

Yunxiang Chen^[1], Xiaofeng Liu^[1], Kenneth D. Mankoff^[2]^[1]Department of Civil and Environmental Engineering, Pennsylvania State University, P.A., USA^[2]Department of Glaciology and Climate, Geological Survey of Denmark and Greenland (GEUS), Copenhagen, Denmark

This work is supported by the National Science Foundation (Grant Number: 1503928)

Background

Increasing ice losses of the **Greenland Ice Sheet (GIS)** is contributing almost 1/4 of the global mean sea level rise^[1]. Ice losses due to subglacial ice melting and ice sheet dynamics, however, are still poorly understood due to limited accessibility and thus lack of data. We here show an OpenFOAM-based one-dimensional subglacial conduit model that can be applied to evaluate the diurnal fluctuations and outburst flooding in Greenland ice sheet.

Greenland hydrological system

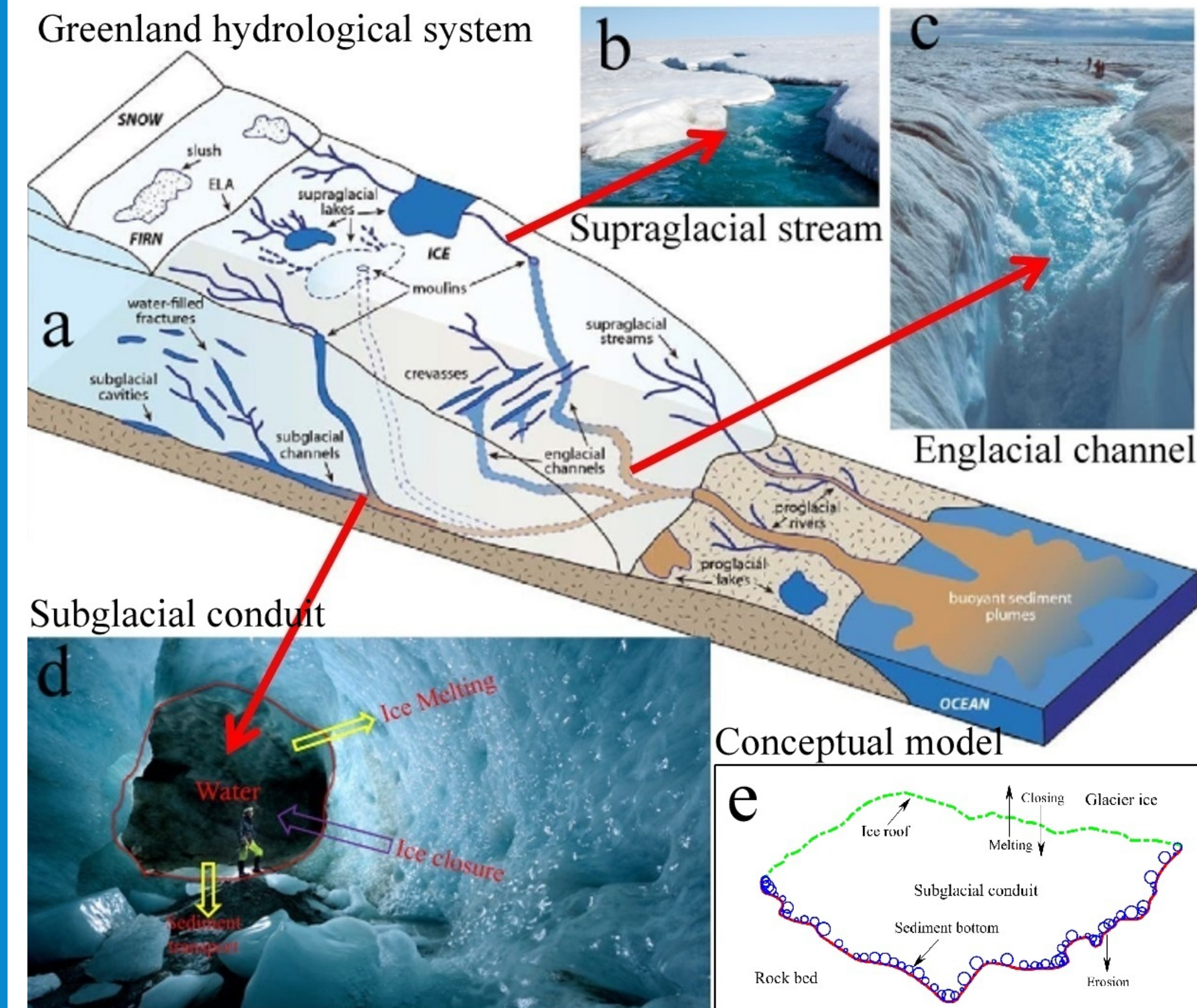


Figure 1: The hydrological system of Greenland ice sheet (a), its three components: supraglacial system (b), englacial system (c), and subglacial system (d), and an conceptual model of key physics in subglacial system (e). Figures (a-d) are modified from Cuffey and Paterson, Los Alamos, Roger Braithwaite, and Robbie Shone, respectively.

Equations & Boundary Conditions (1)

The subglacial model is composed by mass conservation for ice and water, momentum conservation for water, energy conservation, and empirical ice creep-closure model. Entrance pressure boundary is governed by a lake-englacial-subglacial system as shown in Fig. 2(a).

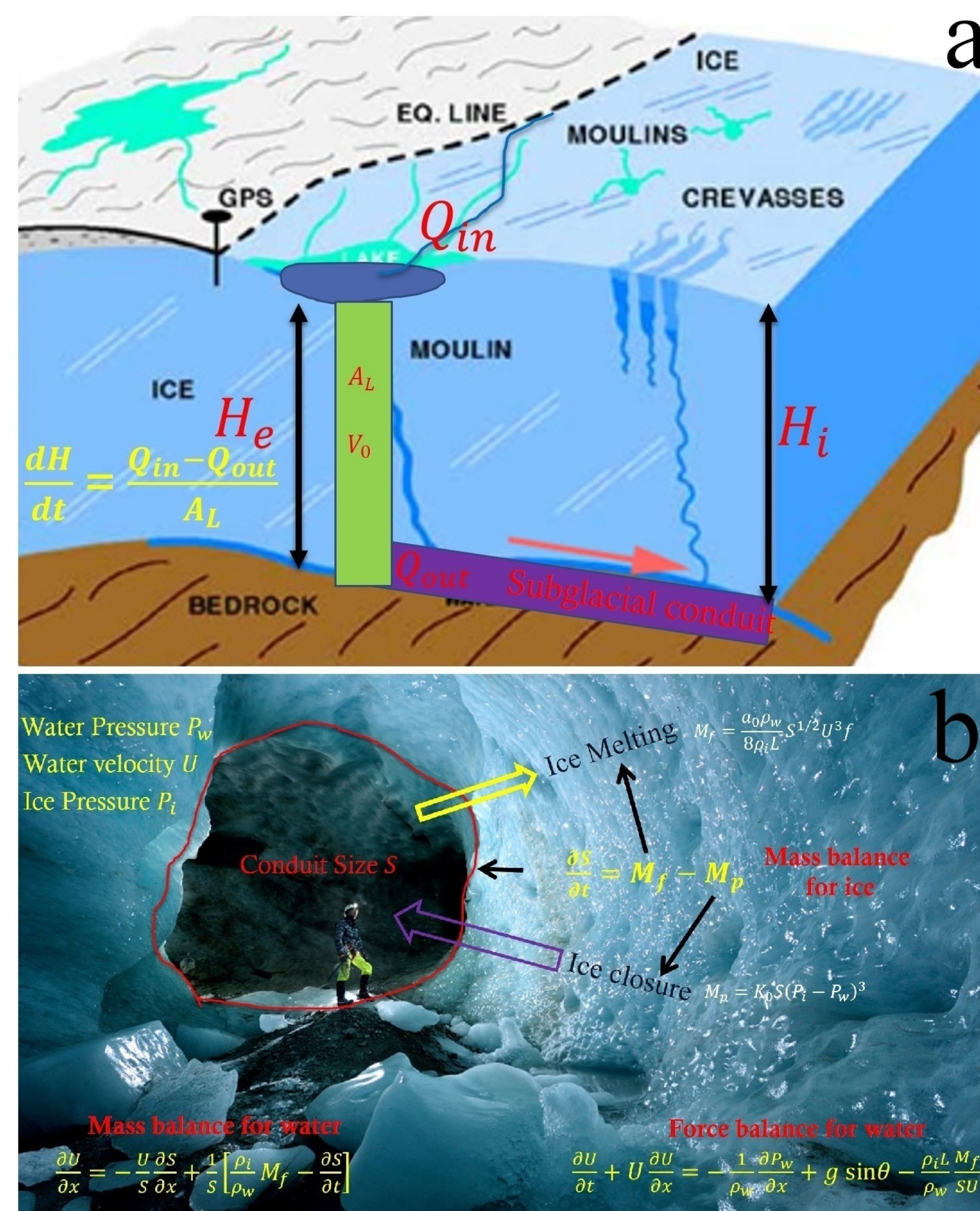


Figure 2: The entrance conditions (a) and mathematical models (b) for subglacial conduits.

Equations & Boundary Conditions (2)

Two simplifications: (a) circular or semi-circular shape, and (b) uniform wall shear stress.

The mass balance for ice, mass balance for water, and momentum balance for water can be expressed by the following equations^[2]:

$$\frac{\partial S}{\partial t} = M_f - M_p \quad (1)$$

$$\frac{\partial U}{\partial x} = -\frac{U}{S} \frac{\partial S}{\partial x} + \frac{1}{S} \left[\frac{\rho_i}{\rho_w} M_f - \frac{\partial S}{\partial t} \right] \quad (2)$$

$$\frac{\partial U}{\partial t} + U \frac{\partial U}{\partial x} = -\frac{1}{\rho_w} \frac{\partial P_w}{\partial x} + g \sin \theta - \frac{\rho_i L}{\rho_w} \frac{M_f}{SU} \quad (3)$$

where S , U , and P_w are time-dependent (t) conduit cross-sectional area, water velocity, and water pressure, respectively. ρ_w and ρ_i are water and ice density, respectively. M_f and M_p denote the ice-melting rate due to viscous friction and the ice-closure rate due to viscous creep, respectively. They are defined as follows:

$$M_f = \frac{a_0 \rho_w}{8 \rho_i L} S^{1/2} U^3 f \quad (4)$$

$$M_p = K_0 S (P_i - P_w)^3 \quad (5)$$

The boundary condition and initial condition are:

BC: $P_w|_{x=0} = \rho_w g H_e$, $P_w|_{x=L} = \rho_i g H_i$, zero gradient for S and U .

IC: Uniform U and S , linear distribution of P_w along channel length x .

Solutions: Numerical & Analytical

Numerical Solution:

(1) Conduit length discretization: $x_i = x_0 + i \Delta x$, $i = 0, 1, \dots, N$, $\Delta x = L/N$.

(2) ICs: $P_{w,i}^0 = P_{w,0}^0 + i/N(P_{w,N}^0 - P_{w,0}^0)$, $U_i^0 = U_0$, $S_i^0 = S_0$.

(3) BCs: $P_{w,0}^j = \rho_w g H_e^j$, $P_{w,N}^j = \rho_i g H_i^j$, $U_0^j = U_1^j$, $U_N^j = U_{N-1}^j$, $S_0^j = S_1^j$, $S_N^j = S_{N-1}^j$. $H_e^j = [V_0 + (Q_0 - Q_{out}) \Delta t] / A_L$, V_0 , Q_0 , Δt , A_L are user-specified parameters, denoting initial water volume, upstream input discharge, iterative time interval, and value column area, respectively. Q_{out} is water discharged from the conduit to the ocean, calculated during the simulation by $Q_{out} = U_0^j S_0^j$.

(4) Solving the coupled water velocity and pressure equations (2) and (3) to obtain $(U_i^{j+1}, P_{w,i}^{j+1})$ using Pressure-Implicit with Splitting of Operators algorithm^[3] based on previous time step value S_i^j , the velocity in the right hand side uses old time step U_i^j .

(5) Updating the conduit size S_i^{j+1} using equation (1).

Numerical Solution:

Assumptions: $M_p \ll M_f$ in the initial stage and M_f is approximately a constant in the quasi-steady stage.

For initial stage:

$$S = S_0 (1 - \beta t)^{-4} \quad (6)$$

$$\beta = \frac{\sqrt{2}}{4} a_0 a_1^{3/2} \frac{\rho_w g^{3/2}}{\rho_i L} S_0^{1/4} K_H \theta f^{-1/2} \quad (7)$$

For quasi-steady stage:

$$S = (\alpha/a_0)^{2/5} g^{-2/5} Q_0^{4/5} \theta^{-2/5} f^{2/5} \quad (8)$$

$$U = (\alpha/a_0)^{-2/5} g^{2/5} Q_0^{1/5} \theta^{2/5} f^{-2/5} \quad (9)$$

For circular shape, $a_0 = \sqrt{4\pi}$, $\alpha = 0.8 \sim 0.9$, θ , Q_0 , f denote conduit slope, upstream discharge, and friction factor.

Application: Diurnal Variations

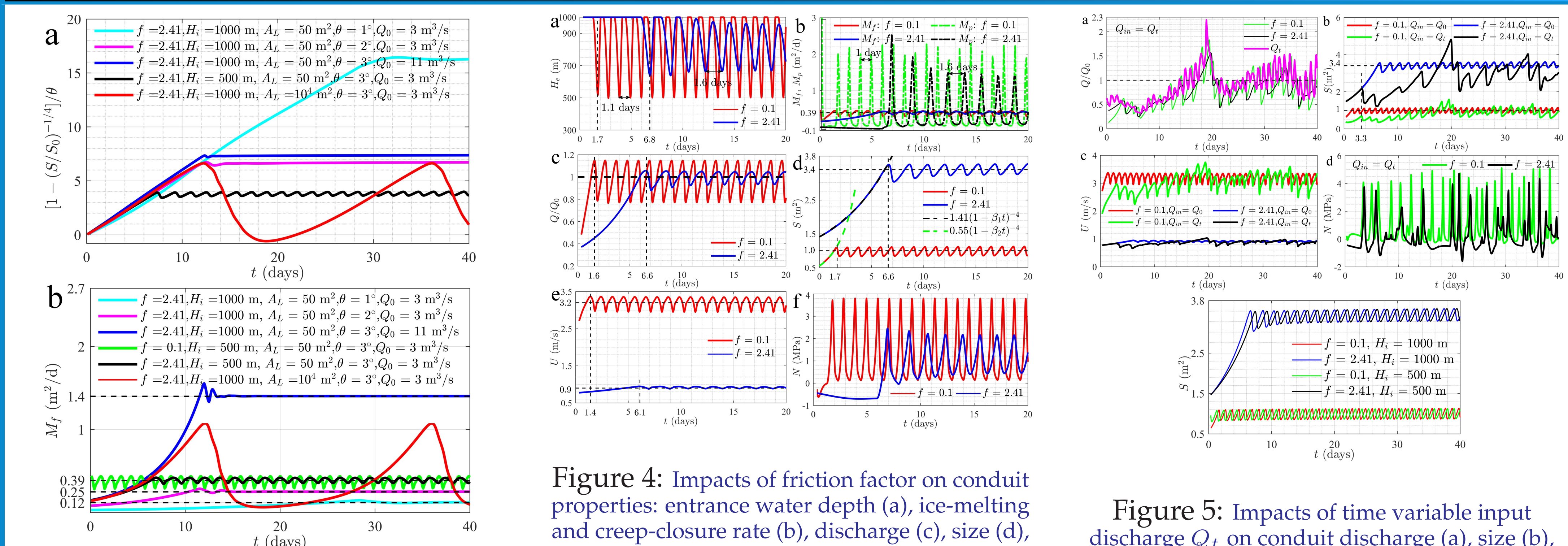


Figure 3: Validating numerical model with analytical solution (a) and test assumption (b).

References

[1] Nerem, R.S., Beckley, B.D., Fasullo, J.T., Hamlington, B.D., Masters, D., MITCHUM, G.T., Climate-change-driven accelerated sea-level rise detected in the altimeter era. Proceedings of the National Academy of Sciences, 2017 114(12), 3128-3133.

Application: Outburst Flood

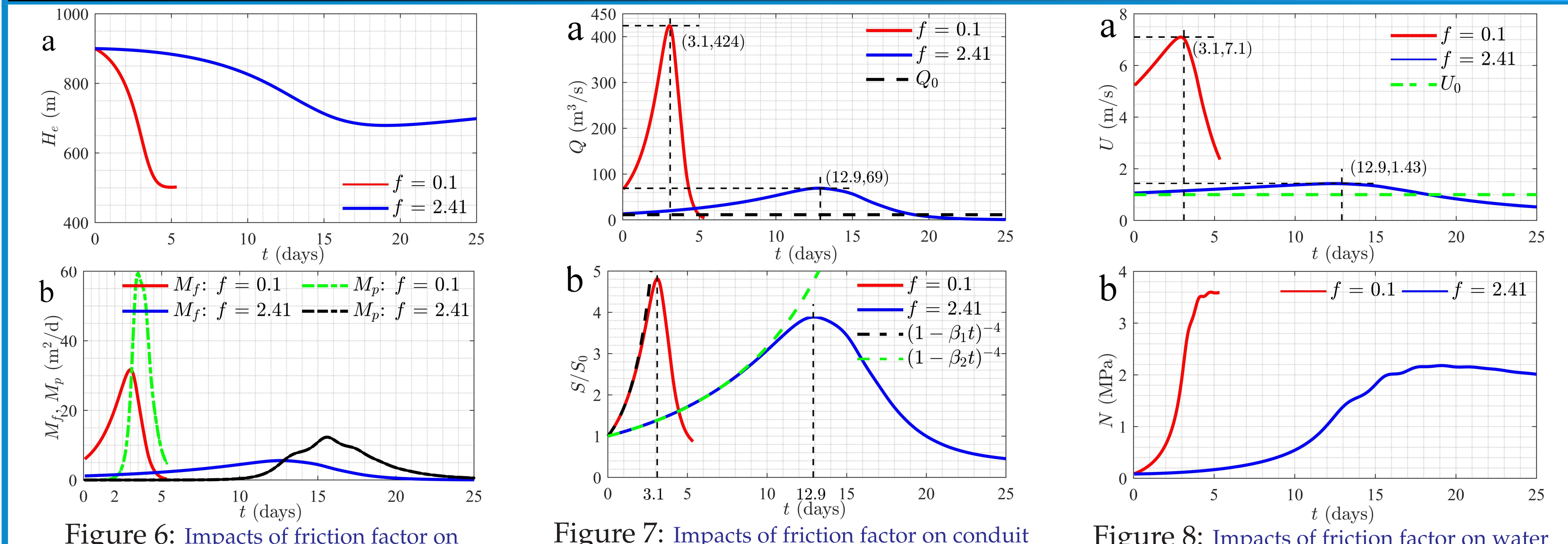


Figure 4: Impacts of friction factor on conduit properties: entrance water depth (a), ice-melting and creep-closure rate (b), discharge (c), size (d), water velocity (e), and effective pressure (f).

[2] Spring, U., Hutter, K., Numerical studies of Jokulhlaups. Cold Regions Science and Technology, 4(3): 227-244, 1981.

[3] Issa, R. I. Solution of the implicitly discretised fluid flow equations by operator-splitting. J. Comput. Phys., 62:40-65, 1985.