

Ice Sheet Modelling
State of the Art & Implications for
Simulation of Glacial Sedimentation

Shawn Marshall
University of Calgary

Garry Clarke, Gwenn Flowers, Dave Hildes
University of British Columbia

Community Sediment Model Workshop
INSTAAR, Boulder CO: Feb 19–22, 2002

Outline

1. Basis of ice sheet modelling
2. Current state of the art – strengths
3. Current state of the art – weaknesses, aka. the current research frontier:
 - Issues re: iceberg calving
 - Issues re: basal hydrology
 - Issues re: basal flow & ice streams
4. Implications for modelling sediment entrainment, transport, and deposition
5. Scale issues: Ice sheets vs. glaciers

1. Basis of Ice Sheet Modelling

- Continuum-mechanical models of ice deformation under gravity
- Several 3D models in the world; resolve 3D velocity, temperature, stress fields as well as ice sheet thickness (t)

ICE SHEET DYNAMICS

Conservation of Mass

$$\frac{\partial H}{\partial t} = -\frac{\partial (uH)}{\partial x} - \frac{\partial (vH)}{\partial y} + b$$

**Ice Flux
Divergence**

**Mass
Balance**

ICE SHEET DYNAMICS

Conservation of Momentum

$$\frac{\partial \sigma}{\partial x_j} = -\rho g$$

Glen's Flow Law

$$\dot{\epsilon}_{ij} = B(T) \left[\sum_2' \right]^{\frac{(n-1)}{2}} \sigma_{ij}'$$

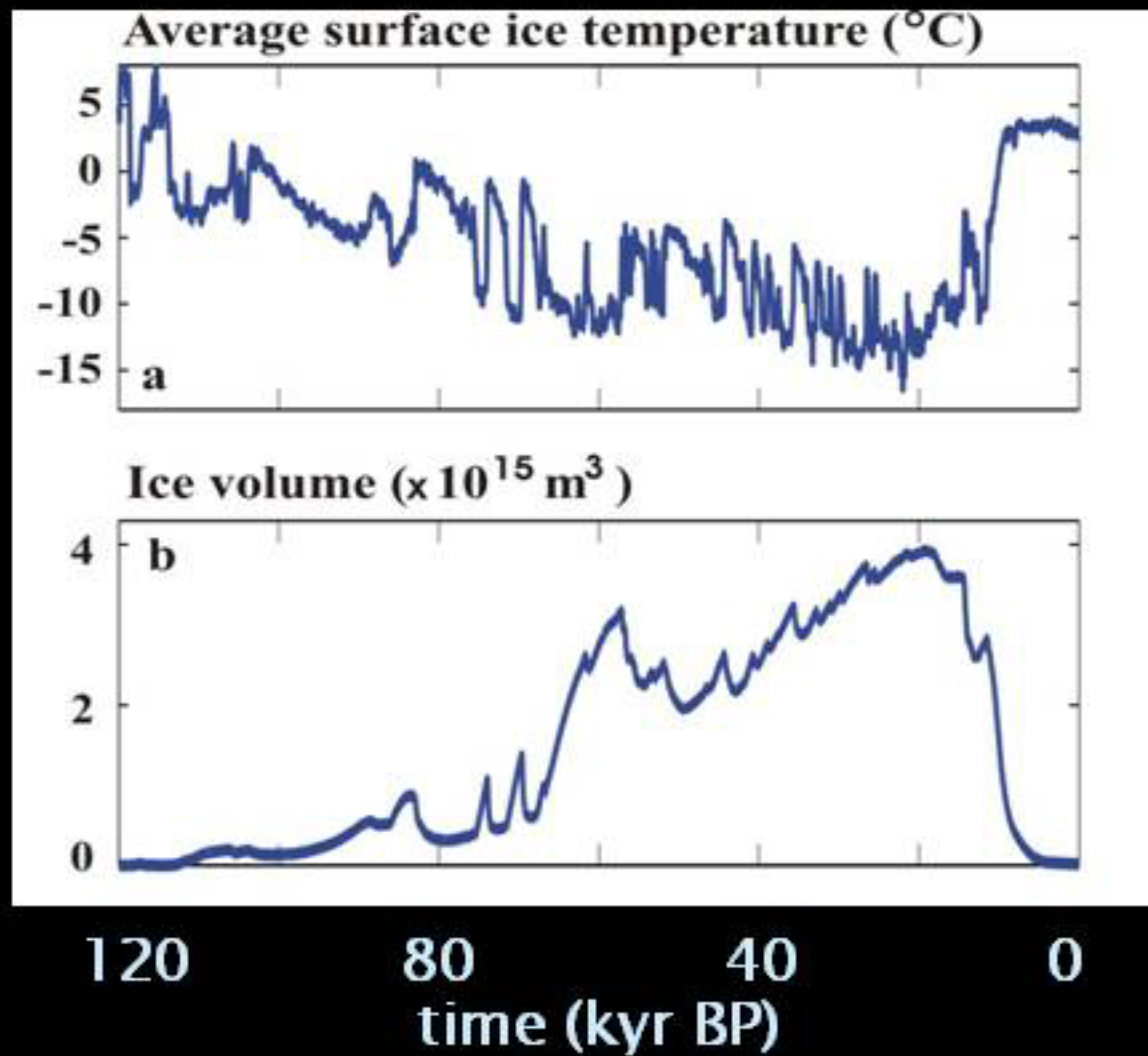
ICE SHEET THERMODYNAMICS

❄ Conservation of Energy

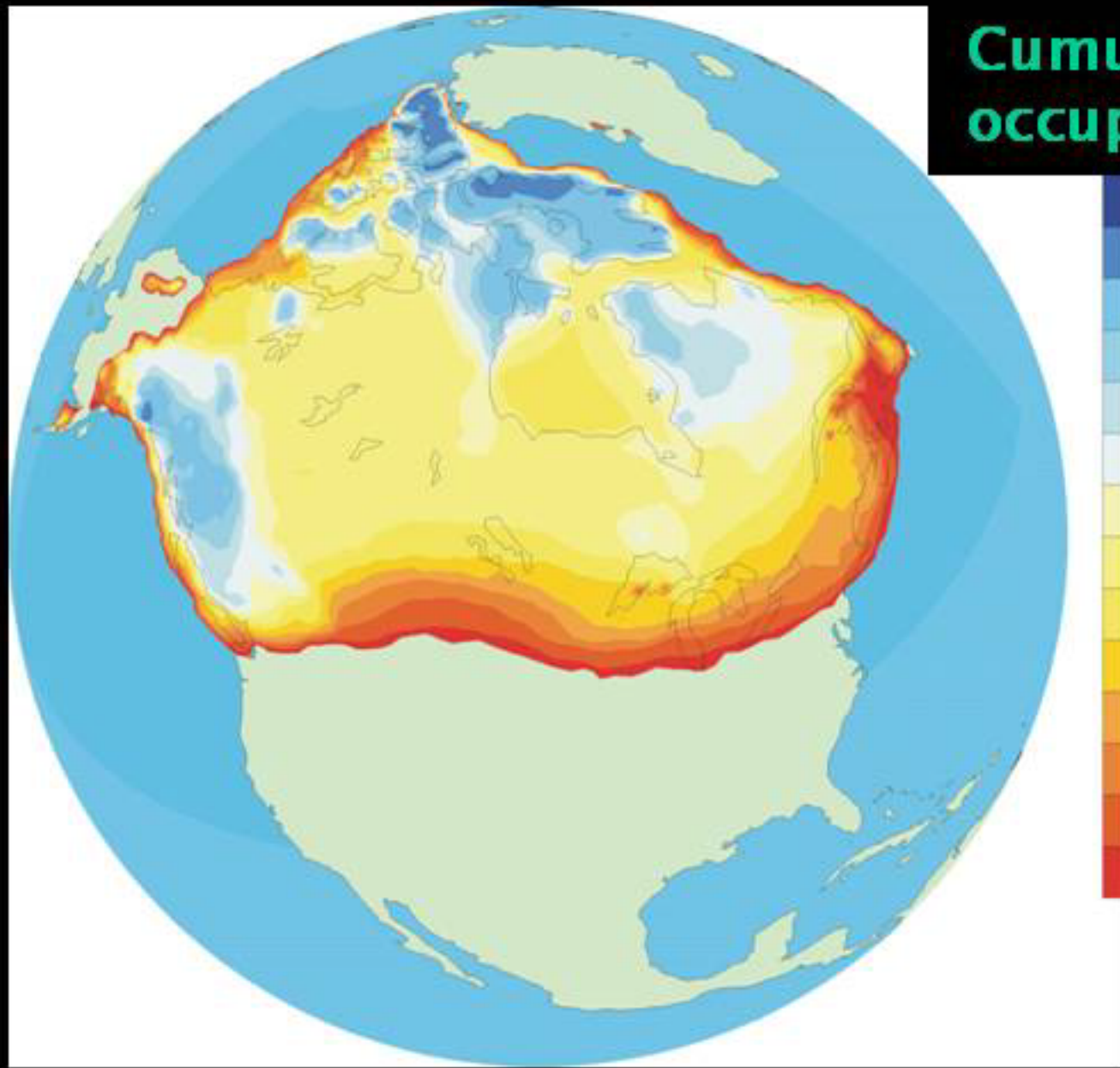
$$\frac{\partial T}{\partial t} = \underbrace{-v_k \frac{\partial T}{\partial x_k}}_{\text{Advection}} + \underbrace{\kappa \frac{\partial^2 T}{\partial x_k^2}}_{\text{Diffusion}} + \underbrace{\frac{\Phi}{\rho c}}_{\text{Strain Heating}}$$

$$\Phi = \dot{\epsilon}_{ij} \sigma'_{ij}$$

The Last Glacial Cycle in North America



Cumulative ice occupation (kyr)



100

80

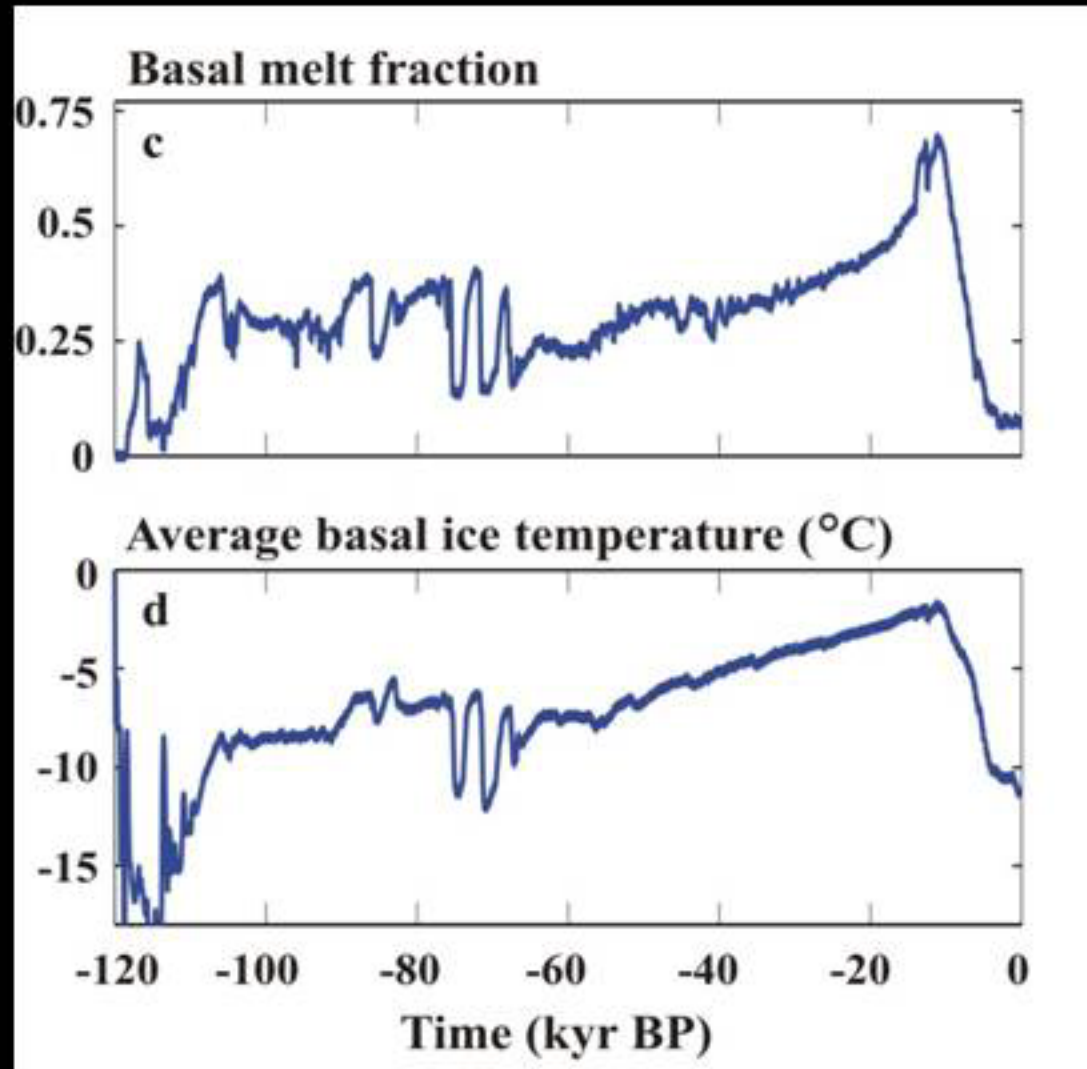
60

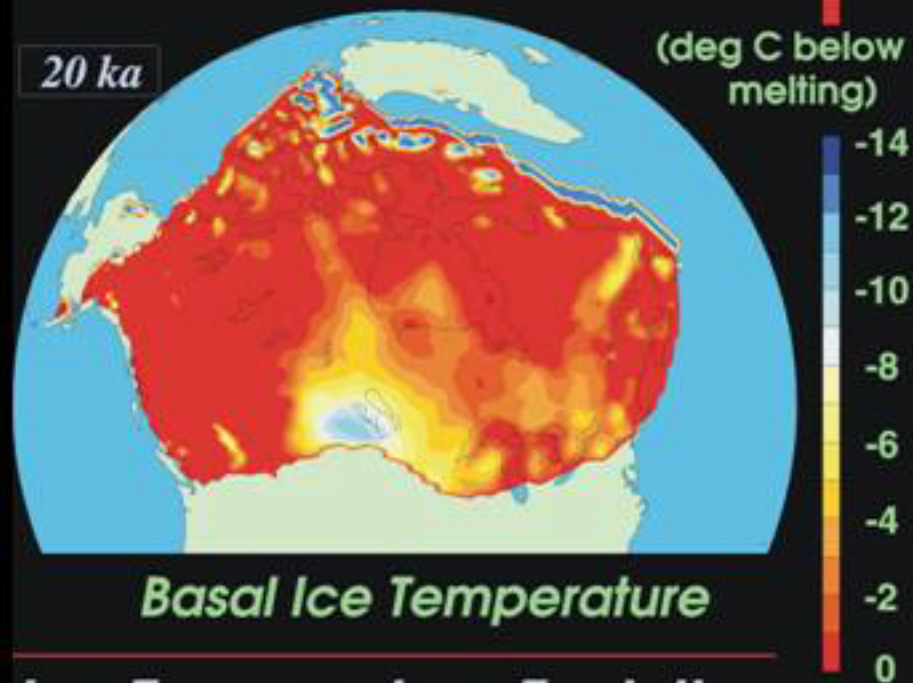
40

20

0

The Last Glacial Cycle in North America





Ice Temperature Evolution

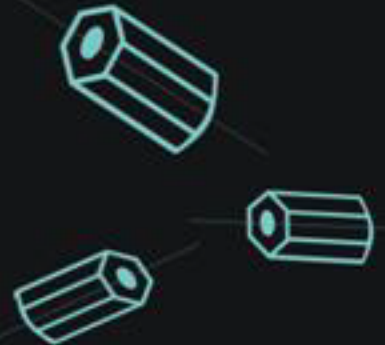


Uncertainties: Ice Rheology

Glen's "Enhanced" Flow Law

$$\varepsilon_{ij} = EB_0 \exp\left(\frac{-Q}{RT}\right) \sum_2' {}^{(n-1)} \sigma_{ij}'$$

Complications: Anisotropy
Impurities
Water content

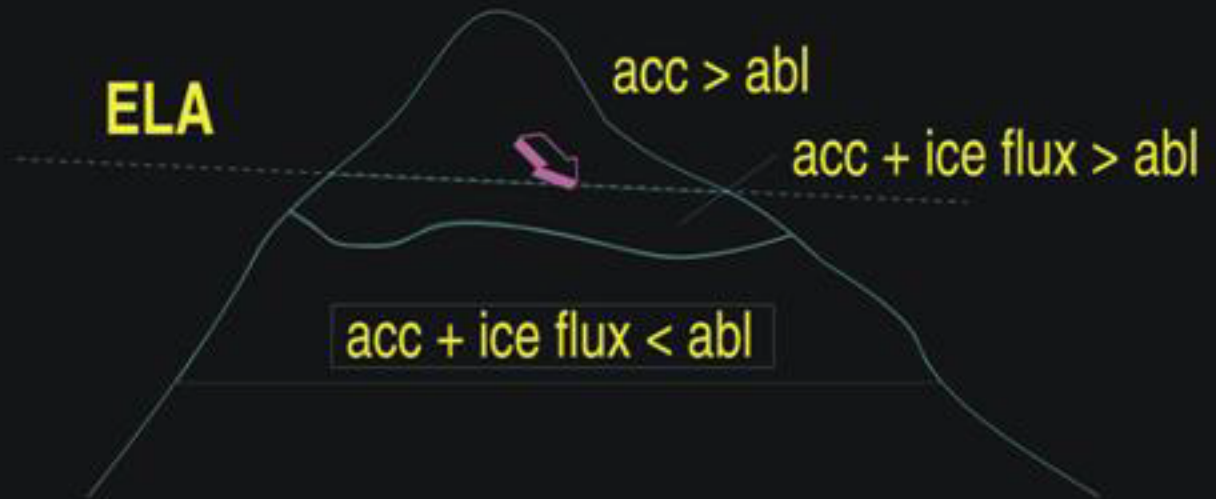




Uncertainties: Mass Balance

Importance:

Ice sheet or glacier nutrition and health



Concerns:

Skill of climate model

Model resolution

Ablation parameterizations: capturing the physics?

3. Current State of the Art – Weaknesses

Current research frontiers:

- Issues re: glacial hydrology
- Issues re: iceberg calving
- Issues re: basal flow & ice streams
- Only just getting to geologic outputs:
e.g. landforms, permafrost, striations,
isotopic labelling of groundwater and
runoff, erosion, dispersal trains,
marine sedimentation...

Glacial Hydrology

Expect large advances:

Has finally caught the interest of many in the ice sheet modelling world

Involves surface, englacial, subglacial, and groundwater routing. Definite challenges in both scale and process understanding

*** The relevant physical fields are relatively well-modelled or available as inputs**

❄ Subglacial Melt/Refreeze

When T_b is at the pressure melting point:

$$\rho L \frac{\partial H}{\partial t} = \rho L \dot{m}_b = \sum Q_b$$

Basal melt or
refreezing rate

Basal heat
fluxes



Subglacial Energy Balance

Same physics in both 3D and 1D models:

$$\frac{\partial T_b}{\partial t} = -v_{bk} \frac{\partial T_b}{\partial x_k} + \frac{k_I}{\rho c} \frac{\partial T_b}{\partial z} - \frac{1}{\rho c} \left[k_W \frac{\partial T_W}{\partial z} \mid k_B \frac{\partial T_B}{\partial z} \right] + \frac{(\Phi_b + \gamma_b)}{\rho c}$$

Basal heat flux

So for basal hydrology:

- OK for calculating sources
- Modelling subglacial drainage is another question. But progress here → Alley, Arnold, Flowers, Fastook and Johnson
 - * need storage & routing considerations

Surface hydrology and runoff:

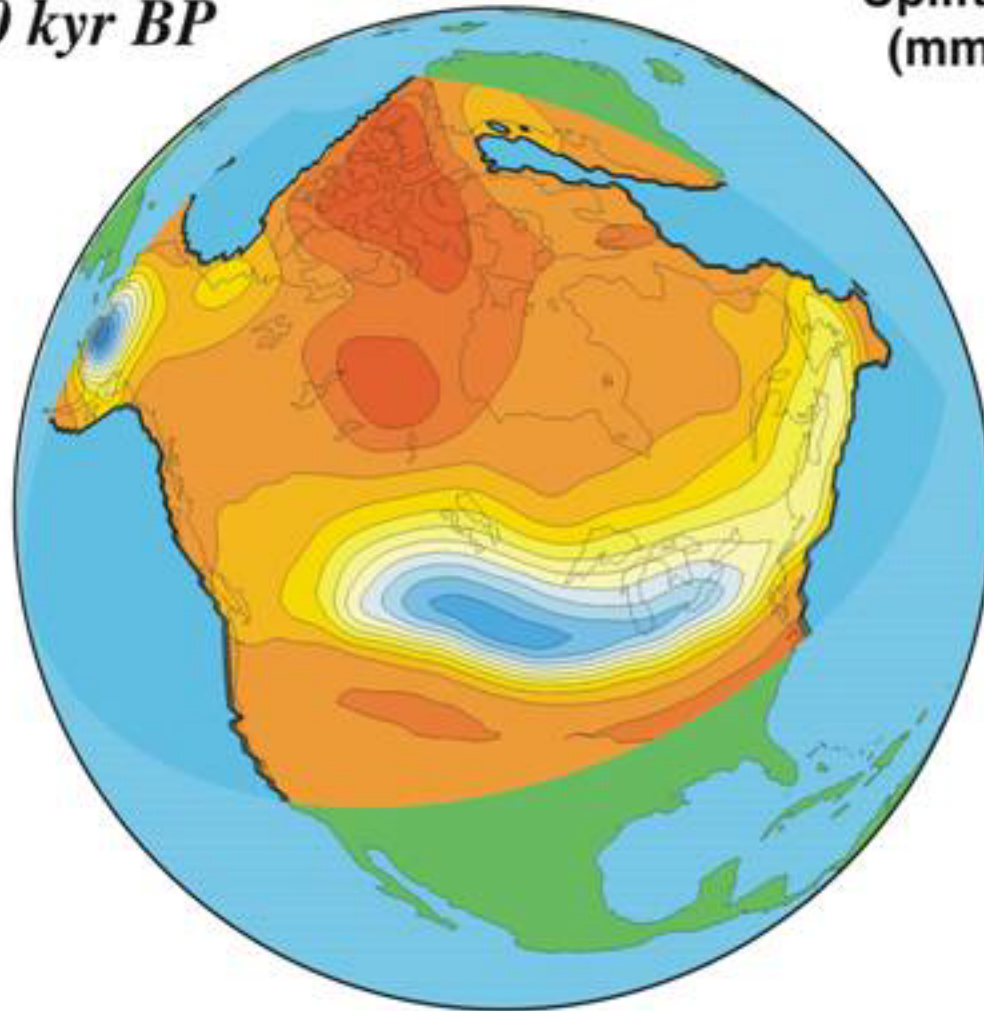
- Actually easier
- Possible to quantify time-evolving “river” flow in a coarse sense → scale issues for more detail

Modelled Isostatic Uplift, LGM

LAST GLACIAL MAXIMUM

21-20 kyr BP

**Uplift Rate
(mm/yr)**



4.5

1.0

-2.5

-6.0

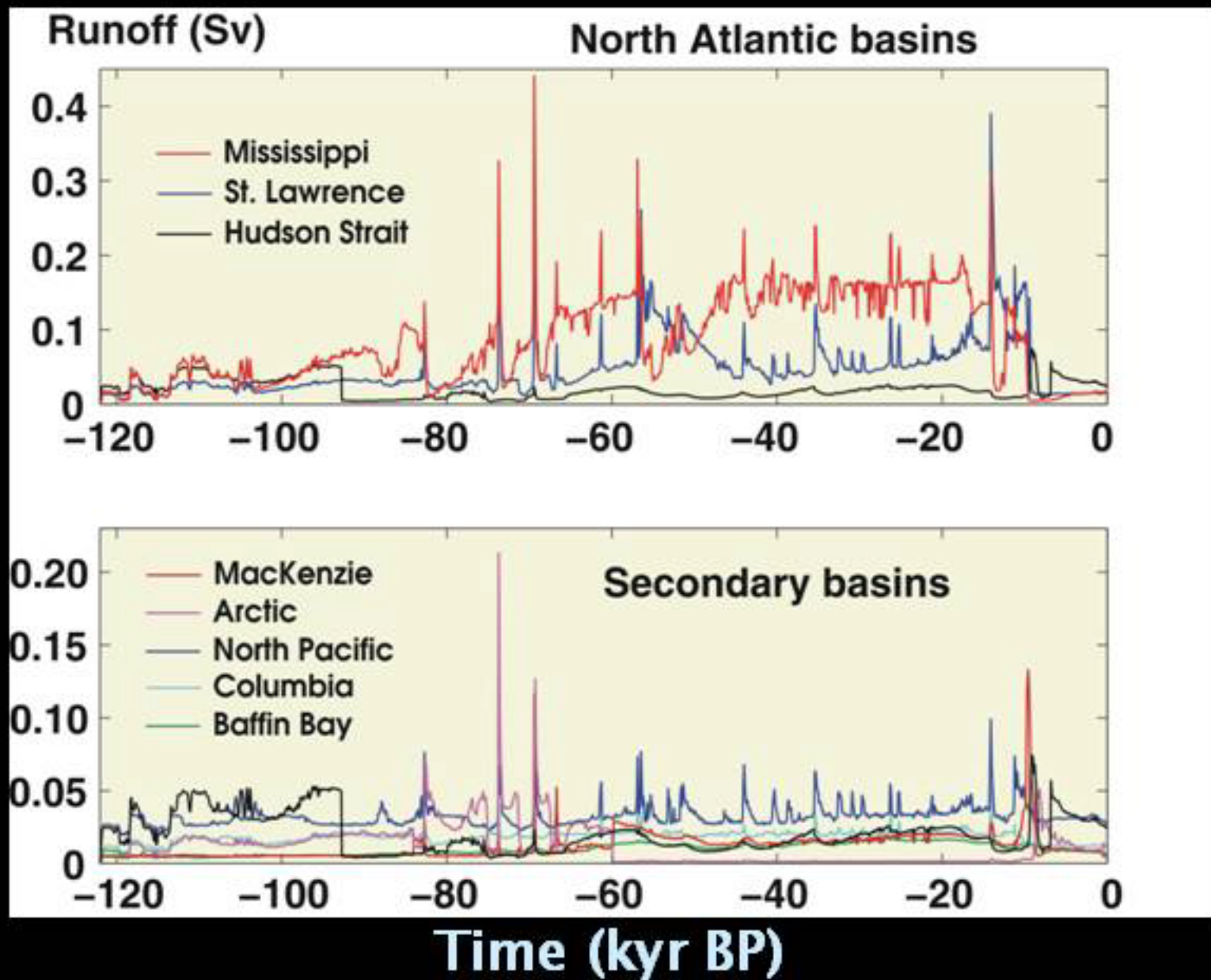
-9.5

-13.0

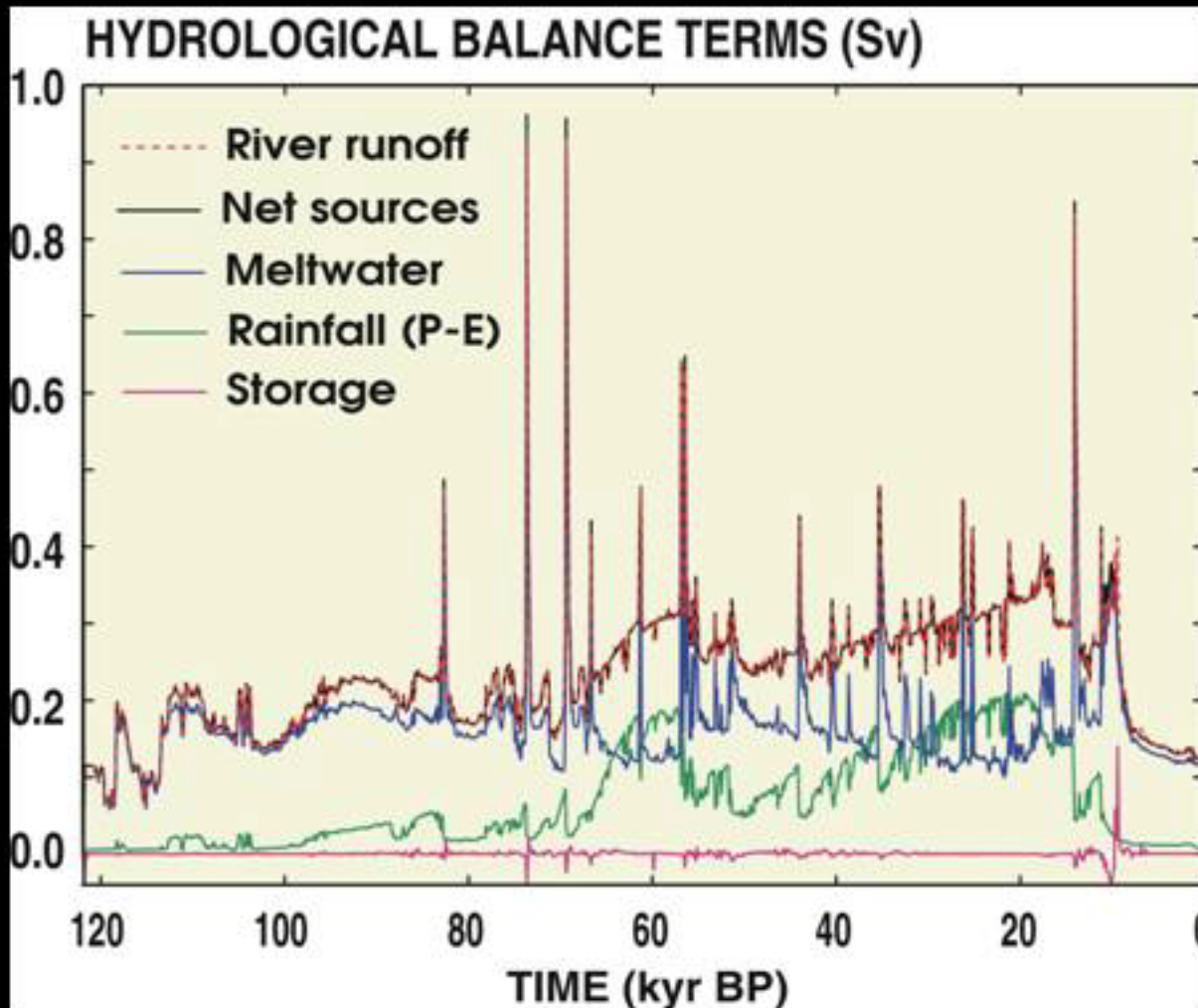
-16.5

-20.0

North American River Runoff, Last Glacial Cycle

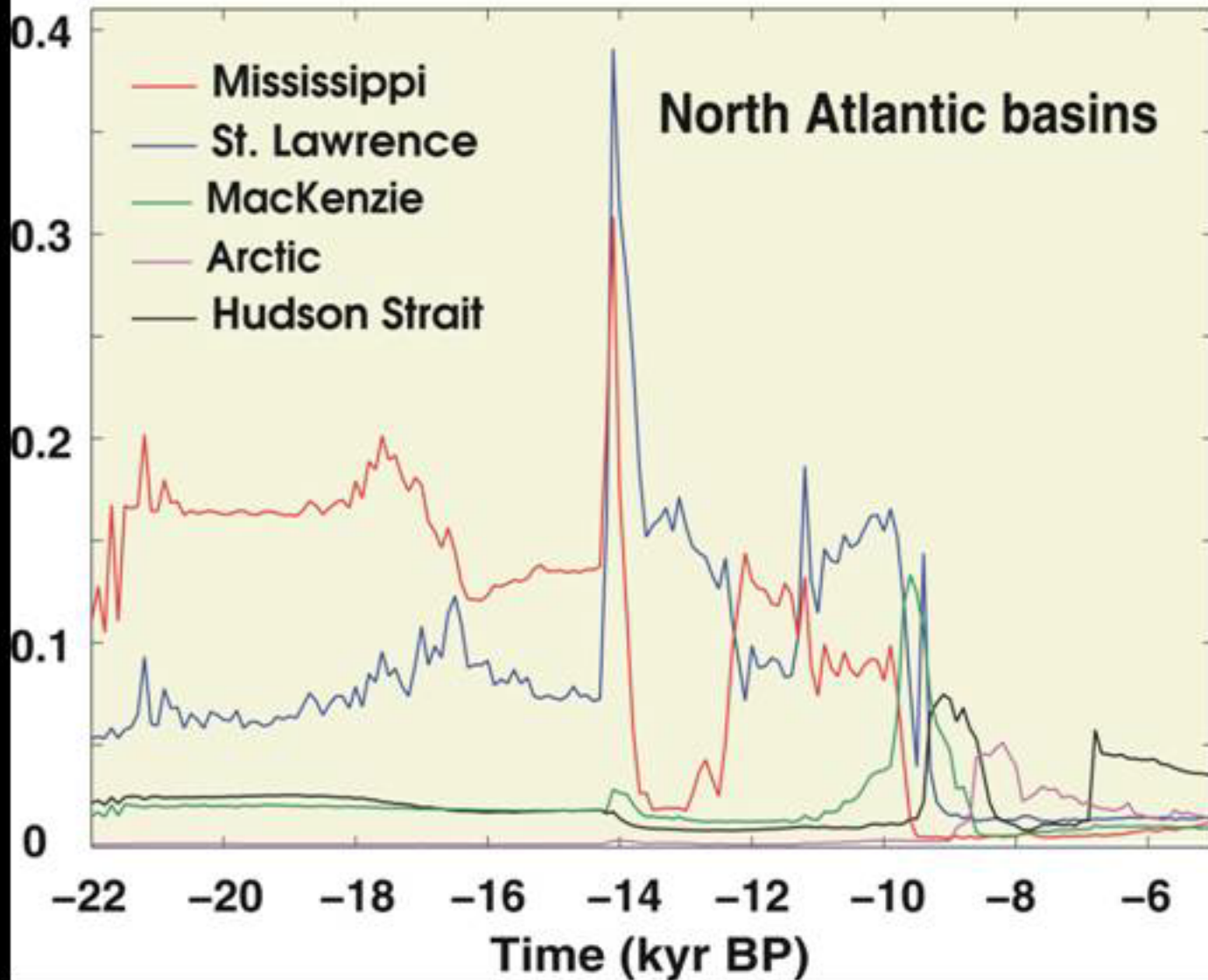


NORTH AMERICAN HYDROLOGICAL BALANCE, LAST GLACIAL CYCLE



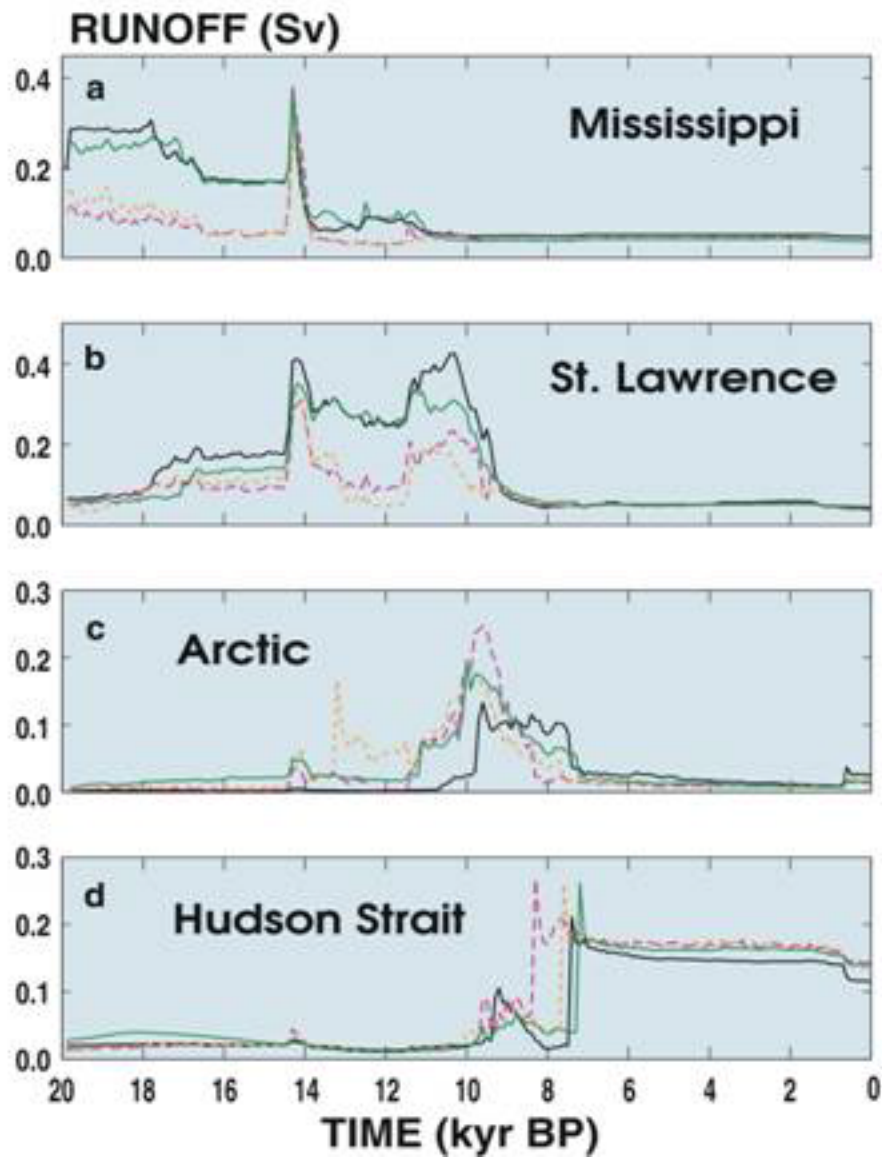
Deglacial Meltwater Pulses

Runoff (Sv)



Staged Runoff Sequence

Sensitivity tests, deglaciation





Iceberg Calving: Model Issues

Importance:

- Dominant mass loss term in Antarctica
- Rapid response to climate change?

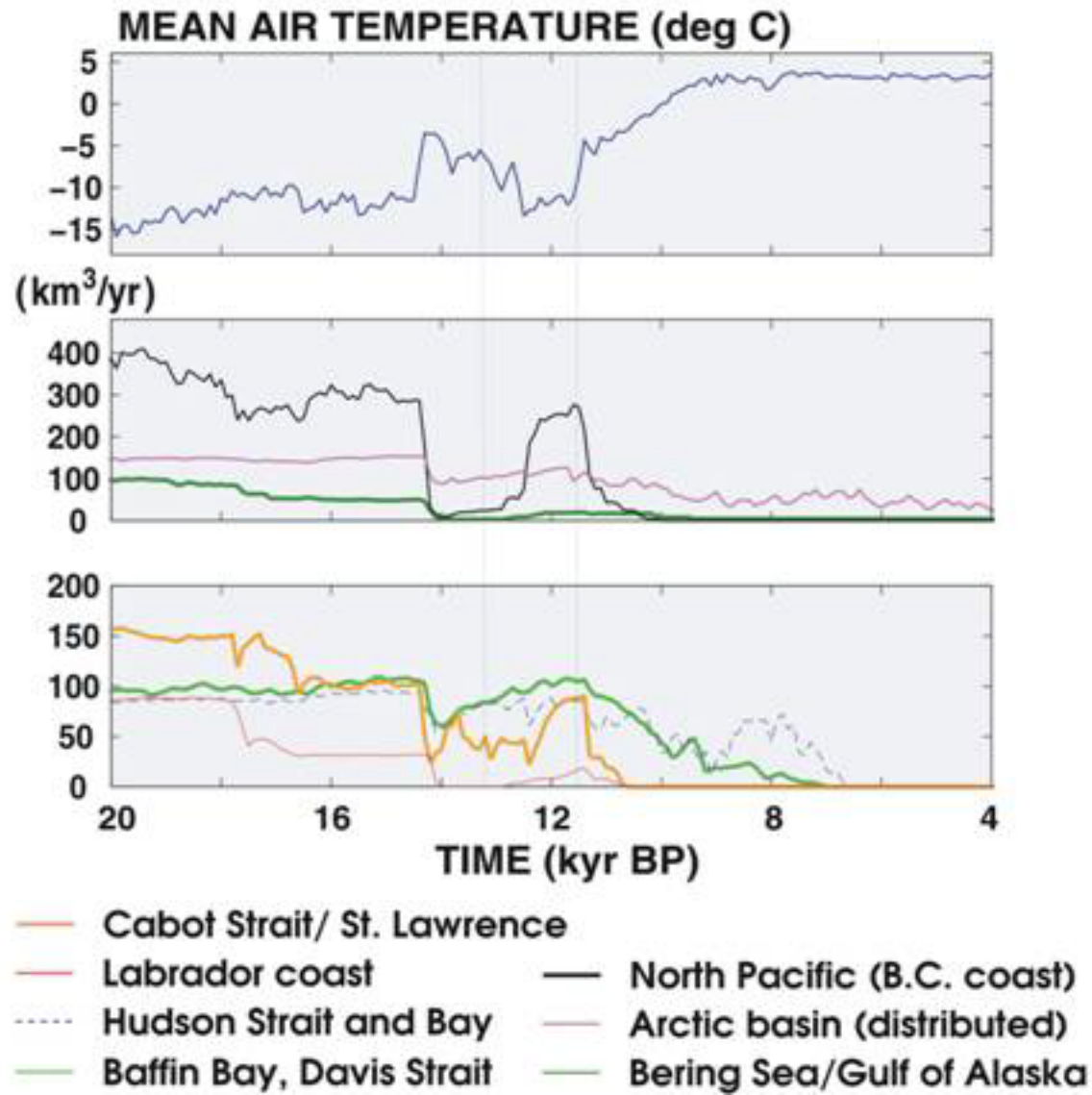


Process Physics:

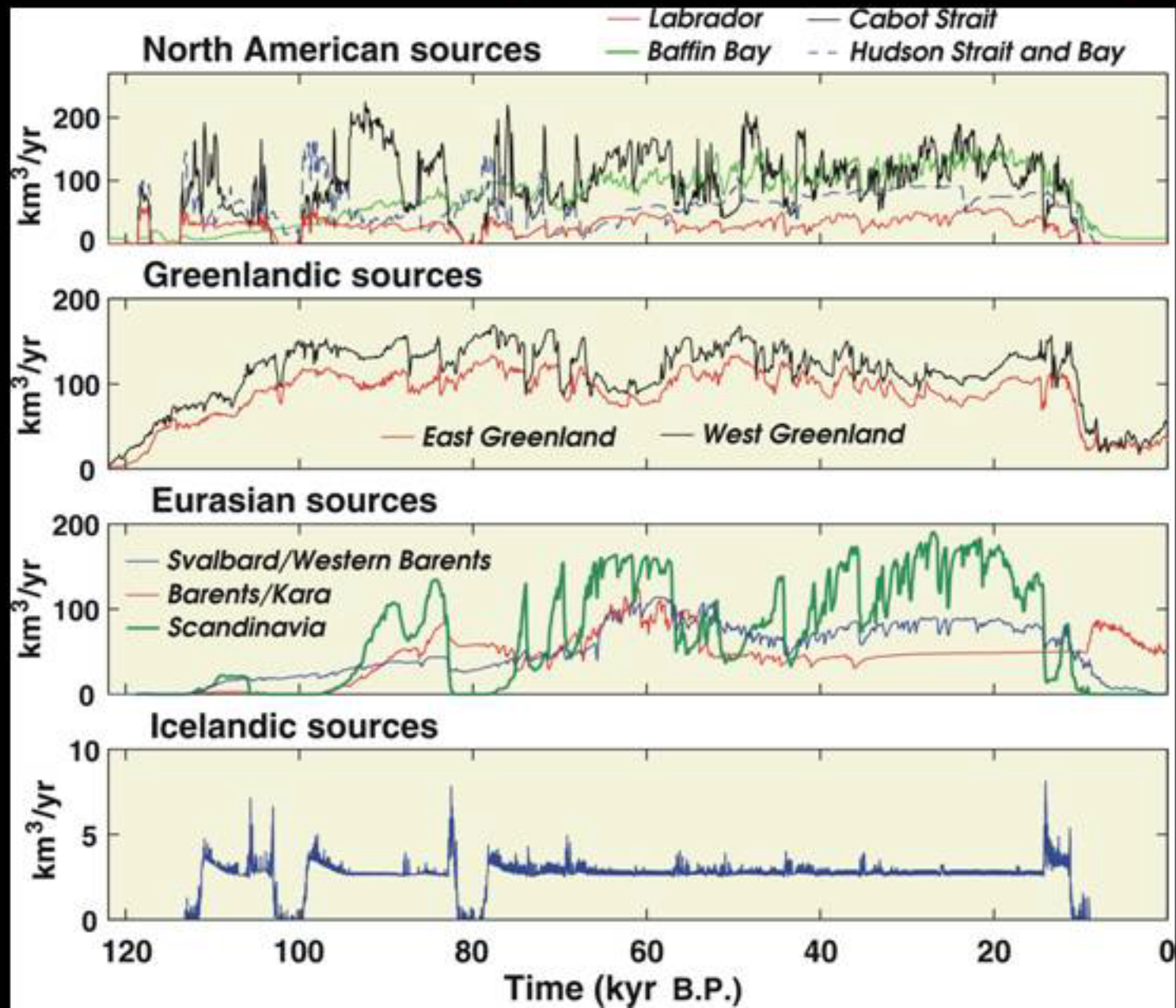
- Not yet quantified, possibly non-deterministic
- Involves 3D stress and temperature fields, fracture mechanics, water depth, open water season, and possibly water temperature, tidal amplitude

Modelled Iceberg Flux: Time Lag with Climate Forcing

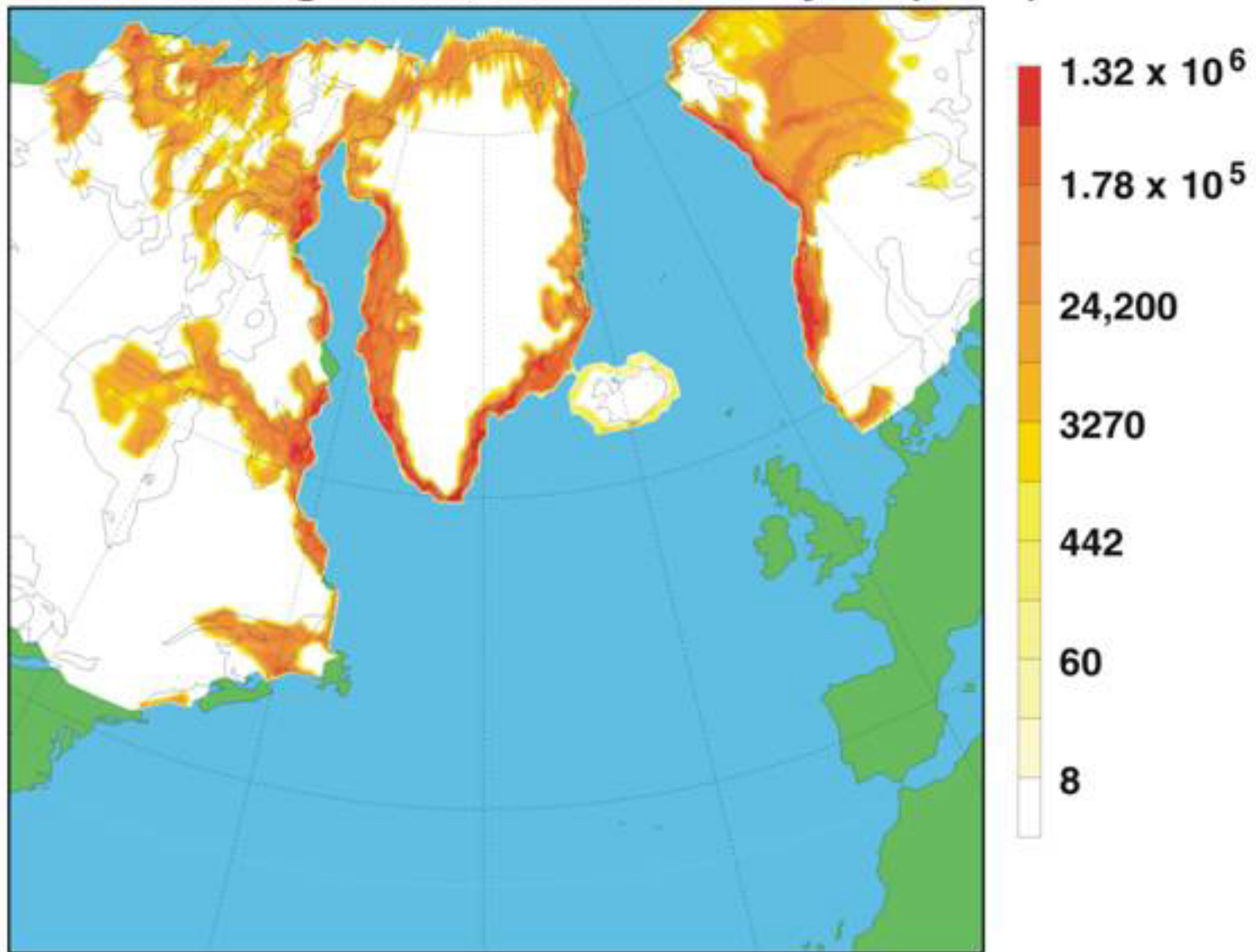
Deglaciation, $t = 20-4$ ka



North Atlantic Iceberg Flux, Last Glacial Cycle



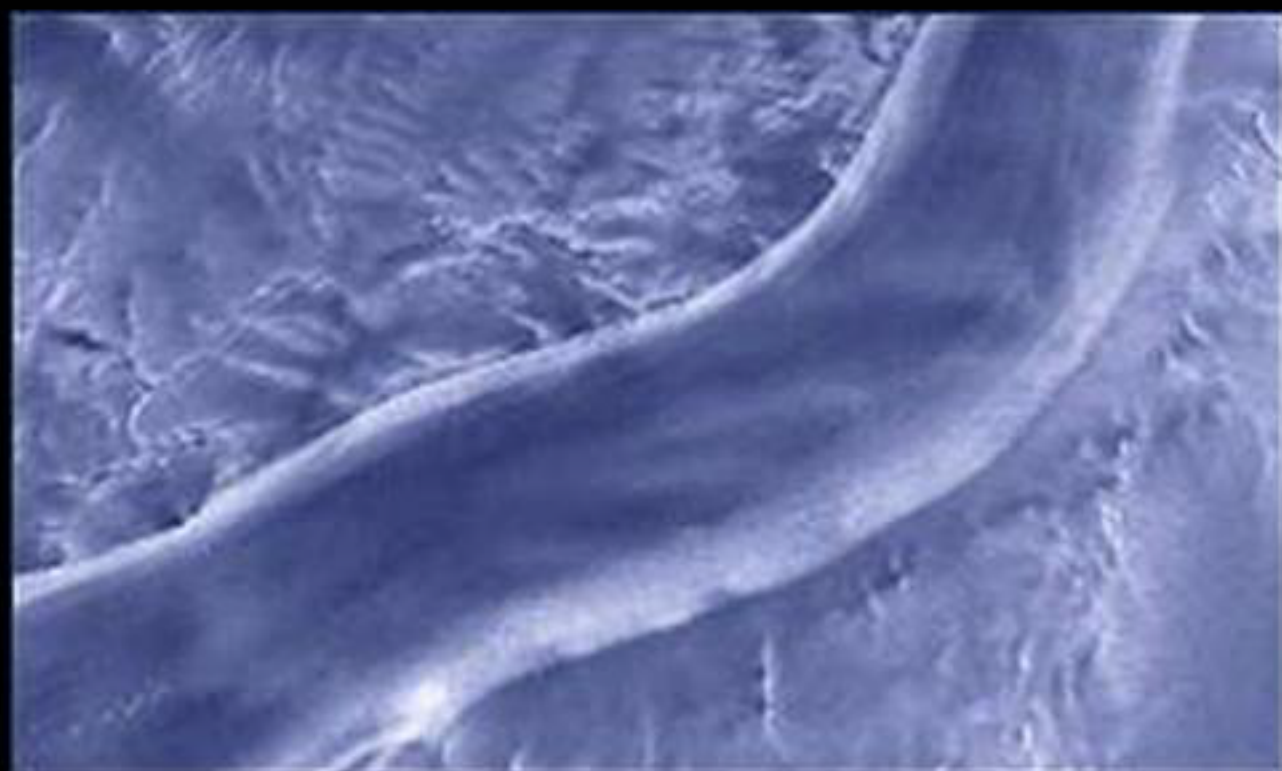
Total Iceberg Efflux, Last Glacial Cycle (km³)



Basal Flow and Ice Streams

Also on the map now

Definite challenges in both scale and process understanding





Uncertainties: Basal Flow

Importance:

Ice streams are the main arteries of an ice sheet

Paleo-fast flow evidence abounds

Main short-term influence on sea level, ocean freshwater balance

Process Physics:

Not yet quantified, probably non-deterministic

Involves ice and bed mechanics, basal temperature regime, subglacial hydrology, bed topography and geology

ICEBERG RAFTED DEBRIS IN A HEINRICH EVENT

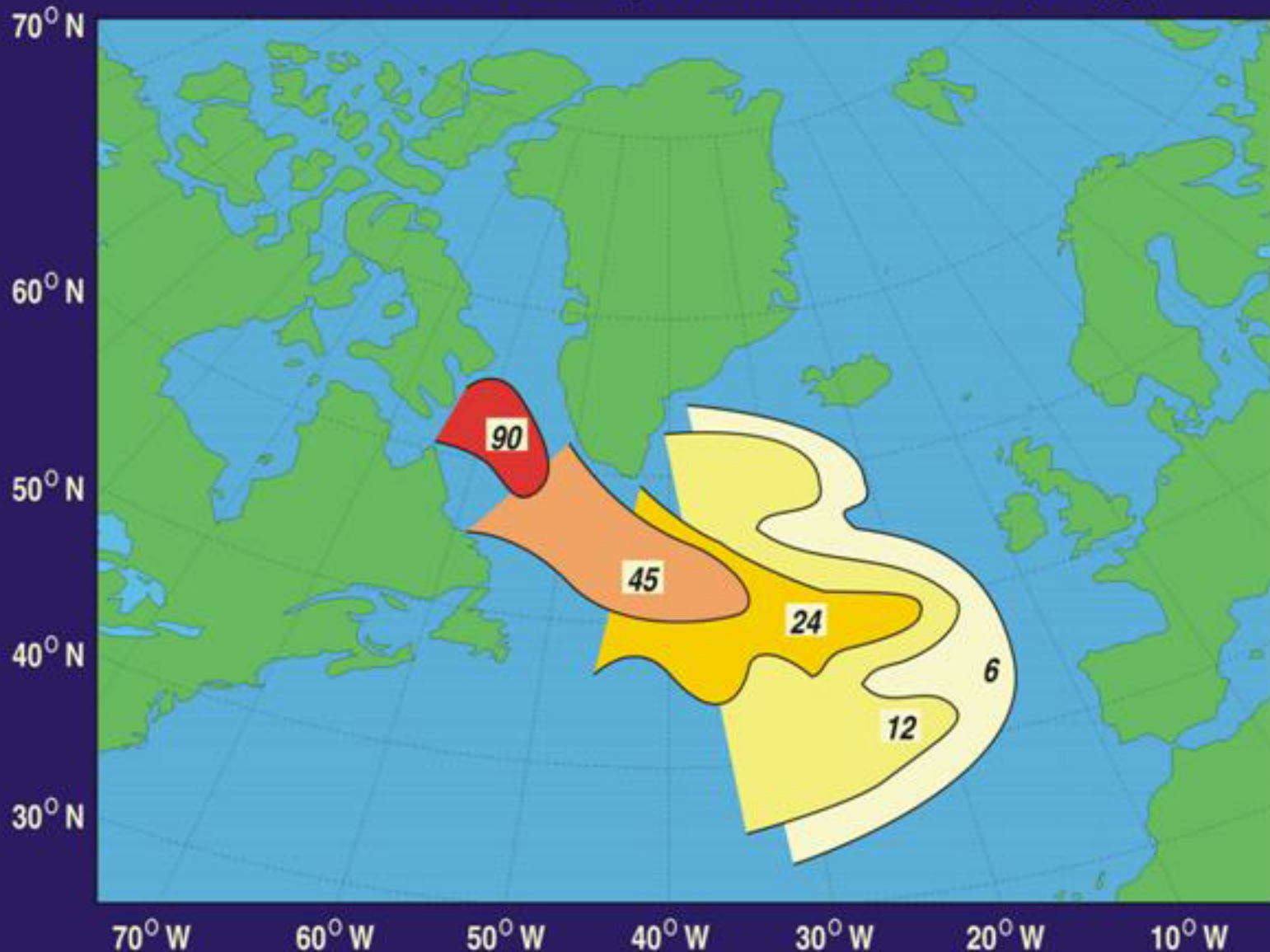
Detrital carbonate layer thickness (cm)



After Dowdeswell et al. (1995)

ICEBERG FLUX IN A HEINRICH EVENT

Freshwater-equivalent melt rates (cm/yr)



Maximum surge scenario, Marshall and Clarke (1997)

4. Modelling sediment entrainment, transport, and deposition

Again, a lot of the relevant physical fields are currently modelled: basal stress regime, water pressure, temperature, basal flow

4. Modelling sediment entrainment, transport, and deposition

Again, a lot of the relevant physical fields are currently modelled: basal stress regime, water pressure, temperature, basal flow

However, a relatively untouched field:

- geological substrate must be characterized
- process scale \ll model scale
- process physics challenging
- process physics involve some glaciological fields that are not well-characterized, e.g. water pressure fluctuations

Simple approaches still insightful:

e.g. erosion modelling,

$$E(\lambda, \theta, t) \propto E_0(\lambda, \theta) \int f[v_{bj}(\lambda, \theta, t), \tau_{bj}(\lambda, \theta, t)] dt$$

One can experiment with erosion models

$f[v_{bj}(\lambda, \theta, t), \tau_{bj}(\lambda, \theta, t)]$ and consider freeze-on & regelation (T, P) entrainment processes, etc.

→ Advect with the ice or the subglacial drainage

Last Glacial Cycle in Quebec/Labrador

Potential Erosion

Easterly Striations



Last Glacial Cycle in Quebec/Labrador

Southerly Striations

Westerly Striations

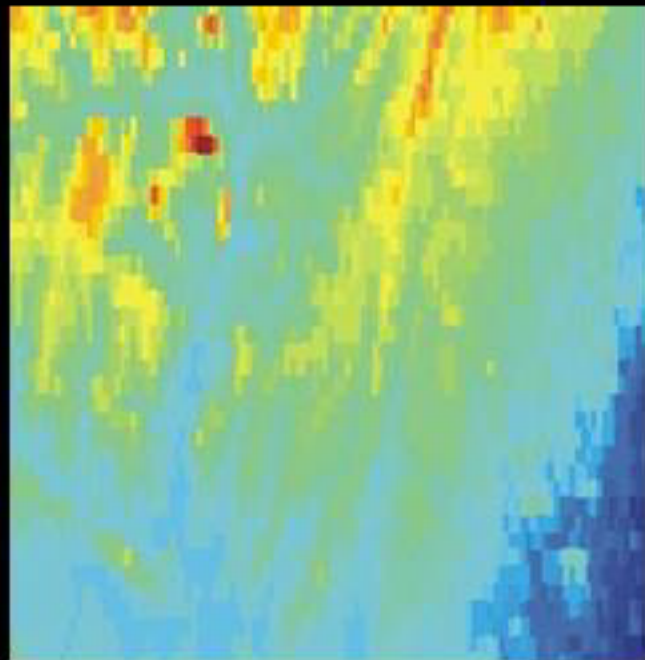


5. Scale Issues: Ice Sheets vs. Glaciers

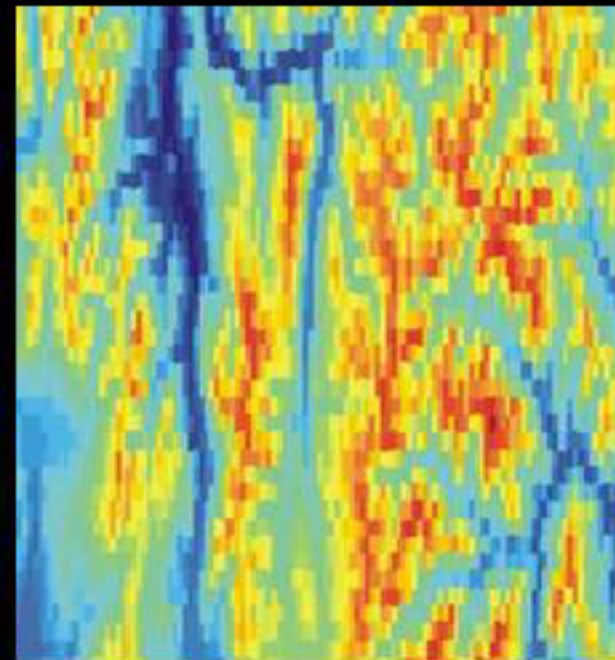
- Limitations to application of large-scale ice sheet models: $\Delta x > \sim 3H$
- Higher-order dynamical models have been developed, esp. for 1d modelling; so we're OK on the scale of individual glaciers.
- Harder: 2d and 3d simulation in complex terrain. Will get there though.

Topographic Demeanor

Wisconsin test point (m)

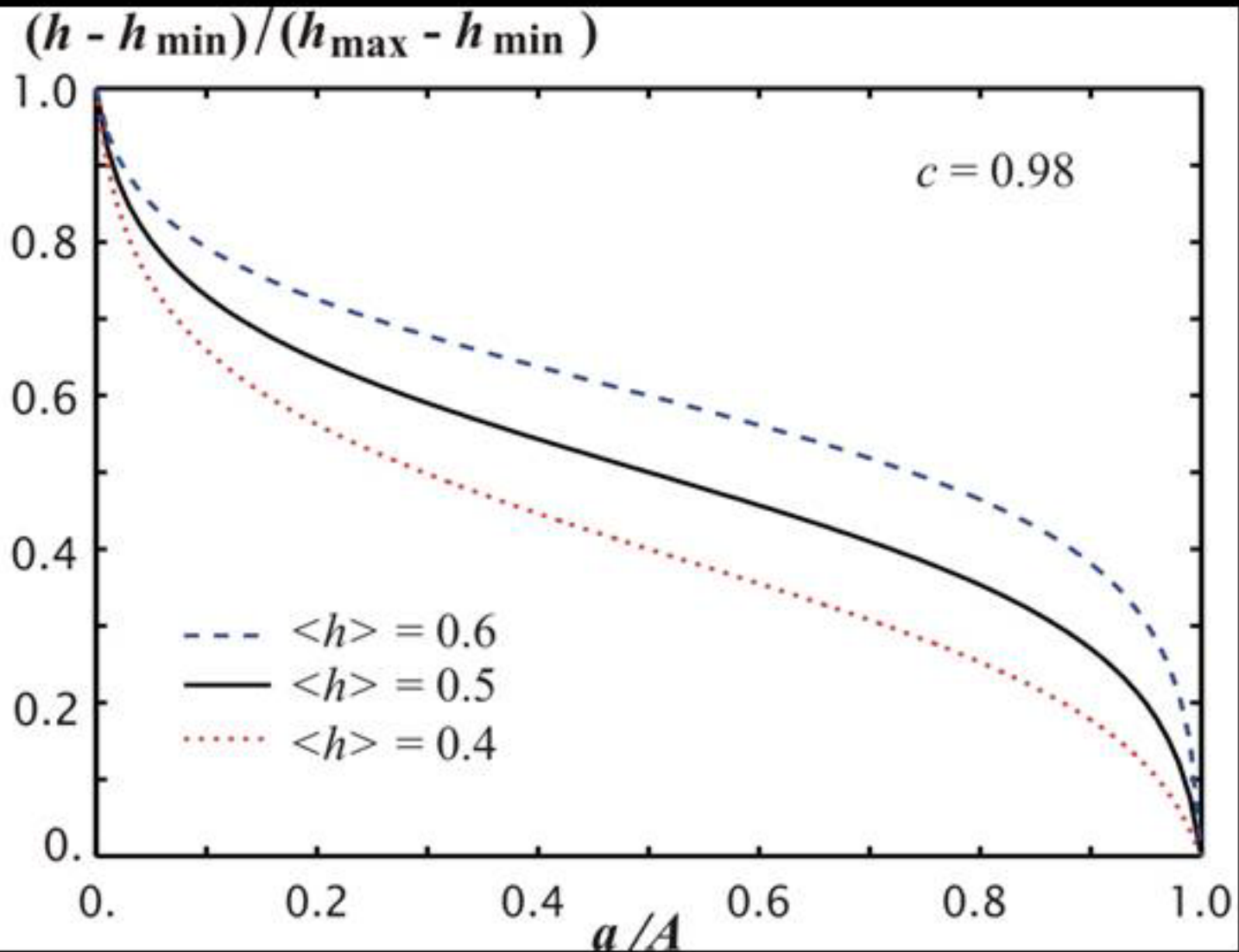


Idaho test point (m)

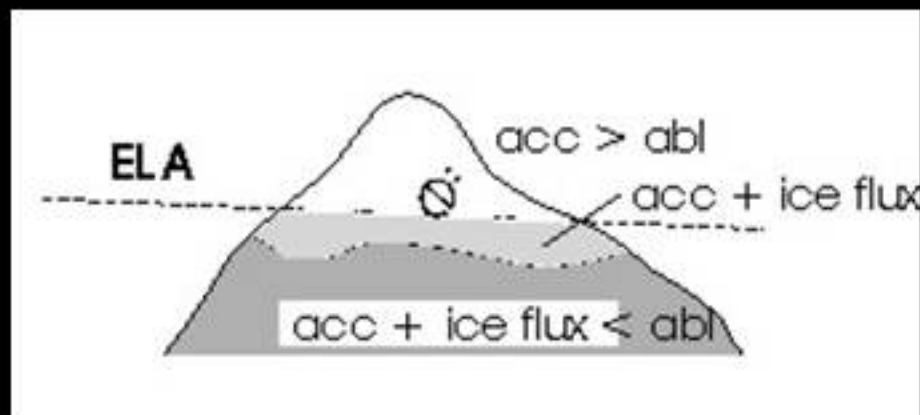


1-Degree Cells

Synthetic Hypsometry Function



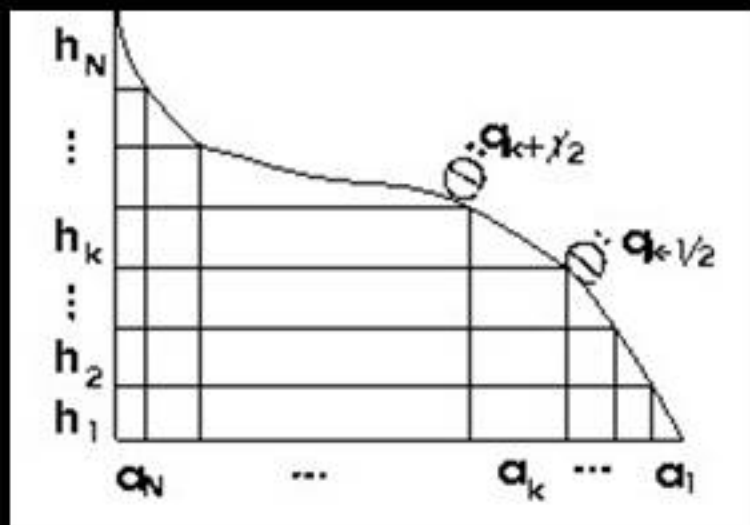
Subgrid Scheme

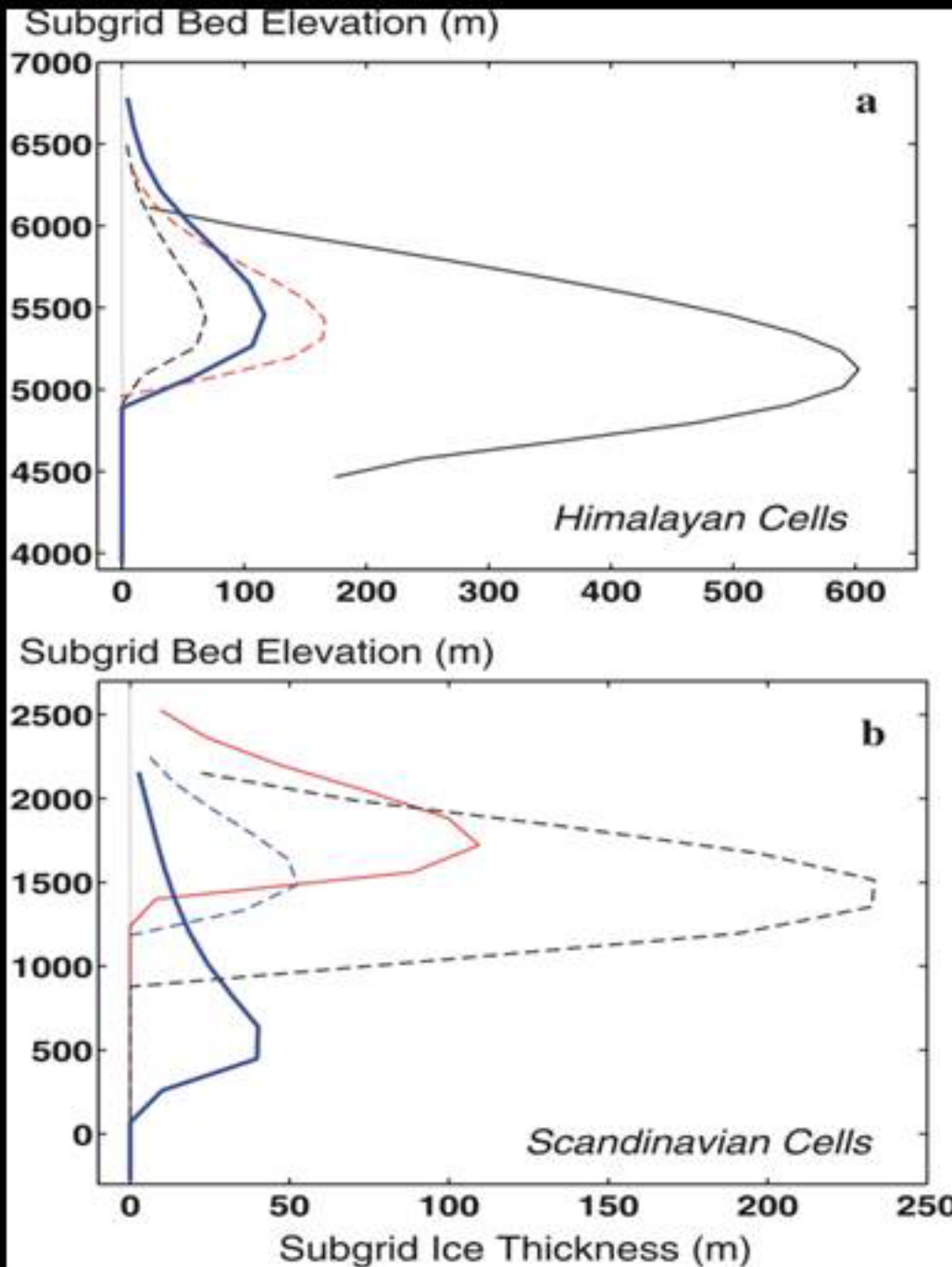


Downslope grid transport in subgrid bins

$$q_k = (H_k + \int b_k dt) / \tau$$

$$[1 - \exp(-(\Delta h_{sk} / L_k)^3)]$$





Subgrid Ice
Distribution

Sample
Synoptic
Grid Cells

Summary

- ⊗ Good timing to incorporate glacial process models into the CSM strategic plan
 - the relevant process considerations and large-scale outputs are on the glacial modellers' module list
- ⊗ As ice sheet models GCMize, there is a new need for top-down co-ordination in the glacial modelling world
- ⊗ Glaciologists need help with this
 - data to constrain, process understanding

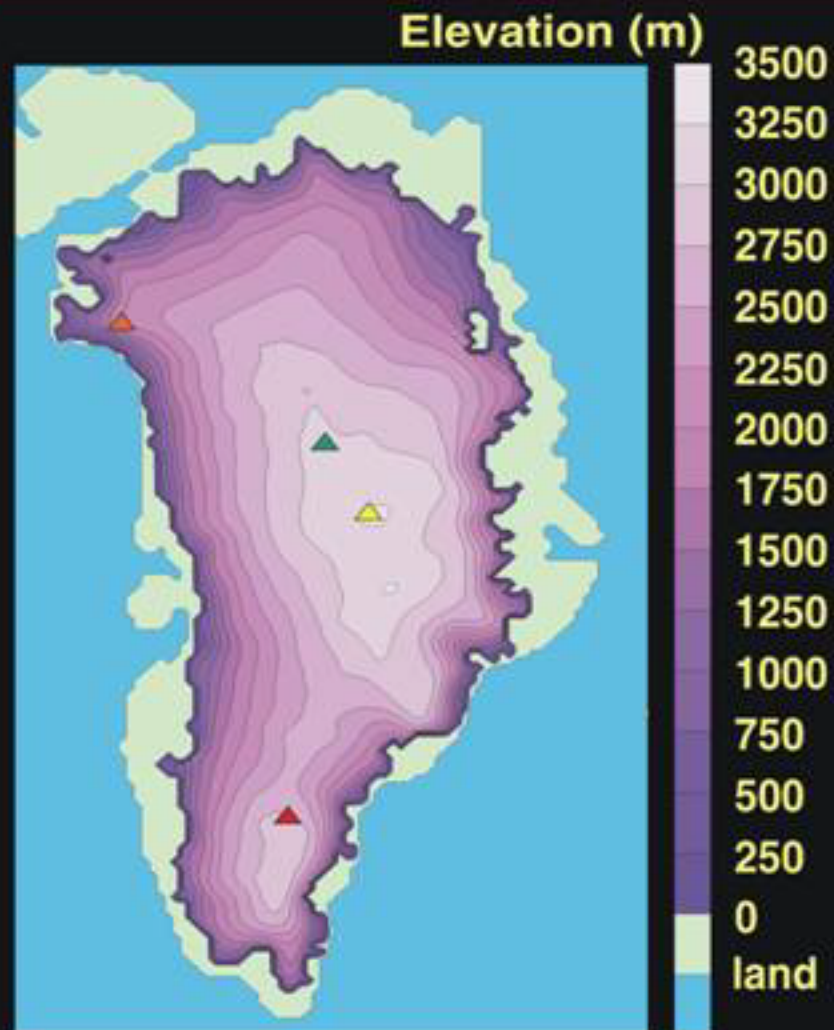
1. Basis of Ice Sheet Modelling

- Continuum-mechanical models of ice deformation under gravity
- Several 3D models in the world; resolve 3D velocity, temperature, stress fields as well as ice sheet thickness (t)
- Typically FD or FE, 5–100 km resolution
- Other approaches: flowline or planform models that permit higher resolution and in some cases, higher-order dynamics

2. Current State of the Art – Strengths

- General success with large-scale ice area and volume (Greenland, Antarctica).
- Good on time evolution, e.g. glacial cycle simulations. Possible to run for O (Ma).
- Well-integrated with climate and isostatic models.
- Considerable experience; intercomparisons and benchmarks established.

Greenland Deep Ice Cores



▲ Summit (GRIP, GISP2)

▲ NGRIP

▲ Camp Century

▲ Dye 3