

# Coupled Modeling of River and Coastal Processes: New Insights about Delta Morphodynamics, Avulsions, and Autogenic Sediment Flux Variability

## Motivation & Research Questions

- Deltas are flat & fertile → densely populated
- Important for agriculture, resources, and transportation
- Inhabitants increasingly susceptible to natural disasters
- Humans have:
  - Decreased sediment supply (e.g. dams)
  - Altered river course (e.g. channelization, levees)
- Relative sea-level rise rate (SLRR) increases → aggradation & backfilling increase (morphodynamic backwater) → avulsions assumed to be more frequent

What key feedbacks between fluvial and coastal processes drive avulsions and delta morphology?

How are delta morphodynamics affected by changing forcings (e.g., sea-level rise) over long time scales?



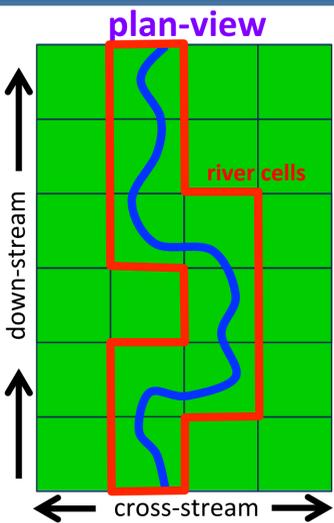
## New Delta Evolution Model

- Need to link both fluvial, deltaic, and coastal systems over multi-avulsion and lobe-building timescales
- Based on couplings using the Community Surface Dynamics Modeling System framework (Basic Model Interface)
- Generalized & scale invariant
- Capable of simulating large space & time scales

River Avulsion and Floodplain Evolution Model (RAFEM)

CSDMS Basic Model Interface

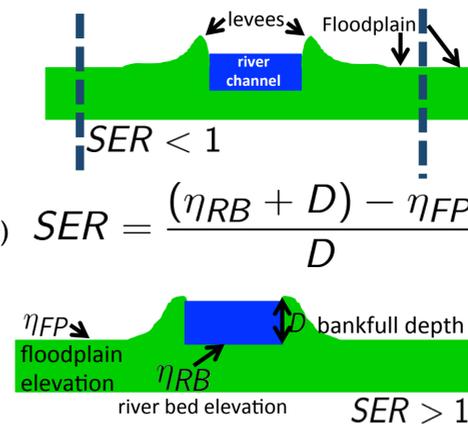
Coastline Evolution Model (CEM)



## River Avulsion and Floodplain Evolution Model (RAFEM)

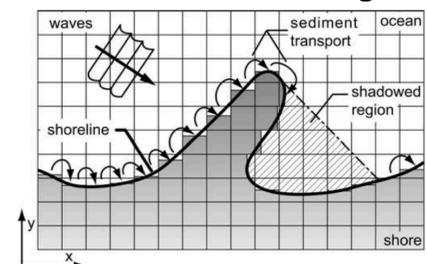
- Cell width  $\gg$  channel width
- Steepest-descent methodology (following Jerolmack and Paola, 2007)
- Diffusion of river profile (Paola et al., 1992; Paola 2000)
- River avulsions triggered by normalized super-elevation ratio (SER) (Mohrig et al., 2000), unsuccessful if not shorter than previous path
- Floodplain deposition = crevasse splay (after 'failed' avulsion; steepest path longer than current course)

## cell cross-section



## Coastline Evolution Model (CEM)

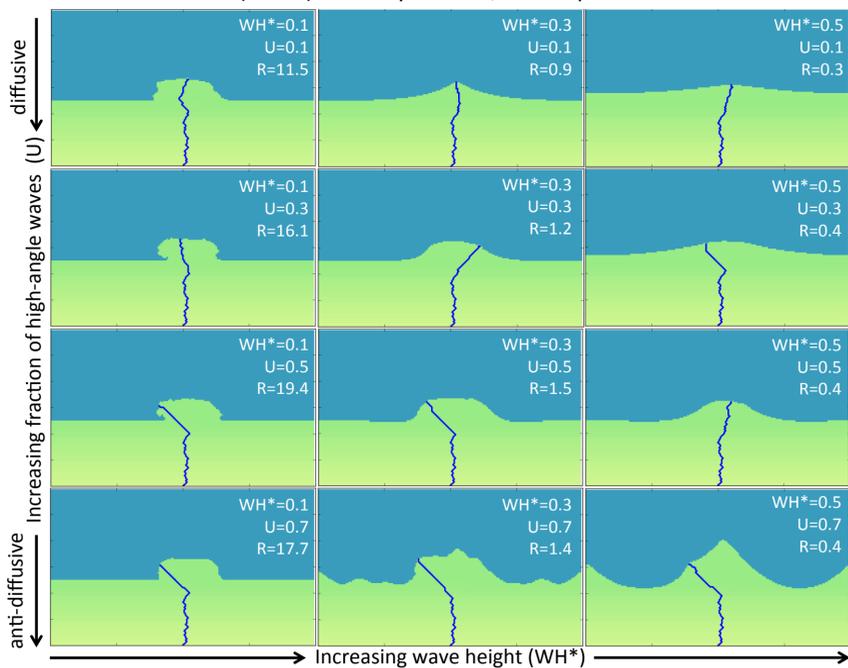
- Shoreline erosion & accretion driven by alongshore sediment transport
- Conserves nearshore sediment
- Wave climate and shadowing



Ashton and Murray, 2006

## Wave climate diffusivity affects morphology

- low wave height: sign of wave climate diffusivity doesn't matter; waves too low to affect shape
- higher wave height: sign does matter, affects morphology & avulsion time scales
  - diffusive ( $U < 0.5$ ) → flat shorelines, progradation inhibited
  - antidiffusive ( $U > 0.5$ ) → locally smooth, but cusped shorelines



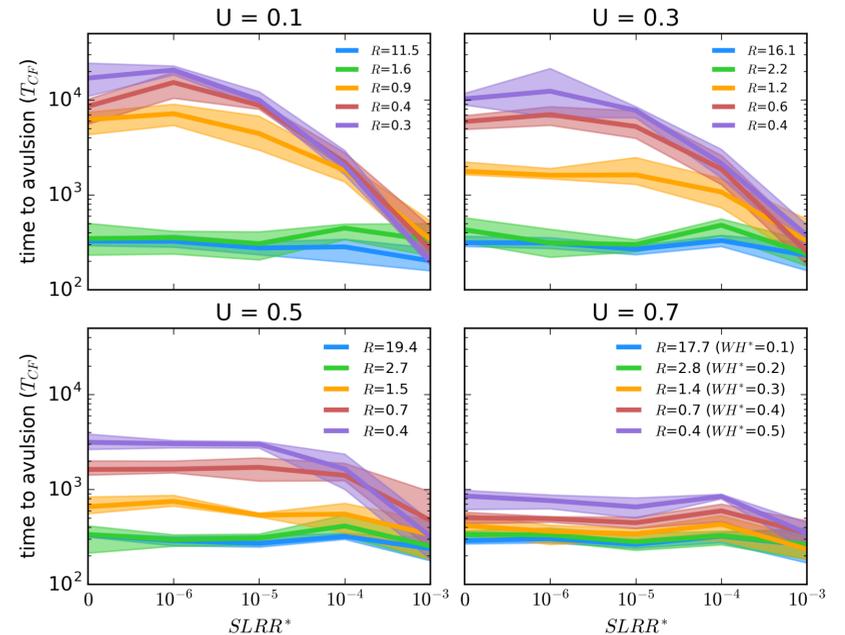
fluvial dominance ratio R expresses how quickly fluvial sand is delivered to shore vs. how quickly it is reworked by waves (Nienhuis et al., 2015):

$$R = \frac{Q_r}{Q_{s,max}}$$

R > 1: river-dominated  
R < 1: wave-dominated

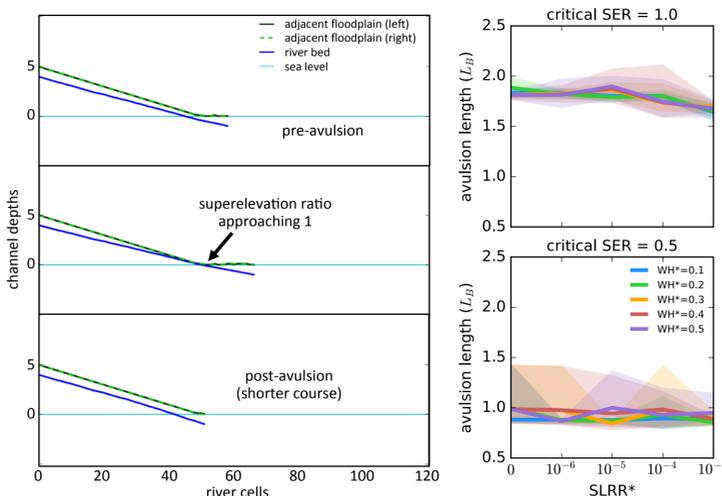
## Avulsion time scales

- diffusively wave-dominated: progradation slow, avulsions take longer to occur
- river-dominated or  $U > 0.5$ : progradation not inhibited, avulsions happen quickly
- Increasing SLRR\* only decreases avulsion time scales for wave-diffused deltas!  
→ In river-dominated or  $U > 0.5$  cases, lateral (transgressive) movement of shoreline counteracts base-level driven aggradation, no net effect on avulsion timing



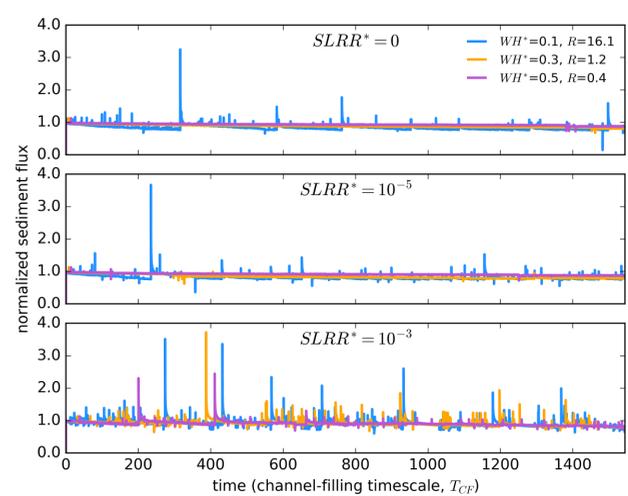
## Avulsion length scale

- Preferential length scale is a function of geometry (not varying discharge)
- over long-term, profile diffuses more rapidly than surrounding floodplain, becomes super-elevated at terrestrial concavity
- length scales with critical SER, values scale well with lab and field data



## Sediment flux variability

- Peaks represent avulsions
- River-dominated: larger cyclical peaks in flux
- Wave-dominated: less autogenic variability
- Increasing SLRR\* → more avulsions and variability for wave-diffused deltas (b/c base-level driven aggradation causes avulsions to occur more frequently)



## References

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## Acknowledgements

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