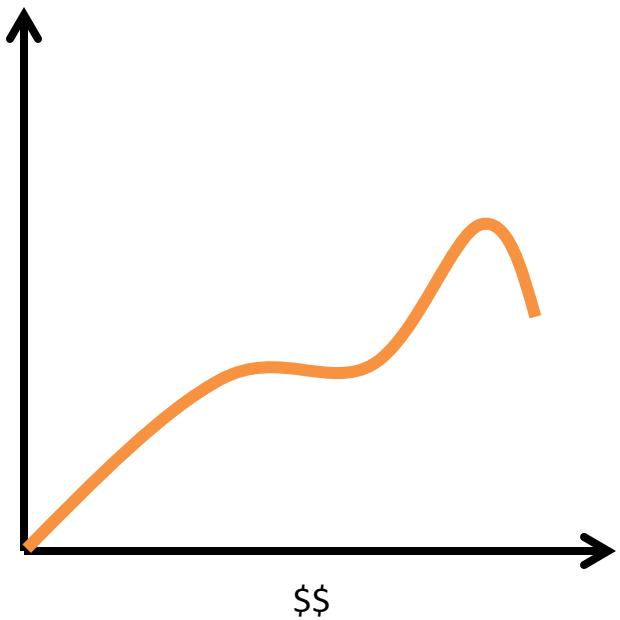


A Reduced-Complexity Channel-Resolving Model for Delta Formation

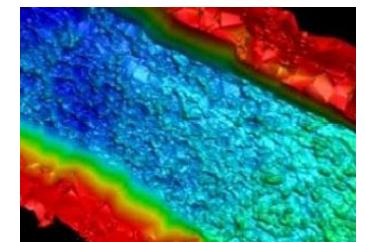
Man Liang
Vaughan Voller
Chris Paola

SAFL University of Minnesota

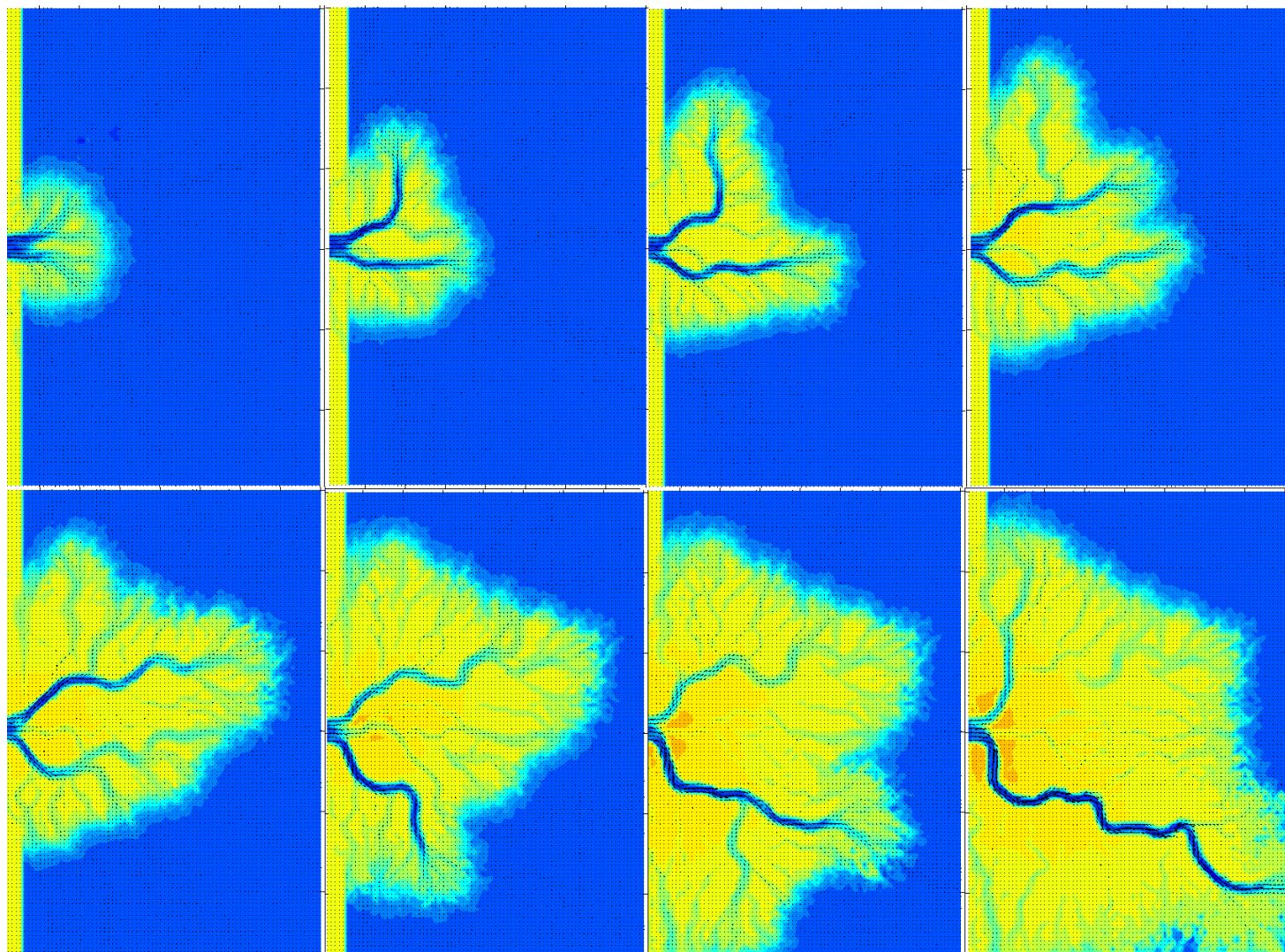
A little philosophy...



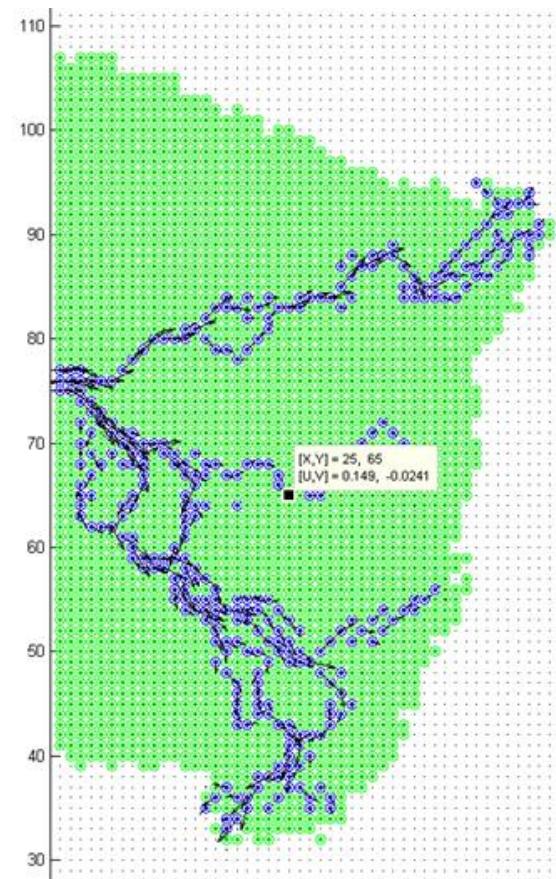
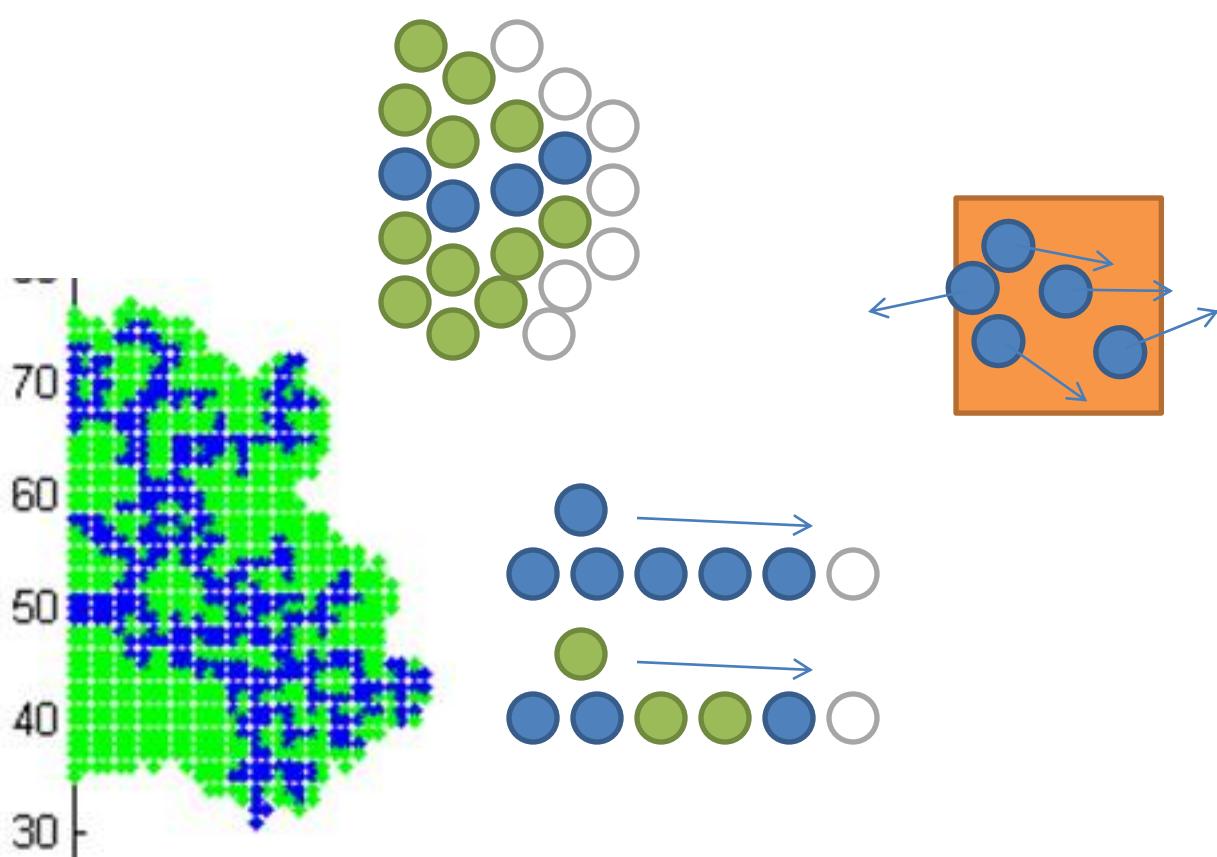
(Kim et al.)

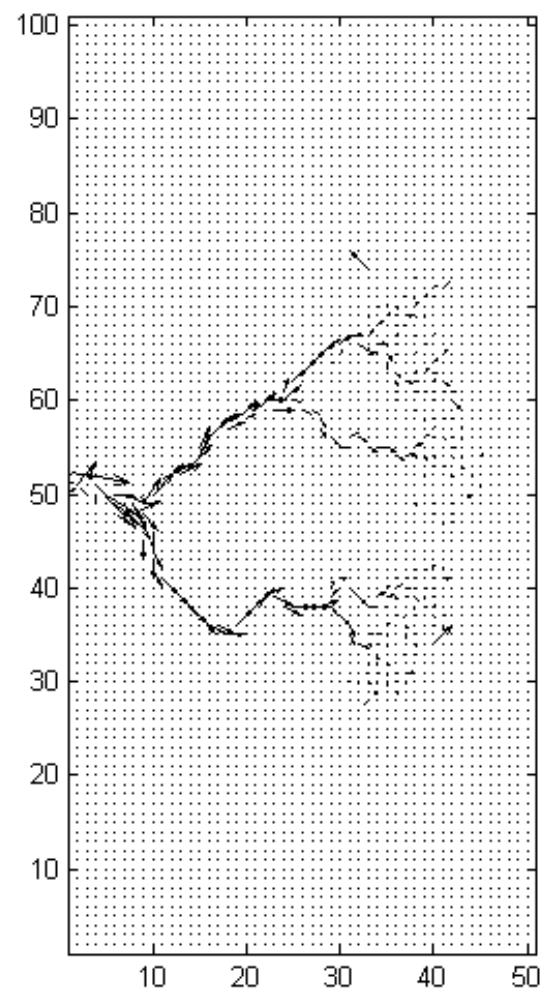
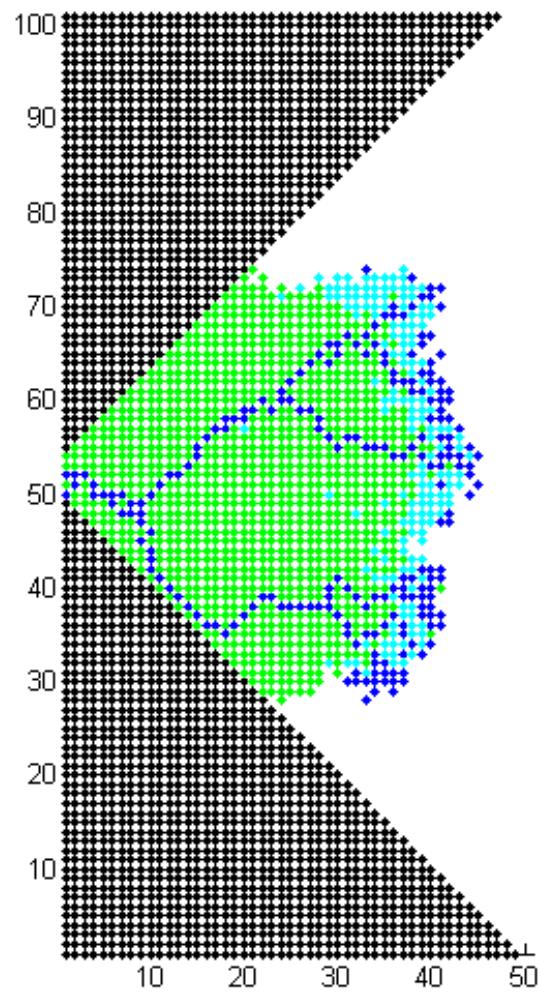


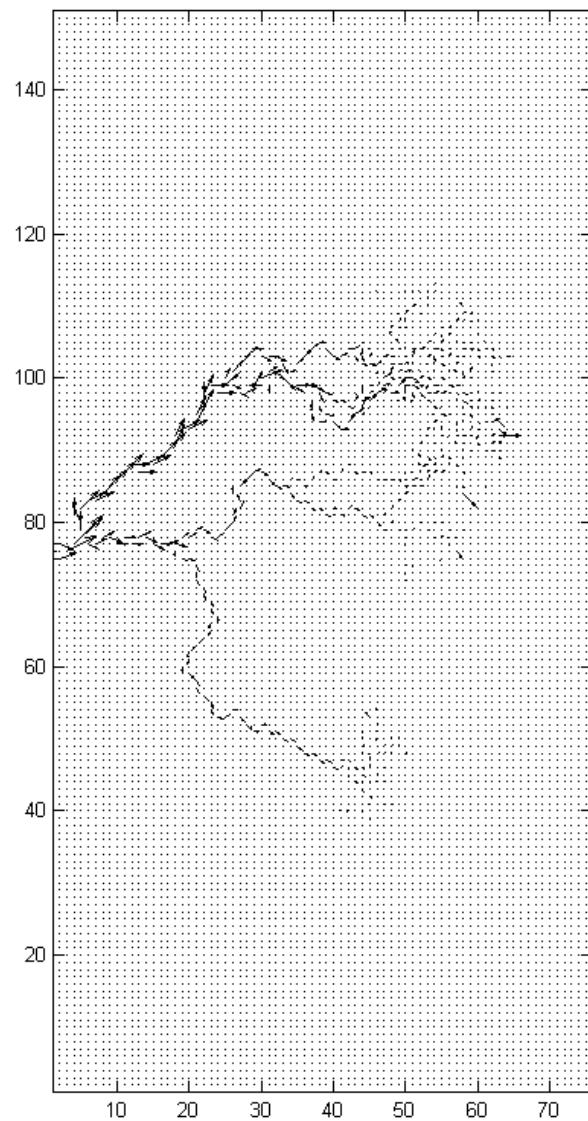
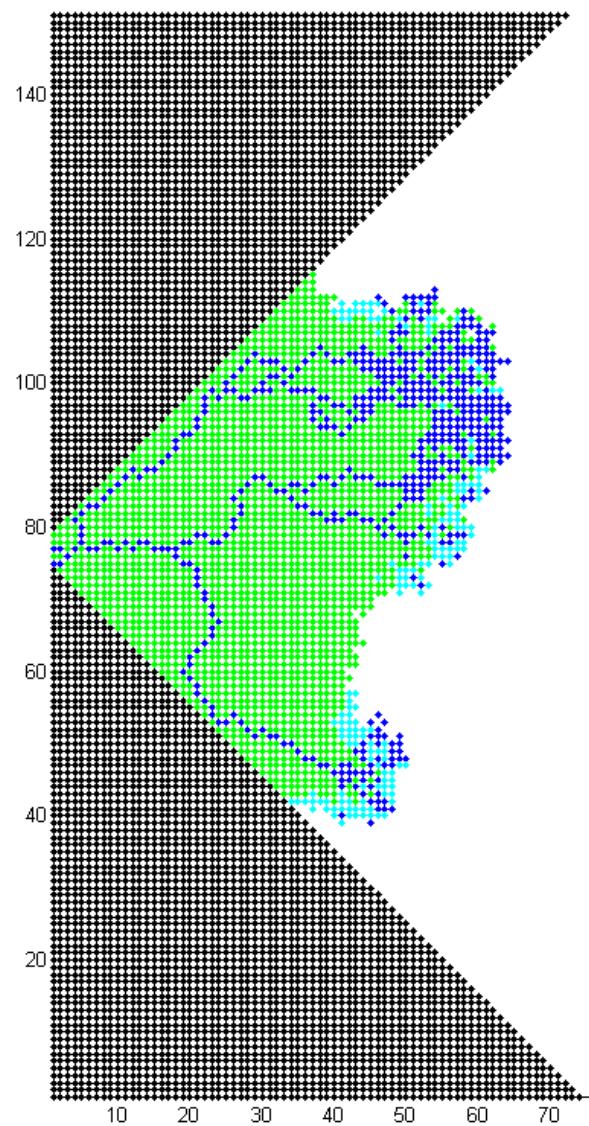
SAFL OSL scan

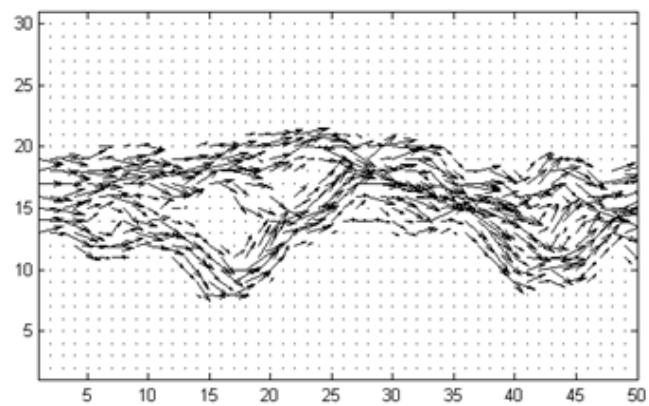
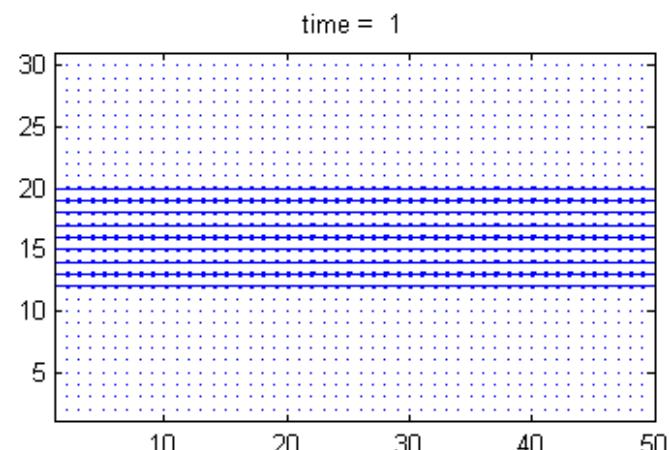
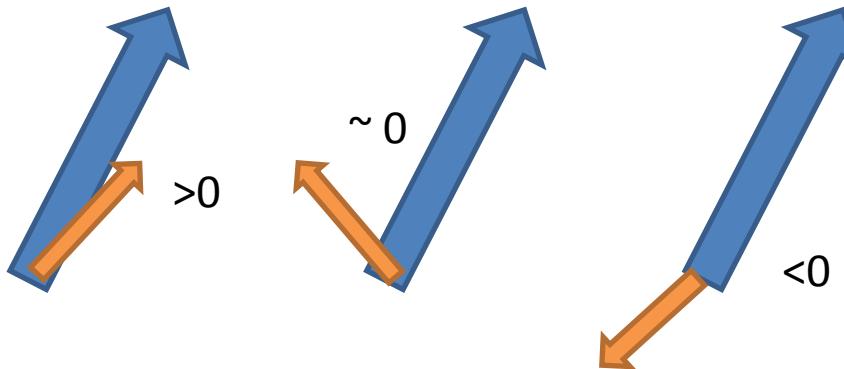
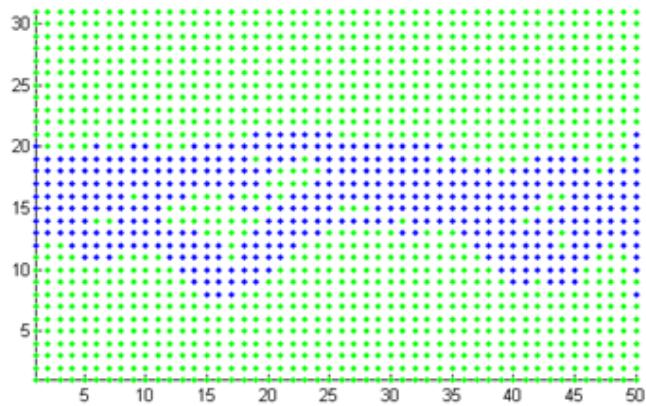
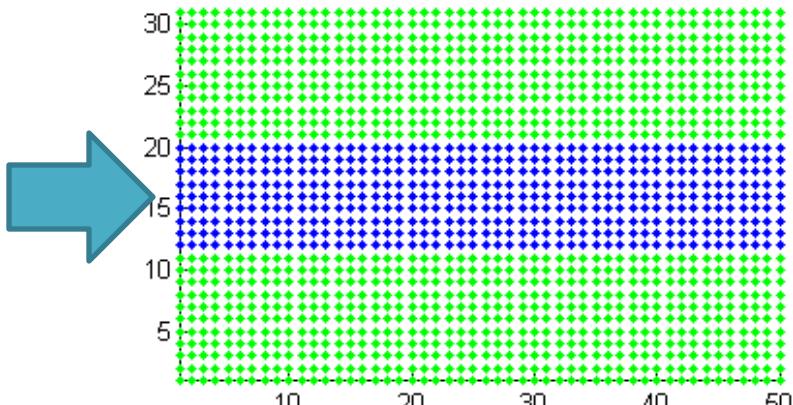
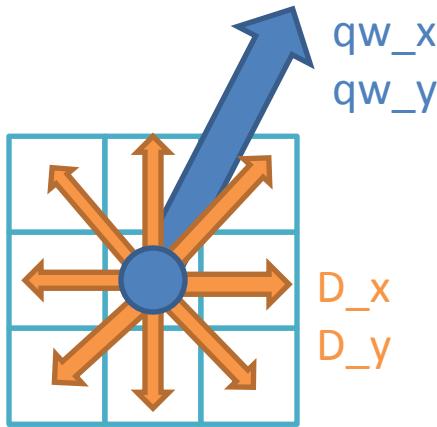


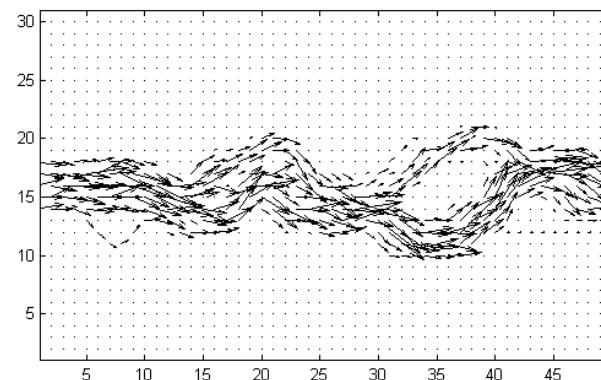
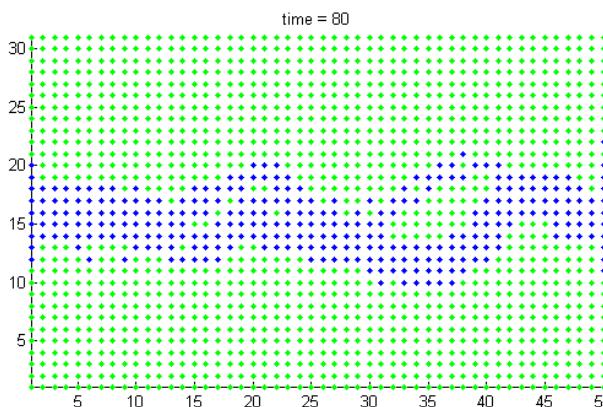
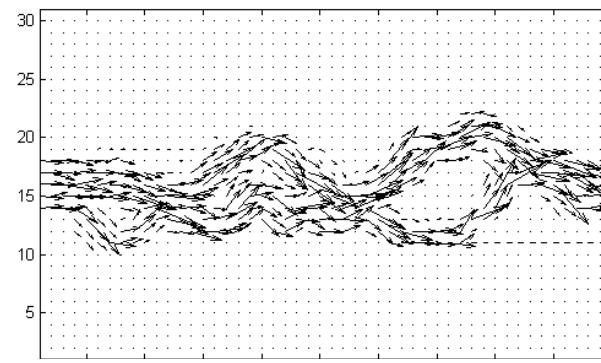
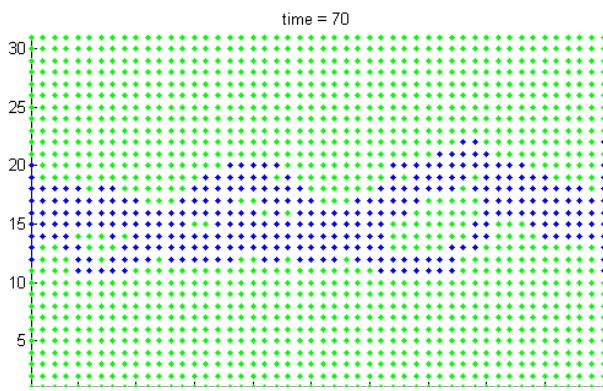
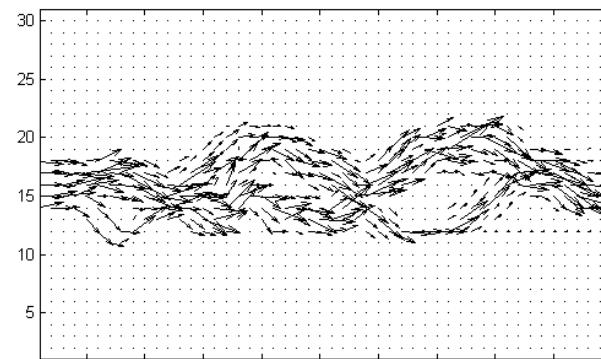
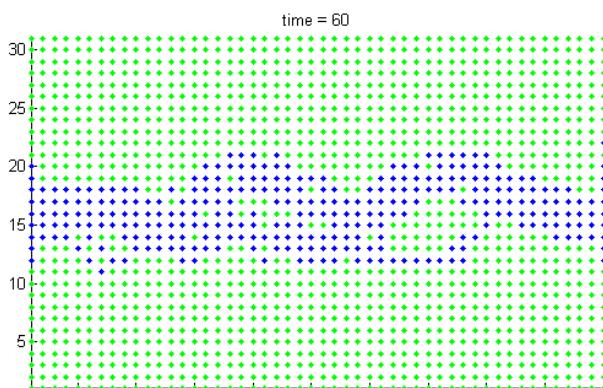
Two particles random walk model

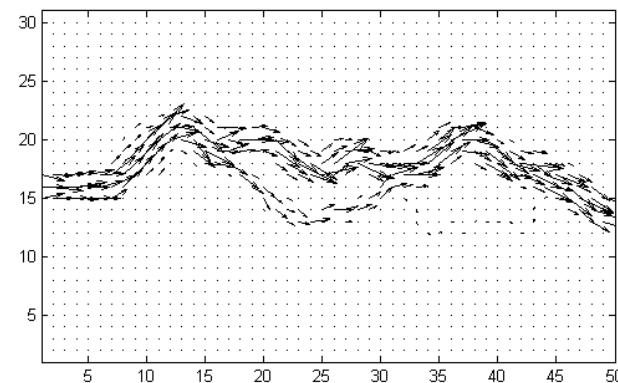
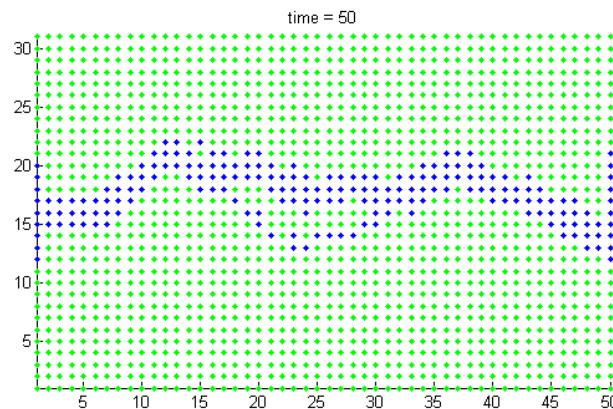
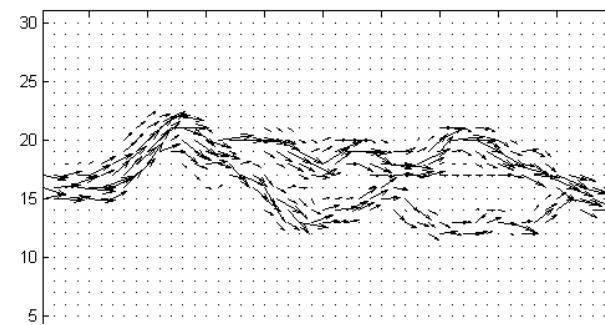
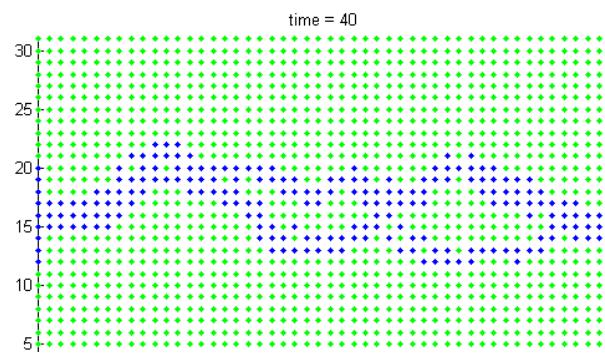
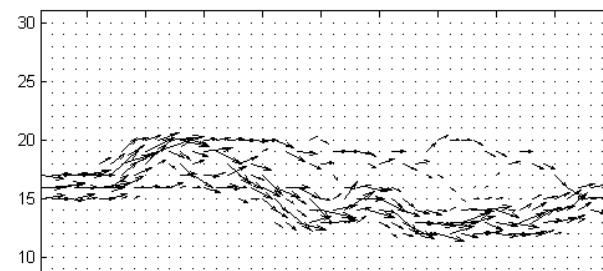
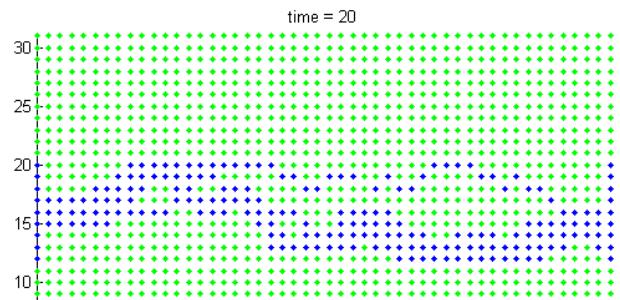


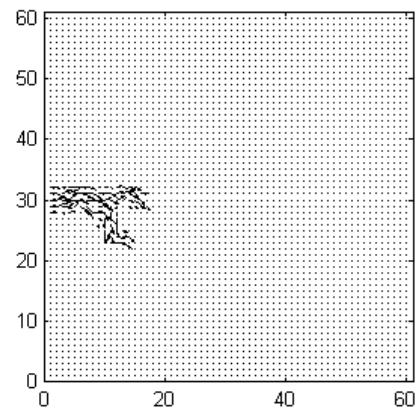
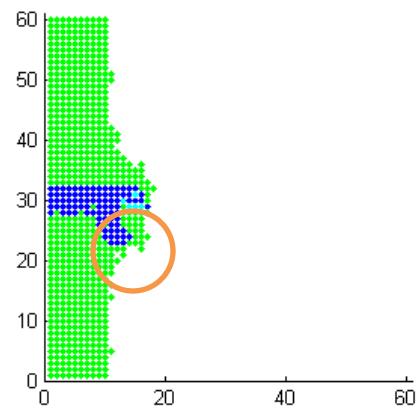
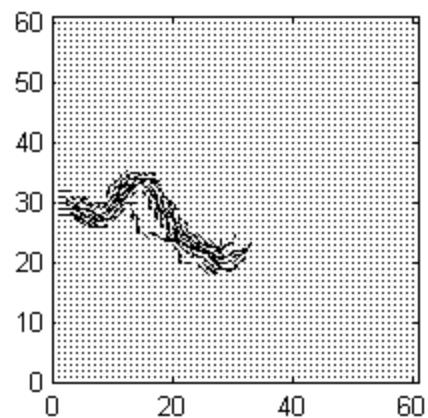
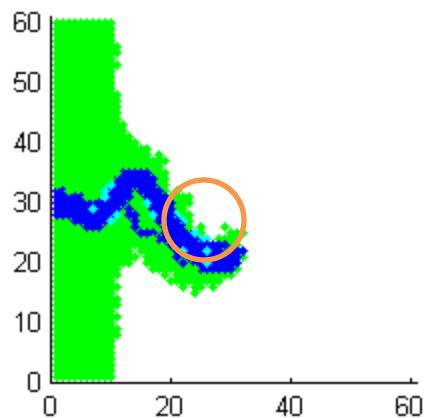


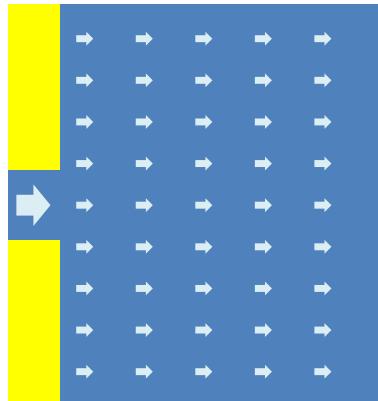




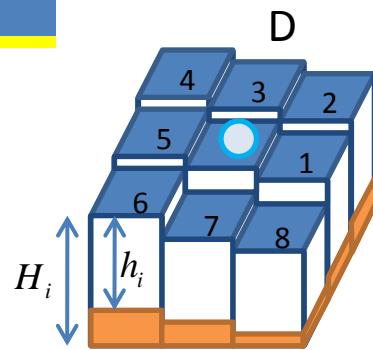




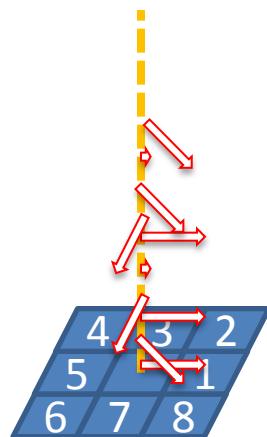




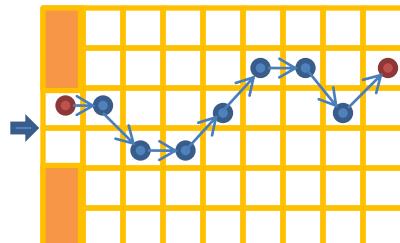
A



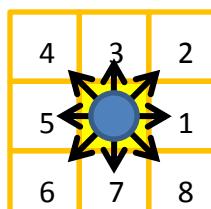
$$\tilde{w}_i^{sfc} = \sqrt{\frac{gh_i^3 |\nabla H|_i}{C_f}} / A$$



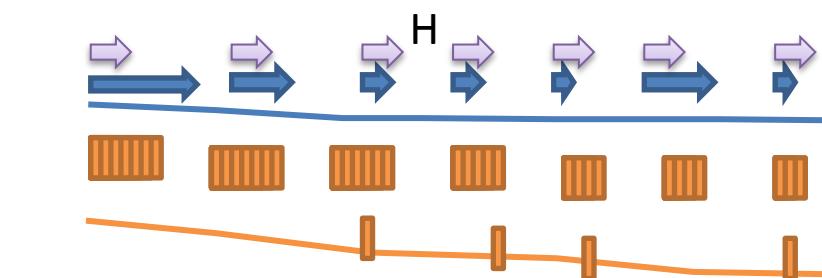
F



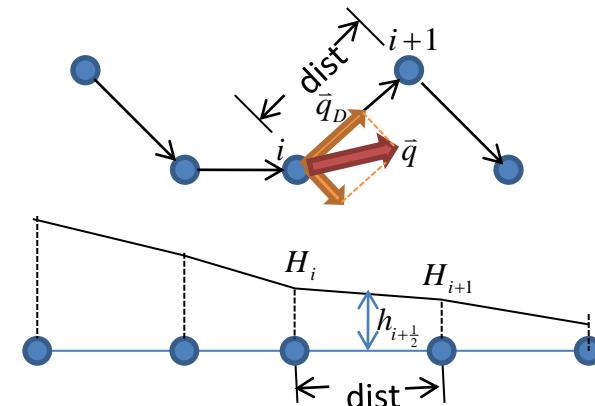
B



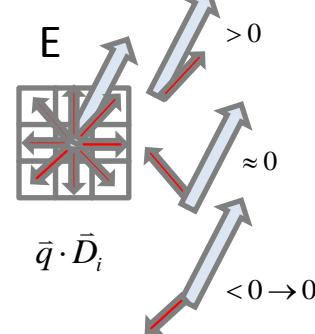
C



$$Q_{dep} = \lambda Q_{res} \frac{U_{dep} - U}{U_{dep}} \quad Q_{ero} = Q_{p_sed} \frac{U - U_{ero}}{U_{ero}}$$

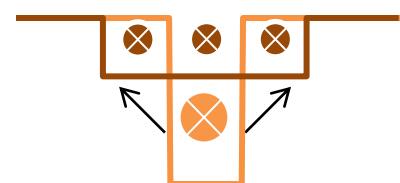


G

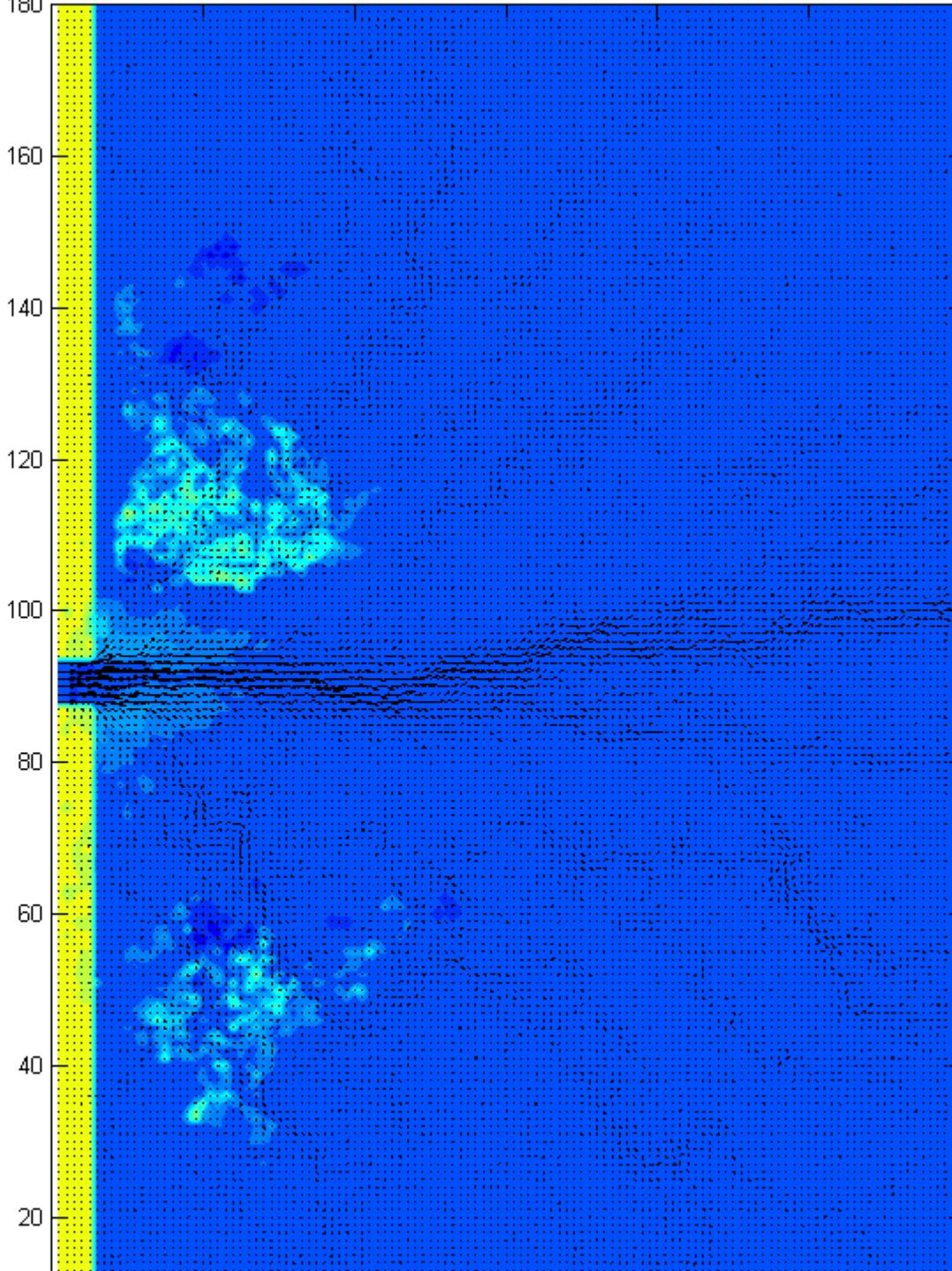


$$\tilde{w}_i^{dxn} = (\max(0, \bar{q}_w \cdot \vec{d}_i))^{\theta_1} h^{\theta_2} / A$$

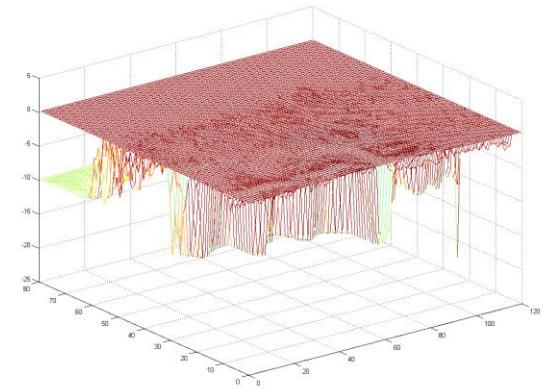
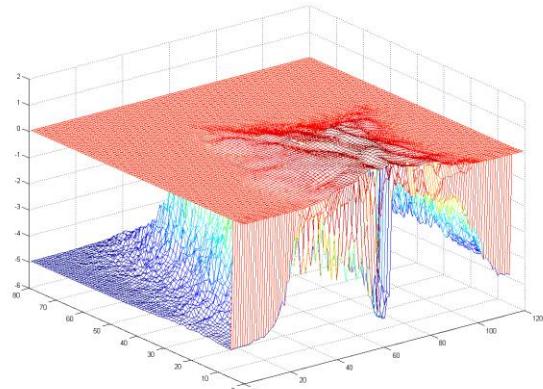
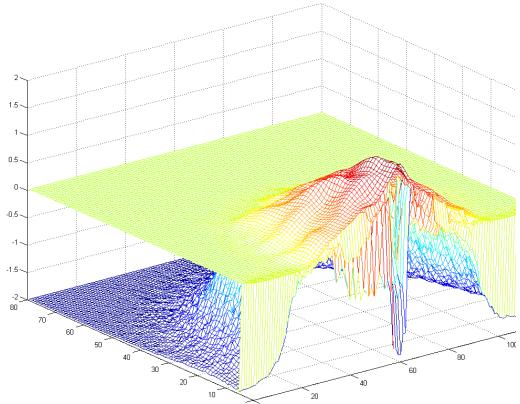
I



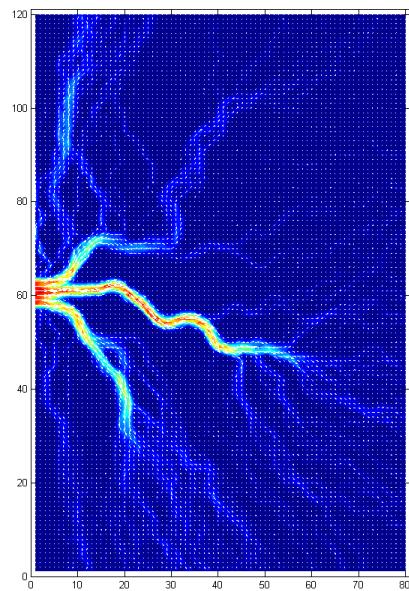
- Intuition is a good first test
- Flow solver is so much simplified, so what do we lose from it?
- Where the randomness comes from?
- Effects of parameters



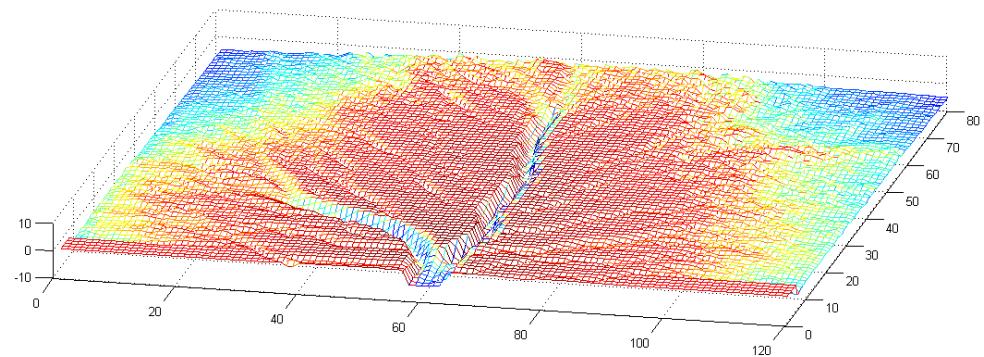
Basic Variables



Free surface

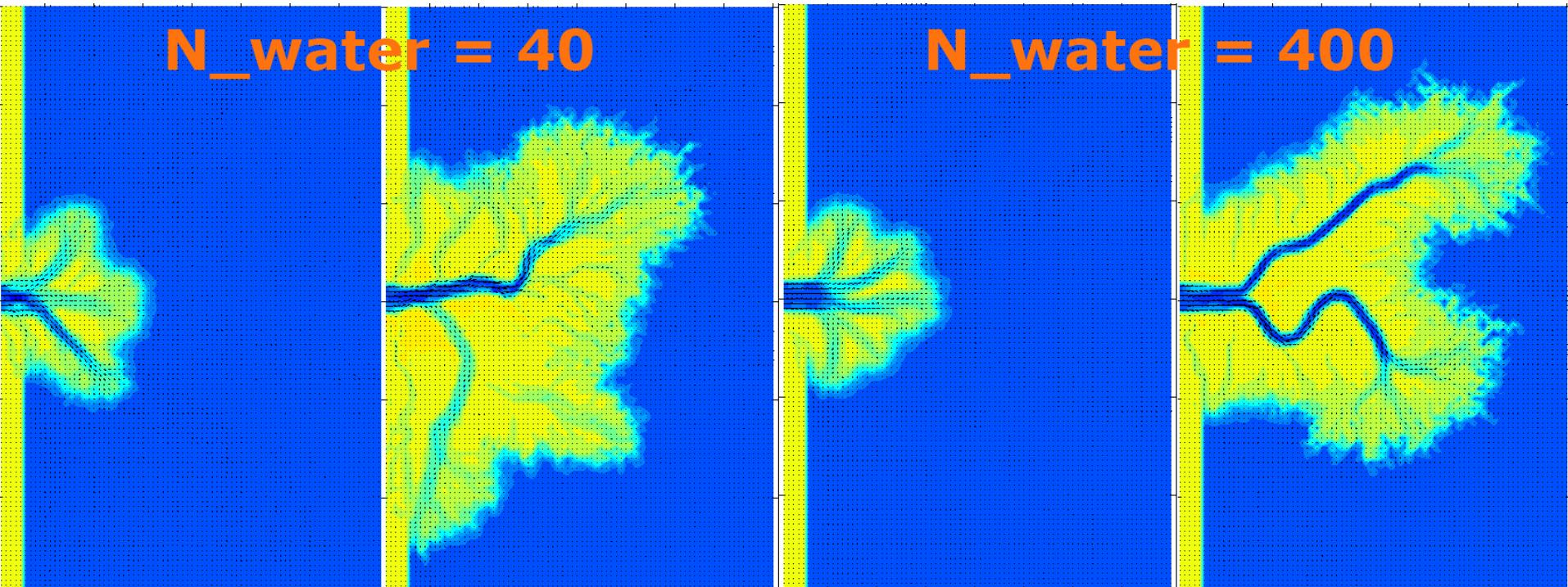


Flow vector

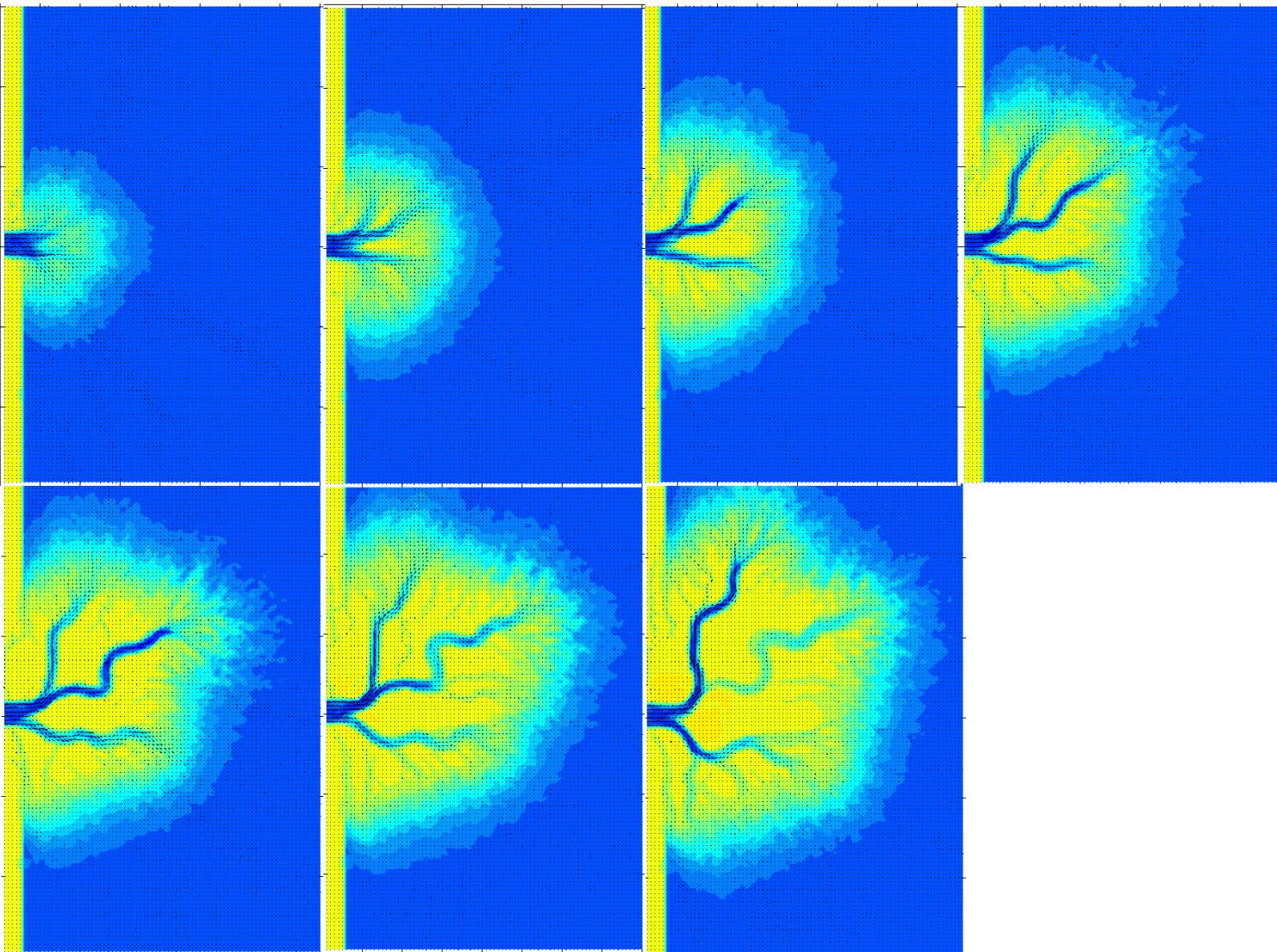


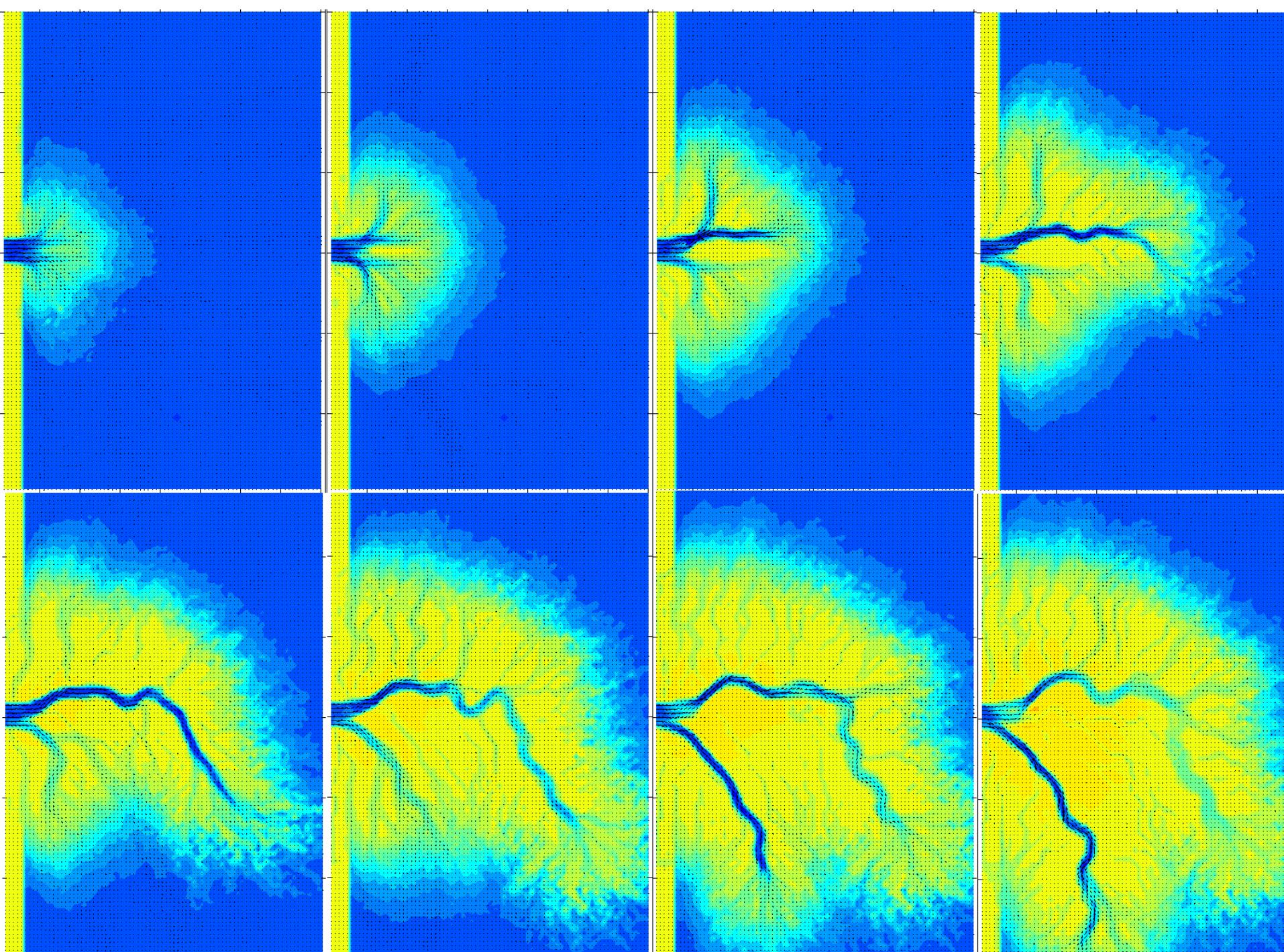
Bed elevation

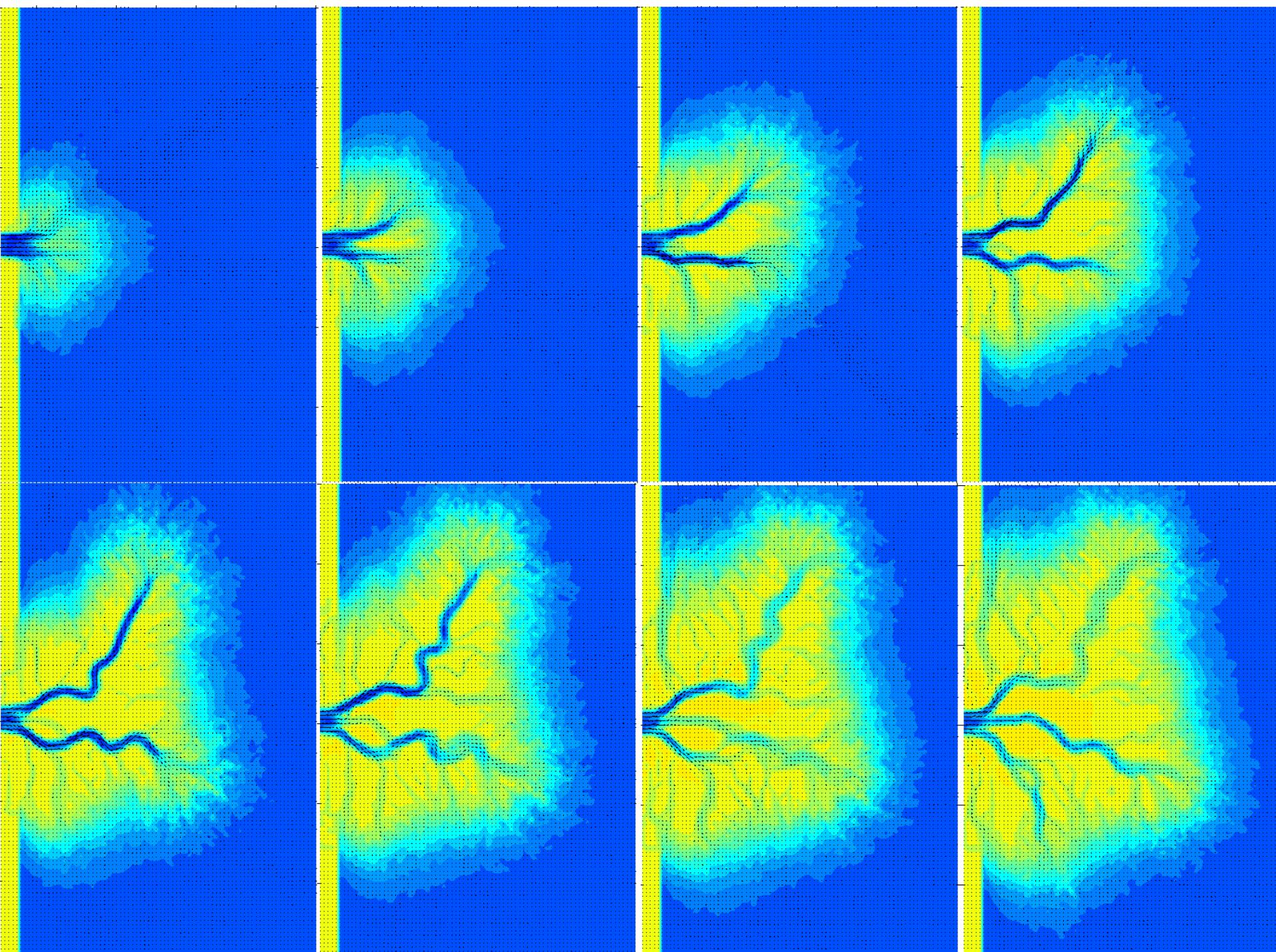
Change # of packages



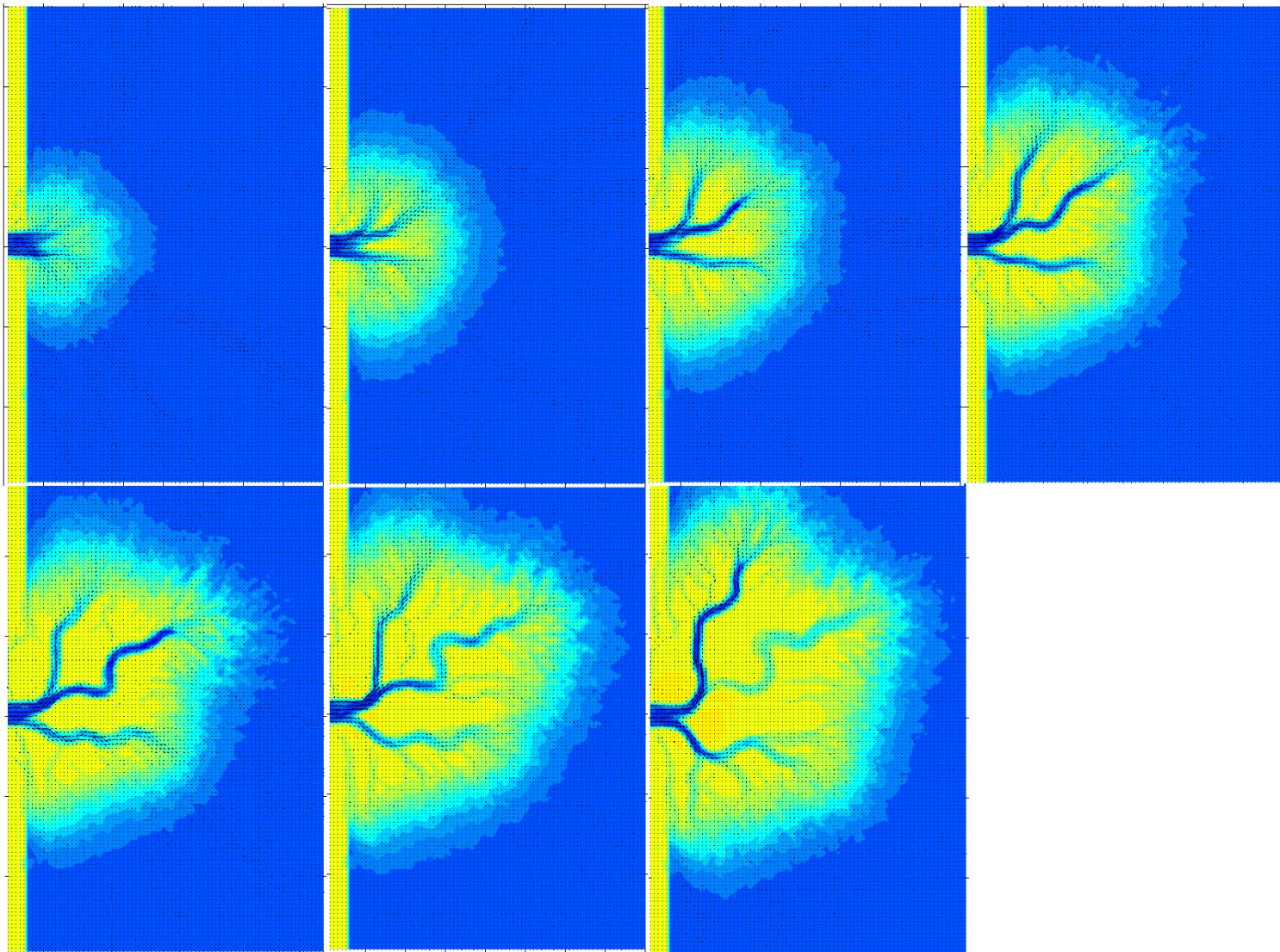
Internal Randomness (variations of the base case)

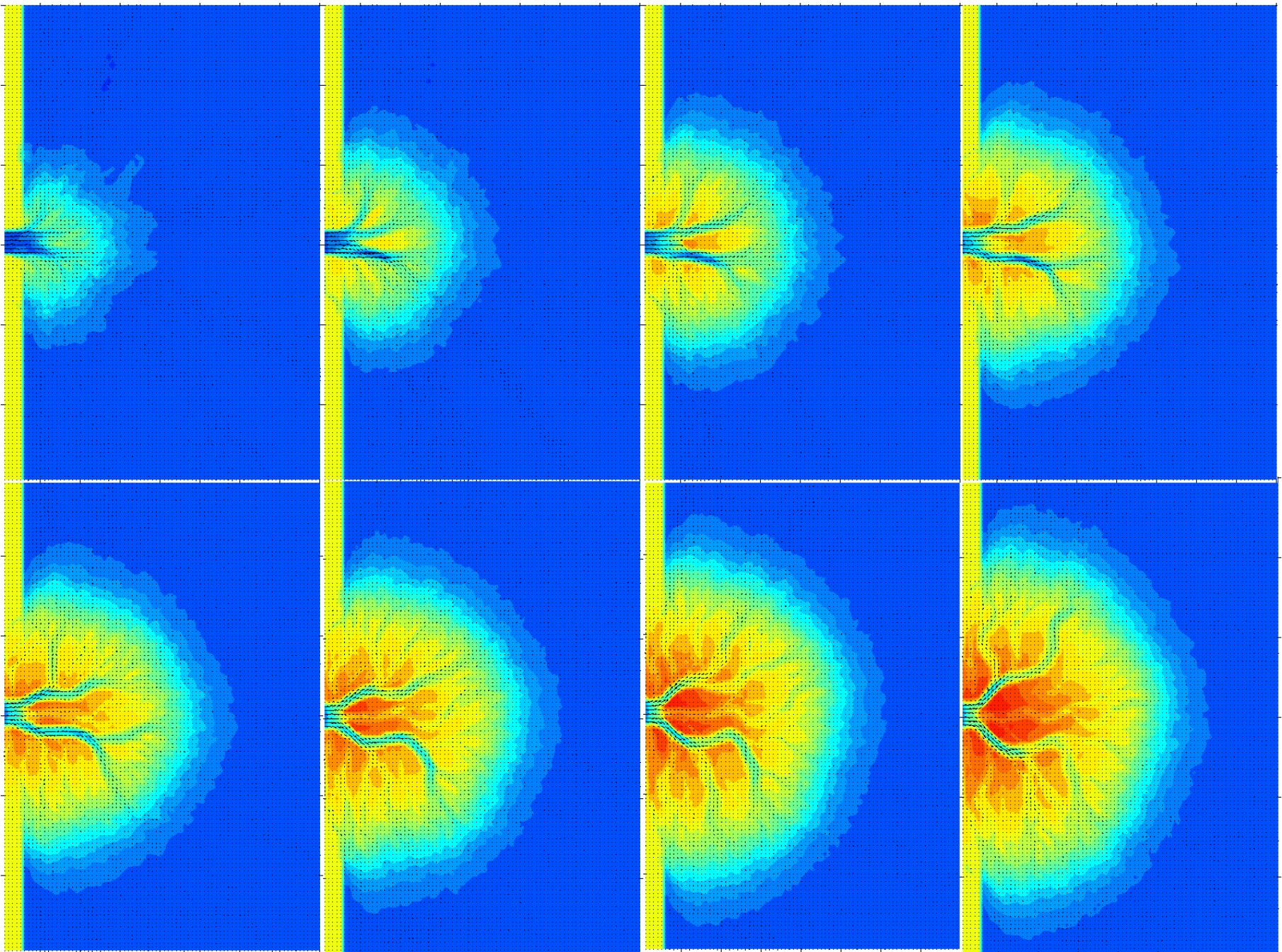


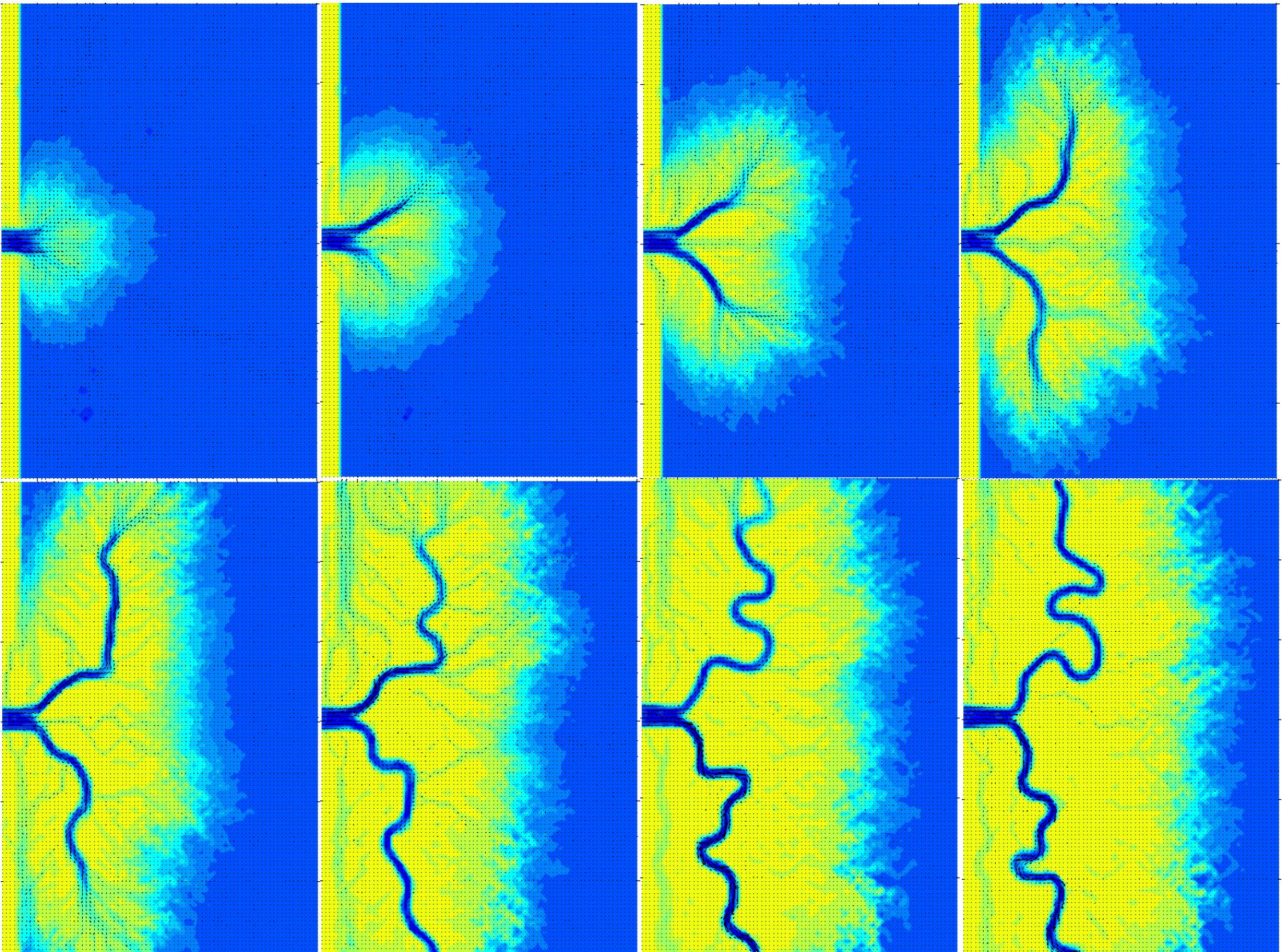


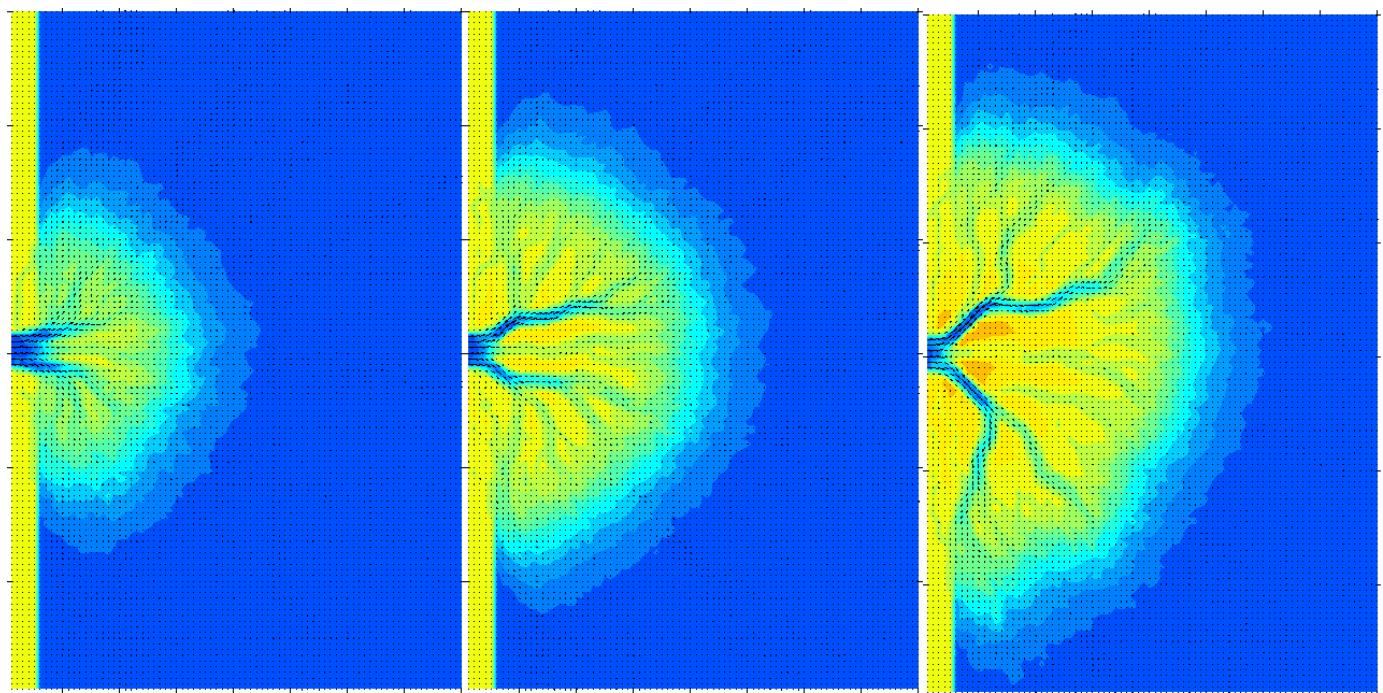
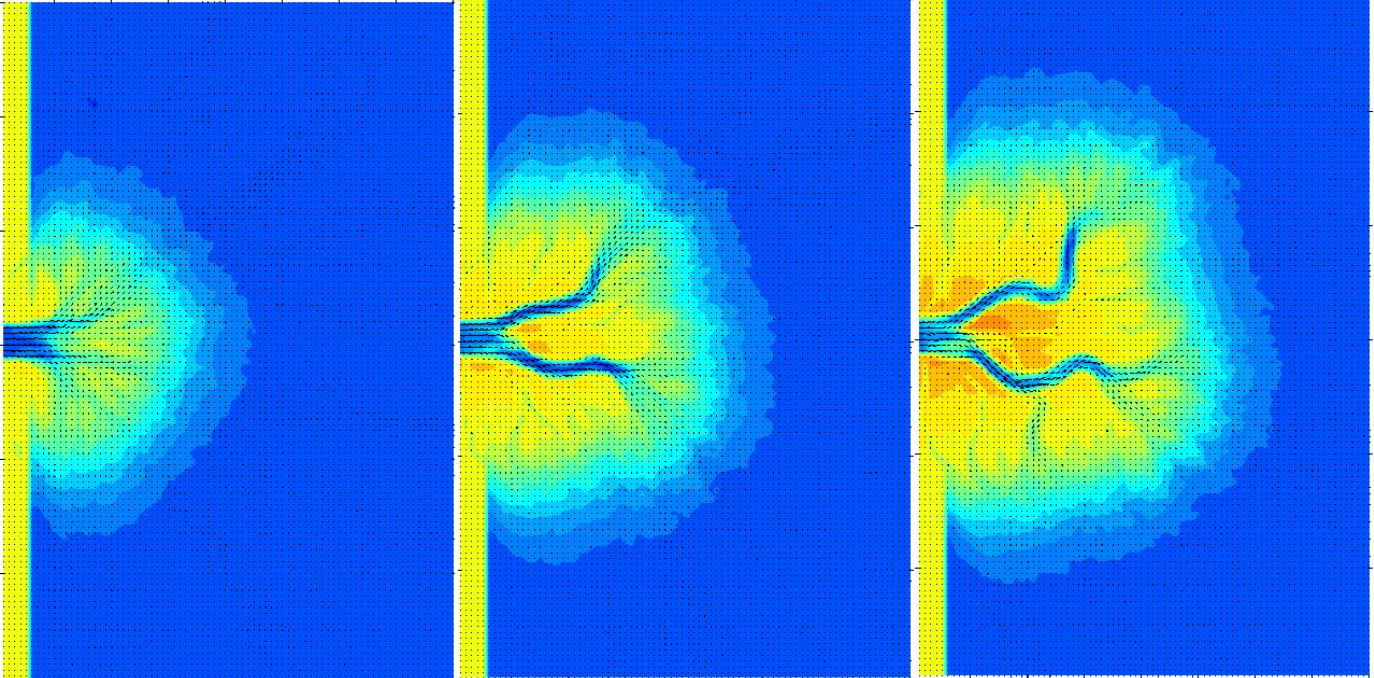


Varing Fr#

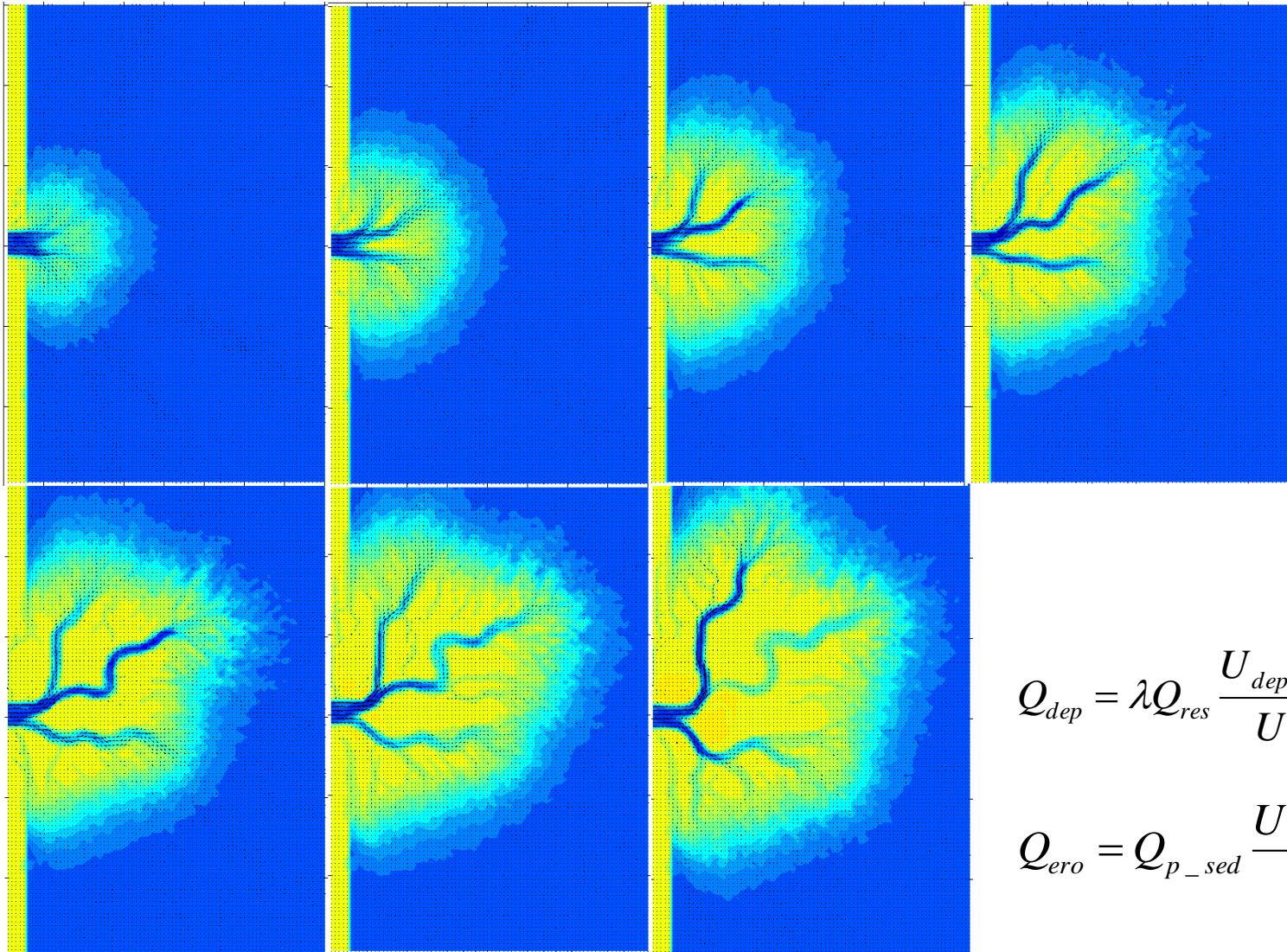






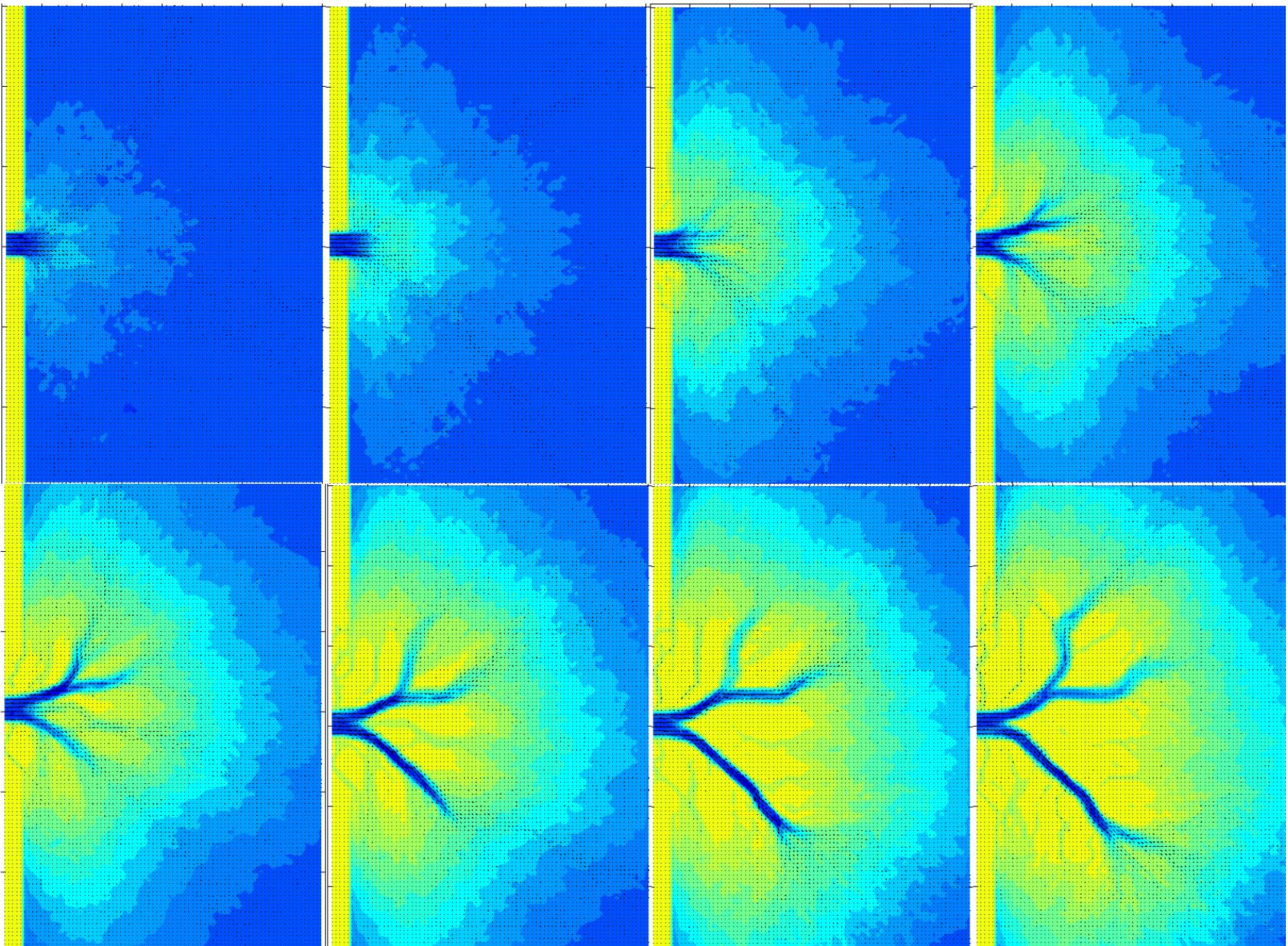


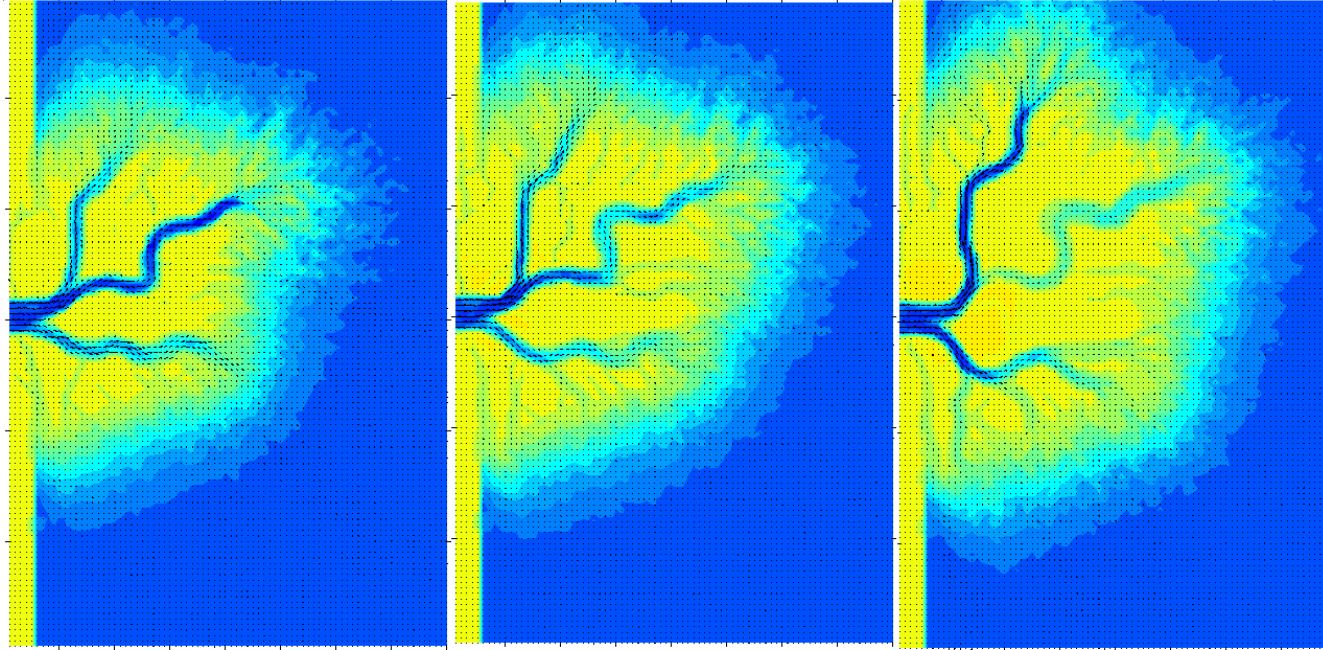
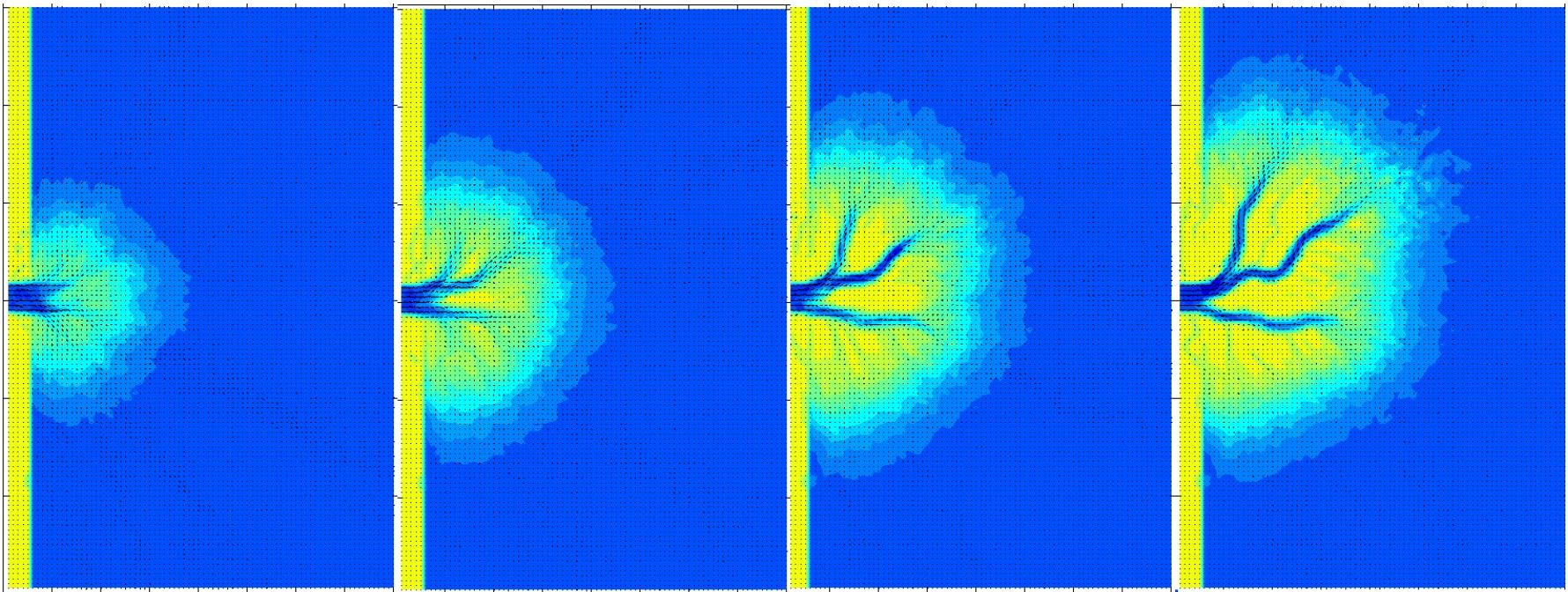
Effects of sedimentation coefficients

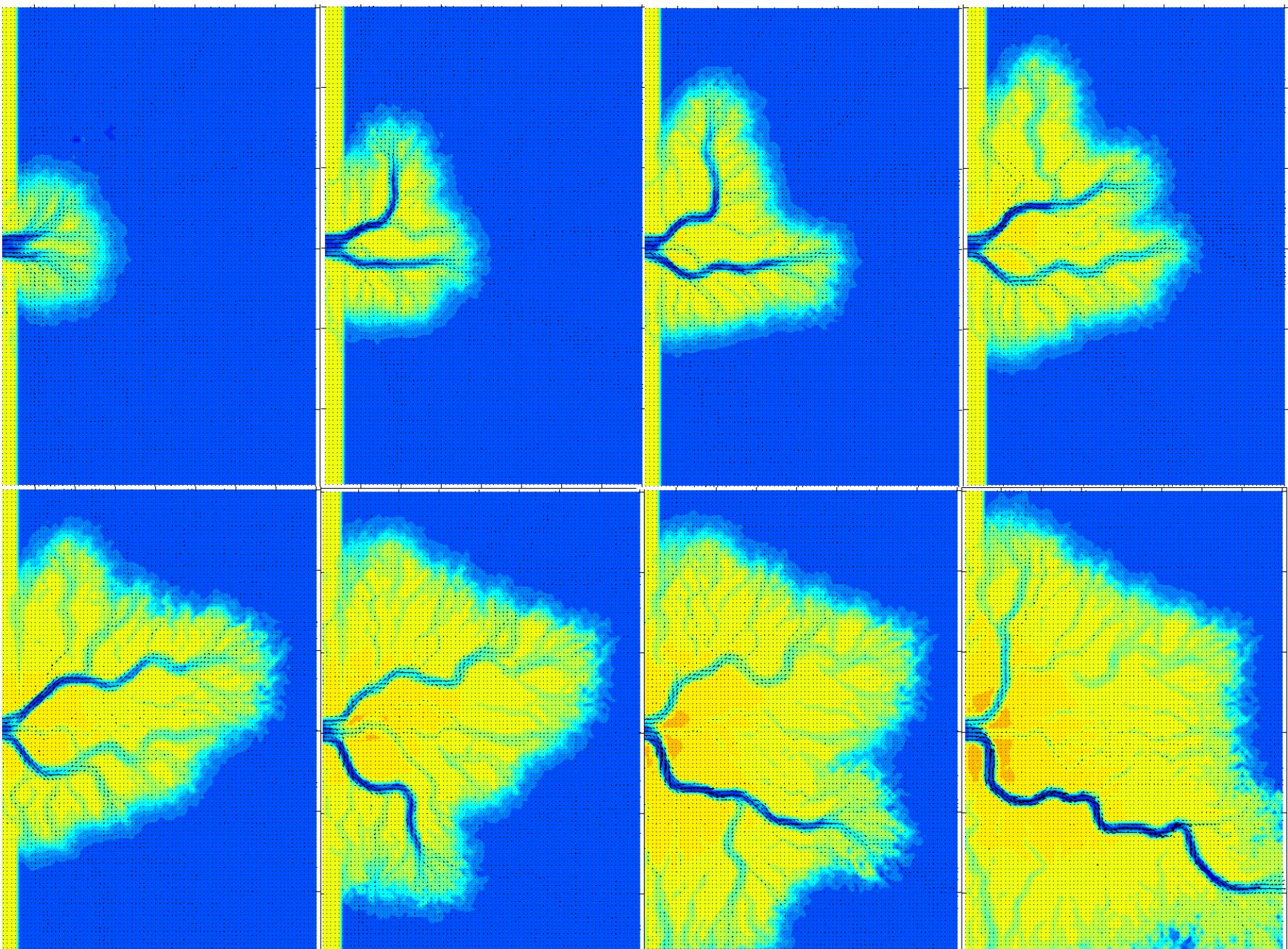


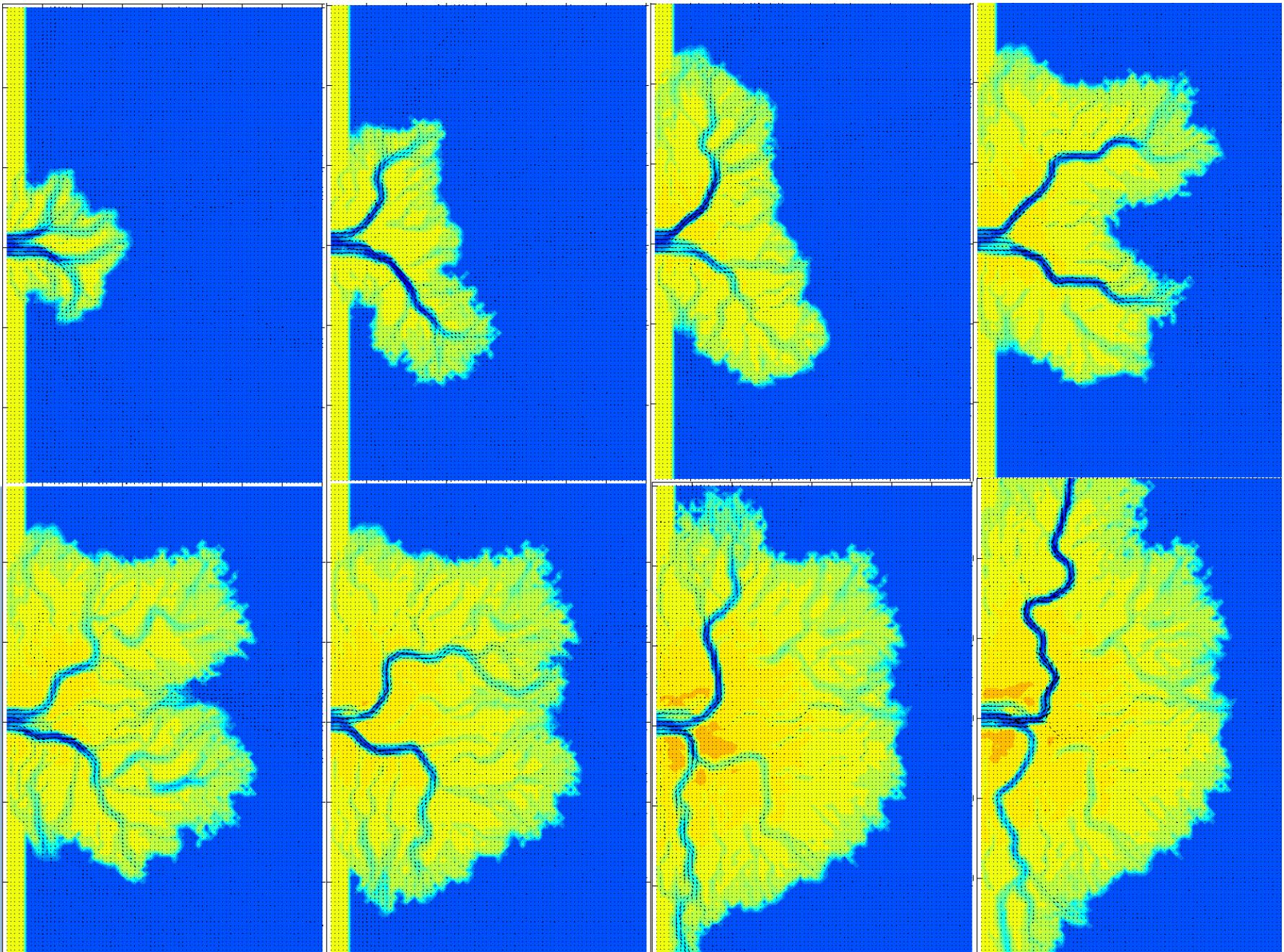
$$Q_{dep} = \lambda Q_{res} \frac{U_{dep} - U}{U_{dep}}$$

$$Q_{ero} = Q_{p_sed} \frac{U - U_{ero}}{U_{ero}}$$

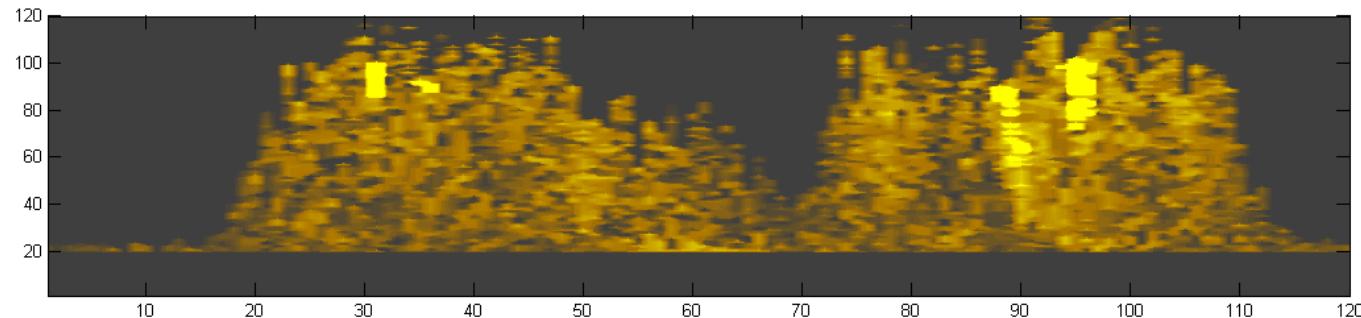
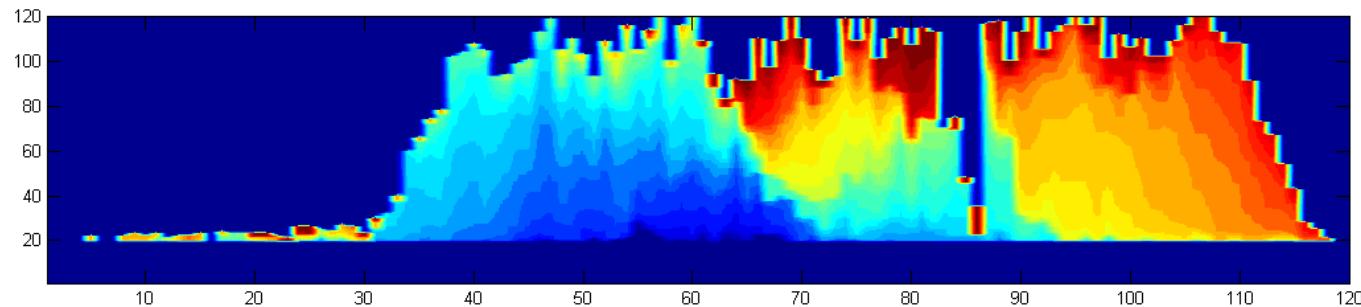






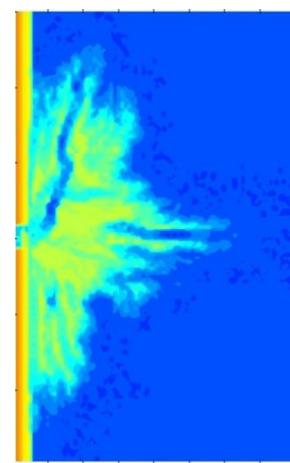
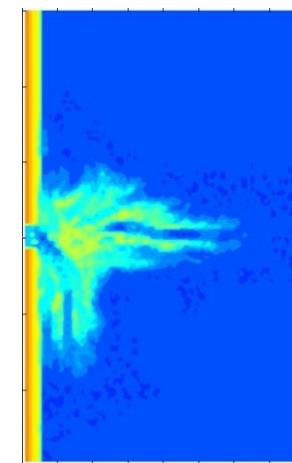
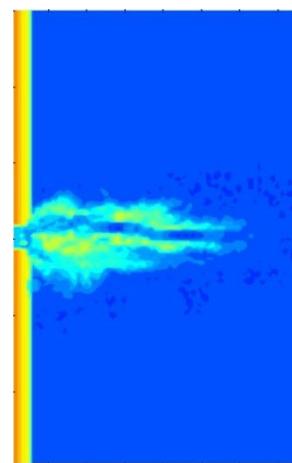
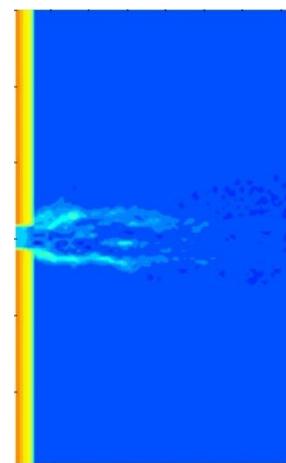
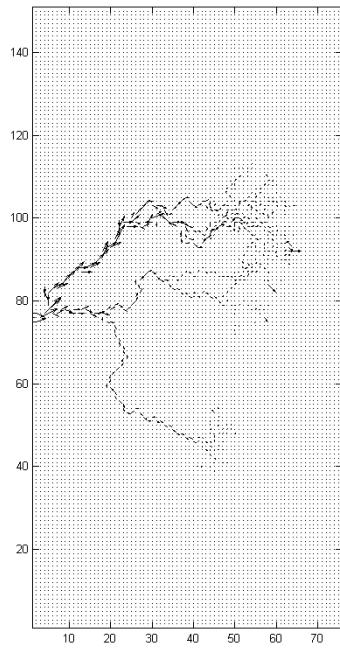
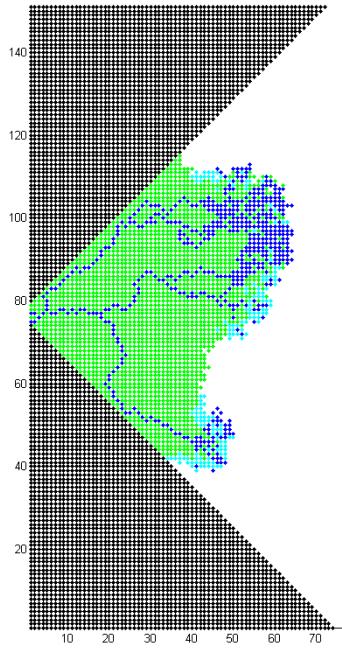


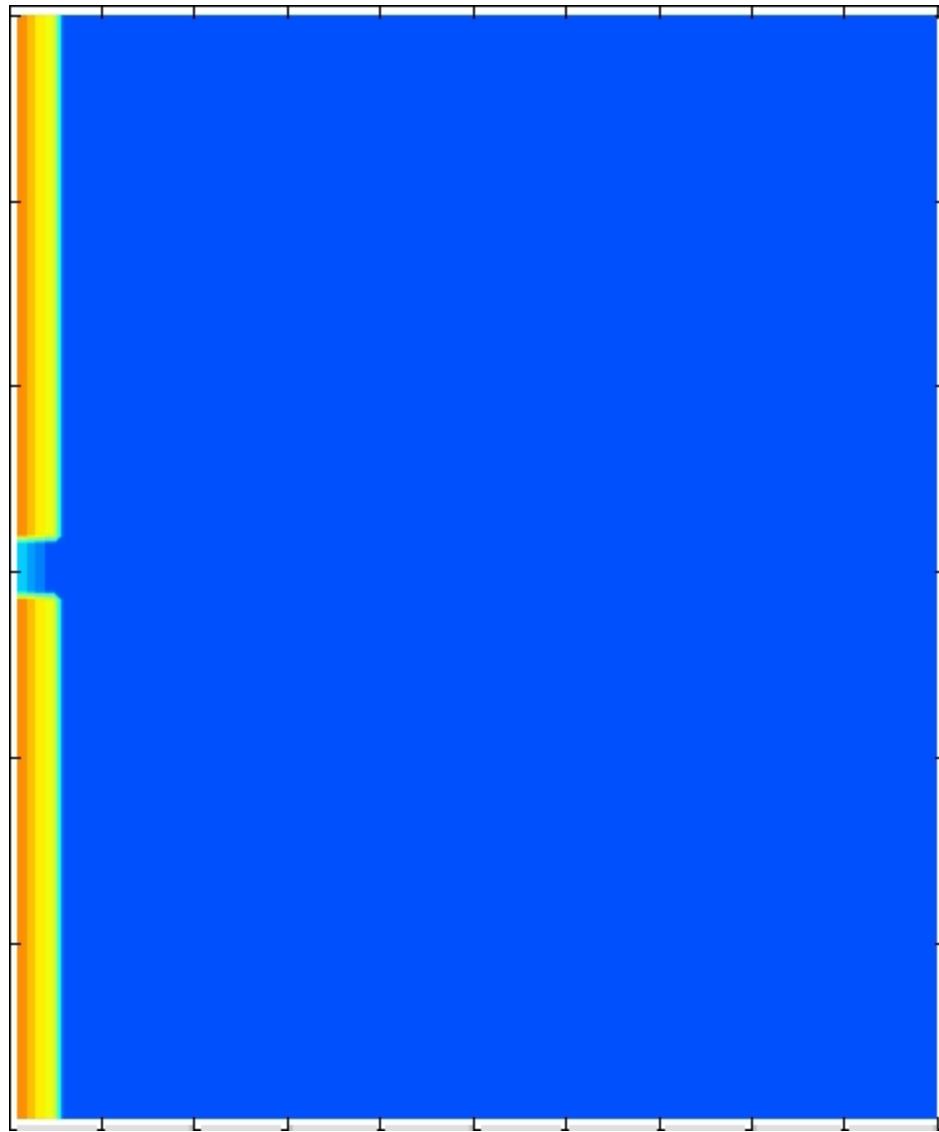
- Stratigraphy potential

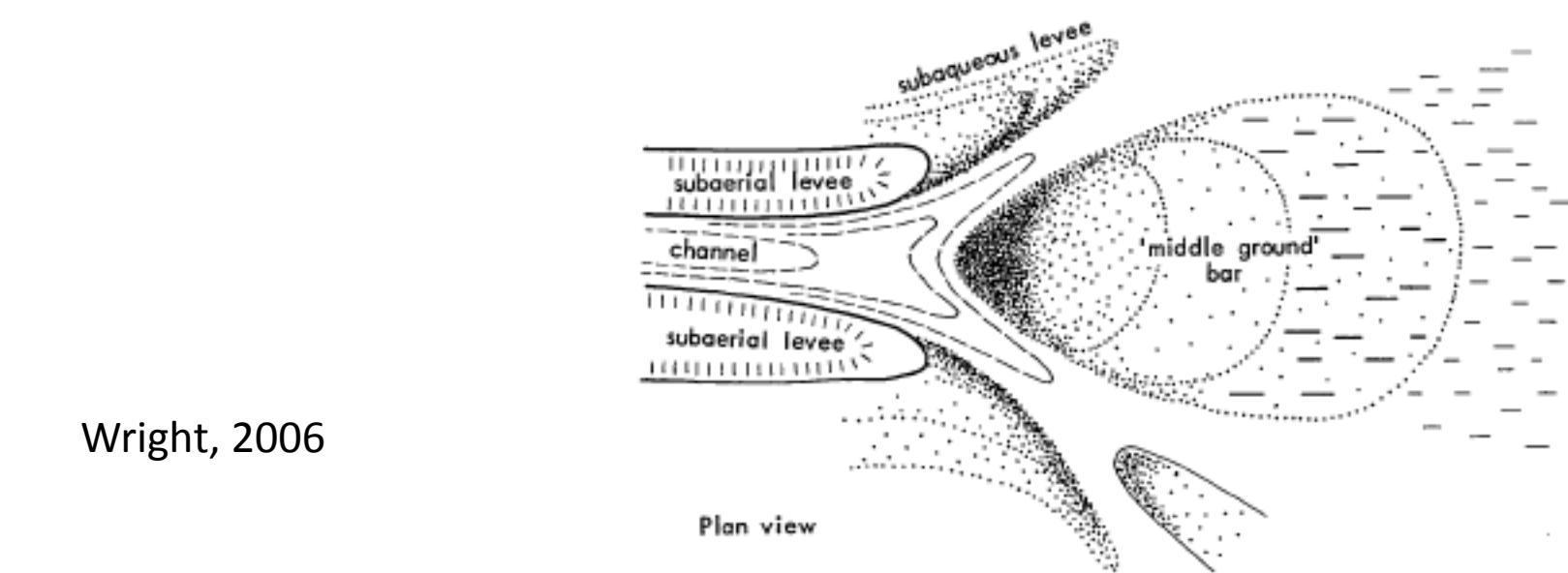
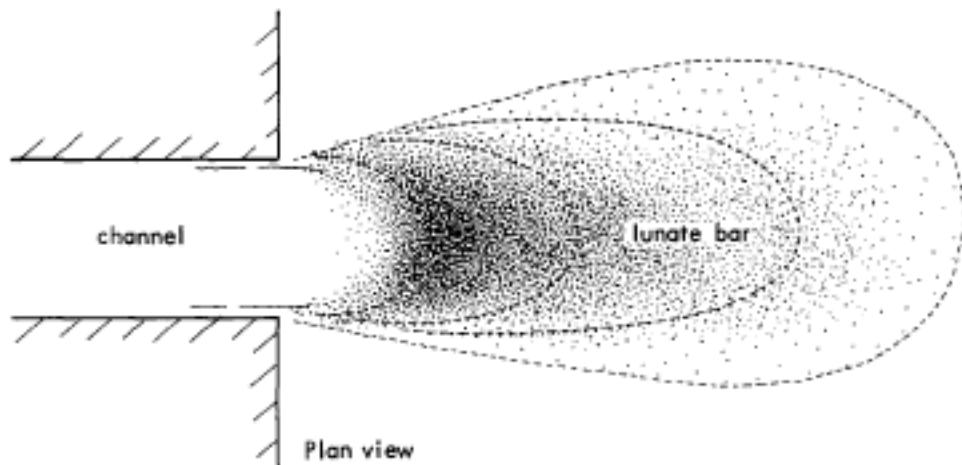


What's Next

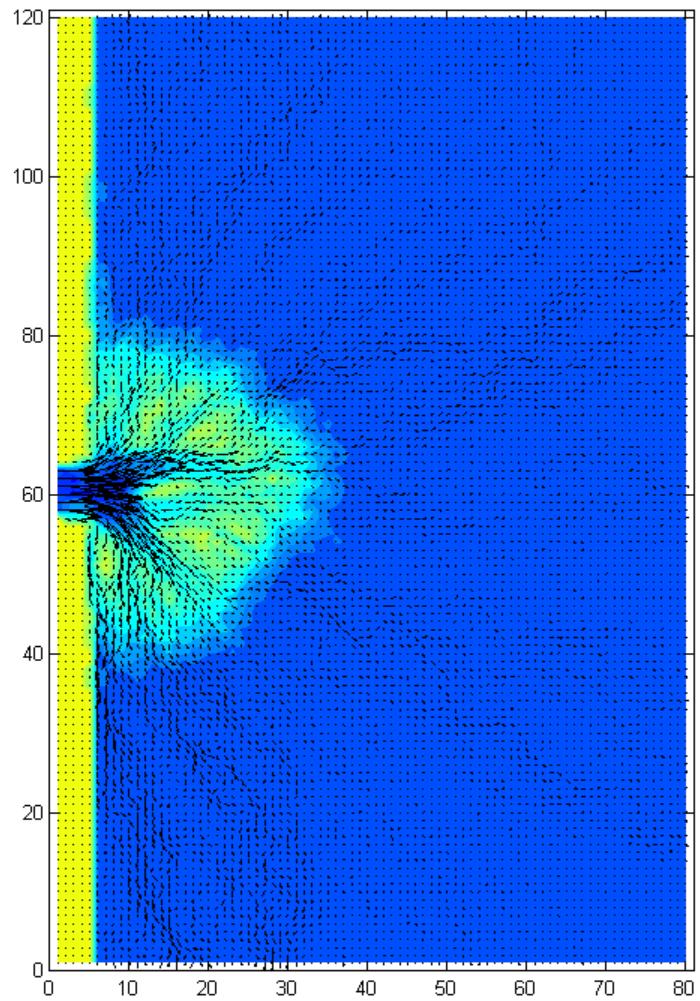
- Improve sediment transport (bed load, grain size, etc.)
- Maybe.. improve flow solver
- Statistical analysis
- Add: wave, vegetation, sea-level, etc.







Wright, 2006



Reduced Hydrodynamic Model

Continuity

$$\nabla \cdot \vec{q} = 0$$

Momentum

$$h\vec{u} \cdot \nabla \vec{u} = -gh\nabla(h + \eta) - C_f |\vec{u}| \vec{u}$$

inertial gravitational
 $\sim \frac{HU^2}{L}$ $\sim \frac{gH^2}{L}$
inertial / gravitational $\sim \frac{U^2}{gH} = Fr^2 \ll 1$



Continuity

$$\nabla \cdot \vec{q} = 0$$

Momentum

$$0 = -gh^3 \nabla(h + \eta) - C_f |\vec{q}| \vec{q}$$

$$|\vec{q}| = \sqrt{\frac{g}{C_f} h^3 |\nabla(h + \eta)|}$$

$$\vec{q} = -\frac{g}{C_f |\vec{q}|} h^3 \nabla(h + \eta)$$

$$\nabla \cdot \left(-\frac{g}{C_f |\vec{q}|} h^3 \nabla h - \frac{g}{C_f |\vec{q}|} h^3 \nabla \eta \right) = 0$$

Define nonlinear diffusivity $K = \frac{g}{C_f |\vec{q}|}$

And linearize the diffusion equation with
Kirchhoff's transformation $\psi = \int_0^\alpha \alpha^3 d\alpha$

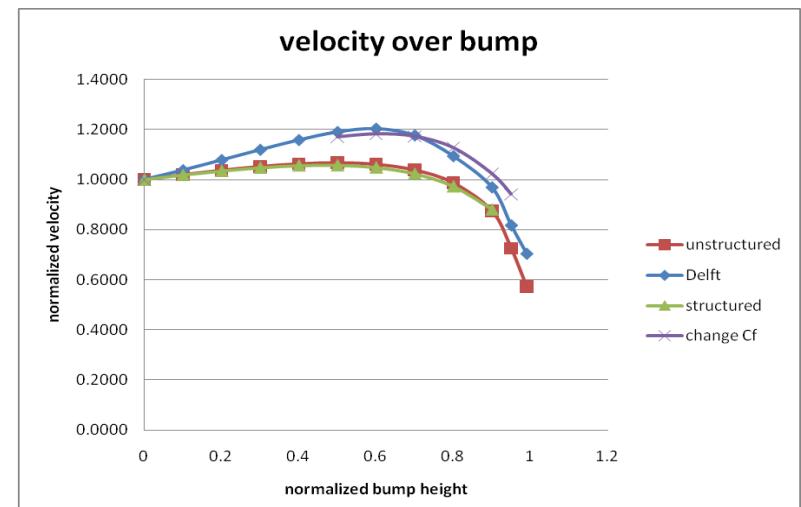
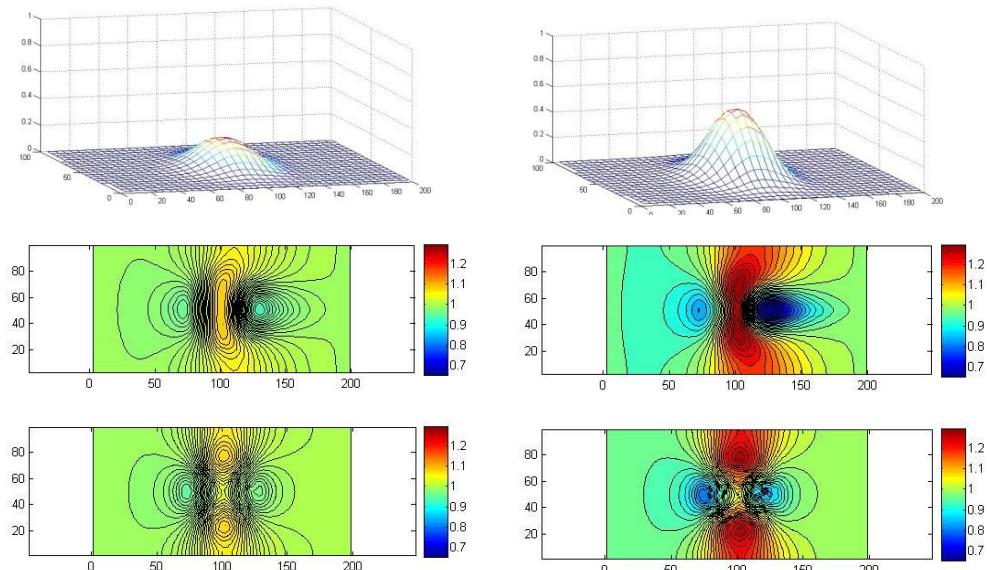
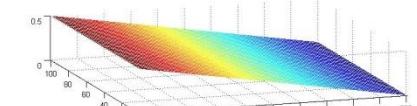
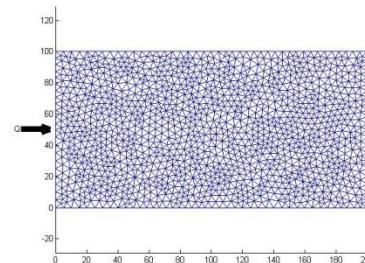


$$\nabla \cdot (K \nabla \psi + S) = 0$$

$$\text{where } K = \frac{g}{C_f |\vec{q}|} = \frac{g}{C_f \sqrt{\left| \frac{g}{C_f} \nabla \psi + \frac{g}{C_f} (4\psi)^{3/4} \nabla \eta \right|}}$$

and $S = K(4\psi)^{3/4} \nabla \eta$ can be considered as a source term.

TEST PROBLEM: the critical mouth bar height needed to divert flow around the bar.



Basic cellular water router:

$$|\bar{q}| = \sqrt{\frac{g}{C_f} h^3 |\nabla(h + \eta)|}$$

$$\varpi_i \propto \sqrt{\frac{g}{C_f} \bar{h}_i^3 |\nabla(h + \eta)|_i} \quad i = 1, 2, 3.$$

$$\sum_{i=1,2,3} \varpi_i = 1$$