



WHAT DO WE KNOW ABOUT DIAGENESIS

1. Generalities relevant to a systems approach

2. Diagenetic processes that affect sediment dynamics

3. Key post-depositional alterations through time and stratigraphic space

David A. Budd
University of Colorado

The True Faces of Diagenesis



DIAGENESIS

- Physical and chemical alteration of sediment and rock in response to changes in its physiochemical regime.
 - Δ fluid flow
 - Δ fluid chemistry (new fluid, reactions within ambient fluid)
 - Δ effective stress
 - Δ temperature
- In carbonates, begins with deposition and extends throughout a rock's history

Carbonate Diagenesis Today

Broad Consensus On

- What diagenetic processes occur in various diagenetic settings
- The conditions that drive most of those processes
- The features that allow us to recognize those processes, conditions, and/or settings
- *Diagenesis is are complex & non-linear with many types of interactions and feedbacks*

Carbonate Diagenesis Today

Less Consensus / Knowledge On

- Microbial influences (catalysis, template)

➡ Lateral spatial patterns in diagenetic products (at scale between one thin section and “regional” patterns)

➡ Presence or absence of self-organization and/or **scale invariance** (hard to evaluate due to location, scales of analysis, & overprinting of diagenetic processes)

- Strength of the linkages (feedbacks) between processes

➡ Thresholds at which changes might “cascade”

Unexplored Spatial Patterns

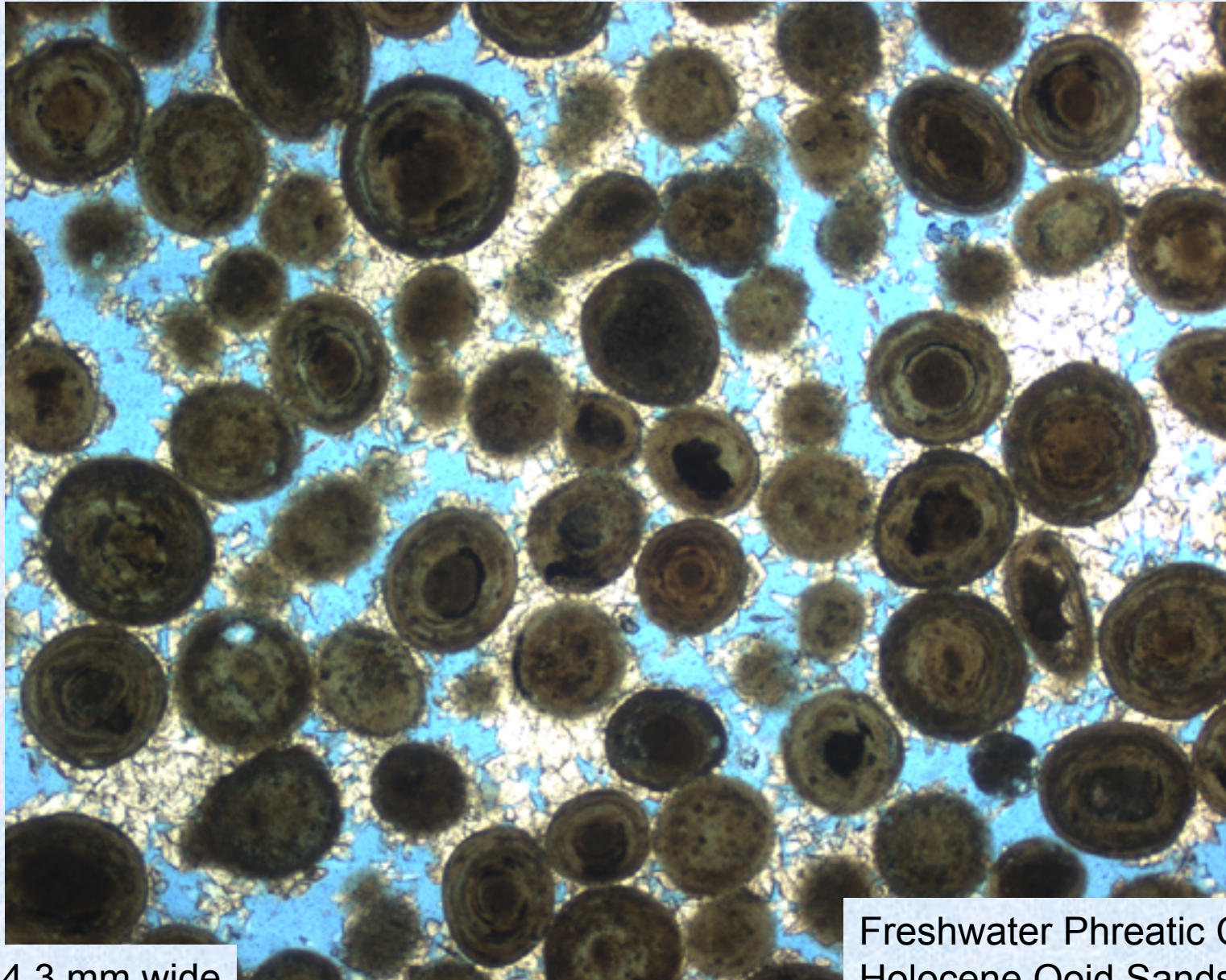


ates

Carbonate nodules at -3m horizon in channel bar sandstone, Cretaceous Frontier Fm., WY (Lee et al, 2007) revealed by GPR

FOV is 80 m wide

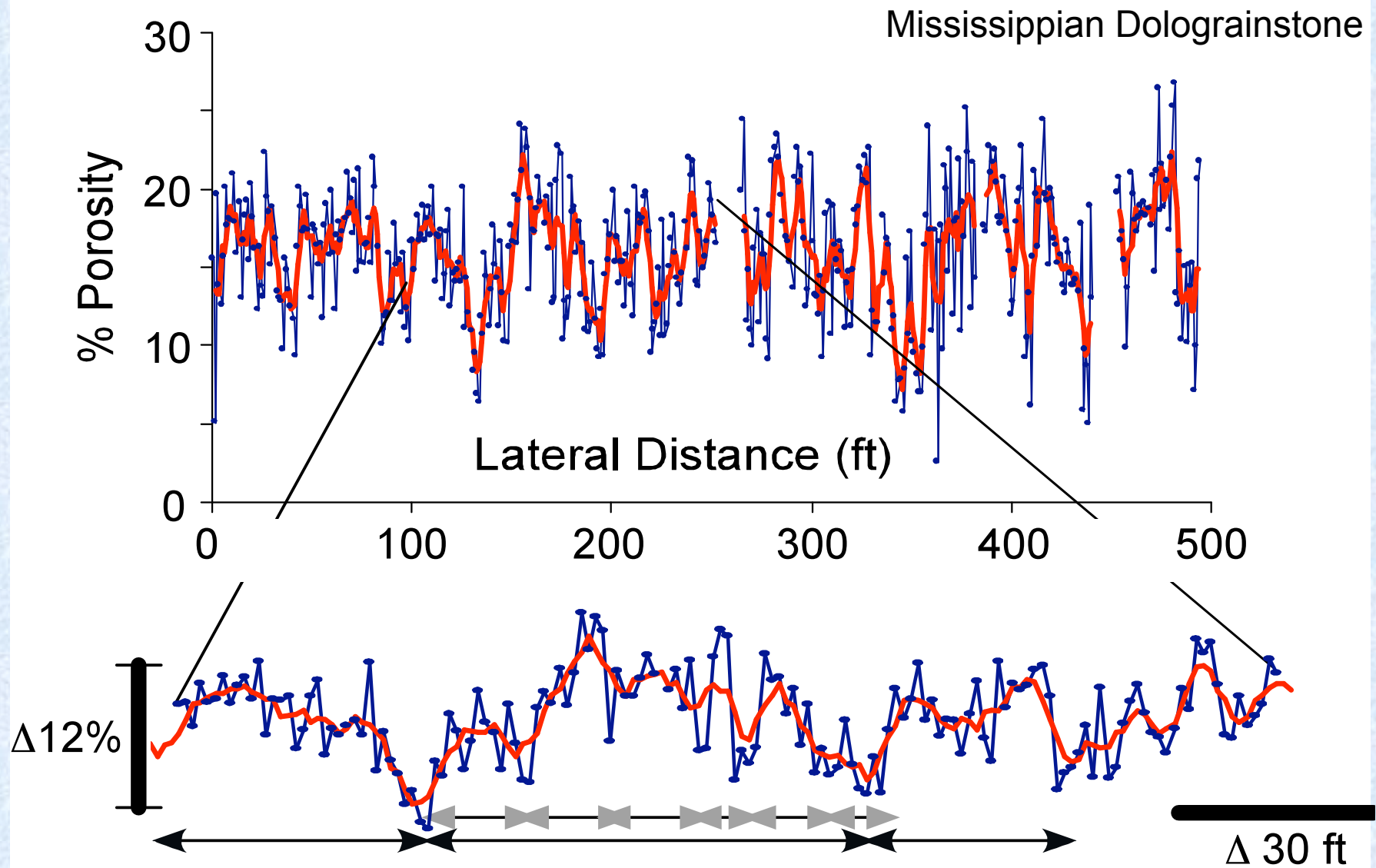
Unexplored Spatial Patterns



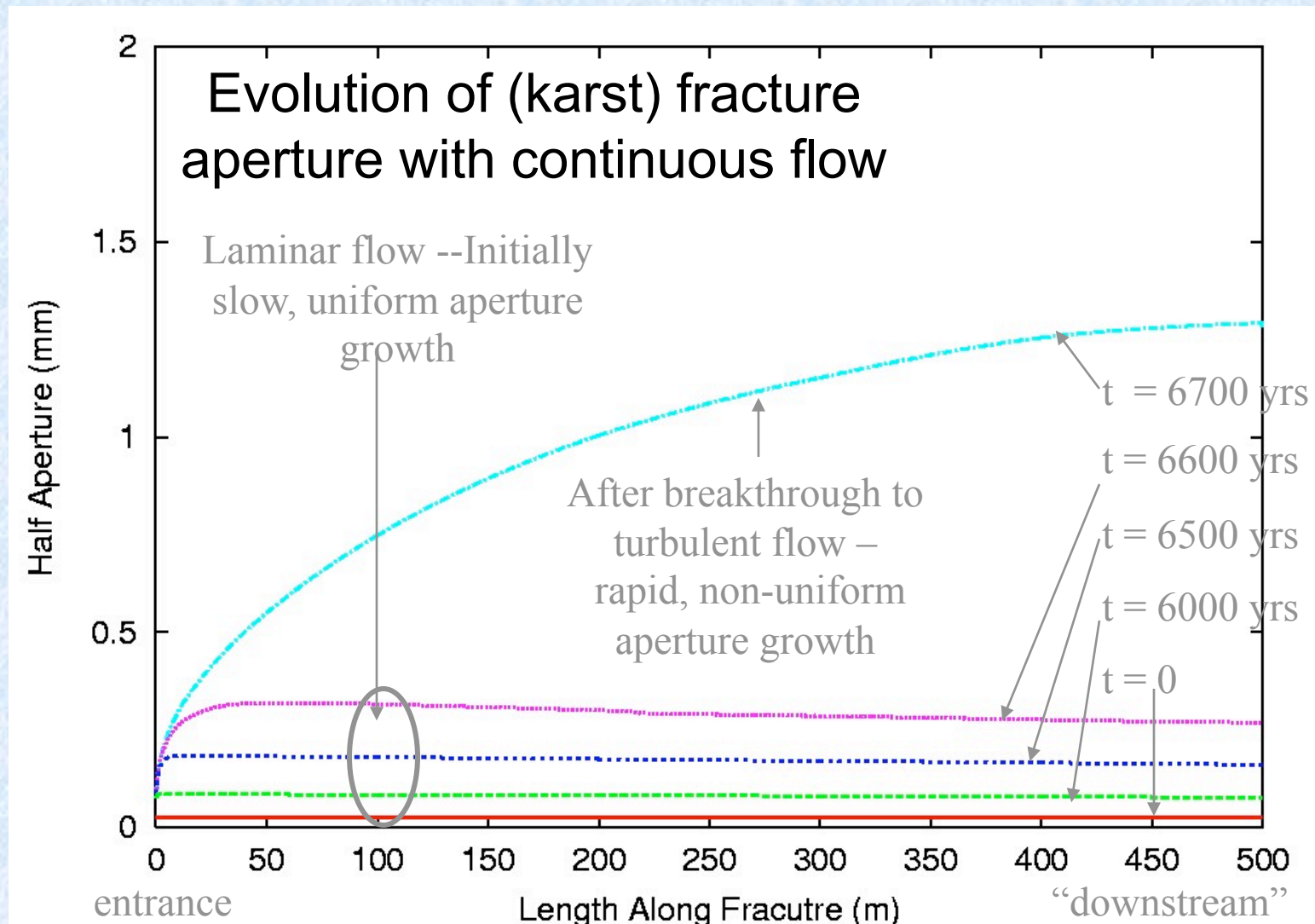
FOV is 4.3 mm wide

Freshwater Phreatic Cements,
Holocene Ooid Sands

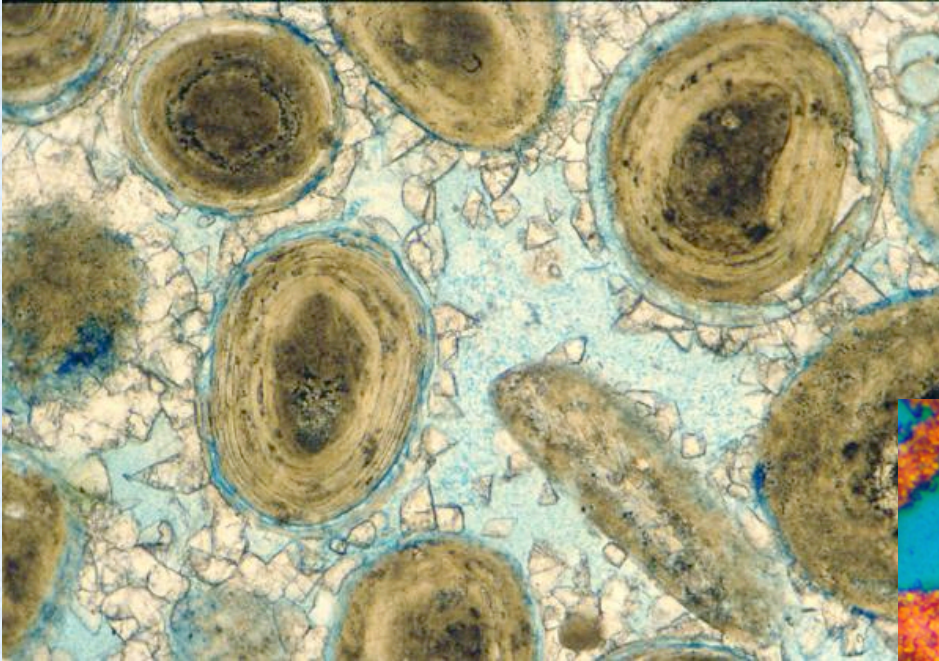
Unexplored Spatial Patterns



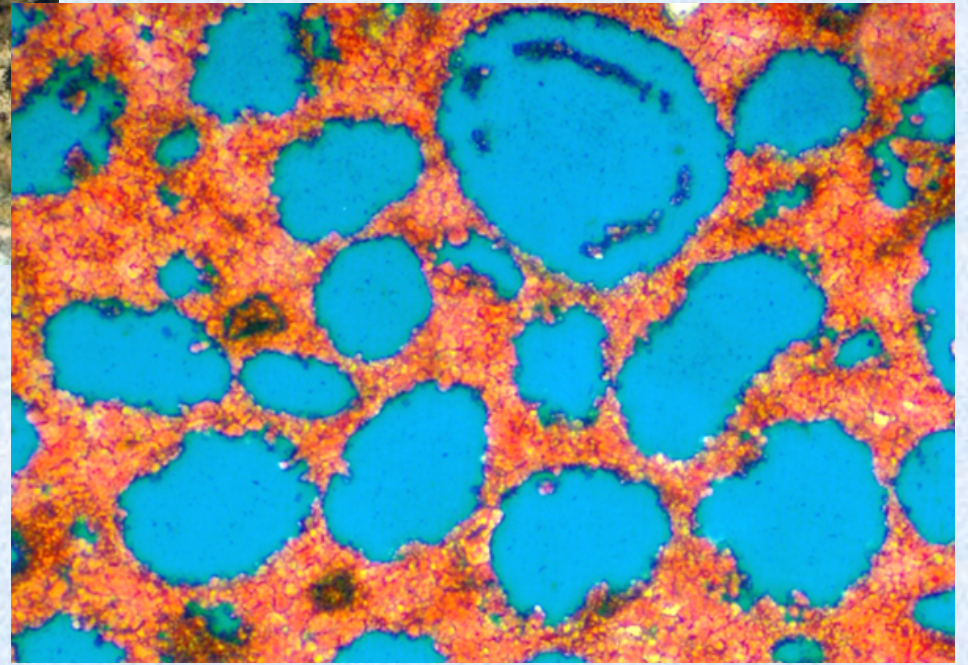
Thresholds to Diagenetic Cascading?



Thresholds to Diagenetic Cascading?



Holocene, Schooner Cays, Bhms



Permian (Scholle & Ulmer Scholle, 2003)

Carbonate Diagenesis Today

Diagenesis is complex & non-linear with many types of interactions and feedbacks,

but lots of uncertainty as to the comprehensive nature of diagenesis as a SYSTEM.

**What is important from a
CSDMS point of view?**

CSDMS Vision - Tools That

- Predict the erosion, transport, and deposition of sediment and solutes in landscapes and their sedimentary basins
- Address the challenging problems of surface-dynamic systems: self-organization, localization, thresholds, strong linkages, scale invariance, and interwoven biology & geochemistry.
- Can be applied to when the earth was abiotic, hotter or colder, when the moon was closer, the oceans more saline or stratified, sea levels rapidly changing, etc
- Are valuable to those working on modern environmental applications, future global warming scenarios, natural disaster mitigation efforts, natural hazard efforts, reservoir characterization, oil exploration, and national security.

Diagenetic alterations that affect concurrent & subsequent sediment dynamics

Post-depositional alterations that affect rock properties through time and stratigraphic space

An aerial photograph of a coastal area, showing a mix of blue water, sandy beaches, and some green vegetation. The water is a deep blue, while the beaches are a lighter, sandy color. The overall scene is a natural, coastal environment.

Alterations that affect sediment dynamics

Sea Floor Precipitation

- hardgrounds, reefs, slopes

Sediment Dissolution

Karstification of Exposed Depositional Surfaces

- surface denudation (epikarst)
- generation of new land surface topography
(karren, dolines, sinks)

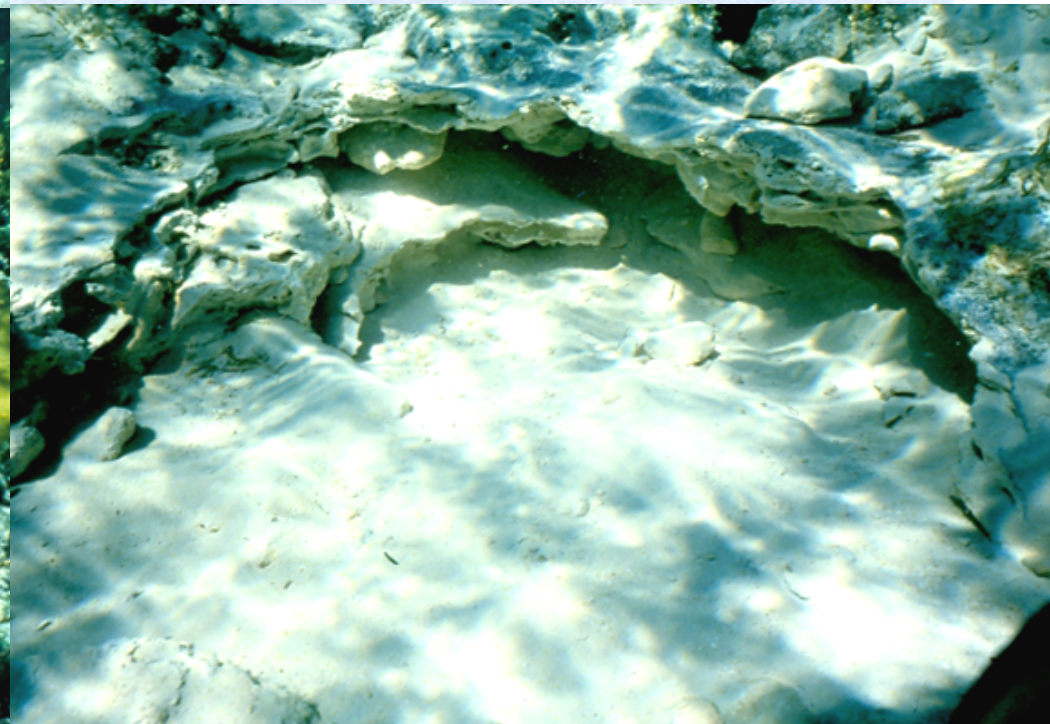
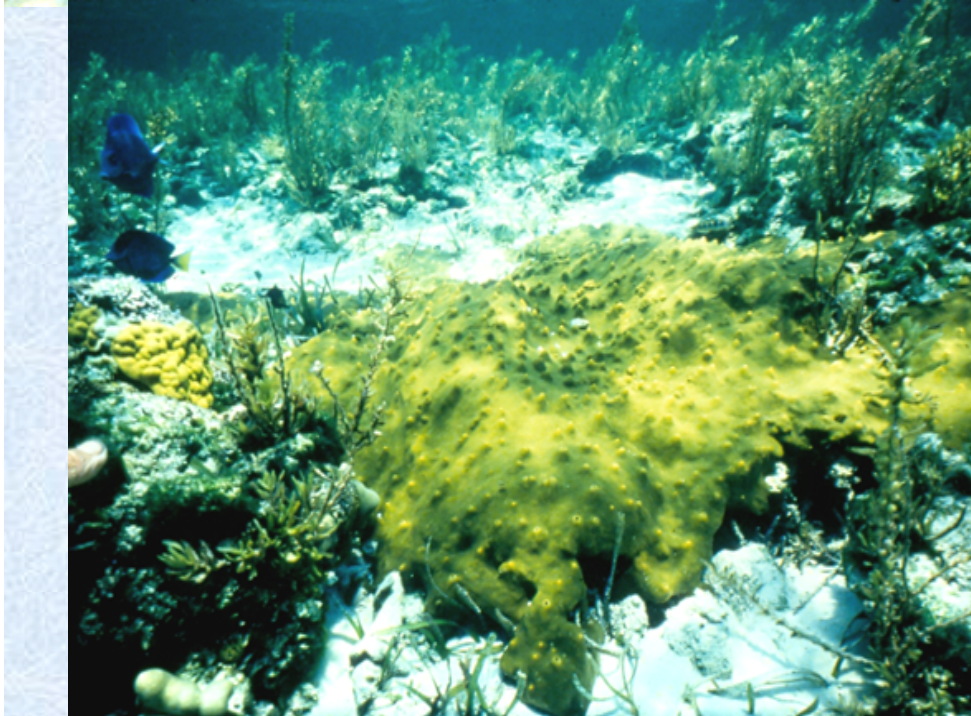
Linkages to Hardground Cementation



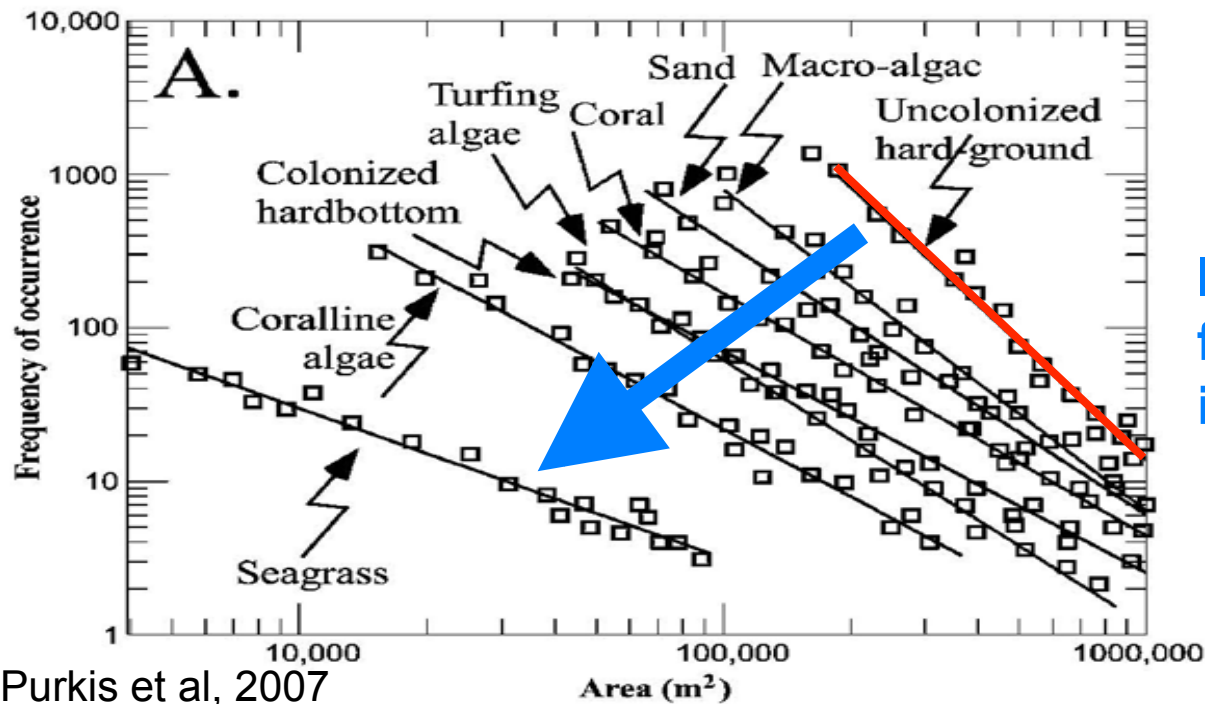
Even minimal cementation can stabilize sediment

Hardgrounds can become surfaces for biohermal development

Multiple or thick cemented layers can become permeability barriers to later fluid flow.

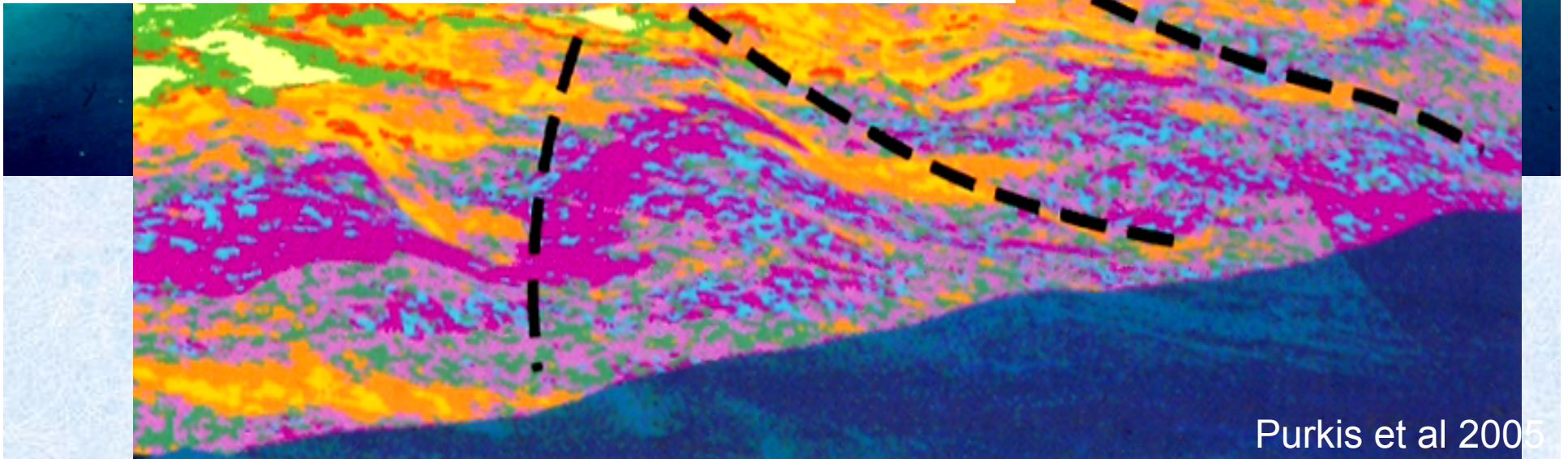


Spatial distribution of hardgrounds



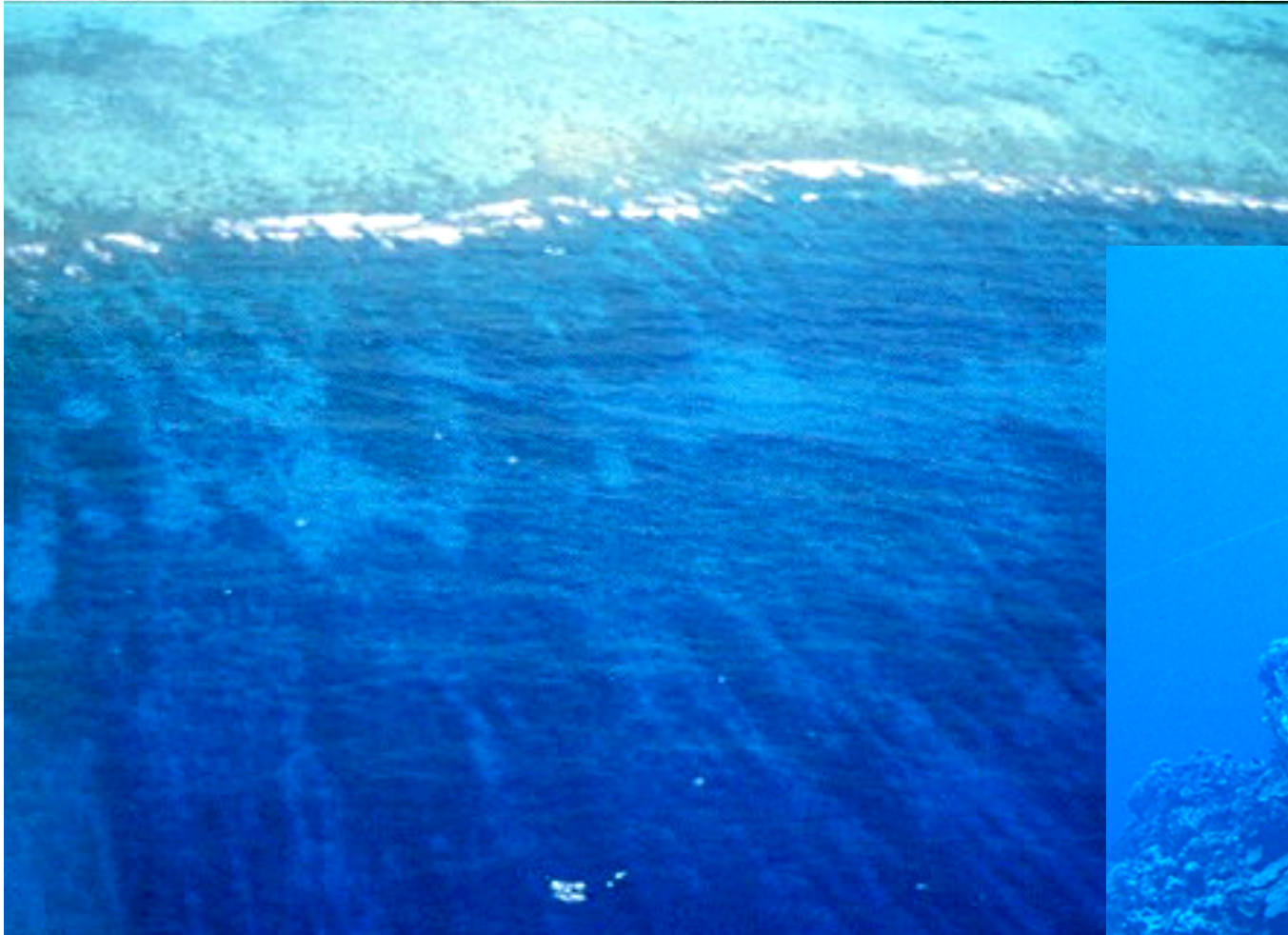
Purkis et al, 2007

Increasing landscape
fragmentation;
increasing complexity

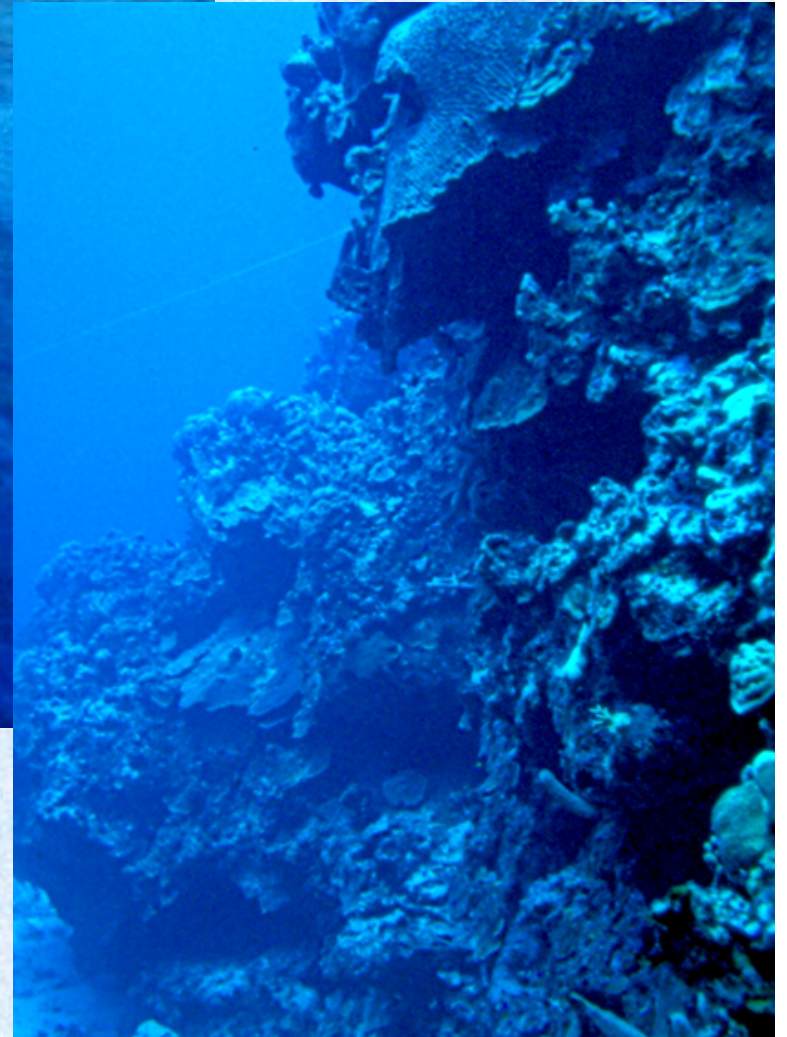


Purkis et al 2005

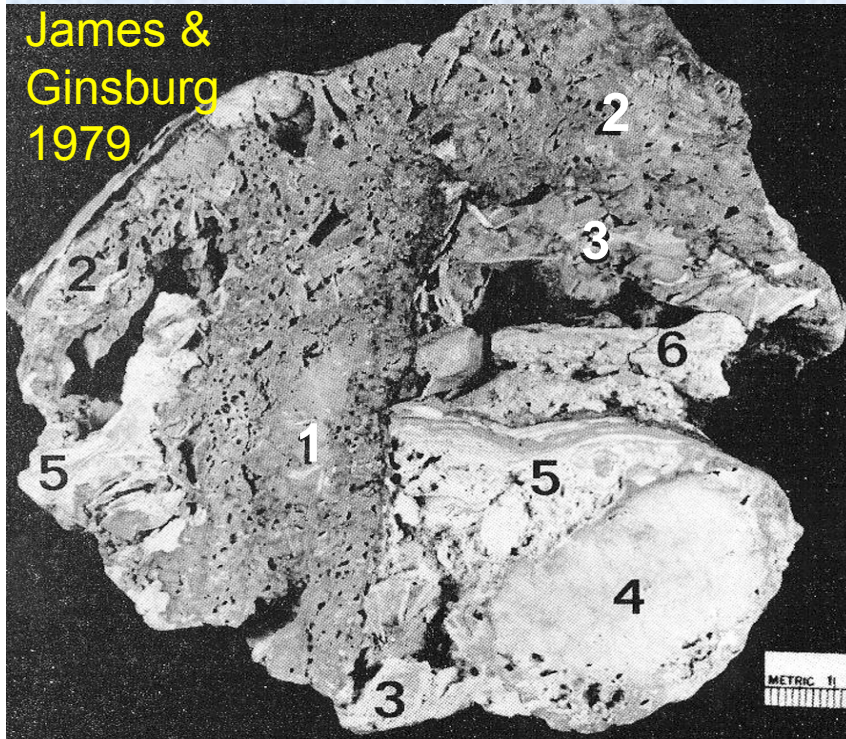
Reef Cementation



Integral component of reef construction, with potential feedbacks to (and from) biota, reef geomorphology, and nature of the eventual sedimentary accumulation

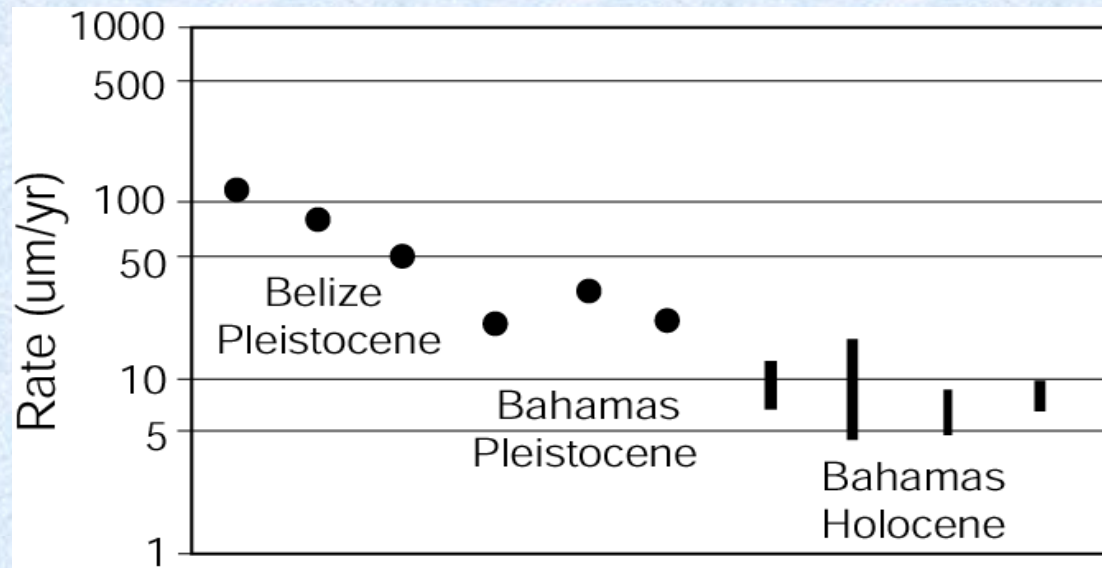


James &
Ginsburg
1979



Reef Cementation – Multigenerational and fast!

- 1 – Original Coral Plate
- 2 – Surrounding Halimeda wackestone (cemented)
- 3 – Boring in cemented mudstone that is filled w/ cemented sediment
- 4 – 2nd cavity in fill of the 1st has a cemented coral fragment at its base
- 5 – laminated sediment that partially fills 2nd cavity; capped by iron oxide crust
- 6 – final grainstone fill of cavity (cemented)

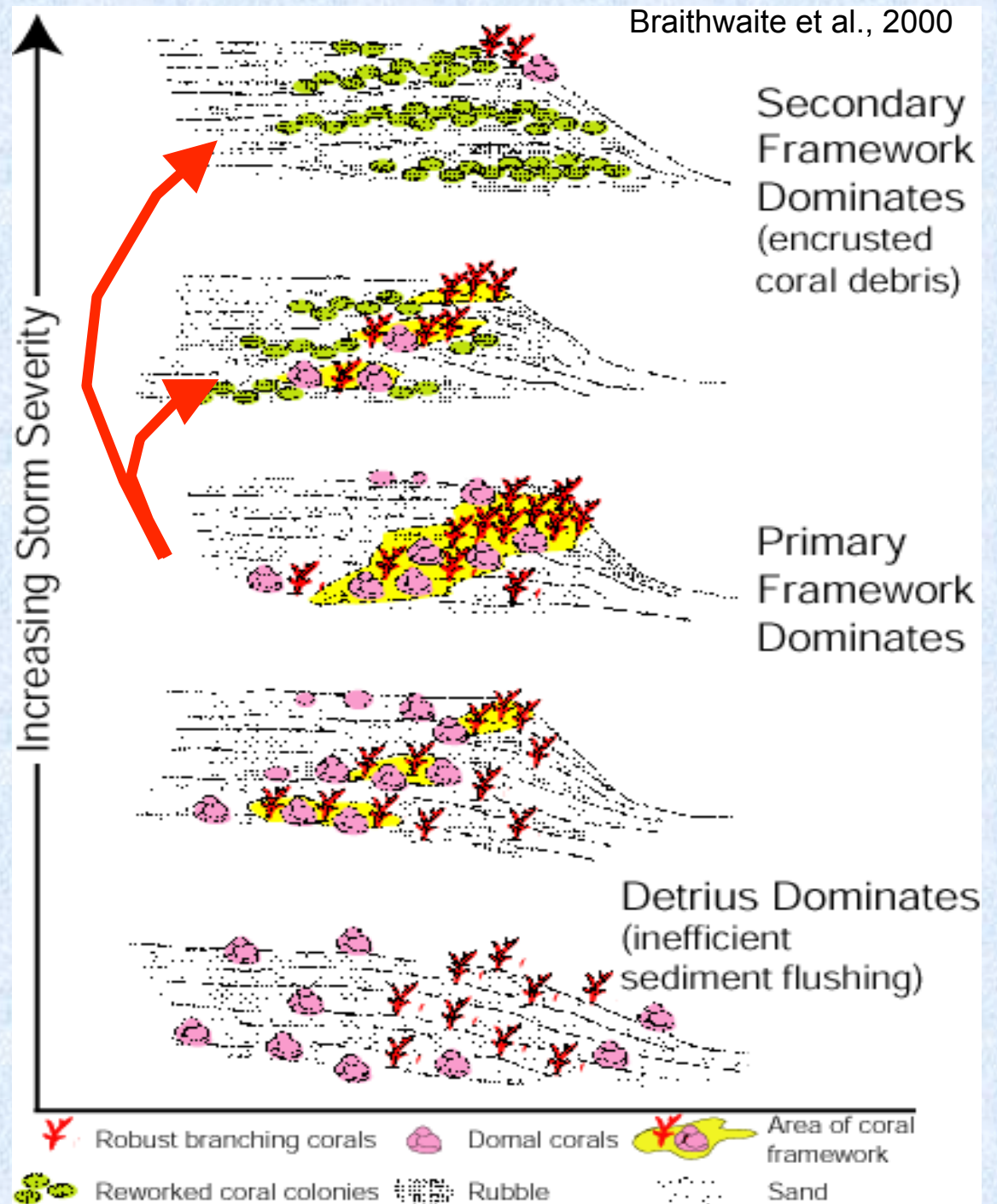


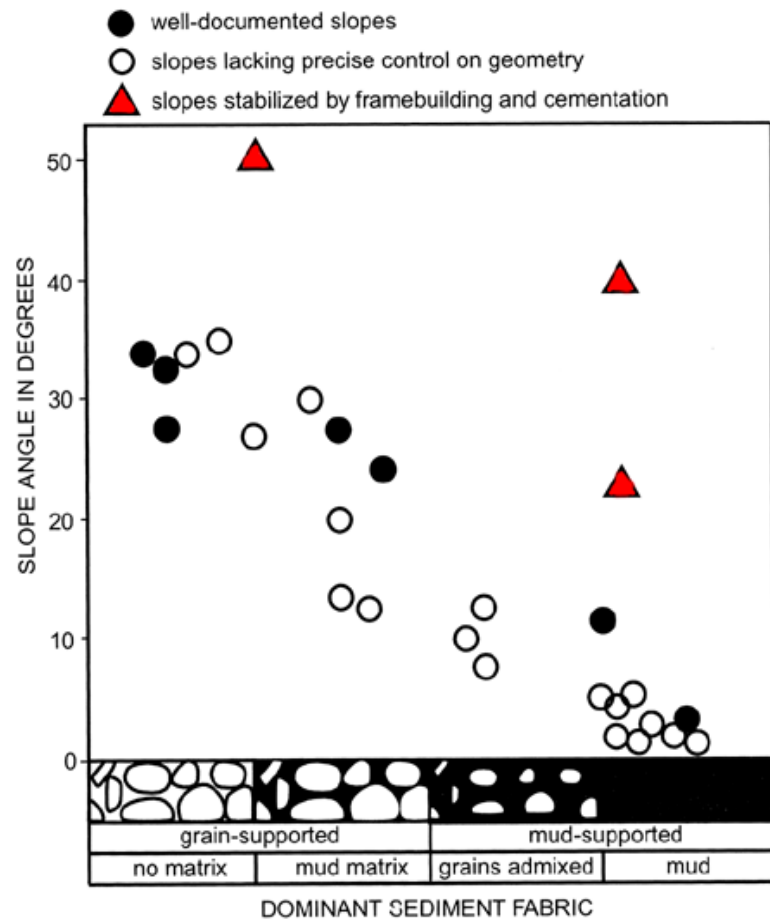
Grammer et al., 1999

Marine cementation
could exhibit rapid
responses
(positive or negative)
to changes in
biogeochemical
drivers of cementation

Fair-weather hydrodynamic energy and severe storms affect the ratio of framework to detrital facies

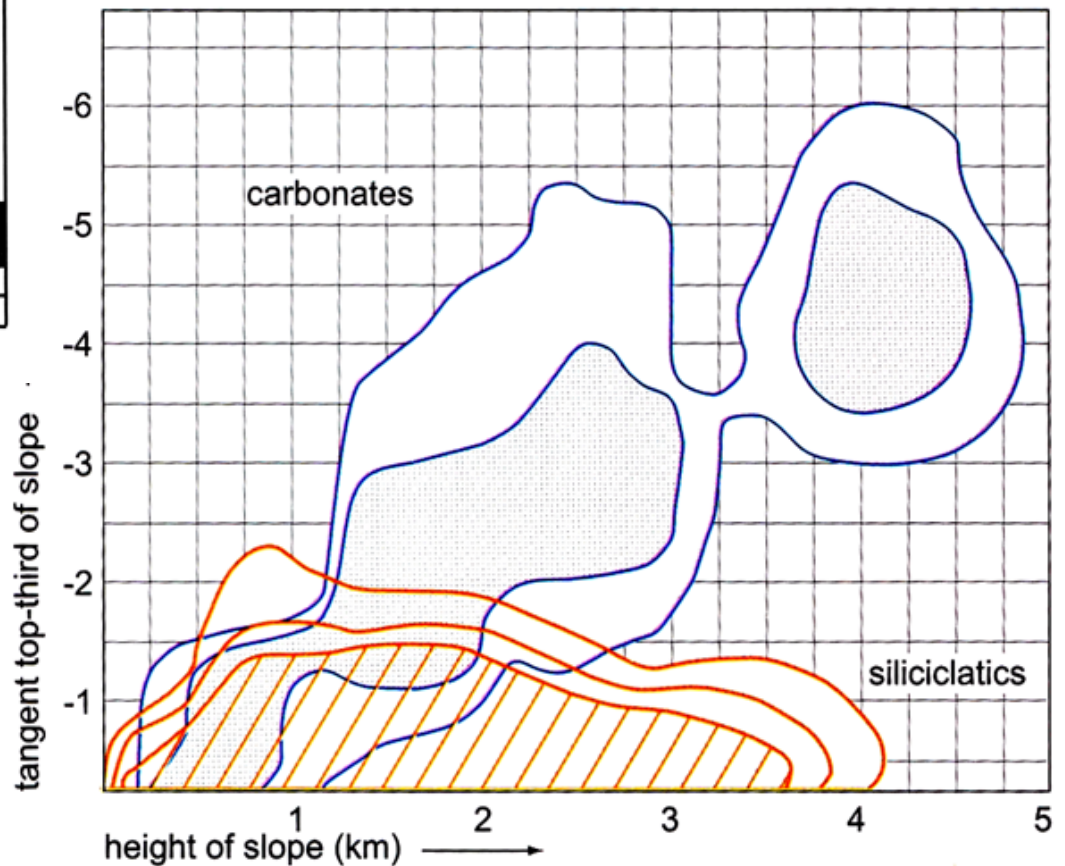
Did higher carbonate saturation states & more cementation determine ultimate reef accumulation patterns in the past?





Slope Cementation

Strong linkages between marine cementation on slopes and slope shape -- and hence platform geometry.



Kenter 1990

Schlager 1989

Sediment Dissolution

Dissolution of sediment was long considered negligible

Since early 1990s, sediment dissolution has been documented

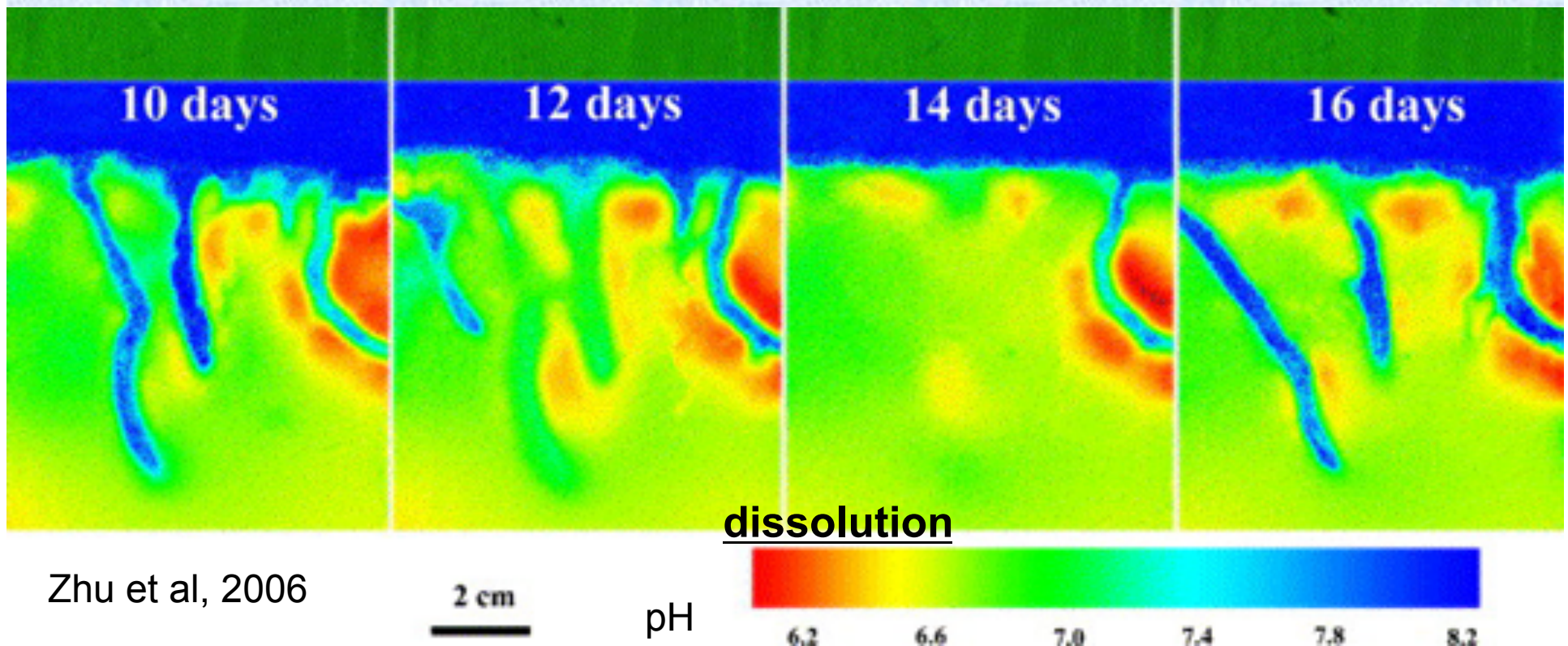
- reaction zone to 10's cm below sediment/water interface
- 0% to >50% of bulk production affected

Affected by

- microbial oxidation of organic matter
- sulphate reduction & oxidation of resultant H₂S
as it migrates upward
- metastability of aragonite in cool-water settings
- carbonate mineralogies and crystal sizes
- diffusive & advective flux rates (bioirrigation beneficial)

Implication of sediment dissolution to modeling carbonate deposition

- if using sediment production functions, one should seemingly consider dissolution
- but if using net accumulation rates conditioned by modern data, dissolution is accounted for



Key linkages if including cementation & dissolution in seascape

Carbonate saturation of seawater (not a constant plus variability from the micro to macro environment)

On-bank seawater fluxes

Wave & tidal energies (fluid fluxes)

Organic matter content of sediments

Bioturbation

Sediment mineralogies, initial petrophysical attributes

But we don't really know spatial patterns by which to evaluate model outcomes

Alterations that affect concurrent & subsequent sediment dynamics

Sea Floor Precipitation

- hardgrounds, reefs, slopes

Sediment Dissolution

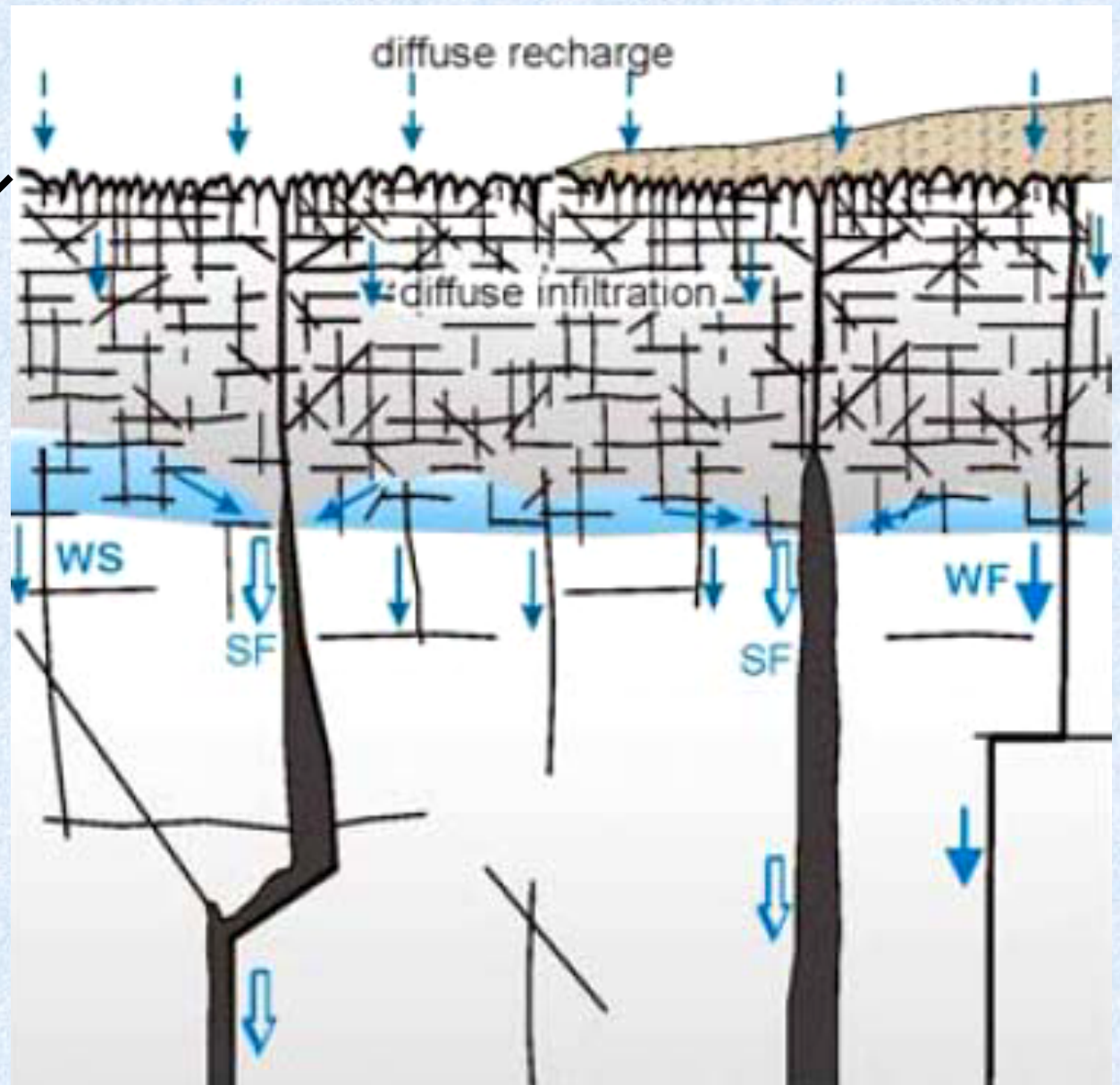
Karstification of Exposed Depositional Surfaces

- surface denudation (epikarst)
- creates topography on landscape

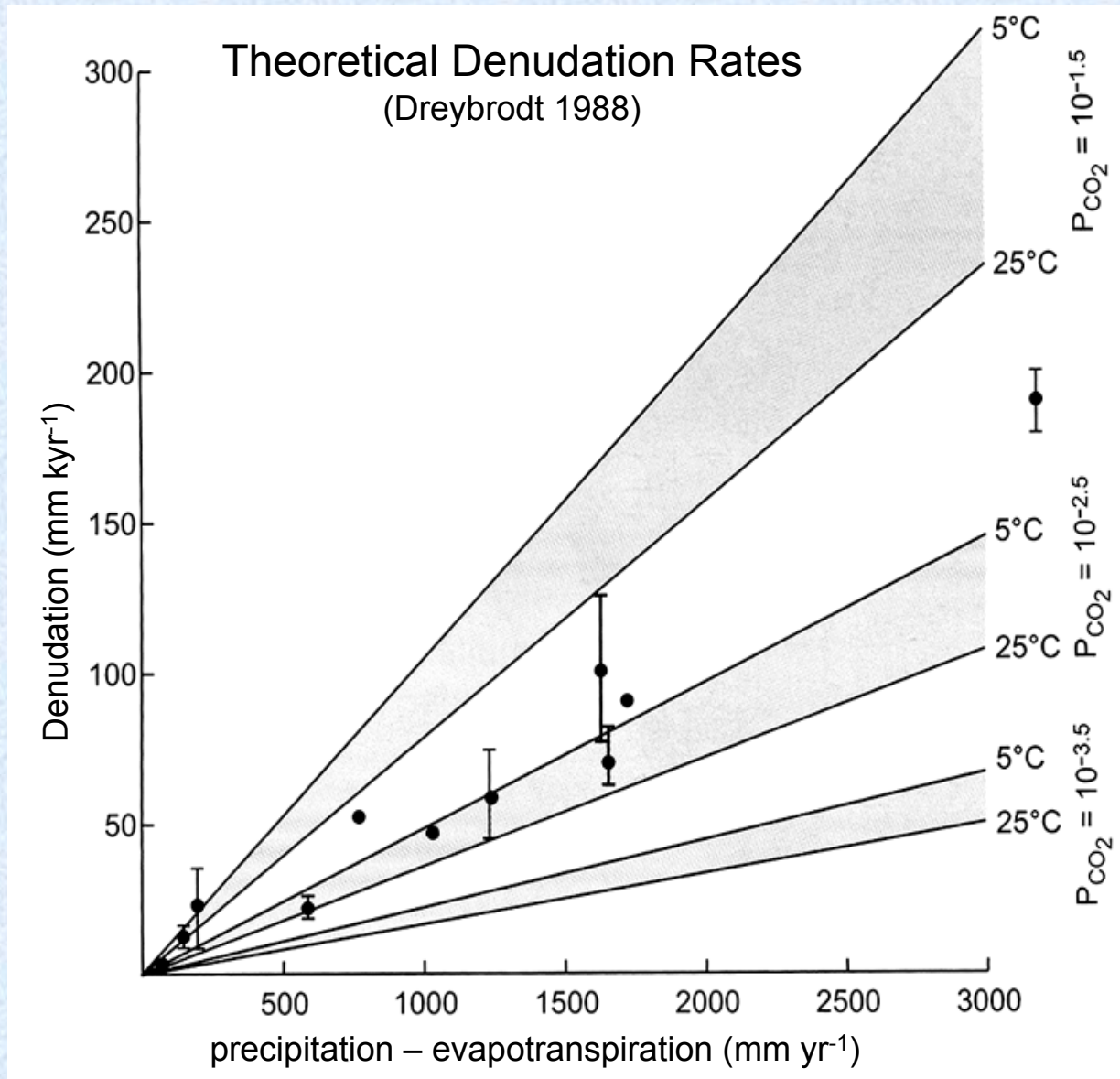
Epikarst zone –
interface between
soil and rock,

between the land
surface and the
vadose zone

Zone of denudation



Klimchouk, 2004



Observed and theoretical rates of surface karst denudation dependent on climate, soil zone PCO_2 , and time.

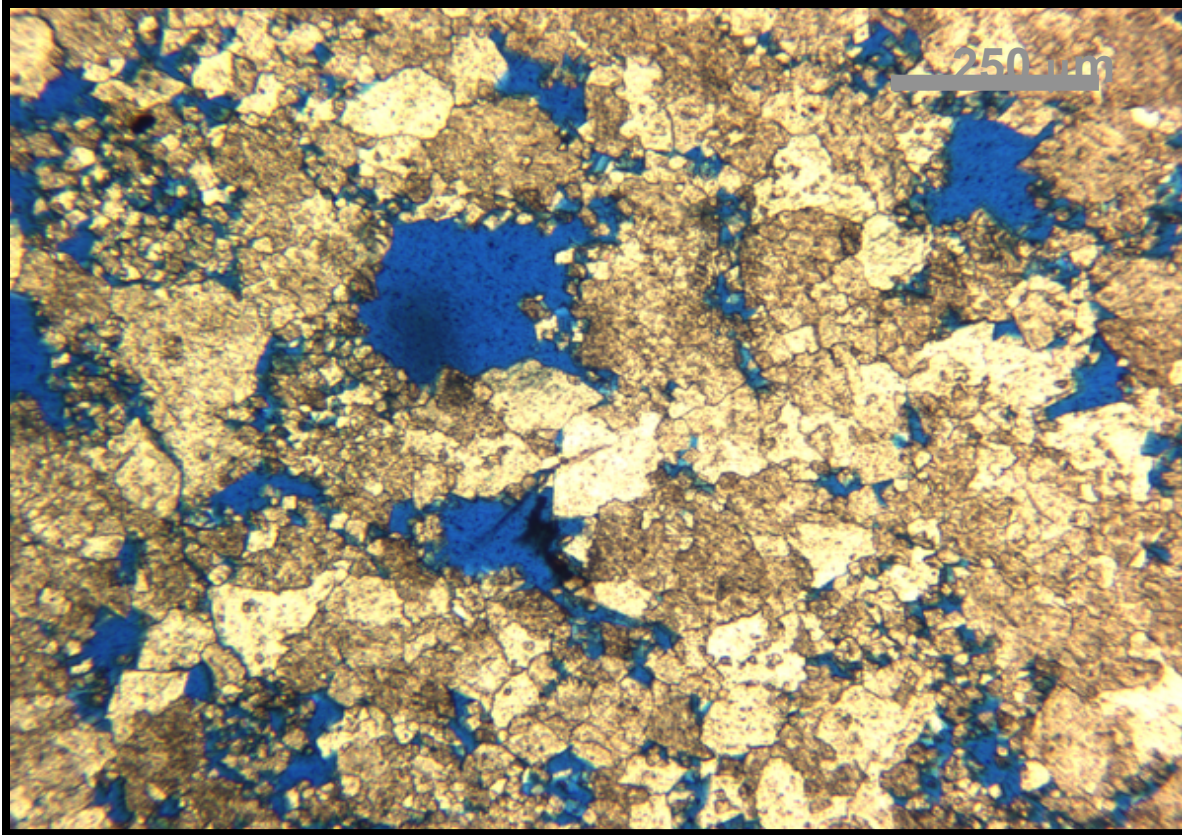
Magnitude of the process:

20 to 100 cm per 10,000 yr

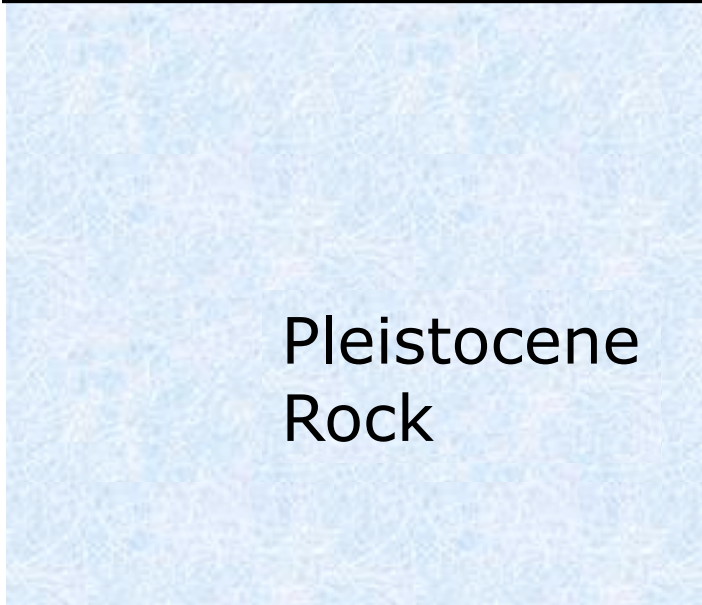
Post-depositional alterations through time and stratigraphic space

Mass transfers that:

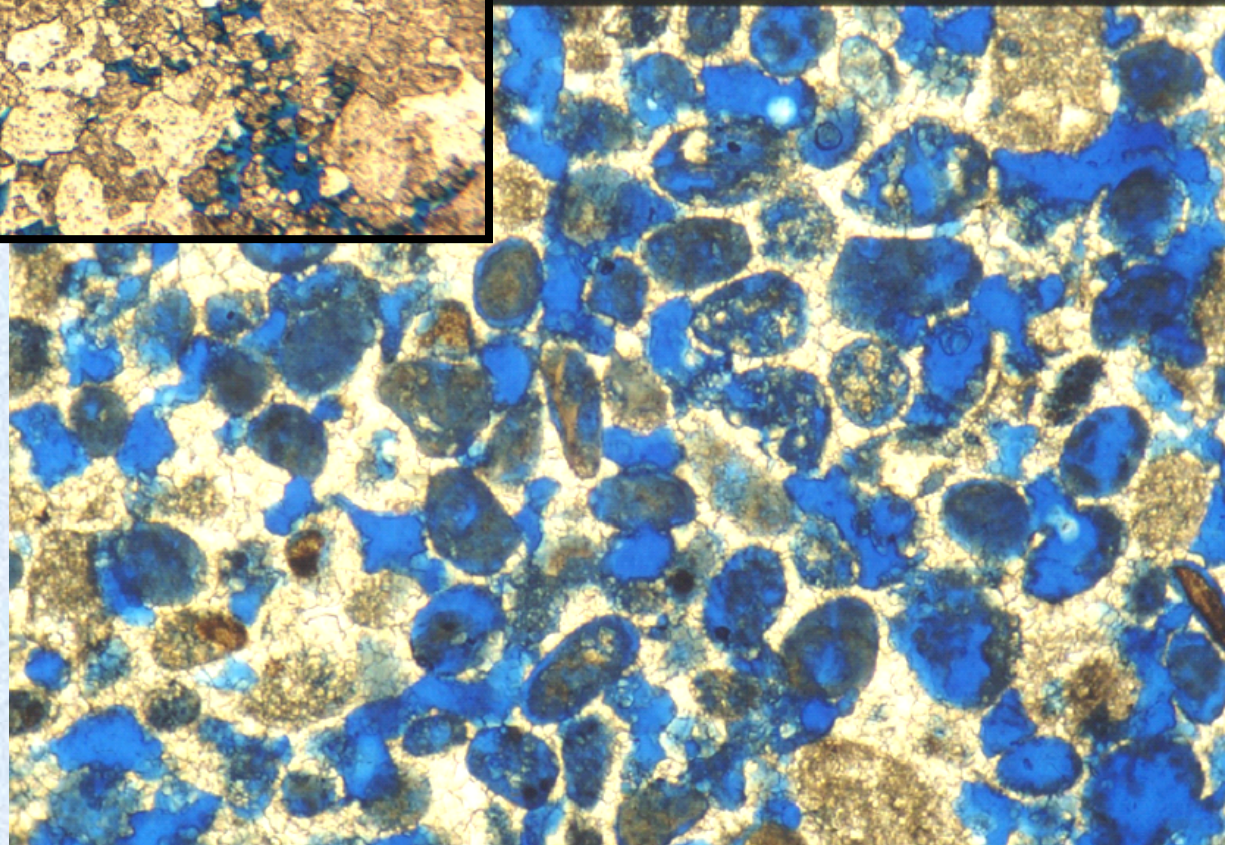
- Change mineralogies
- Alter rock fabrics & textures
- Restructure, and generally reduce, porosity and permeability



Mississippian
Dolostone

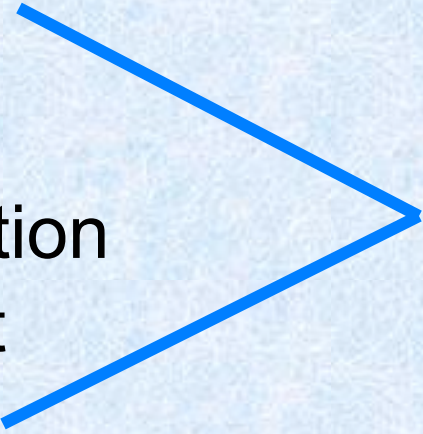


Pleistocene
Rock



Known Diagenetic Processes

- Dissolution
- Cementation
- Recrystallization
- Replacement
- Fracturing
- Mechanical Compaction
- Chemical Compaction



Known from the land surface upon exposure to the deep subsurface

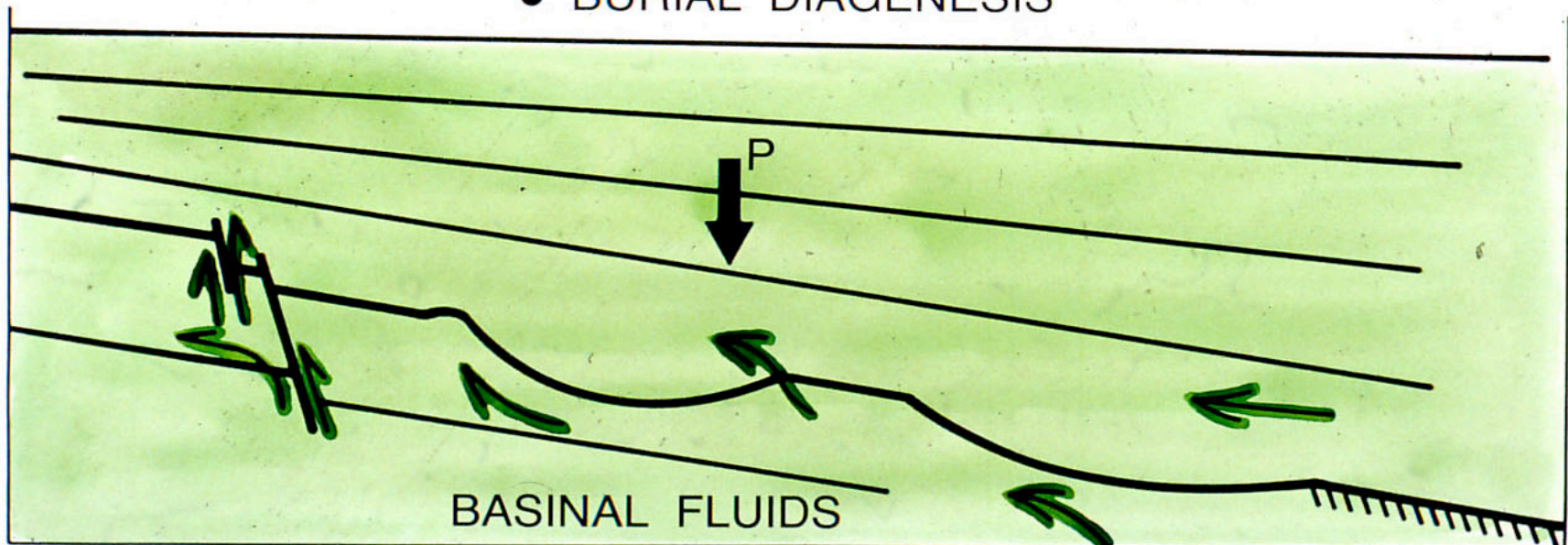
Known from 10s of meters & ~300 m below the surface and downward

Diagenesis historically viewed in terms of hydrochemical regimes that generate internally consistent process-response dynamics

- SURFICIAL DIAGENESIS



- BURIAL DIAGENESIS



Known driving mechanisms for alteration

- **Mineralogical stabilization** (aragonite→calcite; high Mg calcite→calcite; metastable dolomite→dolomite)
- **Mineralogical disequilibrium** (generated by fluid fluxes)
- **CO₂ effects** (CO₂ charging, degassing, effervescence)
- **Microbial mediated reactions** (generate acid)
- **Common ion effects** (anhydrite dissolution; dedolomitization, calcite ppt)
- **Organic diagenesis** (oxidation, decarboxylation)
- **Evaporation**
- **Fluid mixing effects** (freshwater-seawater, freshwater-freshwater, brine-brine)
- **Dehydration reactions** (gypsum→anhydrite, smectite→illite, opalA→quartz)
- **Lithostatic load** (mechanical compaction)
- **Effective stress** (chemical compaction)
- **Tectonic stress**
- **Ostwald ripening**
- **Heat flux**

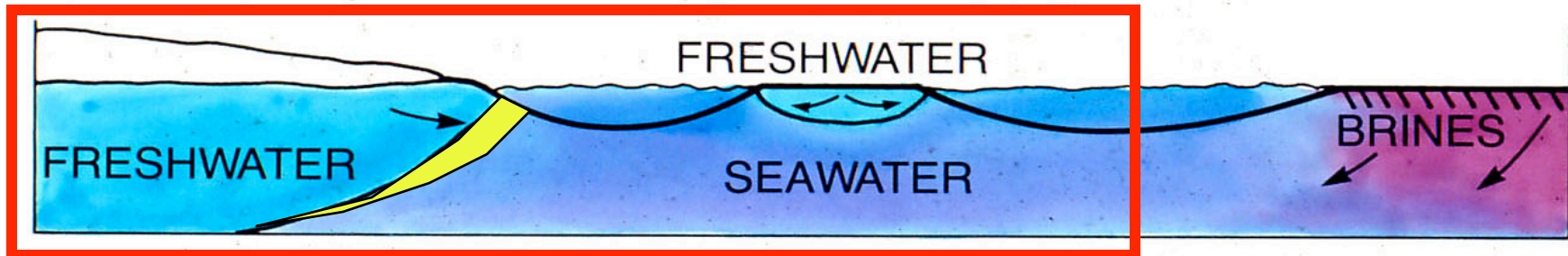
And of course kinetics must be considered ¹

- Nucleation energies
- Reactive surface area (grain size, crystal size)
- Surface-control or transport-controlled reactions
- Crystal-surface contamination & cation-complexing (foreign ion effects, Mg-hydration, PO_4 , SO_4)
- Reaction rates near & far from equilibrium
- Microbial effects (mediation, catalysis, ...)
- Temperature
- Fluid flow rates (relative to reaction rates)

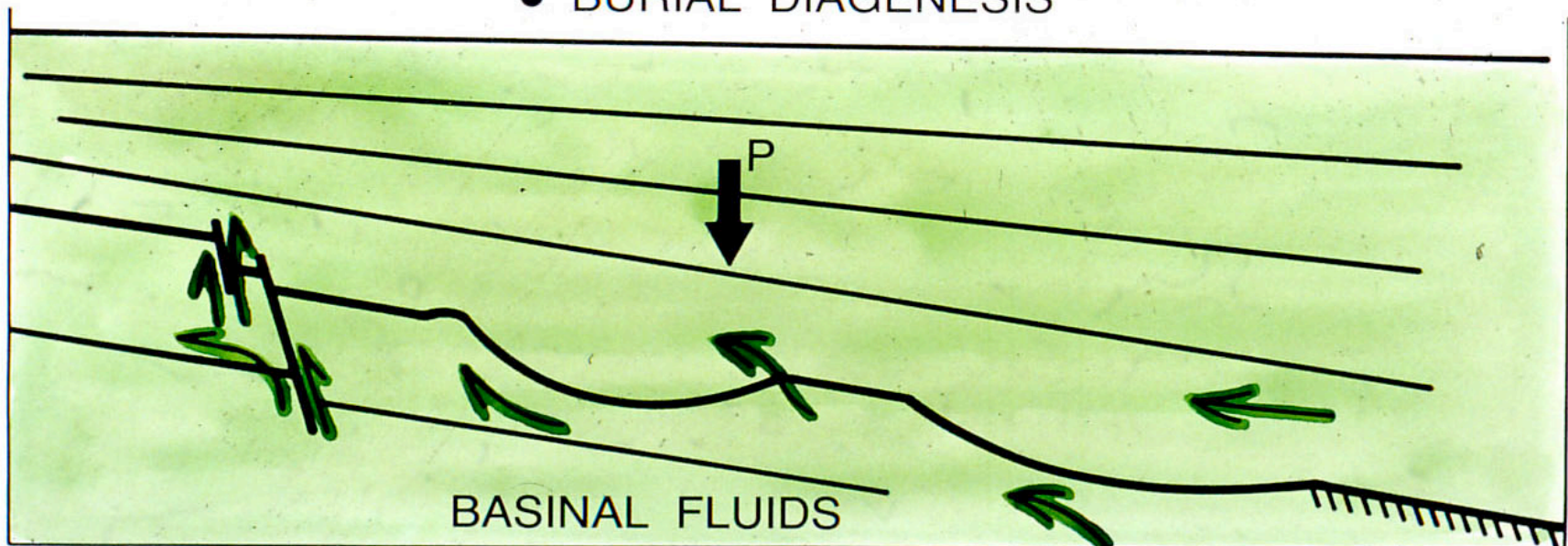
1 – Assuming thermodynamics are also known, but see “dolomite”

Diagenesis historically viewed in terms of hydrochemical regimes that generate internally consistent process-response dynamics

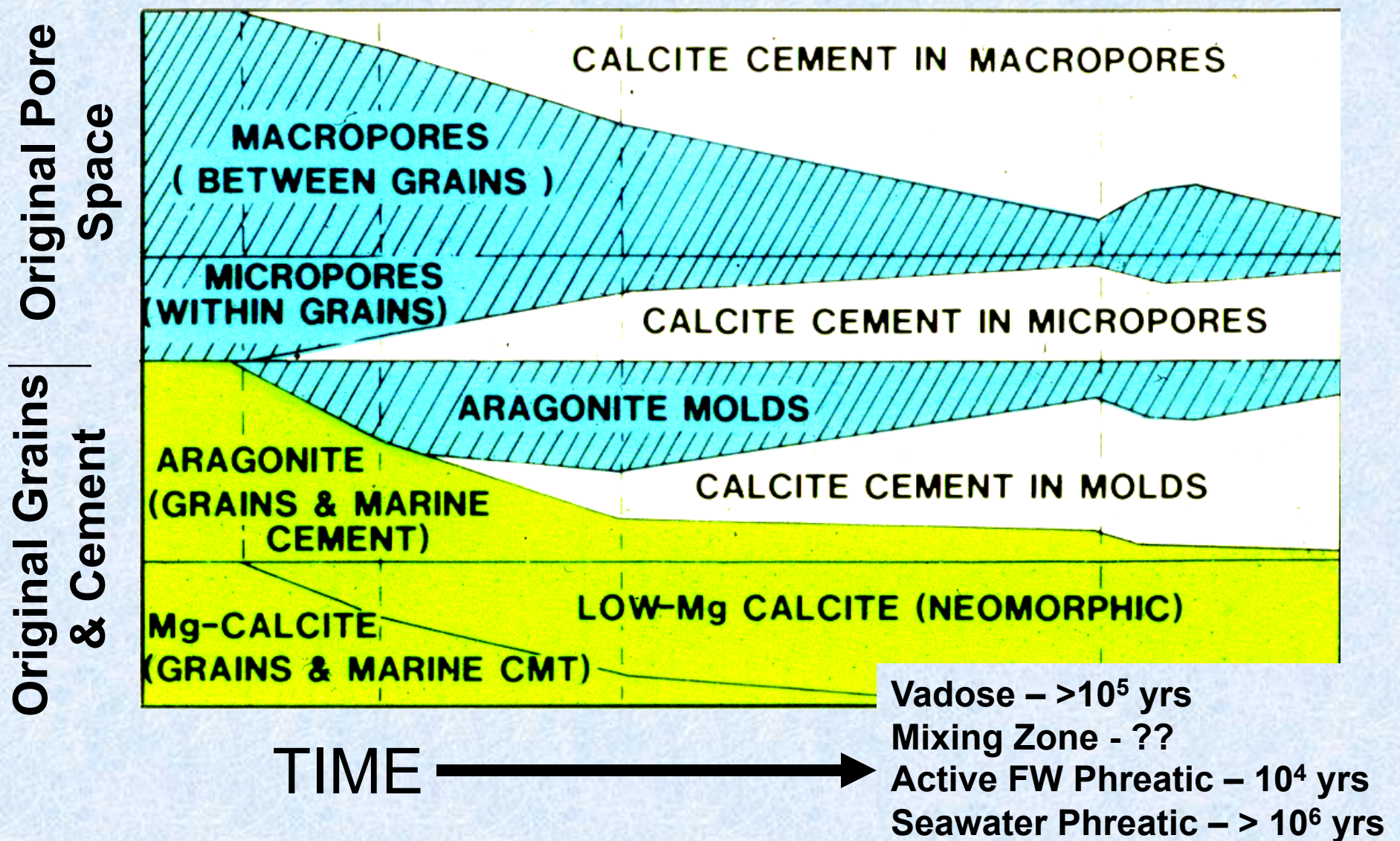
- SURFICIAL DIAGENESIS



- BURIAL DIAGENESIS

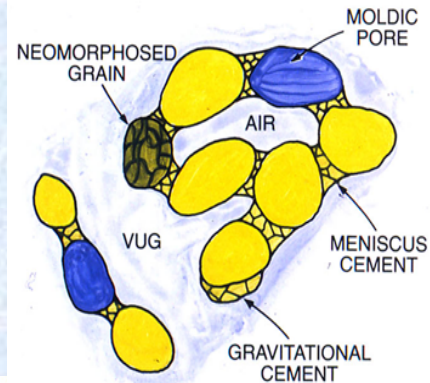


Mineralogical Stabilization

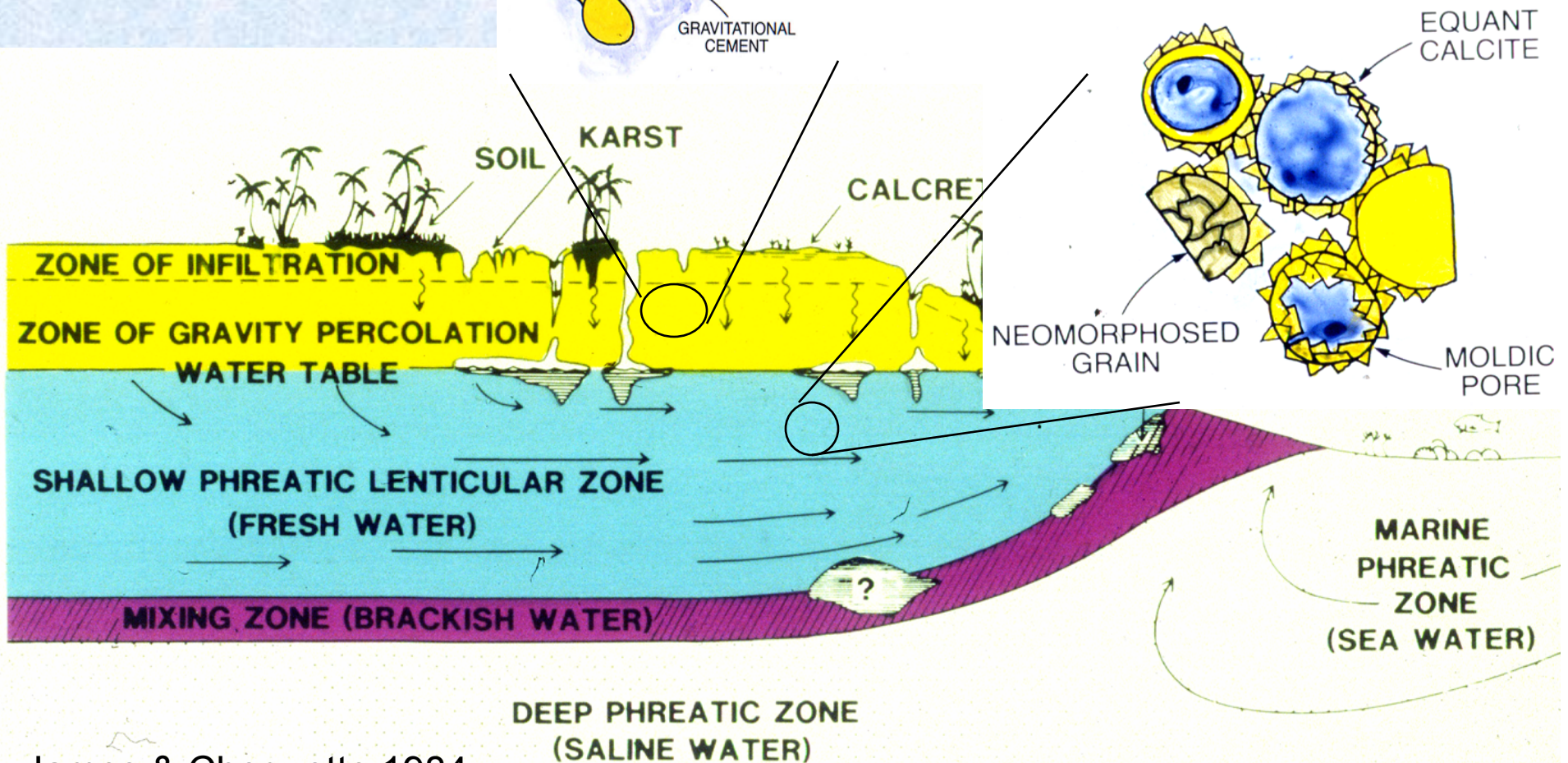


Meteoric Diagenesis

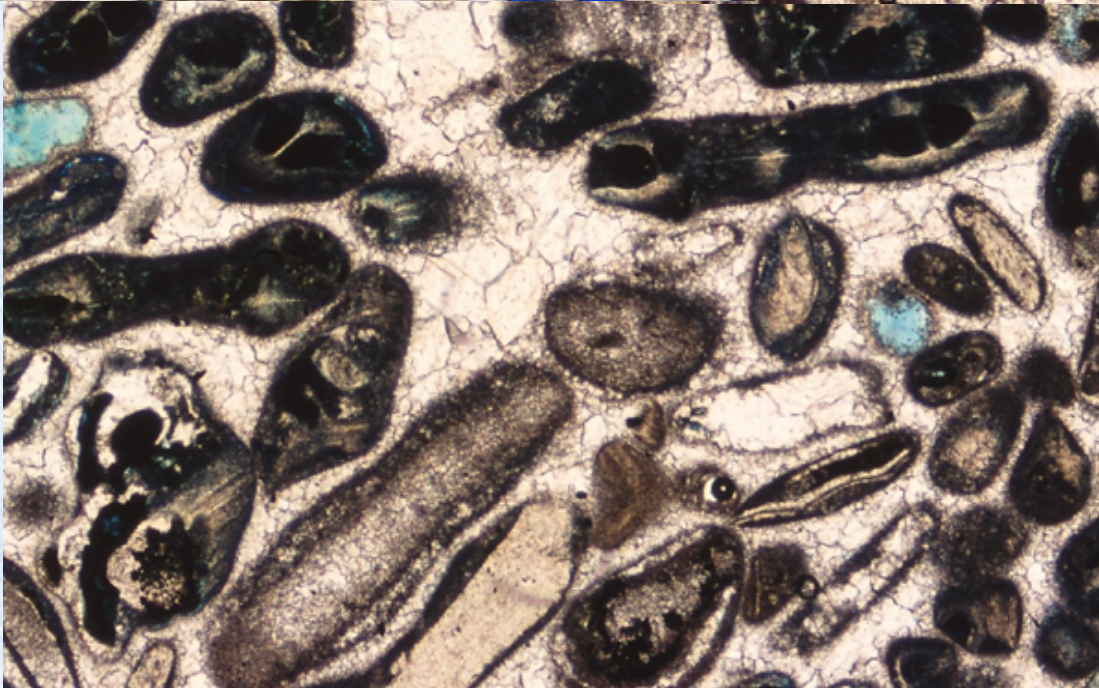
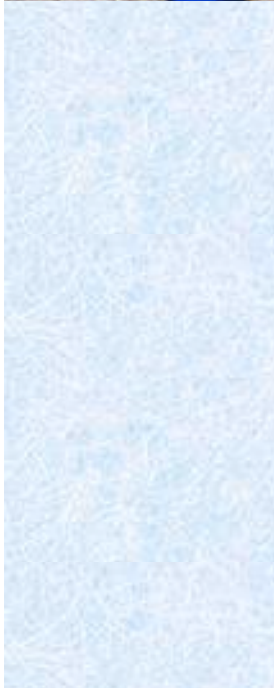
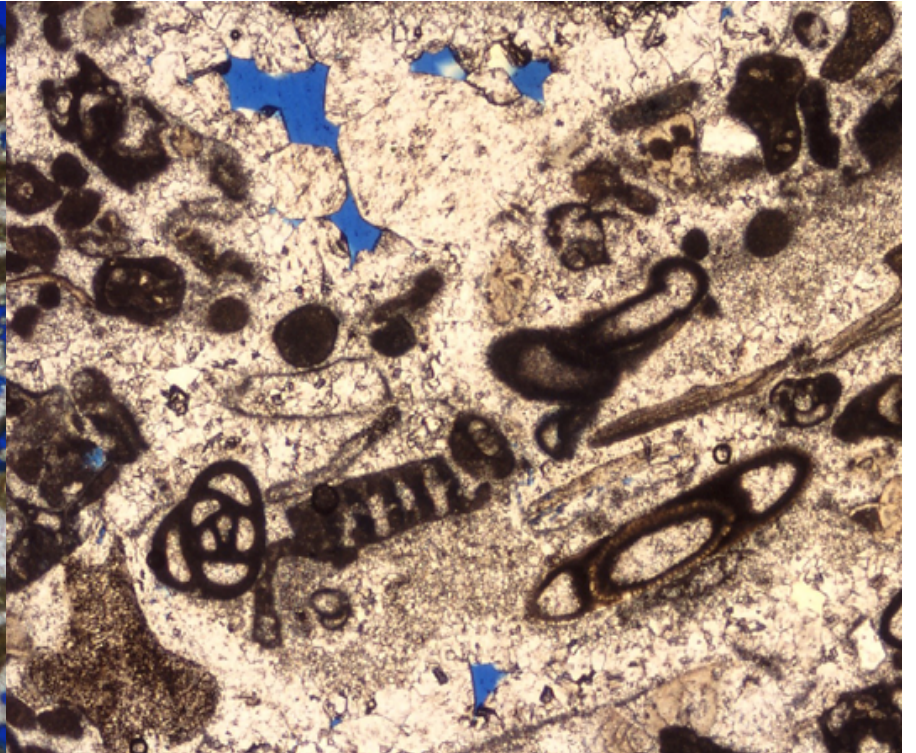
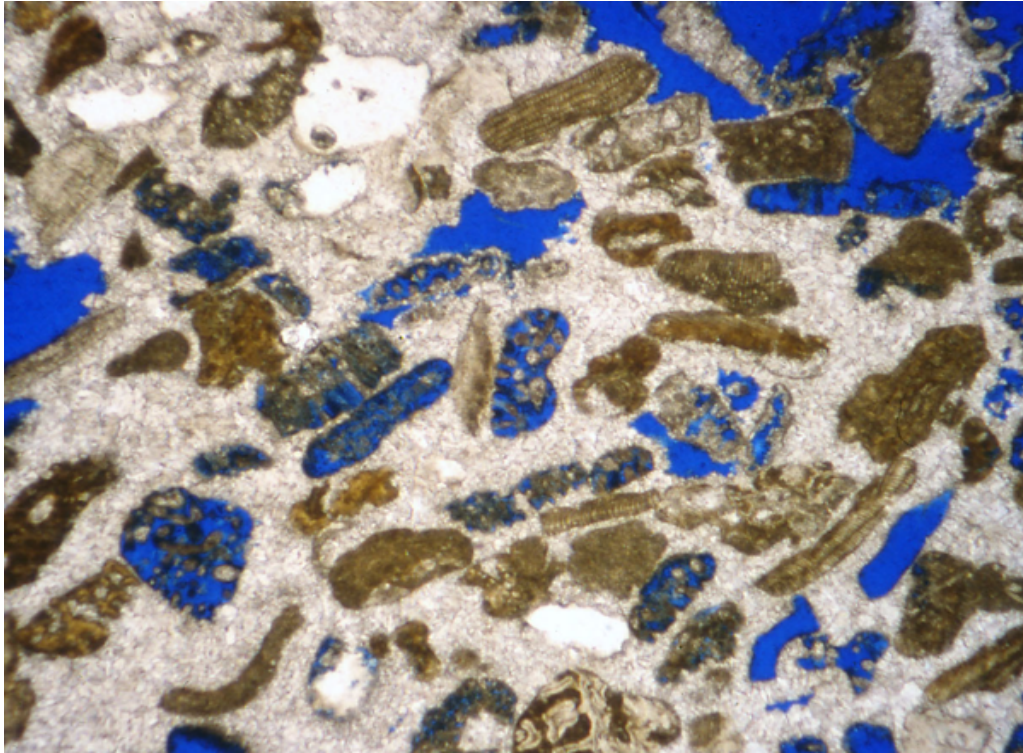
VADOSE-ZONE DIAGENESIS



- MINERALOGIC STABILIZATION
- CALCITE CEMENTATION
 - MENISCUS
 - GRAVITATIONAL
- LEACHING
- SUBARIEL CRUSTS & CALICHE



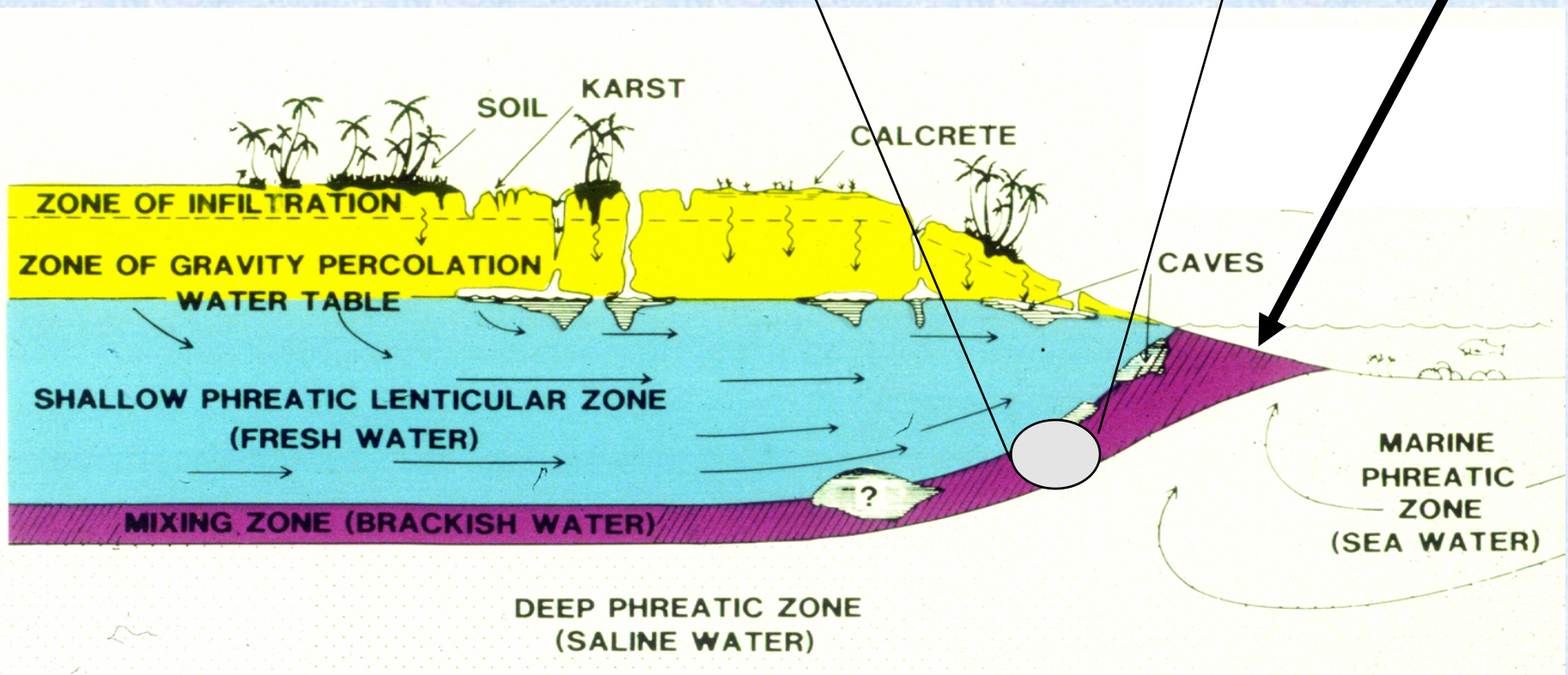
James & Choquette 1984



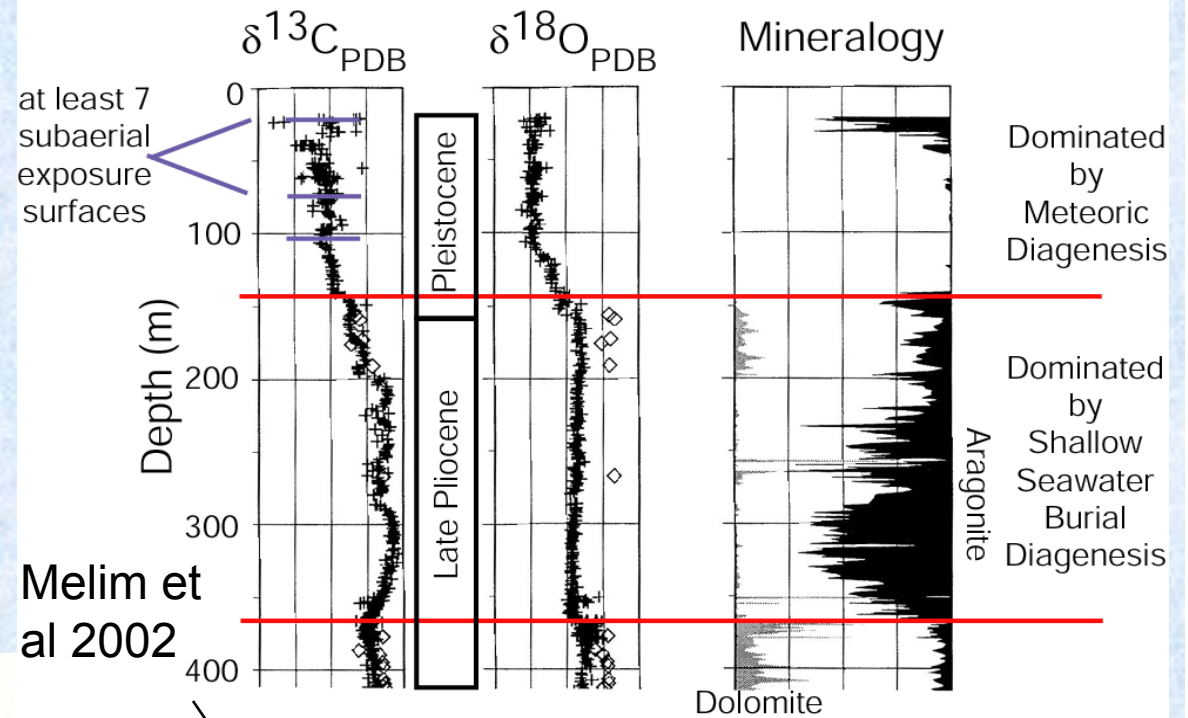
Mixing Zone Diagenesis

Dominated by dissolution.

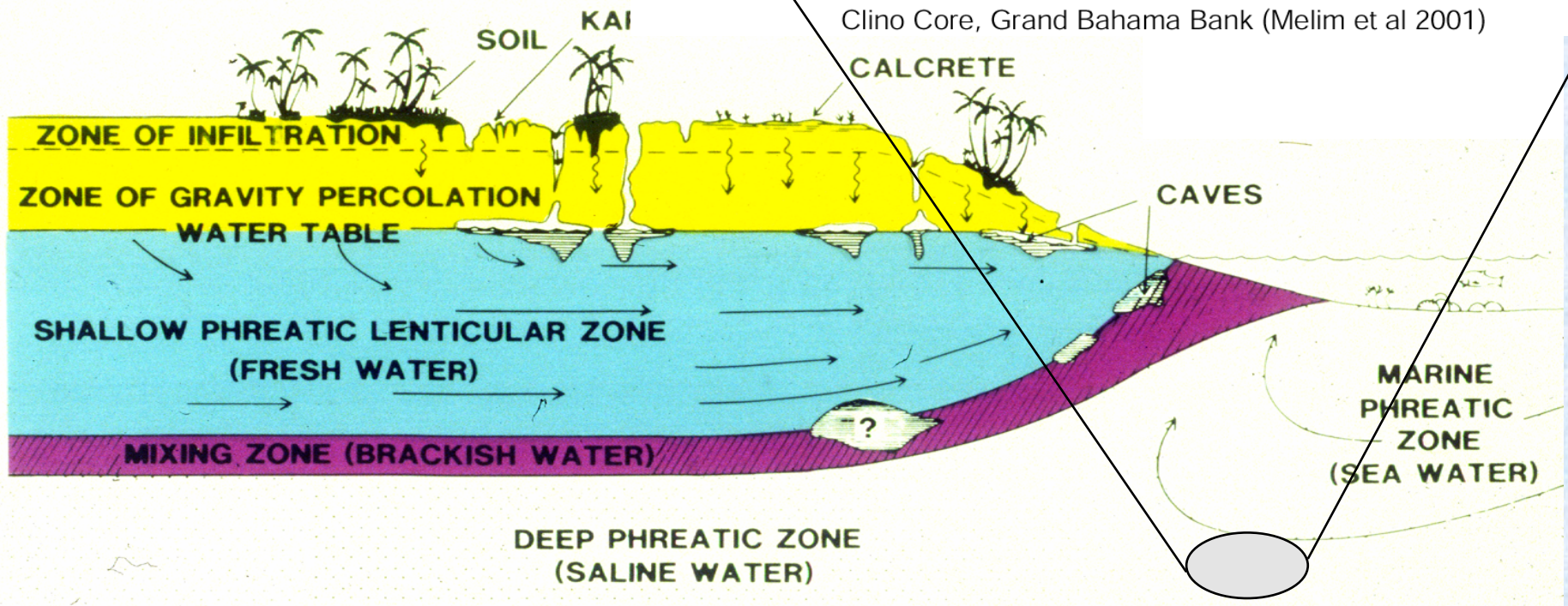
But cementation also noted in hydrochemistry of some modern upper mixing zones (20-30% SW) & implicated in fluid inclusion data from some ancient cements (0-90% SW).



“Shallow” Seawater Burial Diagenesis



Clino Core, Grand Bahama Bank (Melim et al 2001)



Seawater Burial Diagenesis

Unda 443.53 m

molds

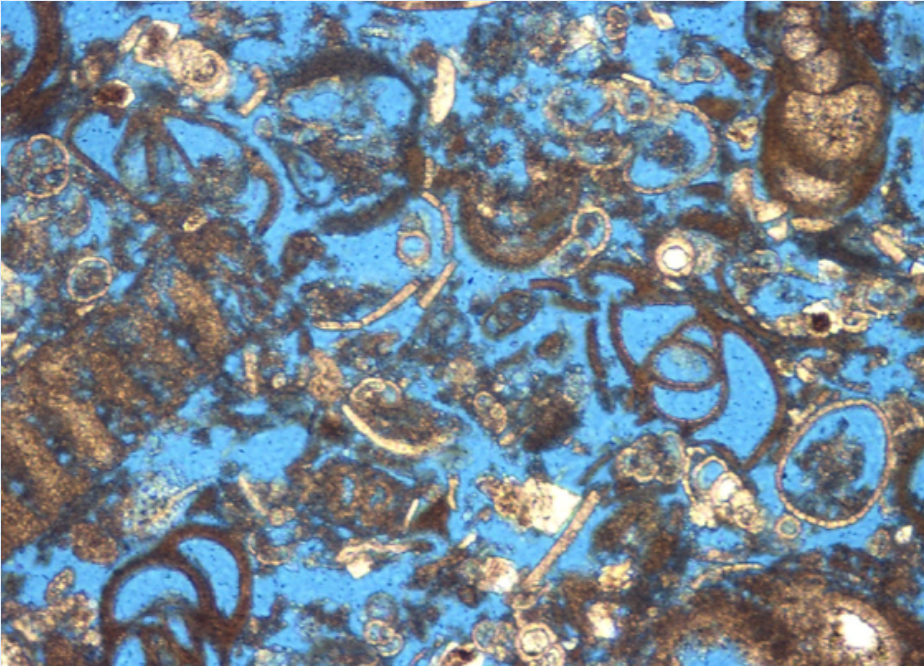
overgrowths

Clino 354.8 m

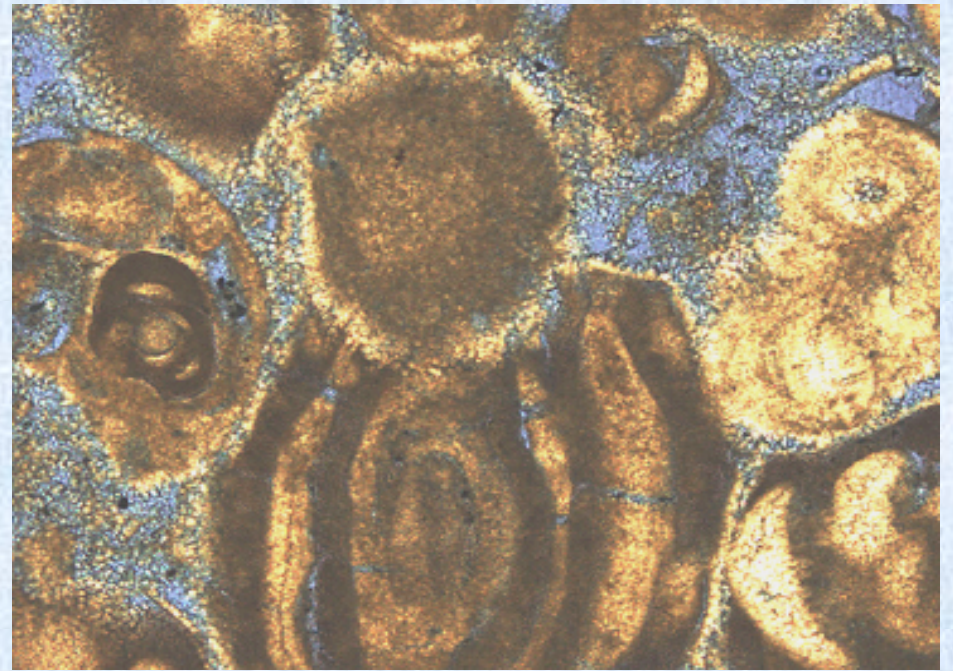
200 μm

Melim et al 2001

Seawater Burial Diagenesis



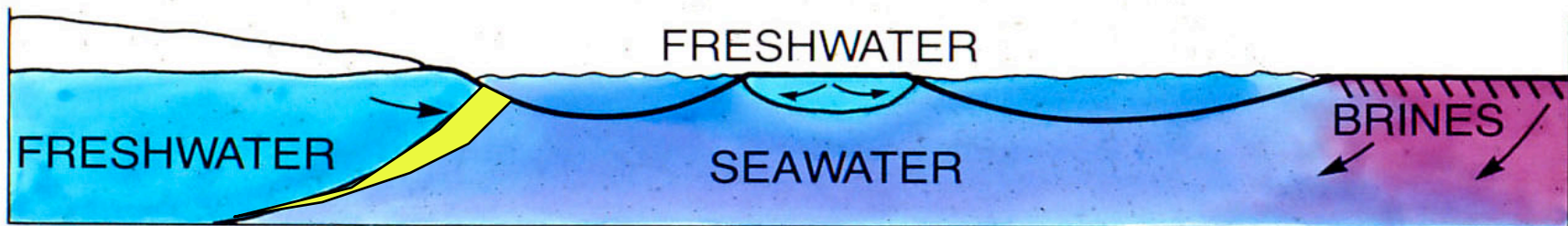
Eocene, Avon Park Fm, 140 m



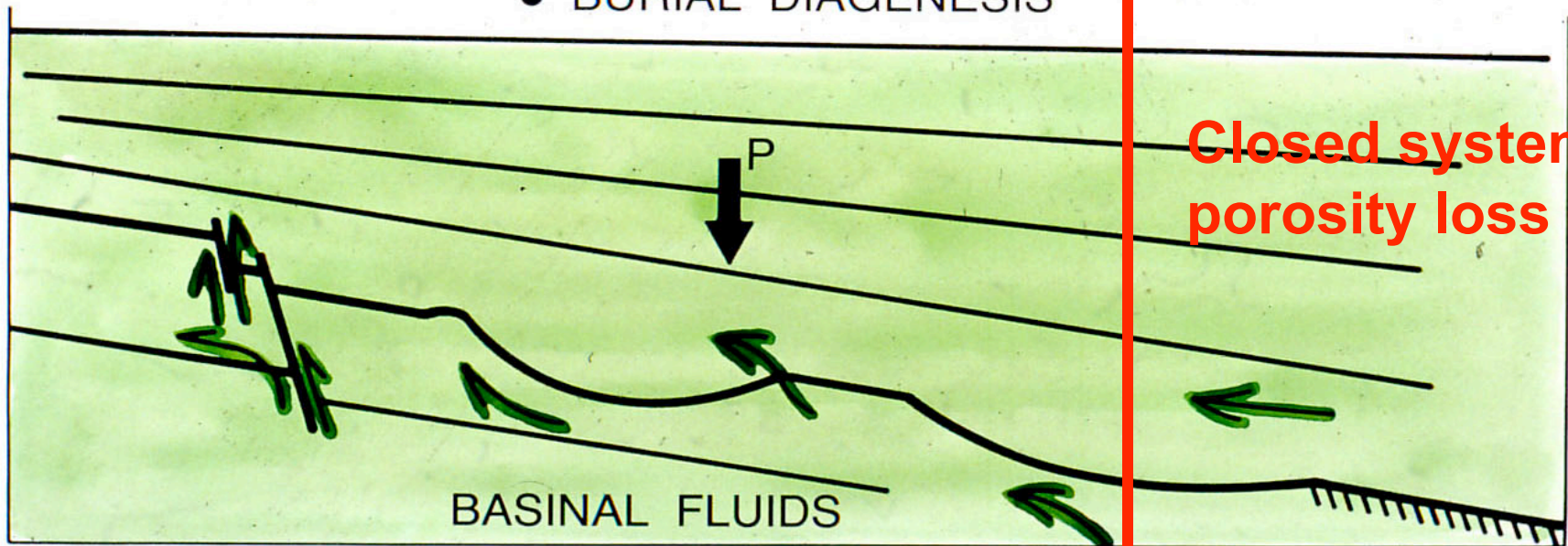
Eocene, Avon Park Fm, 390 m

Diagenesis historically viewed in terms of hydrochemical regimes that generate internally consistent process-response dynamics

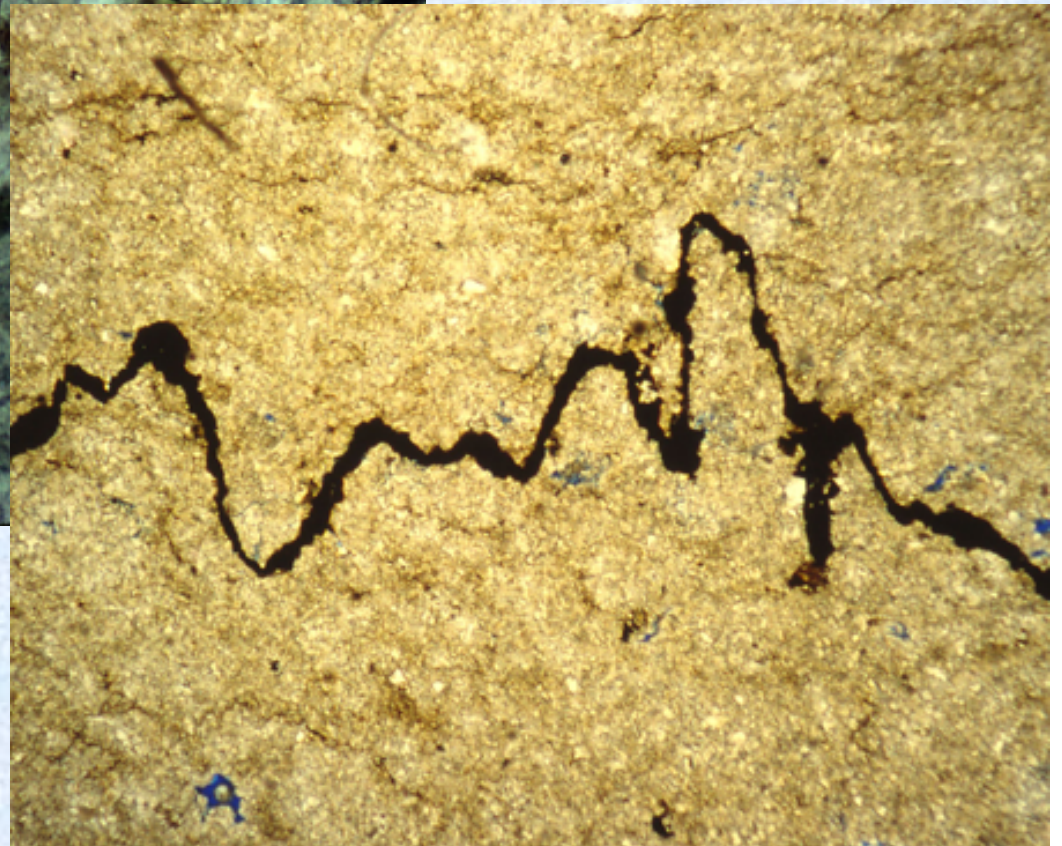
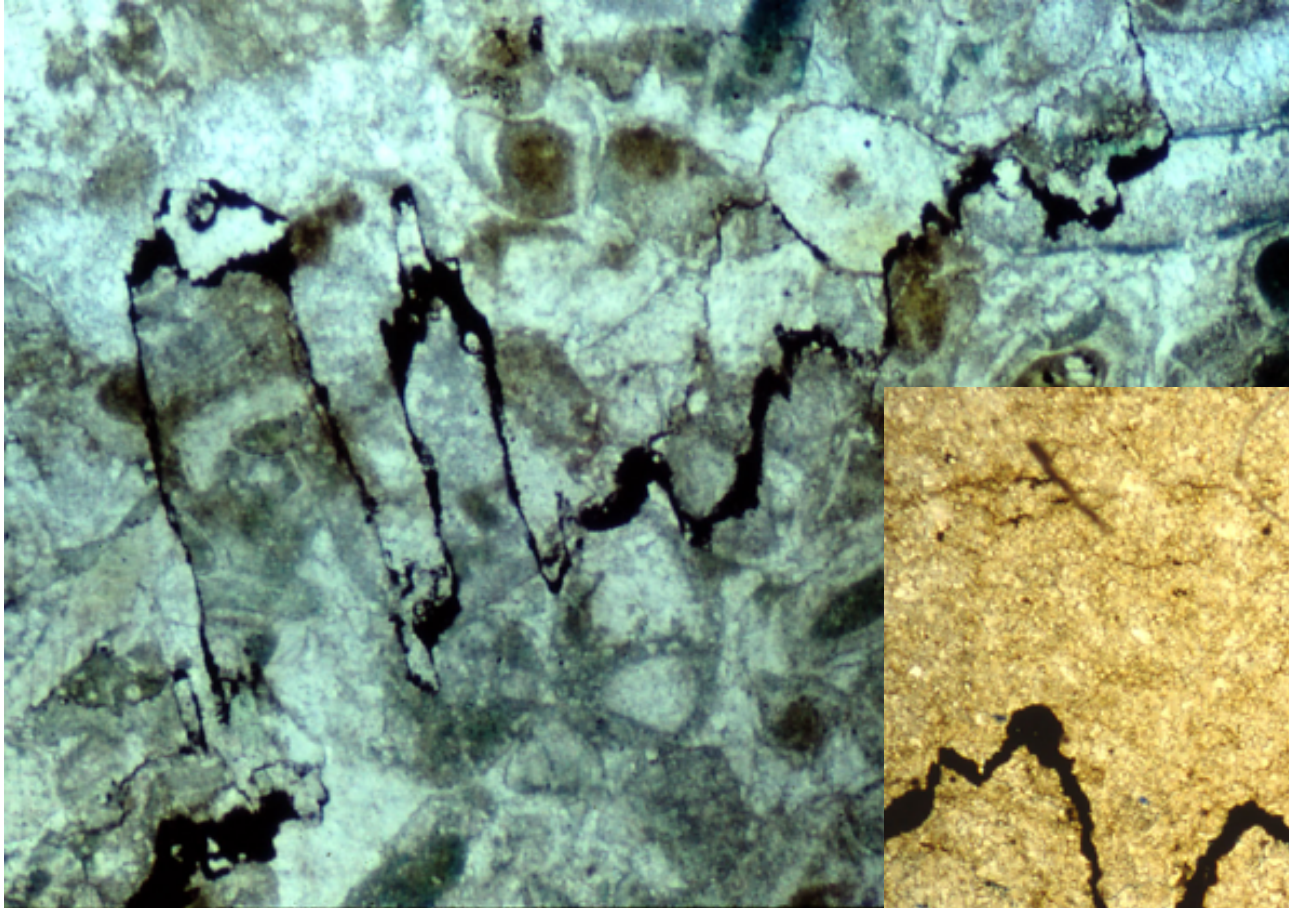
- SURFICIAL DIAGENESIS



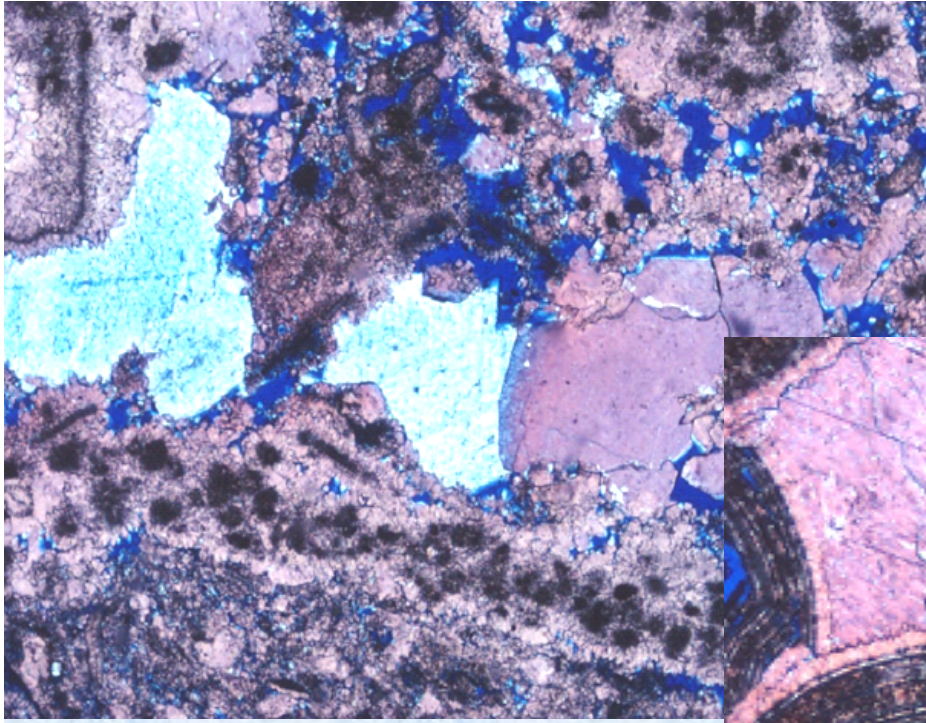
- BURIAL DIAGENESIS



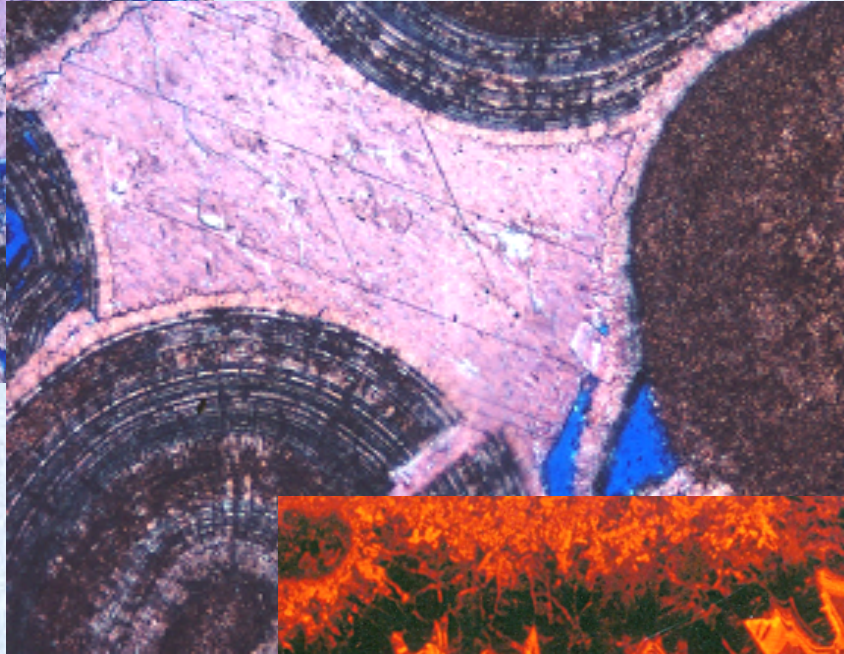
Pressure Solution



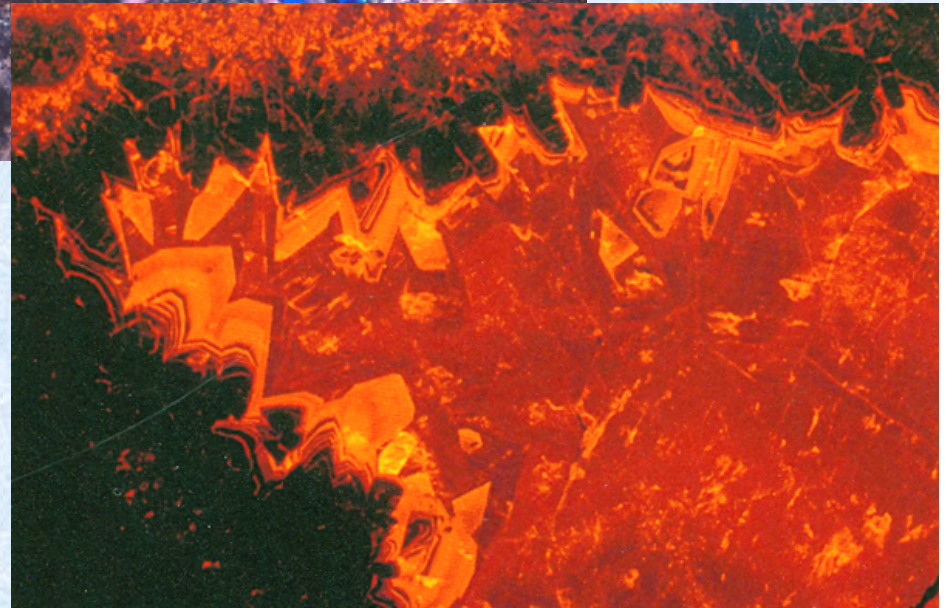
Deep Burial Cement



Pennsylvanian, Paradox Basin

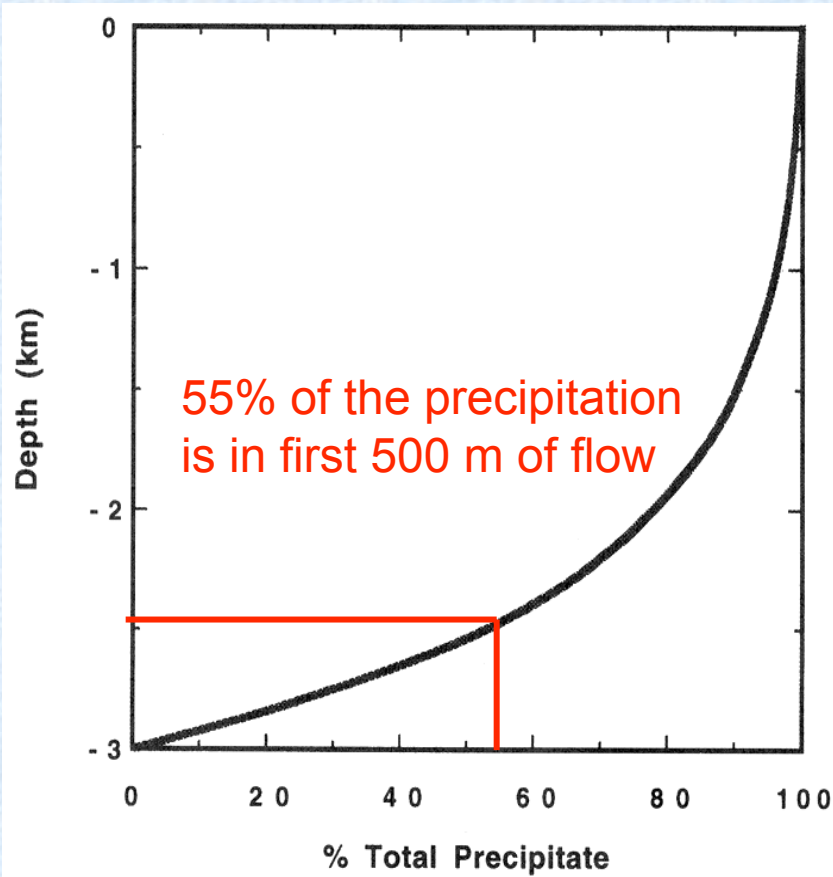


Jurassic, Smackover Fm.

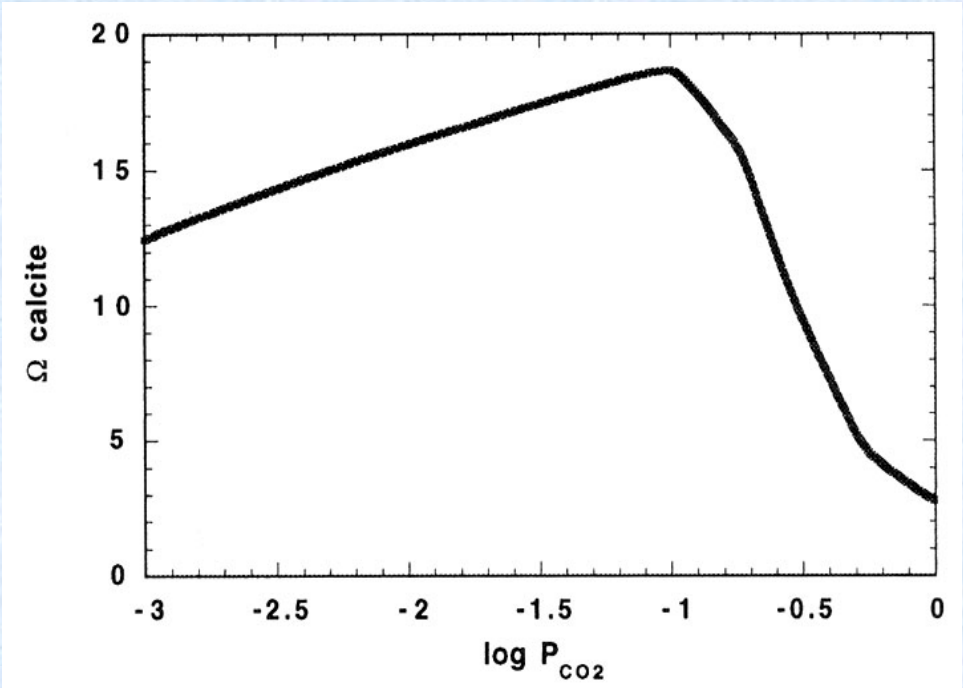


L. Cambrian, Labrador

**Where calcite precipitates as
200,000 mgL⁻¹ brine migrates
from 100°C & 300 bars to STP**



**Supersaturation of a 200,000
mgL⁻¹ brine at 100°C if
pressure drops from 600 bars
to 300 bars**

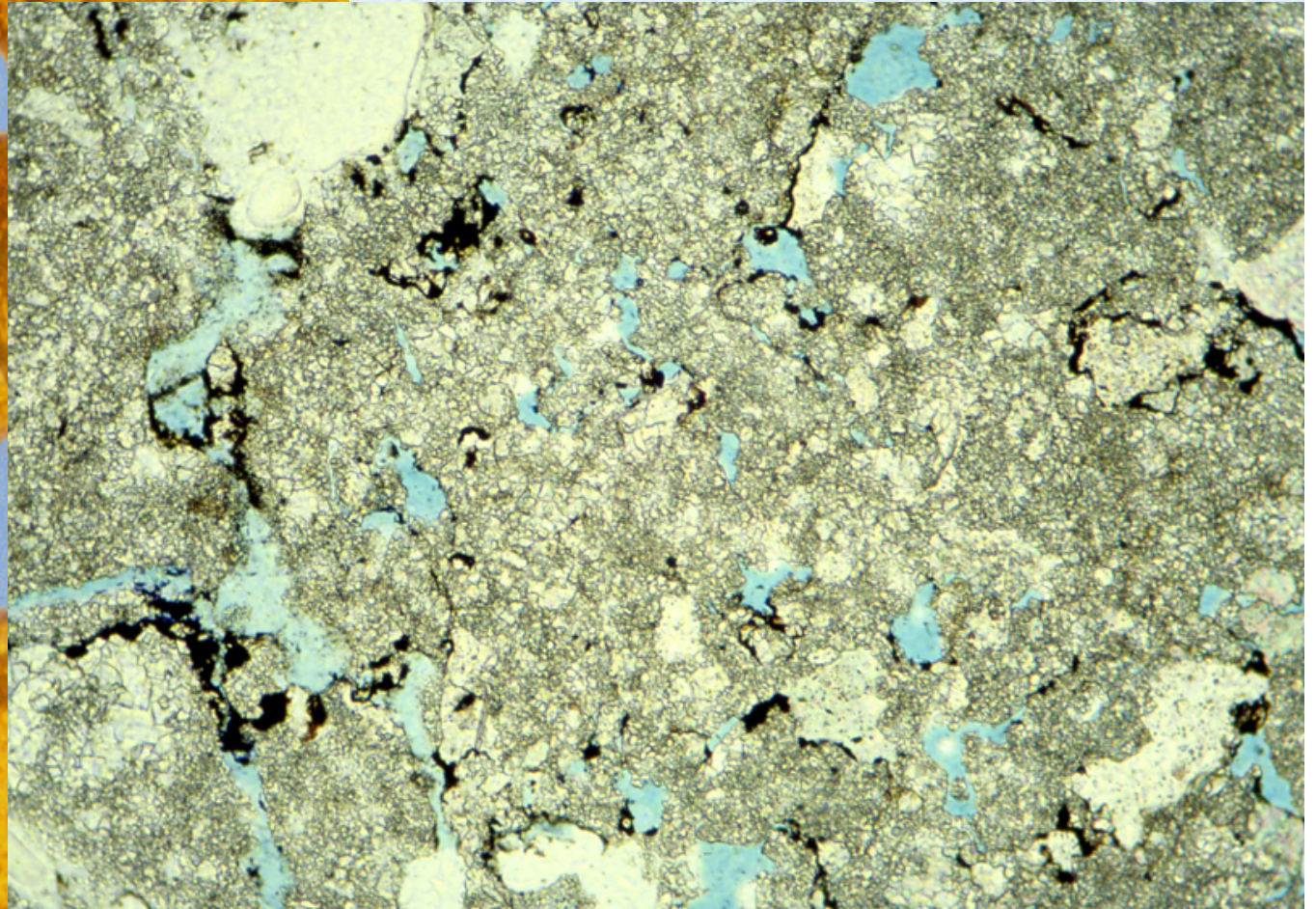
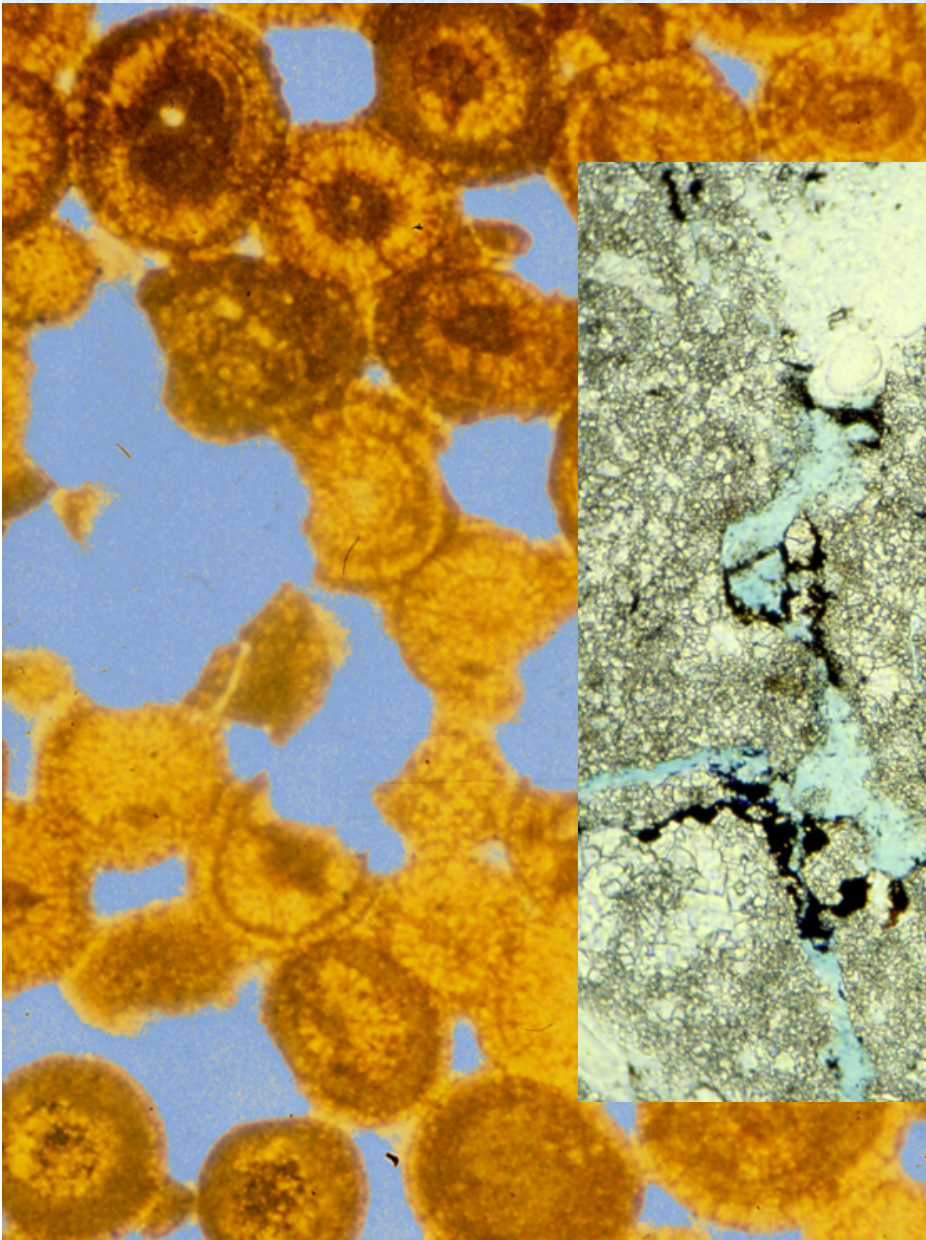


Implication

- deep cementation has strong linkage to abrupt vertical fluid flow**
- also to fluid mixing (convection cells?)**

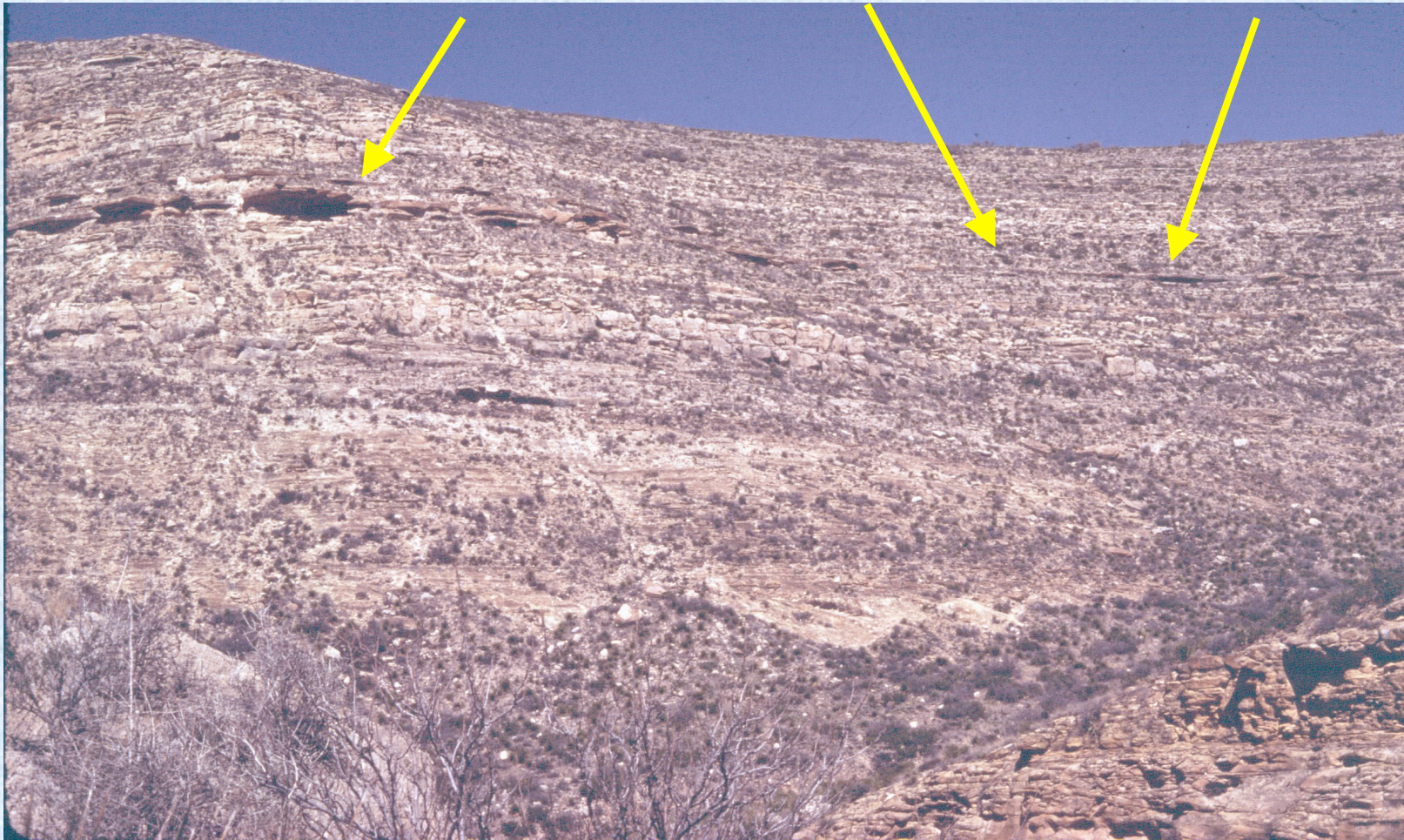
Morse et al 1997

Dissolution

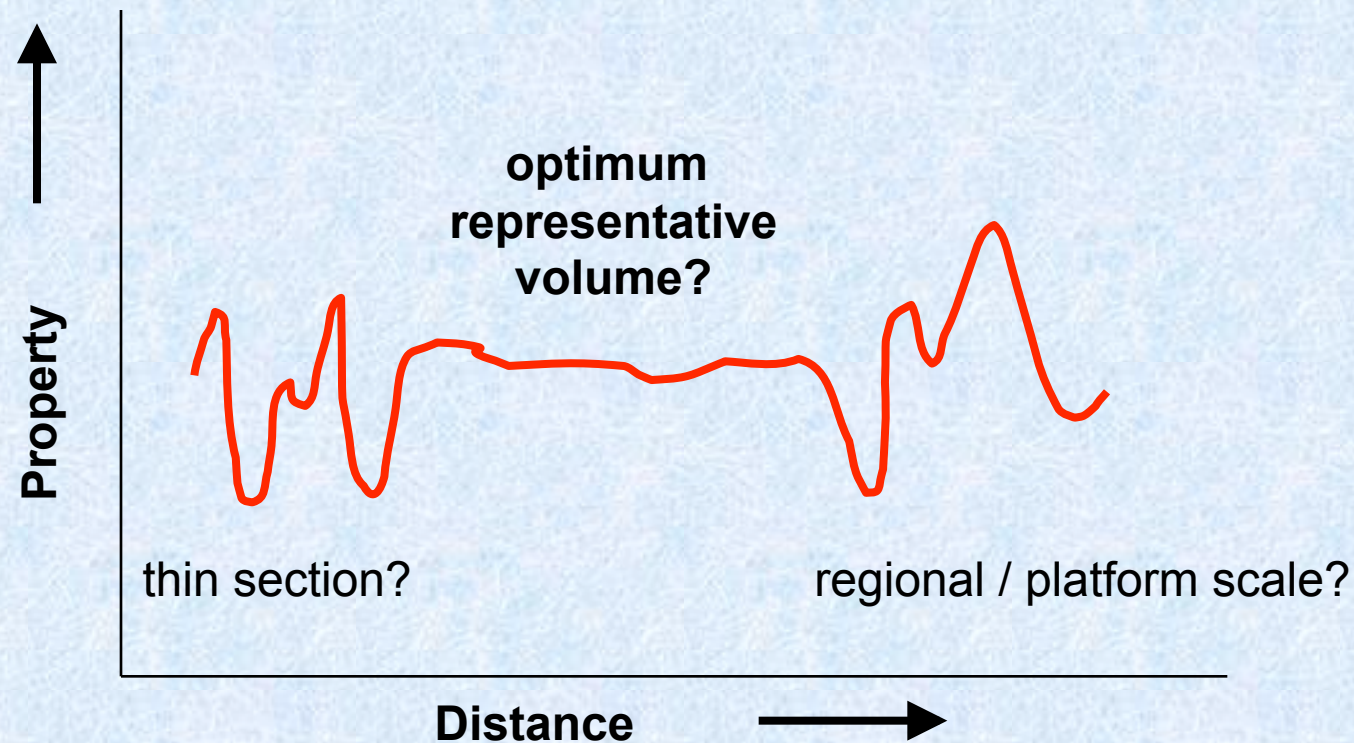


Responsible fluids / reactions still unclear.

*Still unresolved for all these process –
seafloor to deepest burial --- What
happened “there” versus “there”, or “there”?*



Is there an optimum scale at which to approach diagenetic modeling?

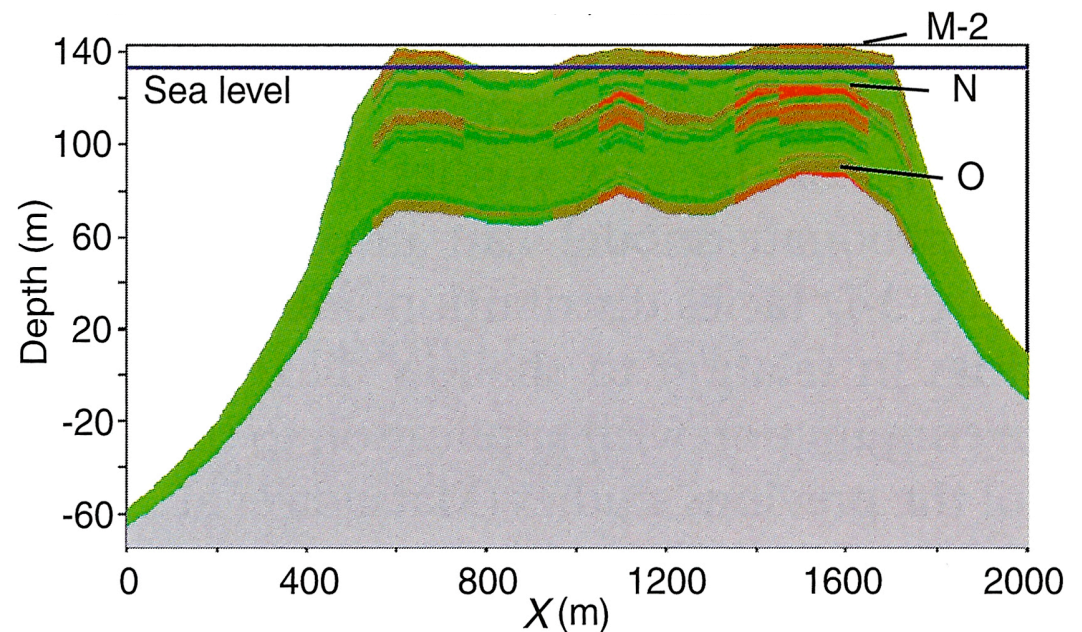
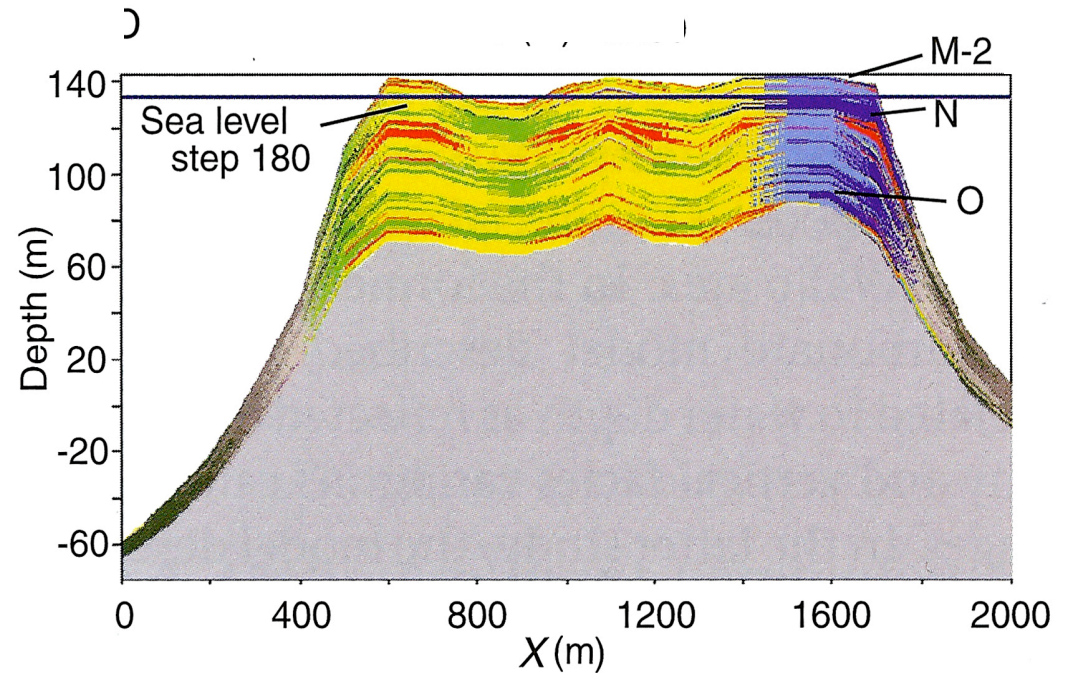


And if so, is that the scale of interest?

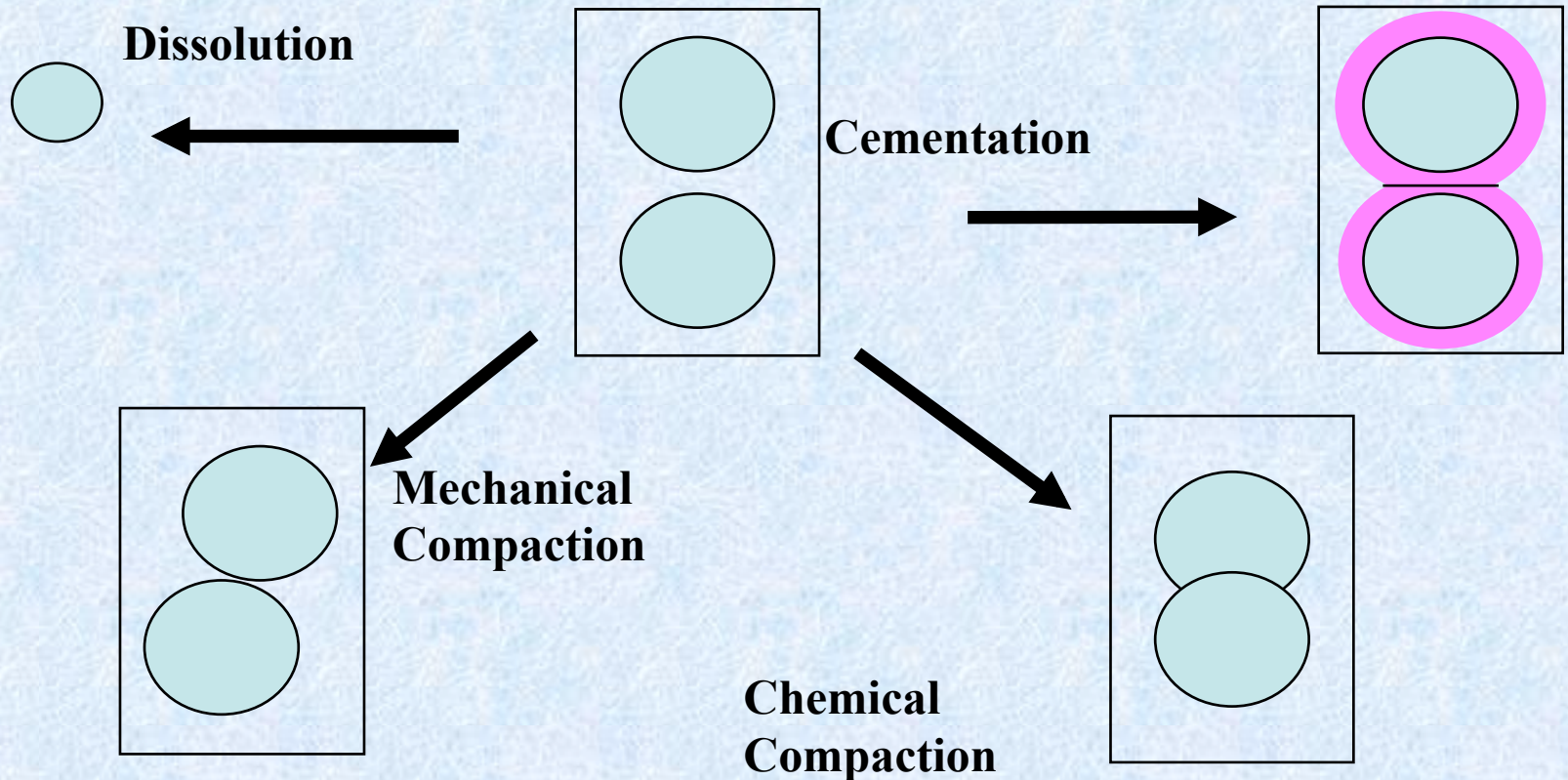
1 Current Modeling Approach

Coupled sediment production and sea level models to predict facies mosaics and exposure surfaces

plus residence time in meteoric hydrozones, and thus diagenetic variability by inference



Can we track fabric dependent processes?



And if so, is that the scale of interest?

Summary Points

We know what diagenetic processes occur in various diagenetic settings

The conditions that drive most of those processes

The features that allow us to recognize those processes, conditions, and/or settings

Diagenetic systems are complex & non-linear with many types of interactions and feedbacks

But there is uncertainty regarding presence/absence of

- self-organization and/or scale invariance
- thresholds in feedback relations

Summary Points

Diagenetic alterations that affect sediment dynamics include
seafloor cementation & dissolution
denudation with exposure

There are many post-depositional alterations that can and do occur through time and stratigraphic space.

Critical questions to their modeling will be
scales of analysis (pore to regional)
linkage of reactions and transport processes
through evolving temperature, pressure, &
petrophysical regimes dependence on fabric relations

And, with the exception of the “regional” and finest spatial scales, we don’t really know lateral spatial patterns by which to evaluate model outcomes