

# Filtering the hydrograph through sediment transport & channel geometry

ST. ANTHONY  
FALLS LABORATORY  
UNIVERSITY OF MINNESOTA  
Driven to Discover™



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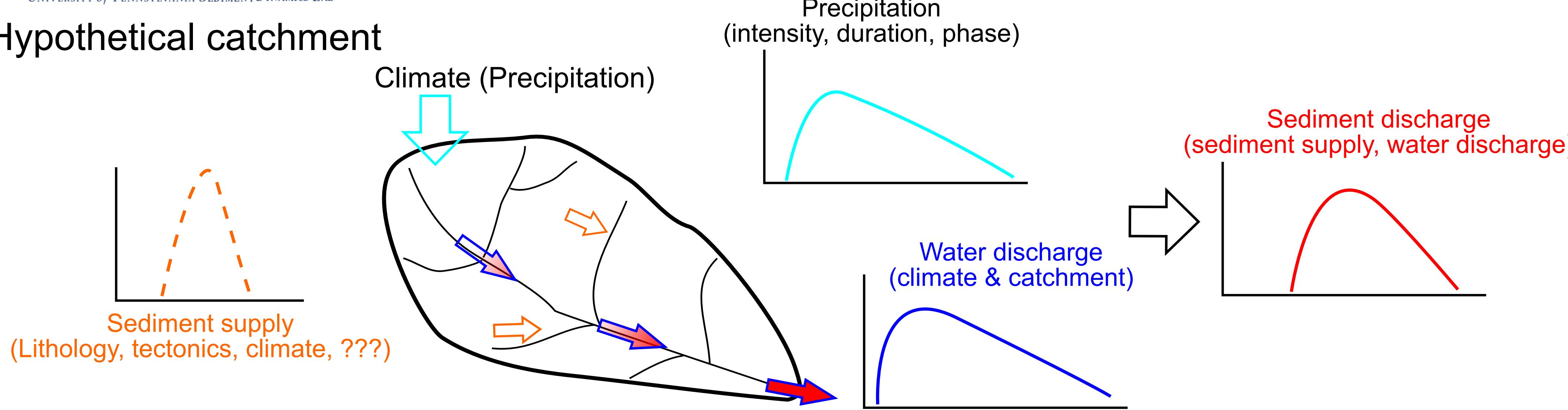
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<sup>4</sup>Earth and Environmental Science, Univ. of Pennsylvania.

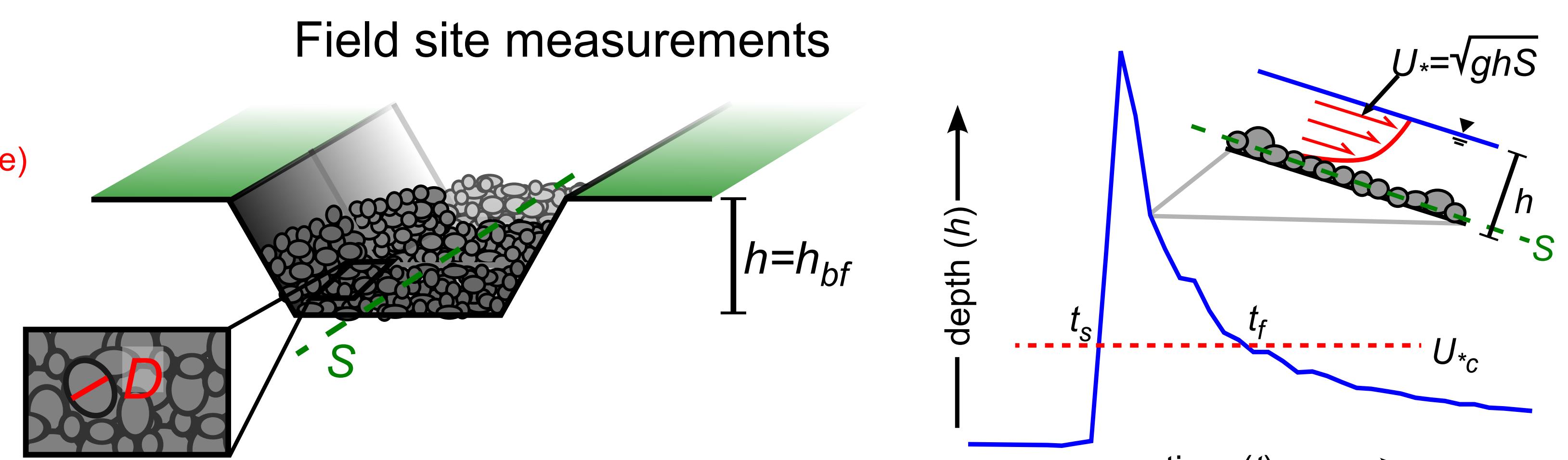
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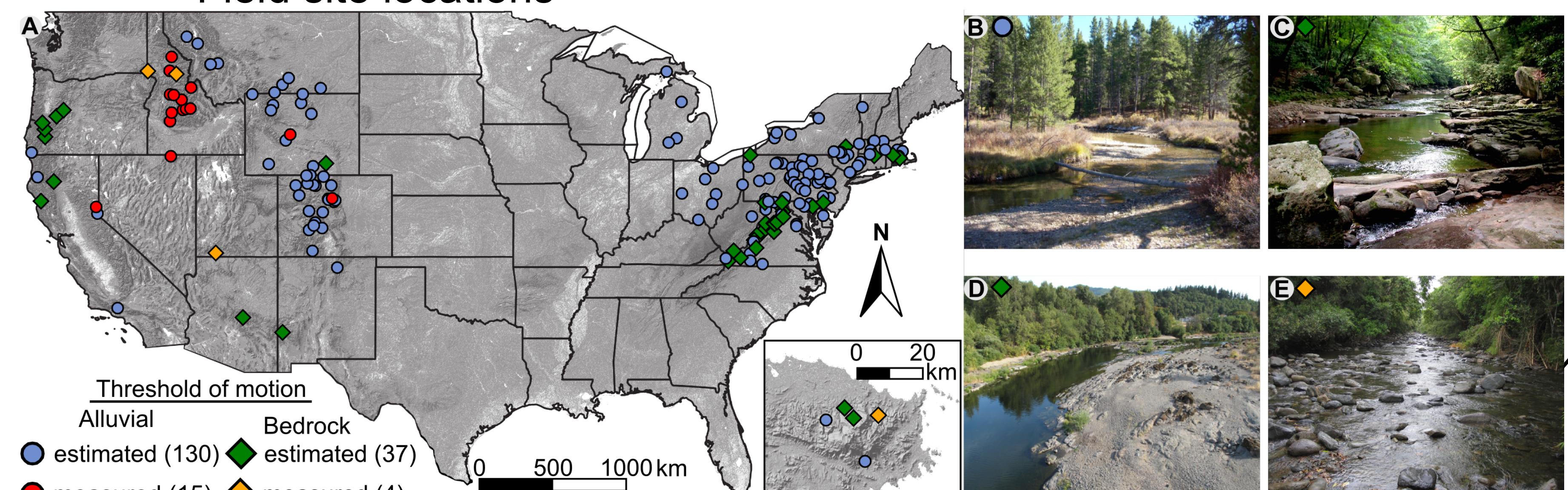
Hypothetical catchment



Field site measurements

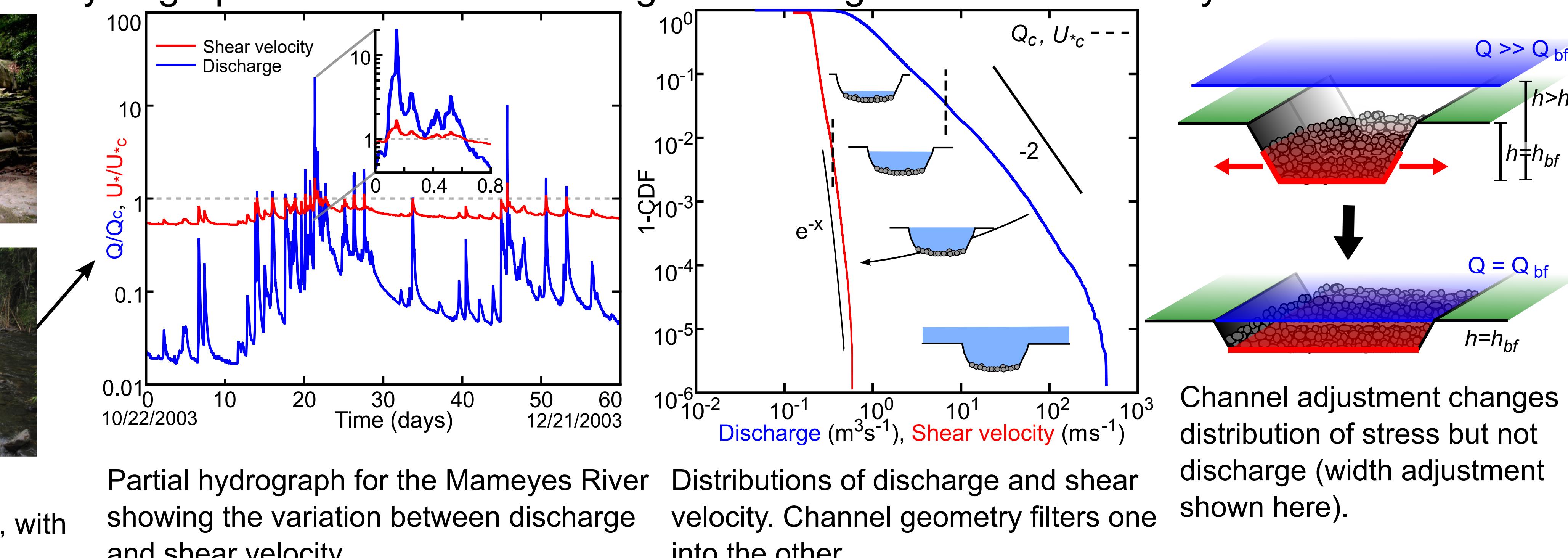


Field site locations

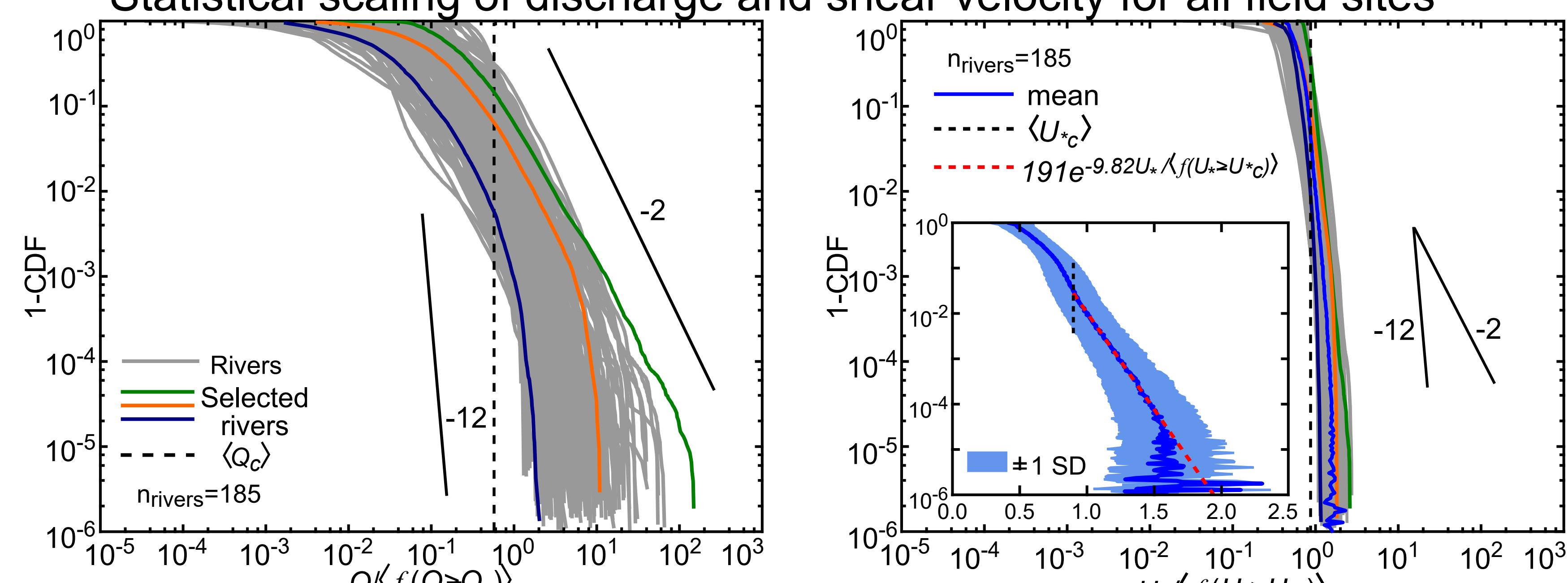


Locations of 186 USGS stream gages analyzed in this study. Field sites span a wide variety of climatic, lithologic, and tectonic regimes. Each field site has 10 years of 15 minute instantaneous discharge data, with channel slope, grain size, and channel geometry measured in the field.

Hydrographs and statistical scaling of Discharge and Shear velocity

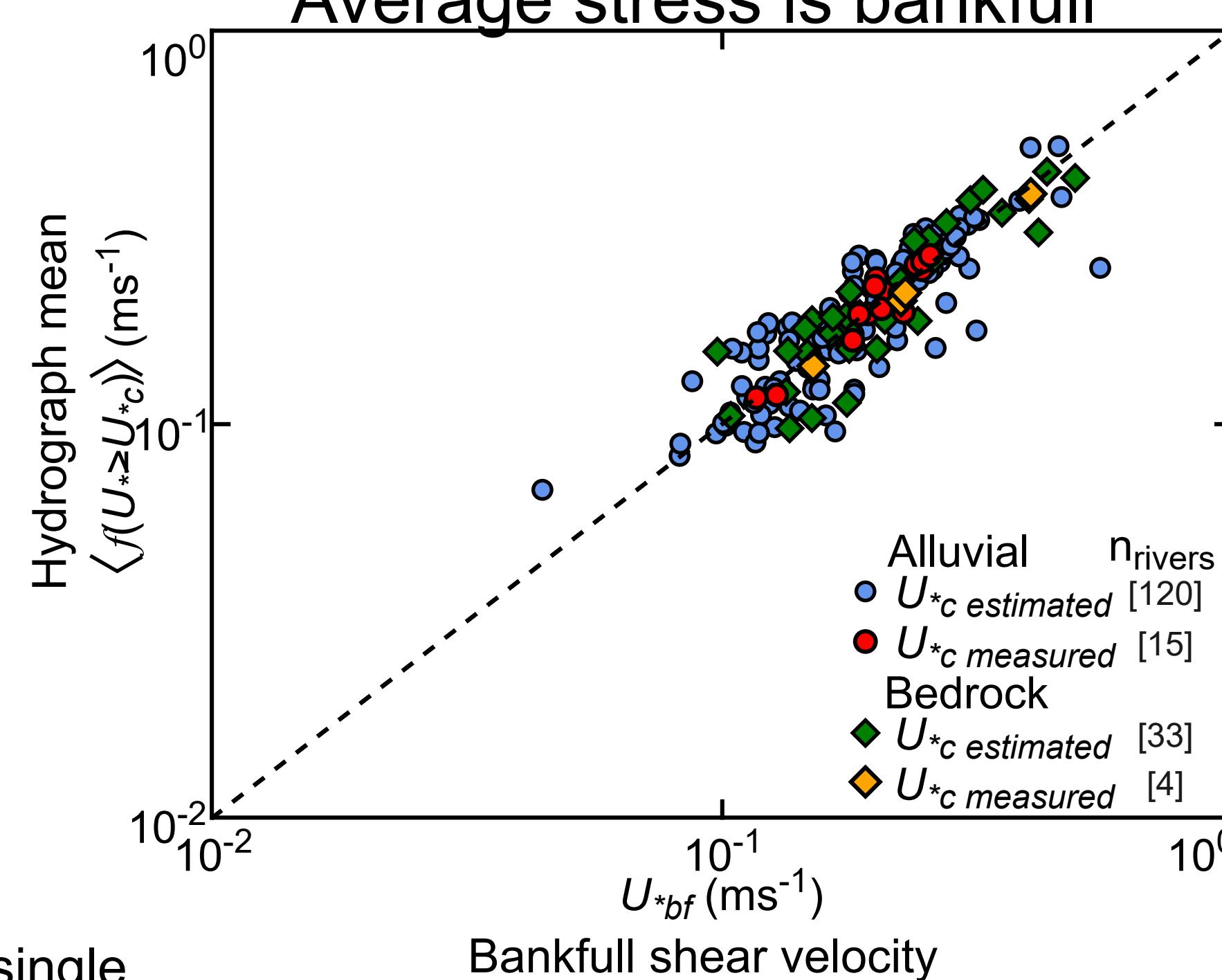


Statistical scaling of discharge and shear velocity for all field sites

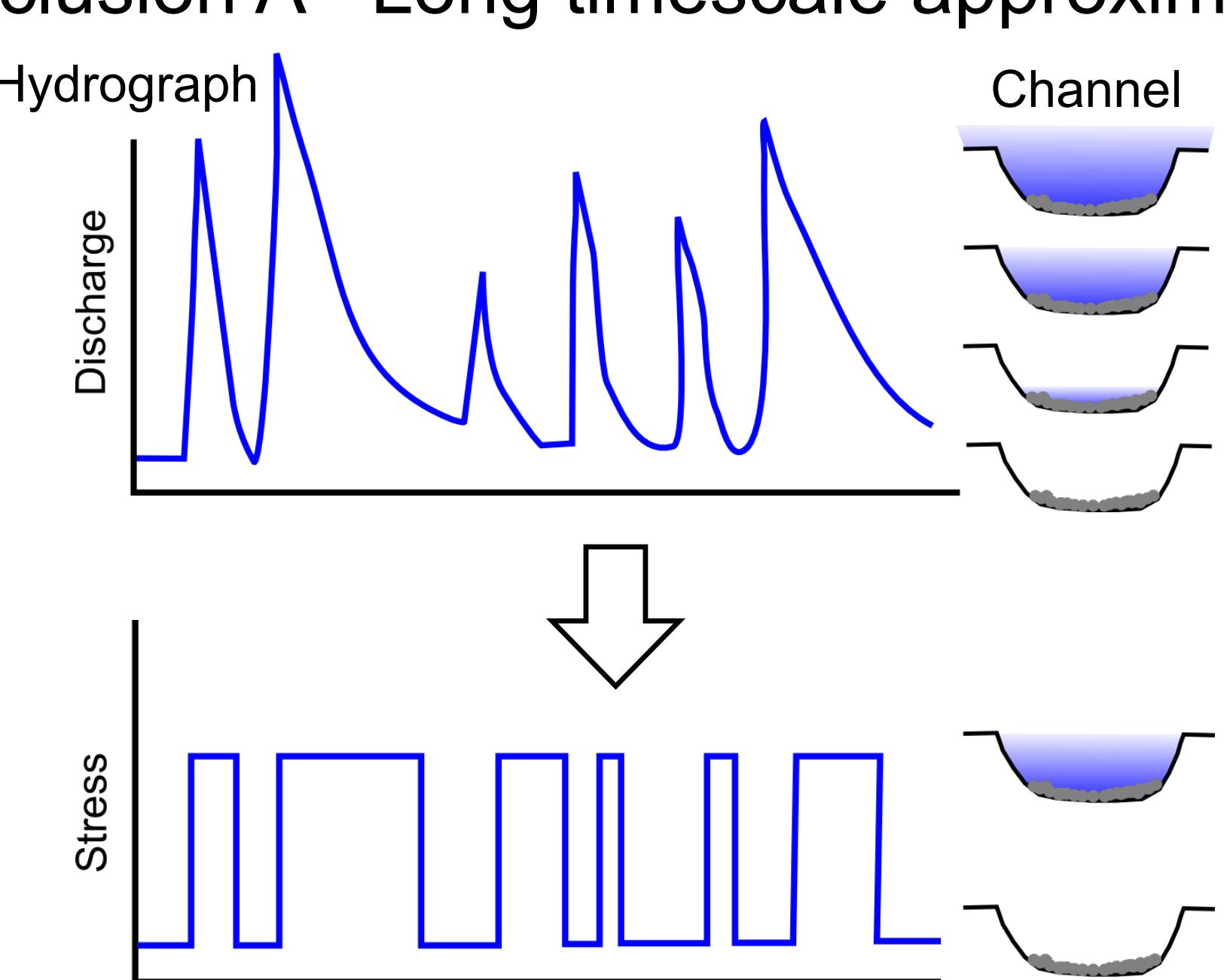


Distributions of discharge and shear velocity for all fields sites. Shear velocity distributions collapse to a single functional form when scaled by the mean shear velocity above the threshold of motion, while discharge does not.

Average stress is bankfull

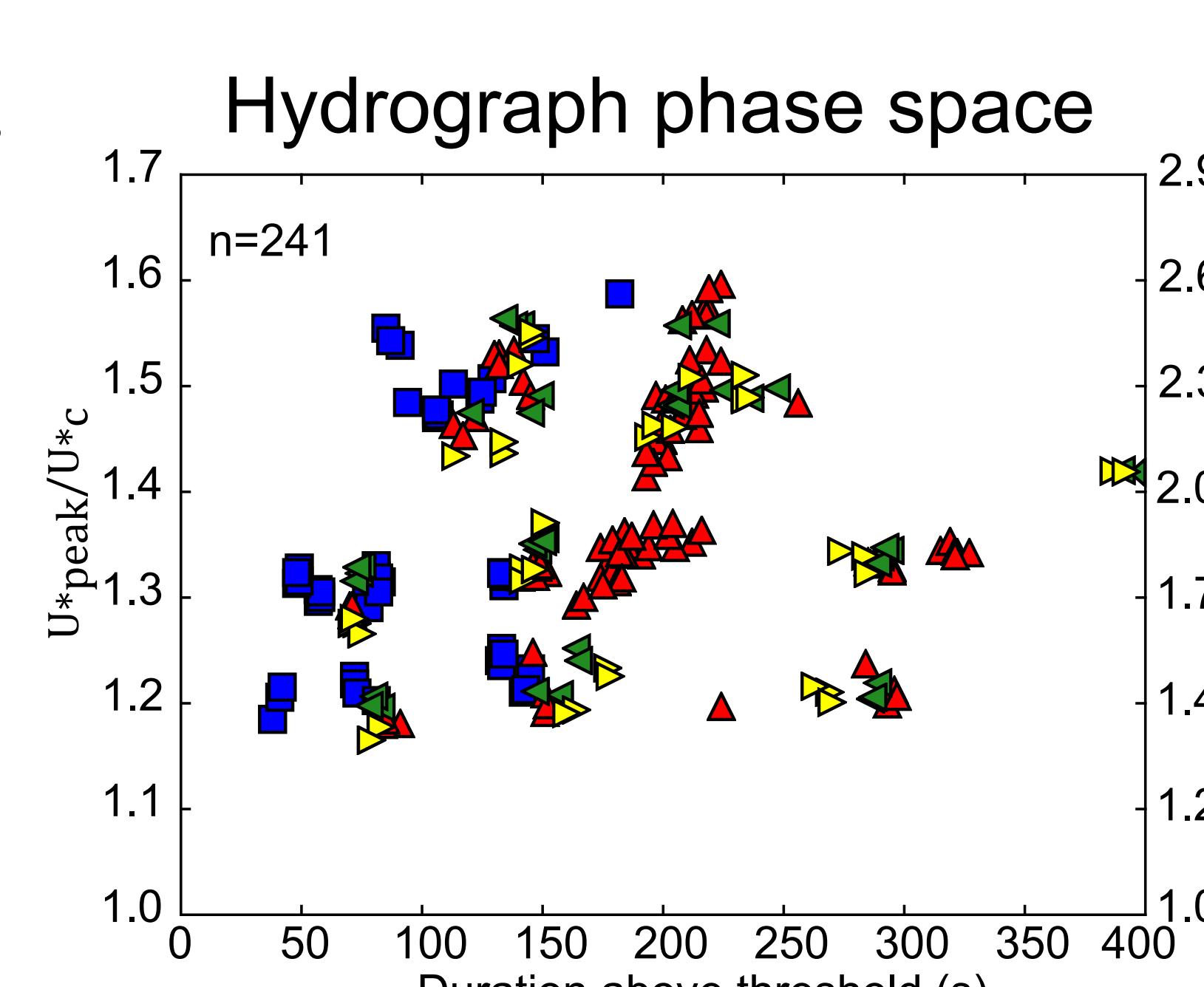
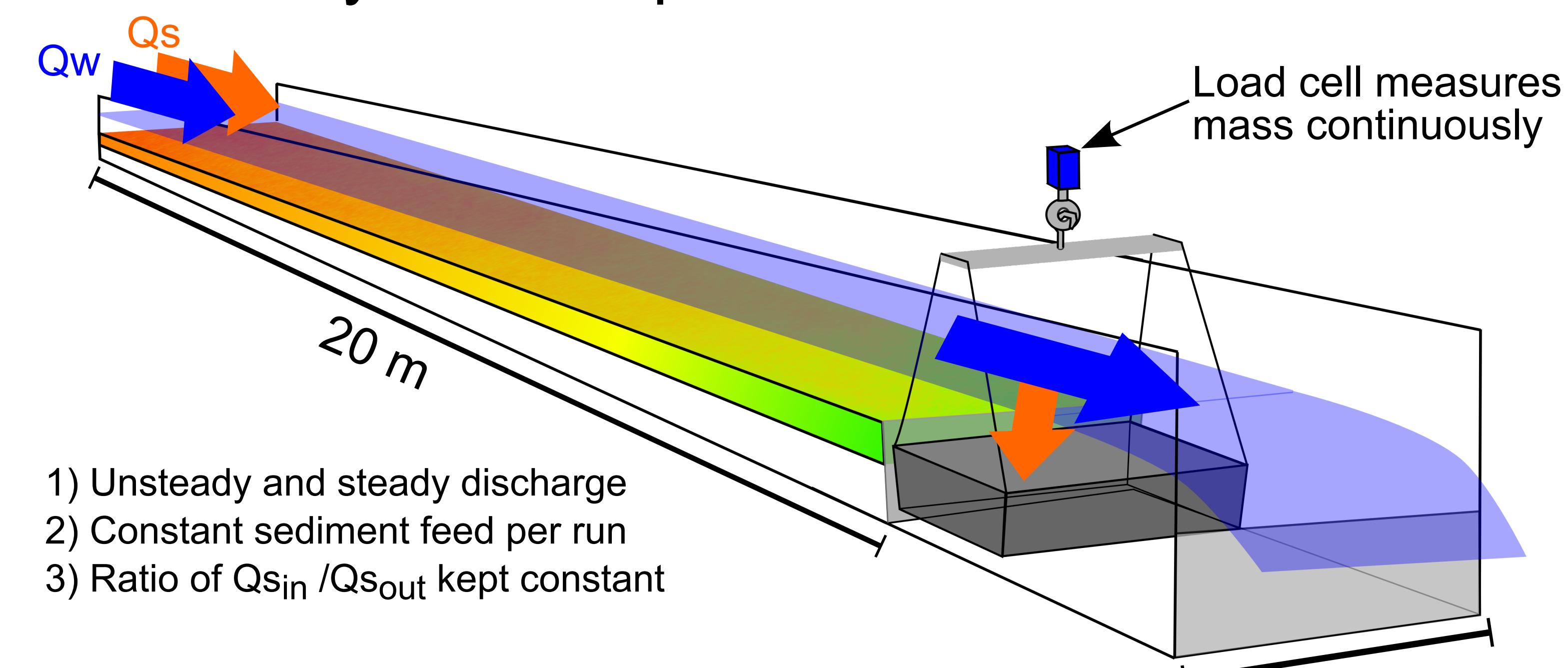


Conclusion A - Long timescale approximation

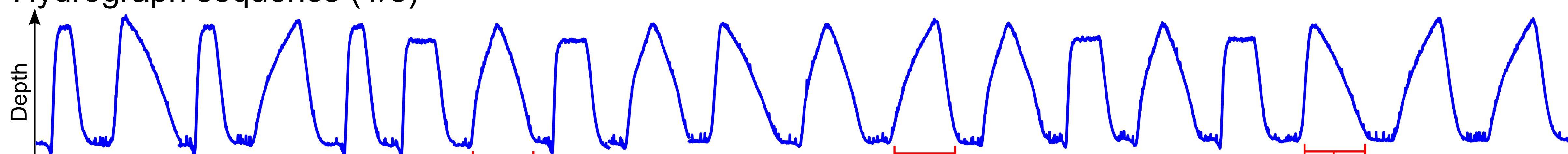


Reduce variable hydrograph to a fixed-magnitude steady flood event with an intermittency factor.

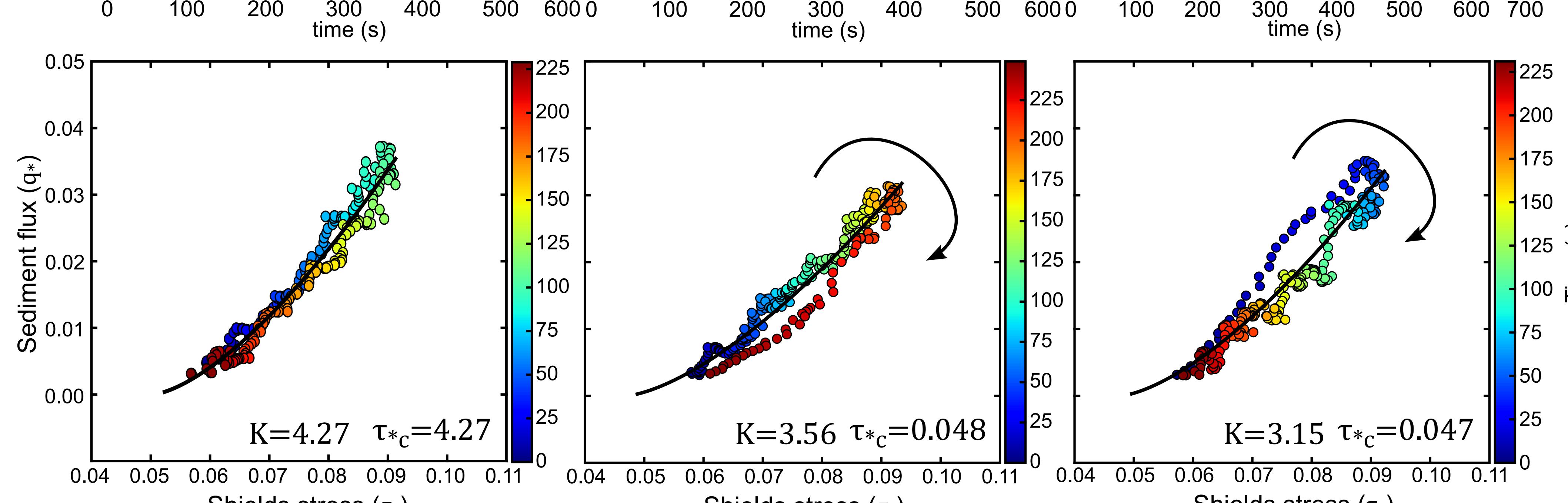
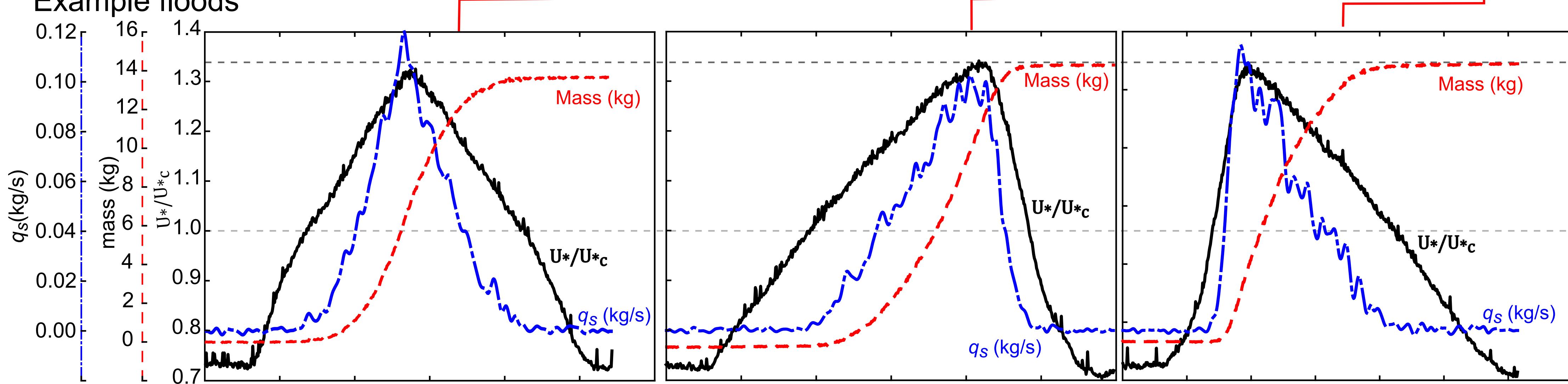
Laboratory flume experiments on flow transience and bed load sediment transport



Hydrograph sequence (1/8)

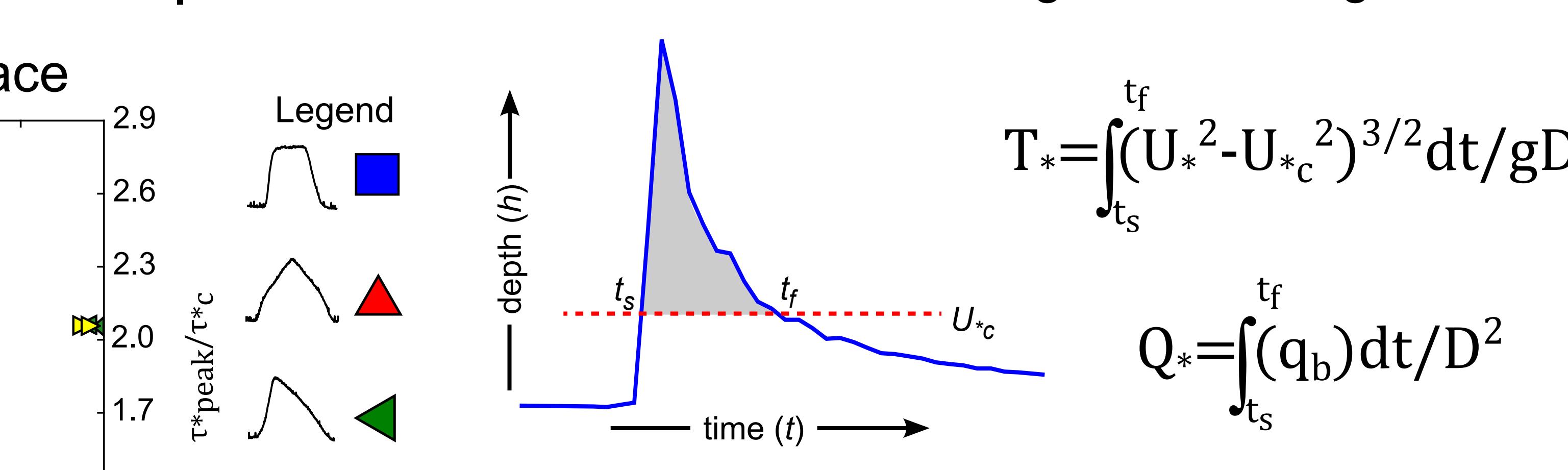
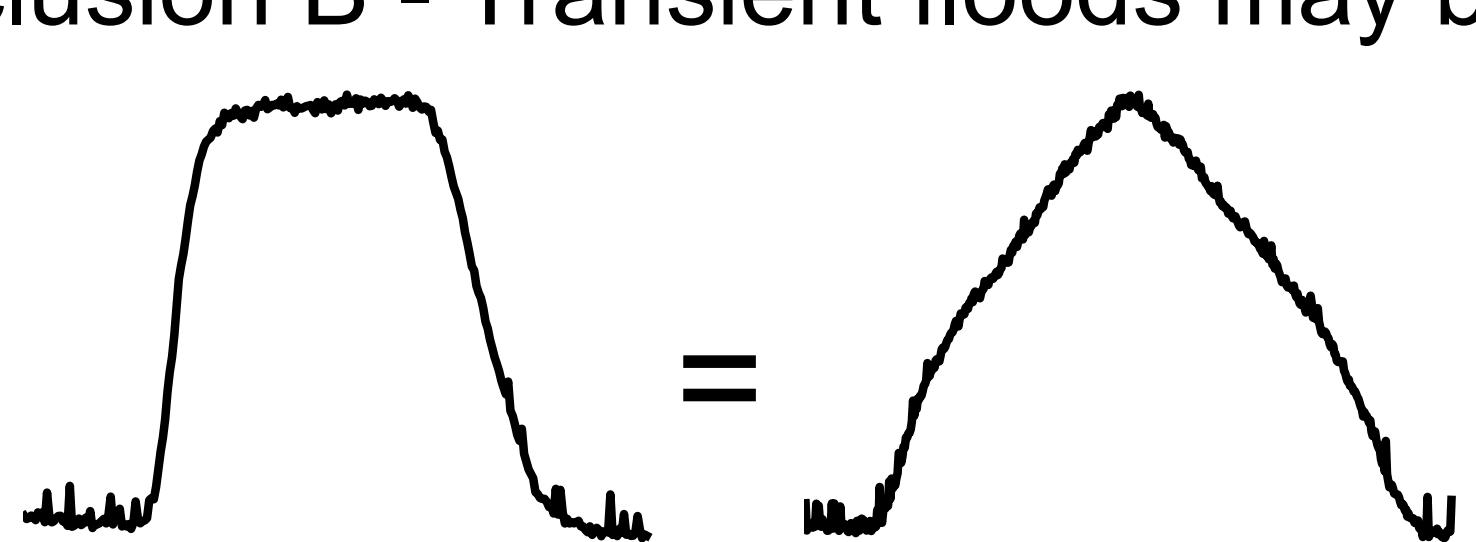


Example floods

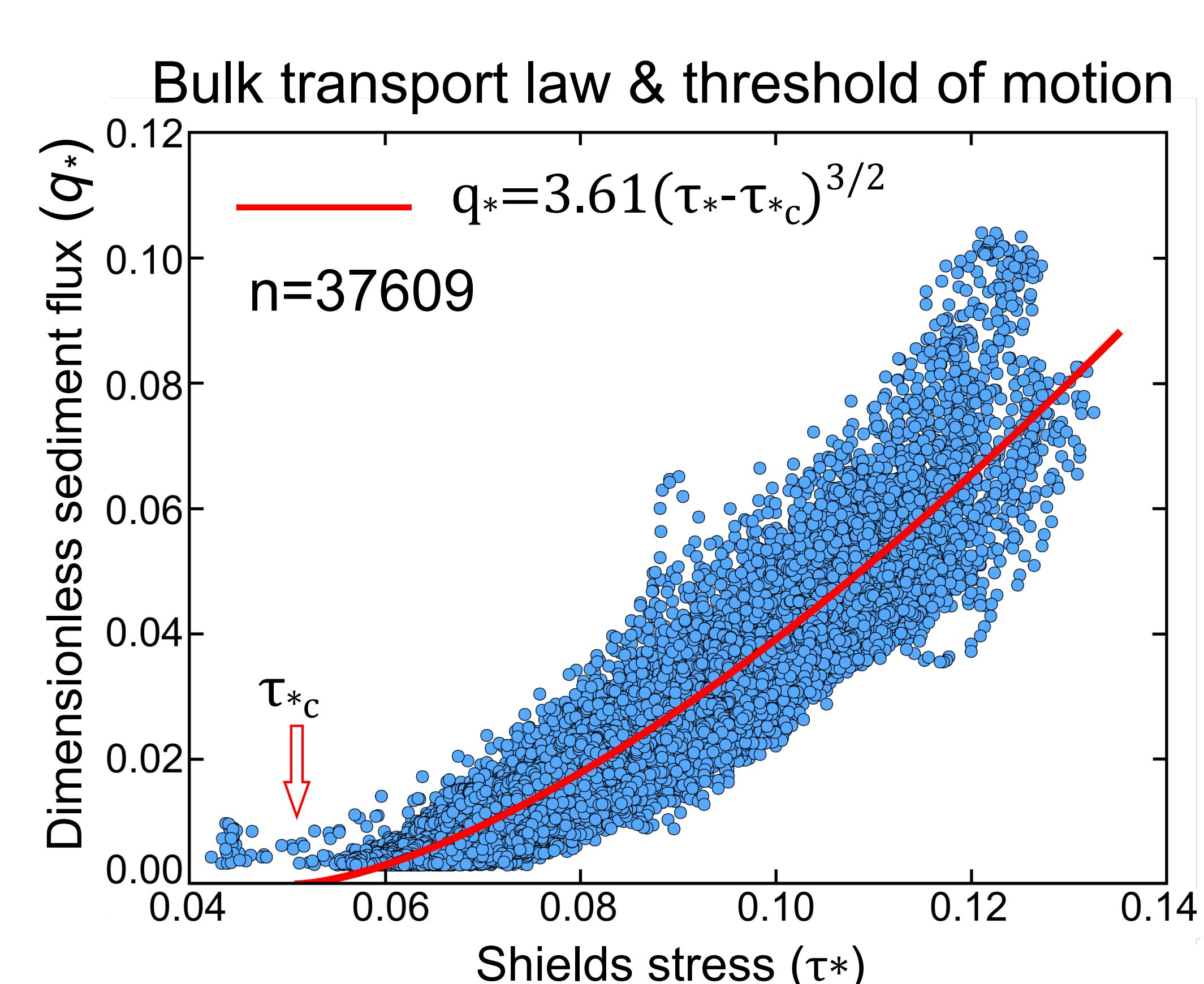


Complex transport phenomena occurs at timescales within floods.

Conclusion B - Transient floods may be approximated by steady flow with equal  $T^*$ .



$$\tau^* = \rho U_c^2 / (\rho_s - \rho) g D \quad q^* = q_b / \sqrt{Rg D^3}$$



Linear scaling between impulse and total flux

