

# The Role of High-Concentration Suspensions in Dispersal of River Sediments

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*with lots of help along the way...*

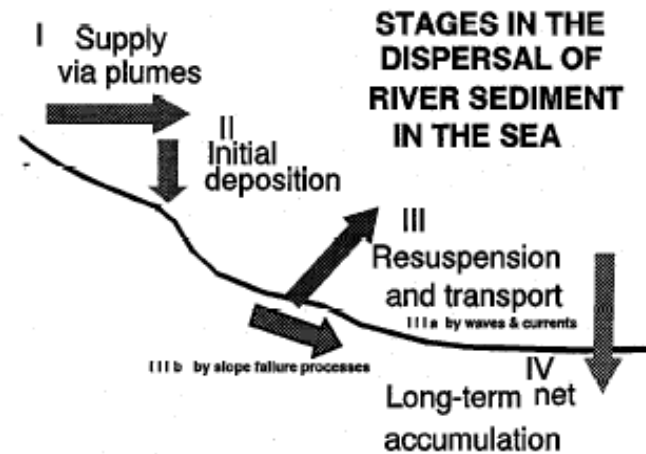


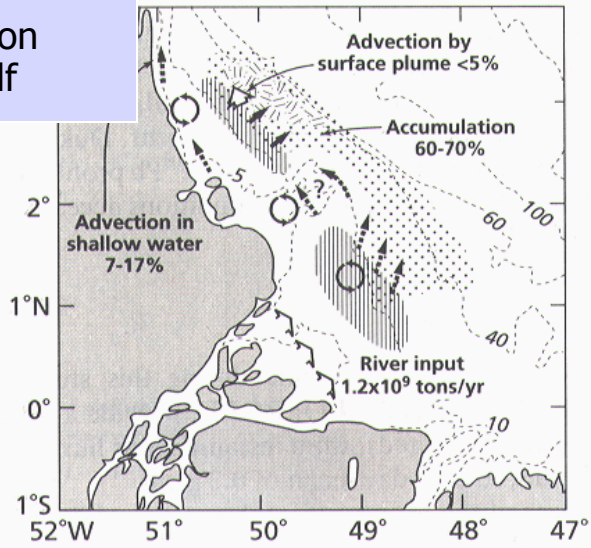
Fig. 2. Conceptual illustration of four major stages in the dispersal of river sediments in the coastal ocean.

Fine Sediment Dispersal in the Coastal Ocean,  
Wright and Nittrouer 1995

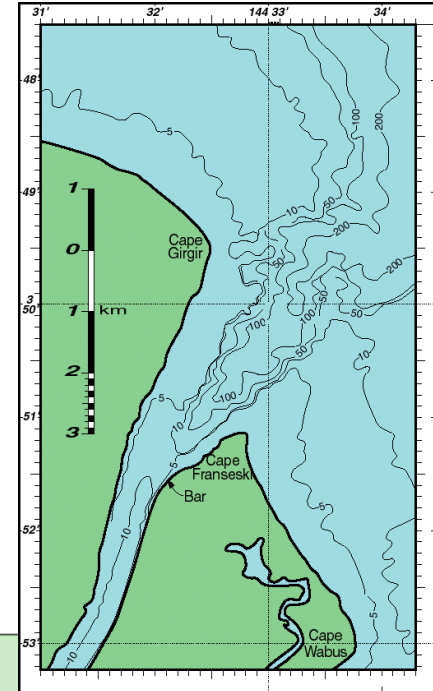
## ***LINKAGES BETWEEN FLUVIAL DISCHARGE AND HIGH-CONCENTRATION FLOWS IN COASTAL ENVIRONMENTS.***

1. Examples where there is a link between high-concentration flows in coastal environments.
2. *Conditions to form high concentration suspensions.*
3. A counter-intuitive example.

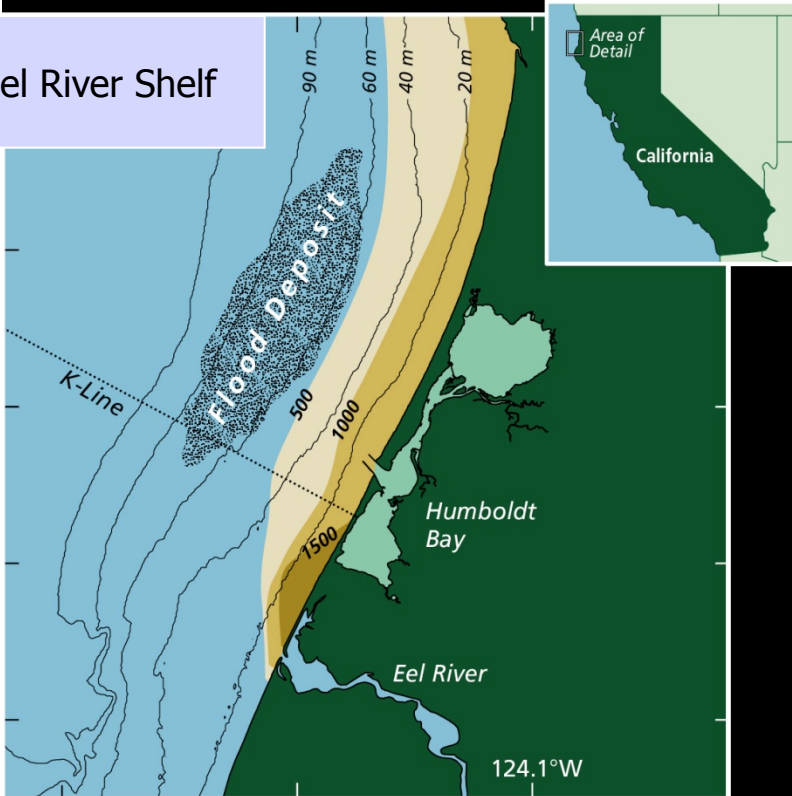
# Amazon Shelf



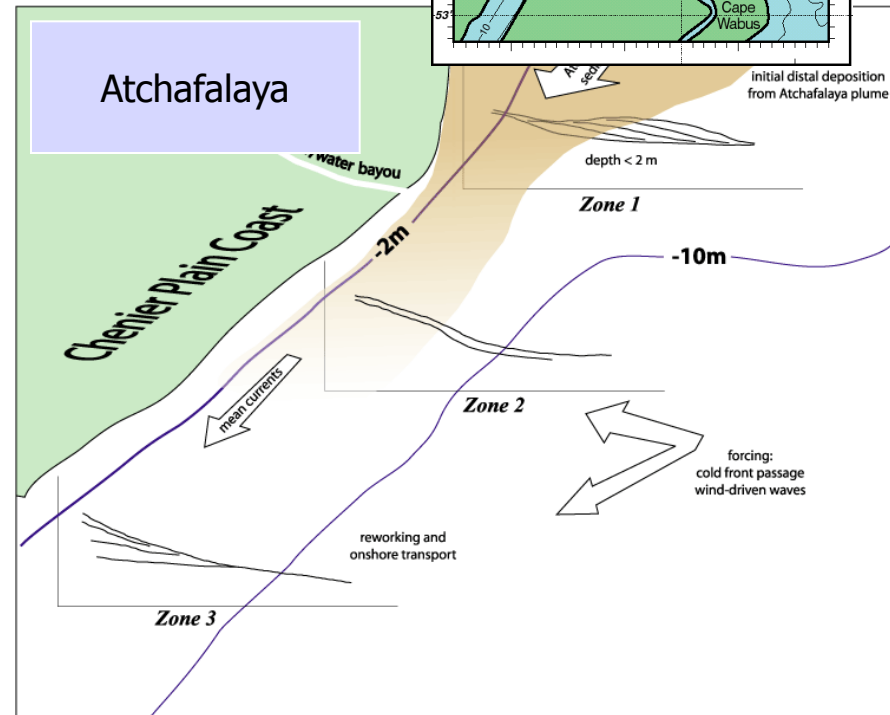
# Sepik River



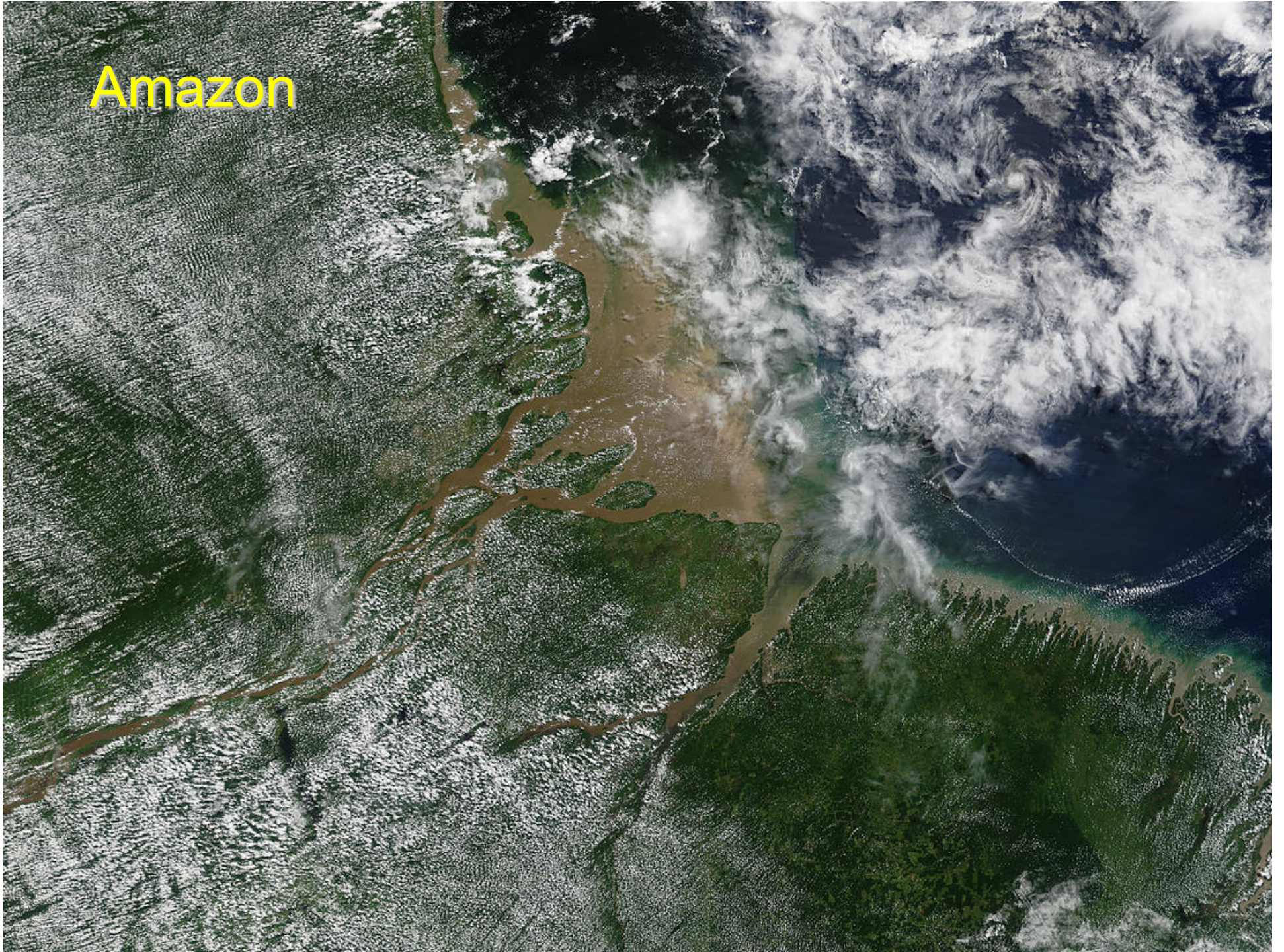
# Eel River Shelf



# Atchafalaya

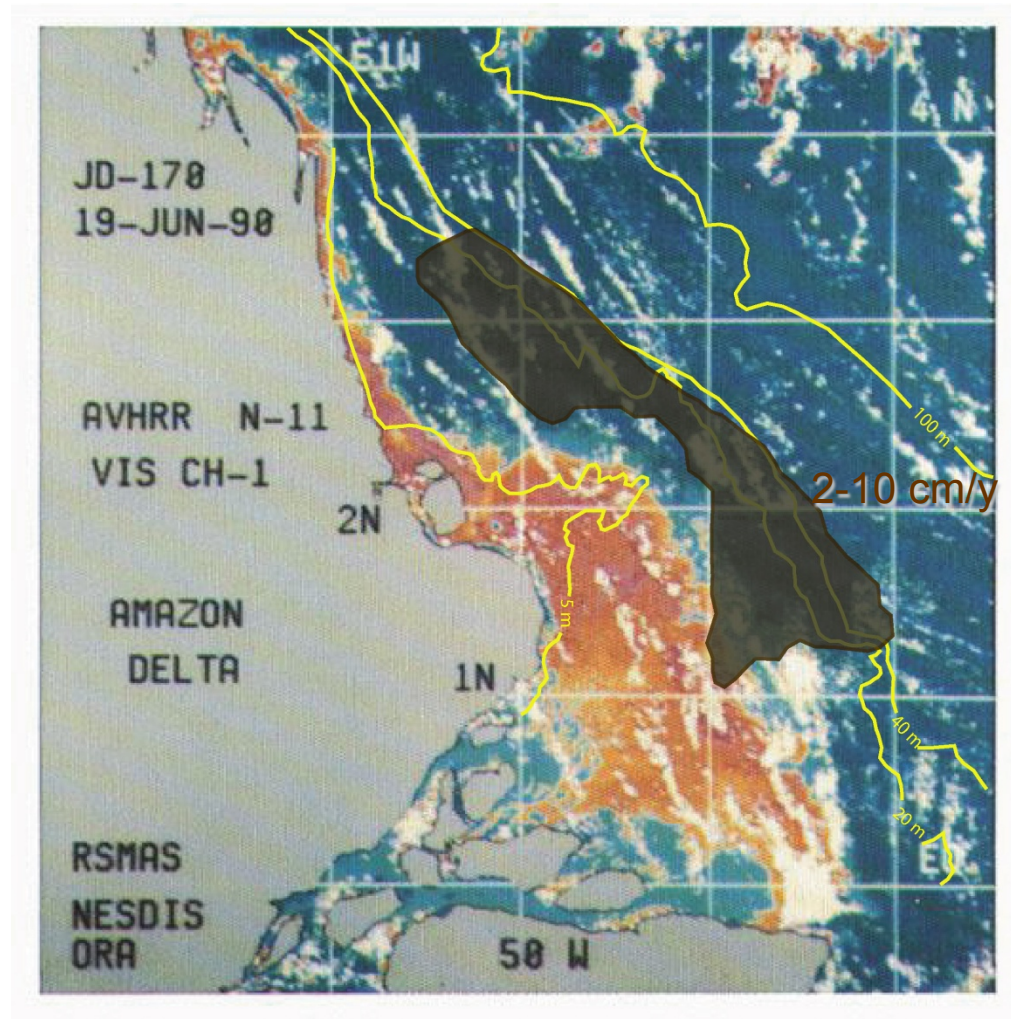


Amazon

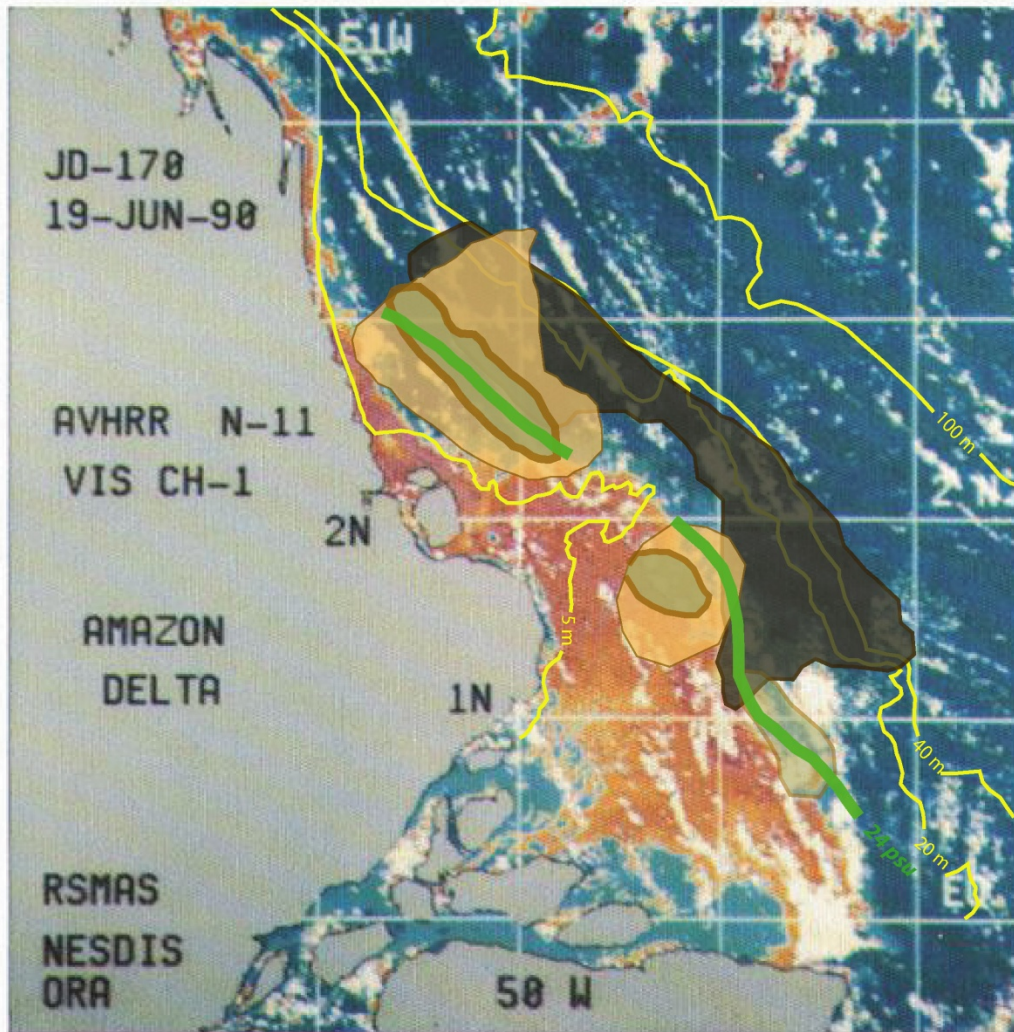


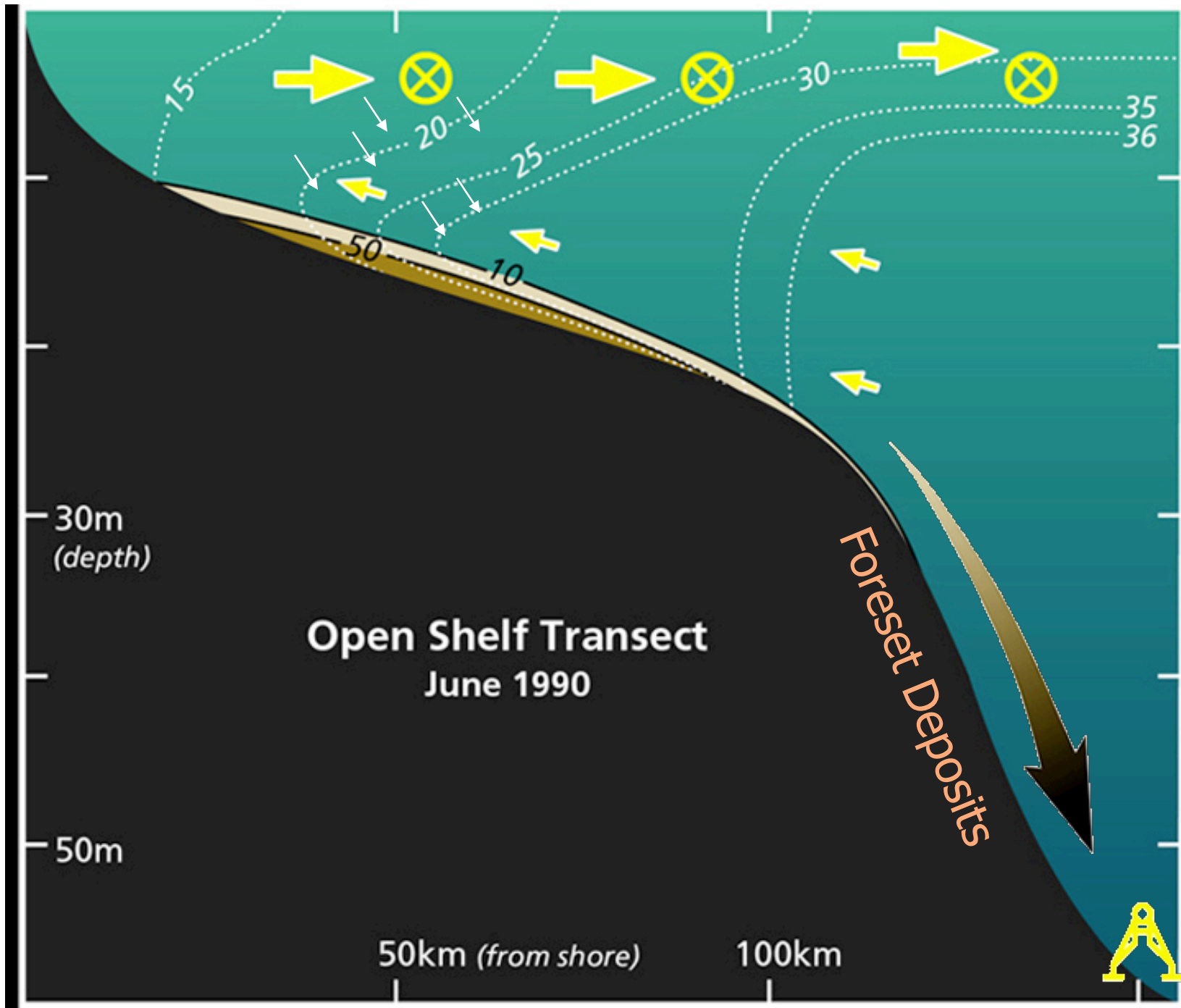
# AmasSeds

## A Multidisciplinary Amazon Shelf SEDiment Study

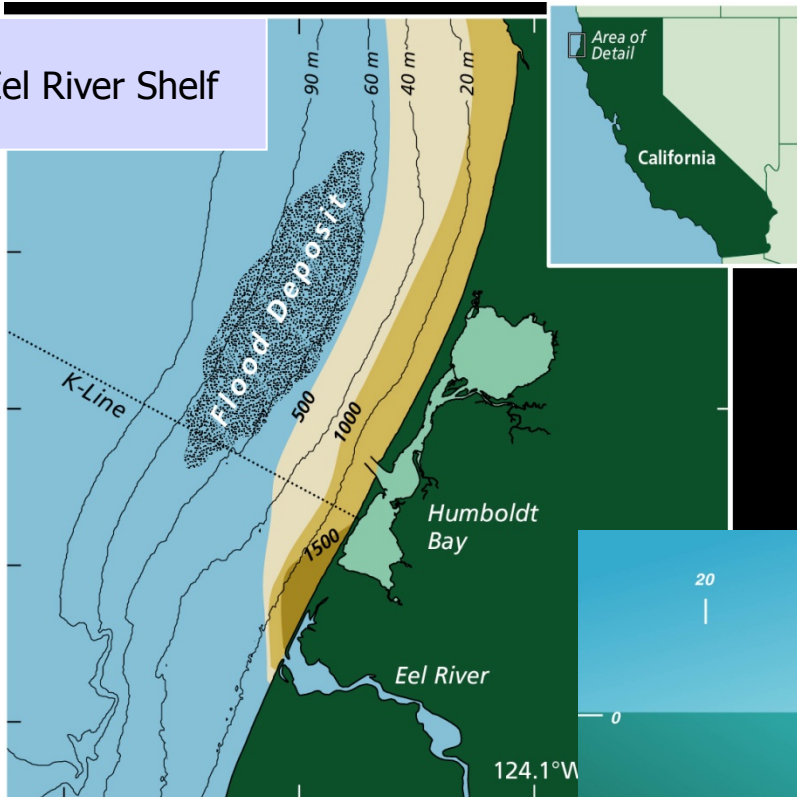


Kuehl & Nittrouer

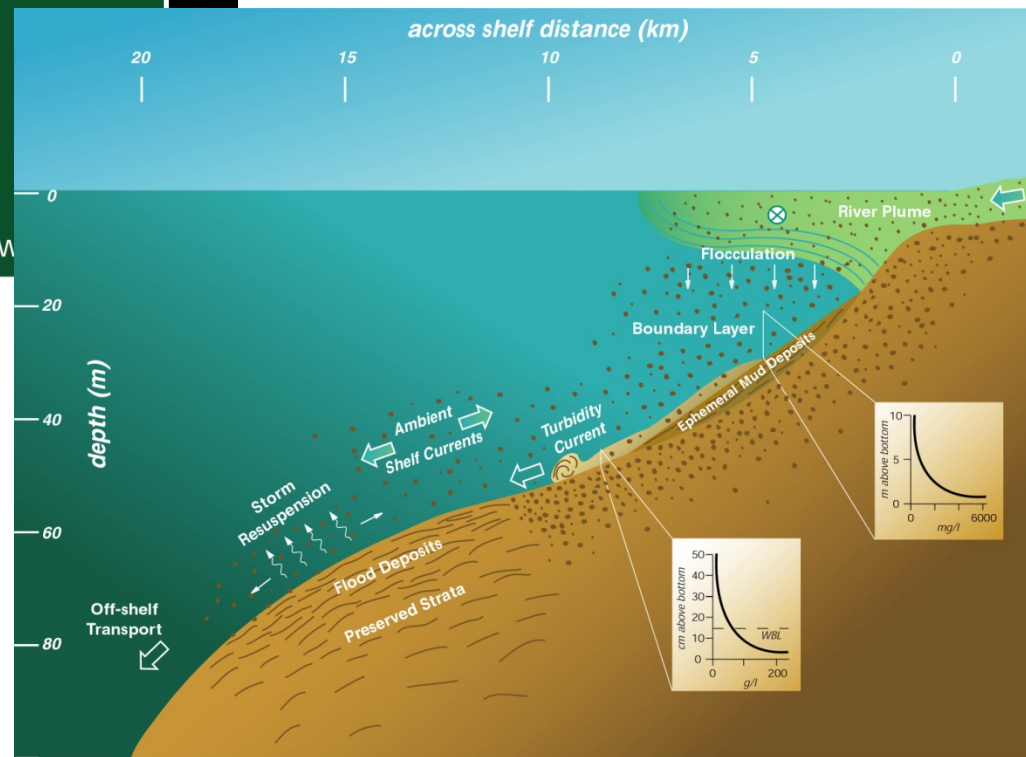




# Eel River Shelf

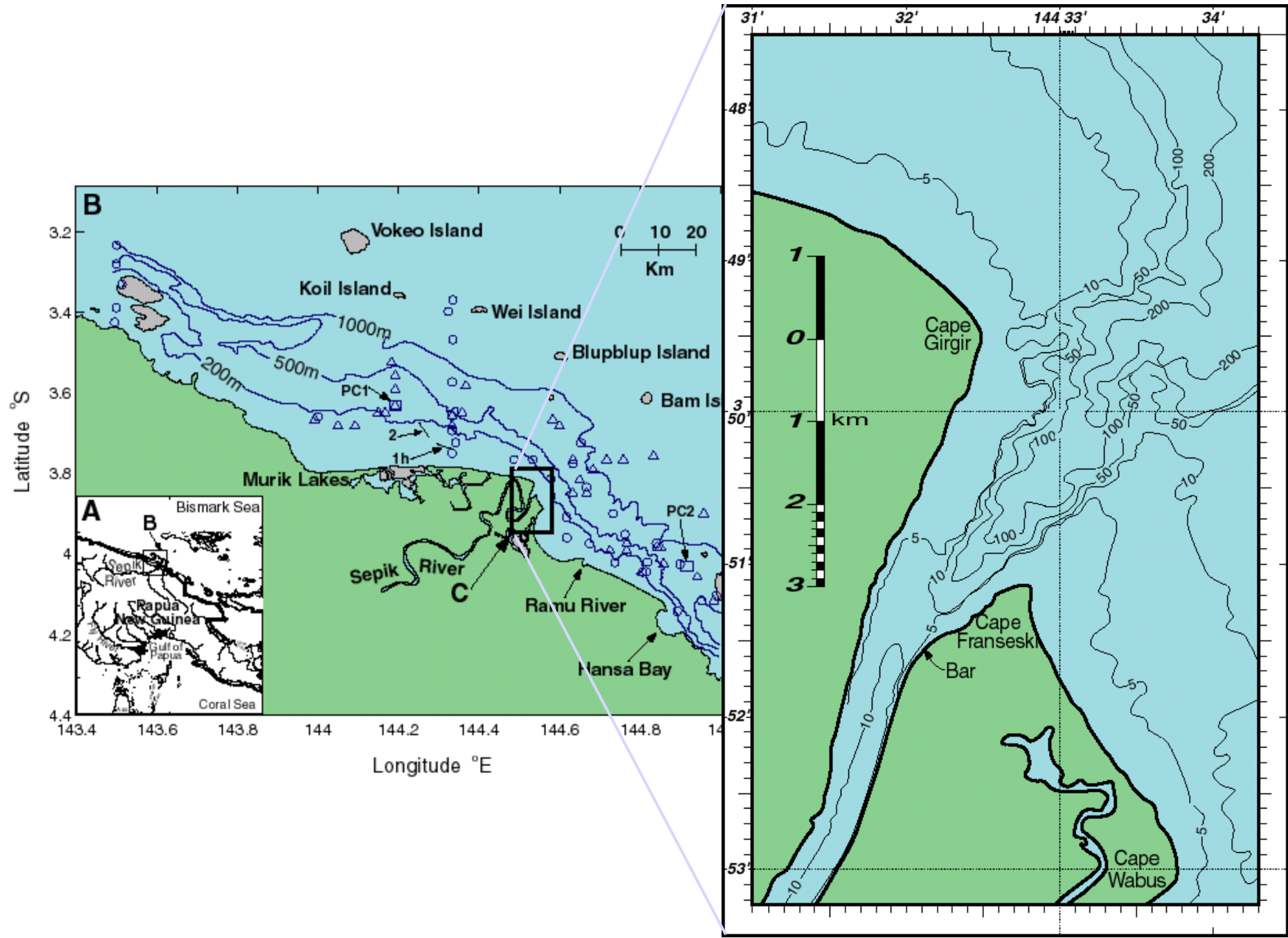


# Wave Supported Gravity Flows

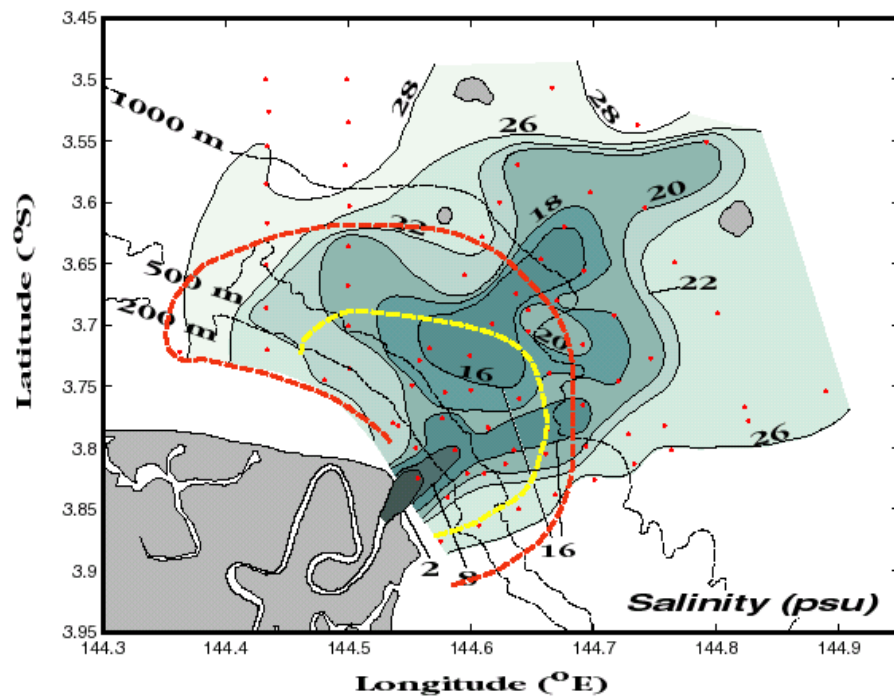
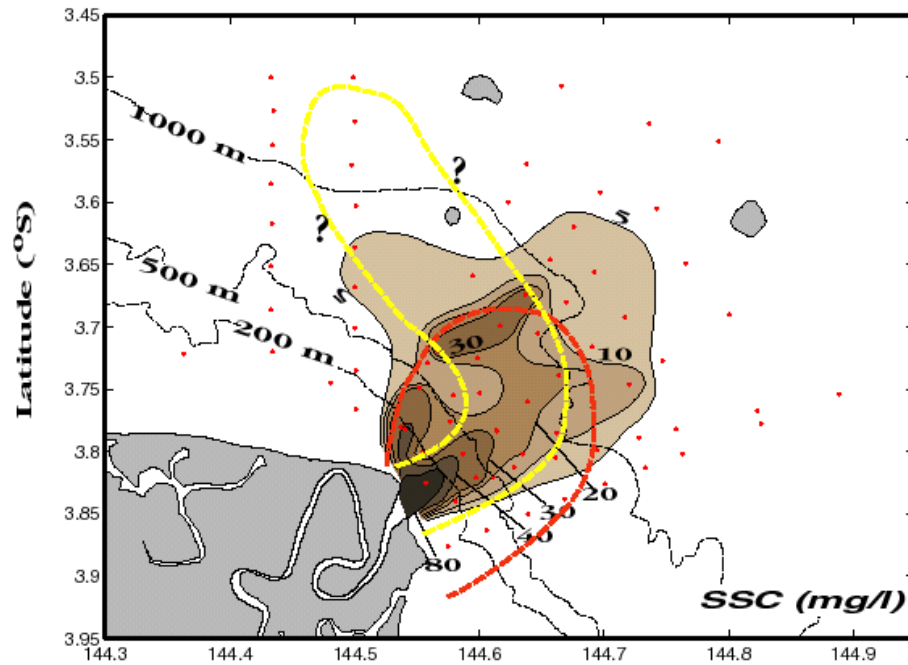


Traykovski, Geyer, Ogston





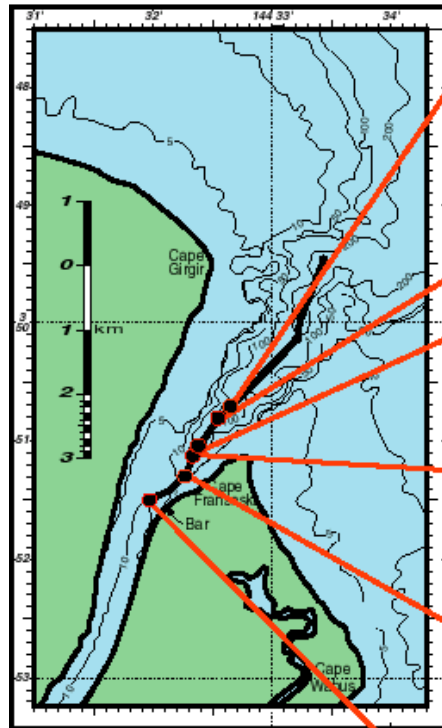
# Surface Plume



## Estimates of sediment flux in surface plume:

- Surface flux exiting river mouth is ~14% input at the bar (integrated over width)

- Implies  $w_s$  is  $O(1 \text{ cm s}^{-1})$  ( $0.3\text{-}2.2 \text{ cm s}^{-1}$ )



→  $Q_s = 0.9 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 0.75 \text{ m s}^{-1}$   
 $C_s = 0.25 \text{ g l}^{-1}$   
 $h' = 4.8 \text{ m}$

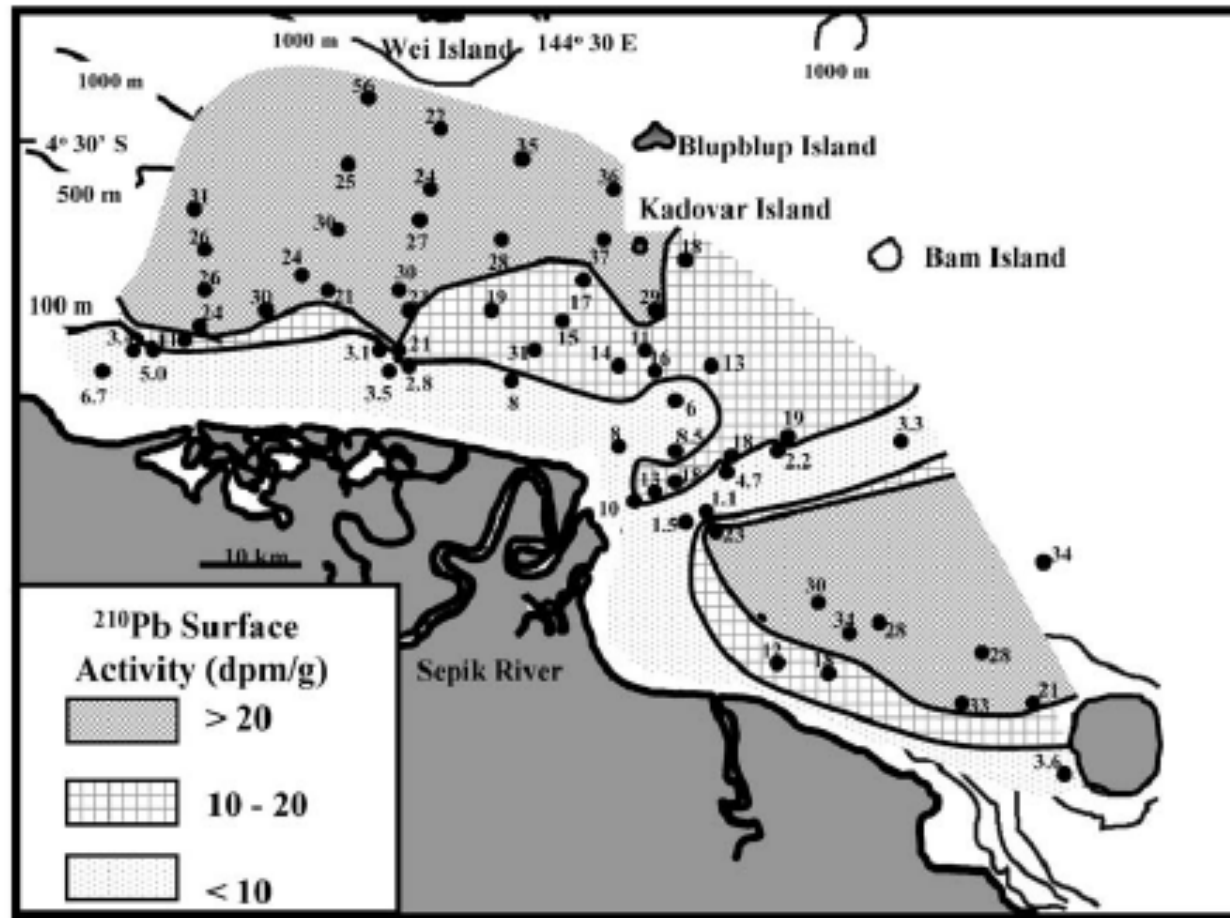
→  $Q_s = 0.4 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 0.47 \text{ m s}^{-1}$   
 $C_s = 0.17 \text{ g l}^{-1}$   
 $h' = 5.0 \text{ m}$

→  $Q_s = 1.1 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 0.80 \text{ m s}^{-1}$   
 $C_s = 0.26 \text{ g l}^{-1}$   
 $h' = 5.5 \text{ m}$

→  $Q_s = 3.1 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 1.26 \text{ m s}^{-1}$   
 $C_s = 0.40 \text{ g l}^{-1}$   
 $h' = 6.0 \text{ m}$

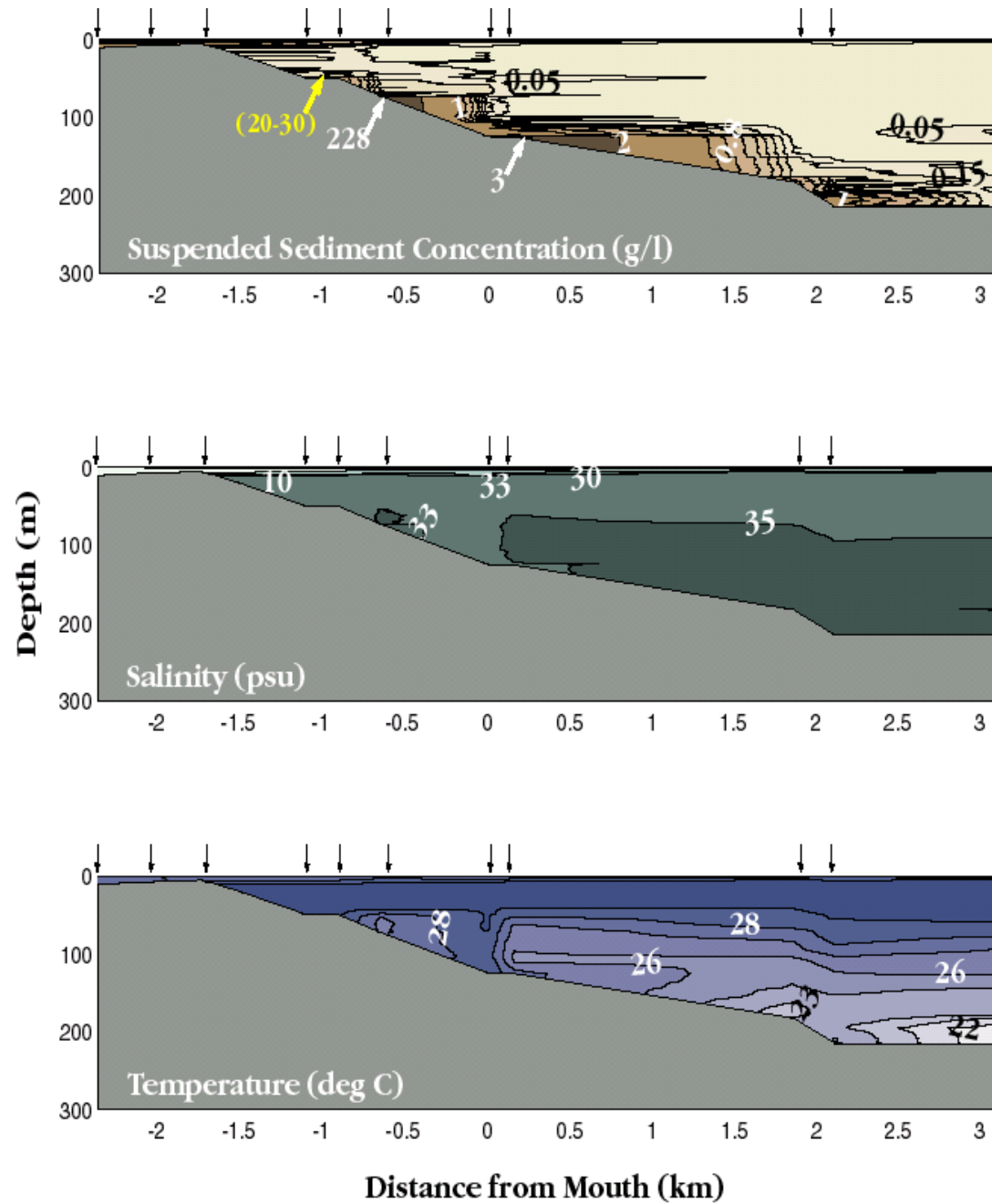
→  $Q_s = 3.7 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 1.37 \text{ m s}^{-1}$   
 $C_s = 0.42 \text{ g l}^{-1}$   
 $h' = 6.5 \text{ m}$

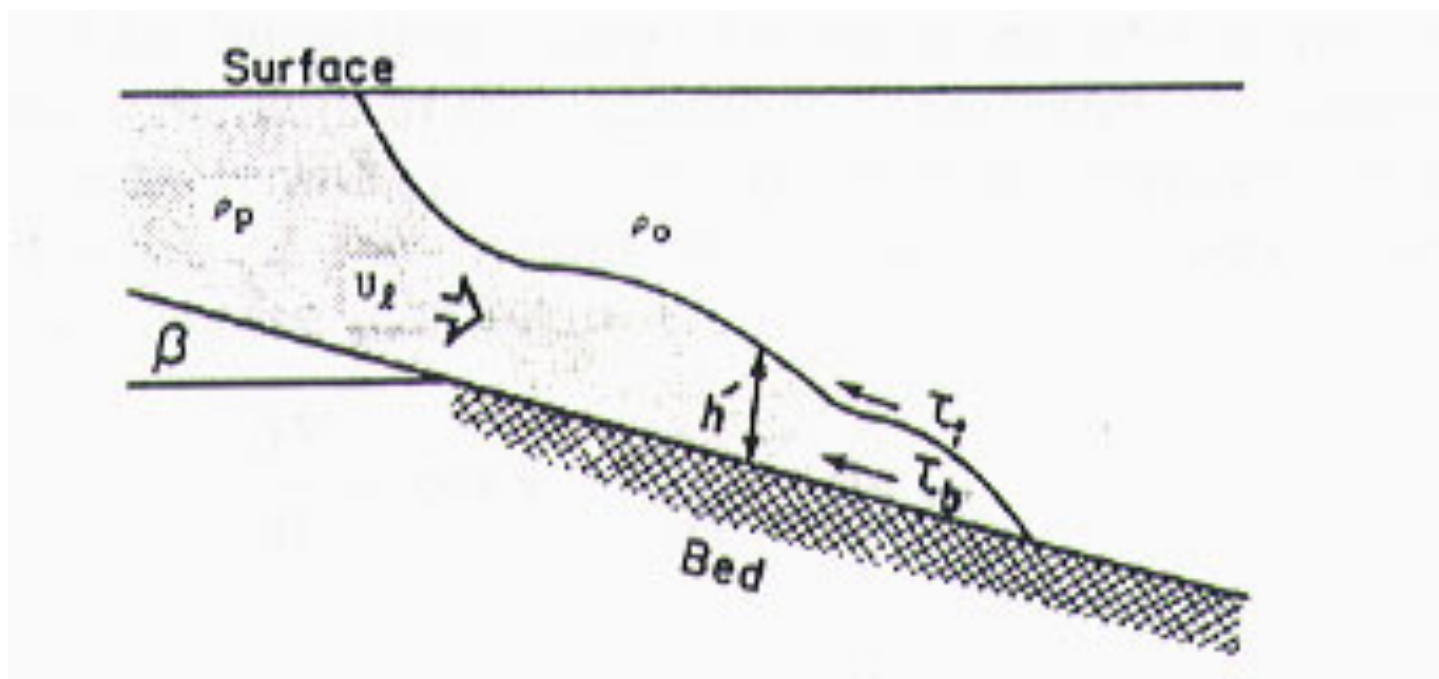
→  $Q_s = 6.2 \text{ kg s}^{-1} \text{ m}^{-1}$   
 $U_s = 1.47 \text{ m s}^{-1}$   
 $C_s = 0.60 \text{ g l}^{-1}$   
 $h' = 7.0 \text{ m}$



Kuehl et al. 2004

April  
1999





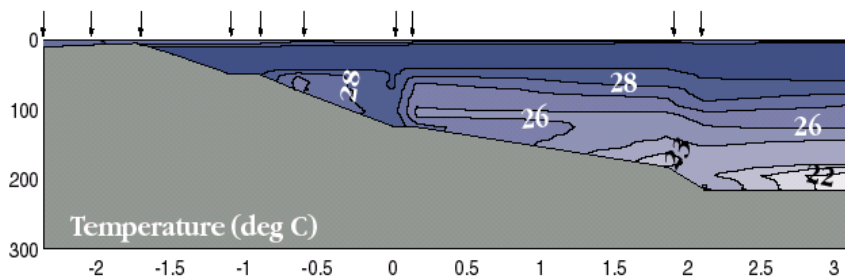
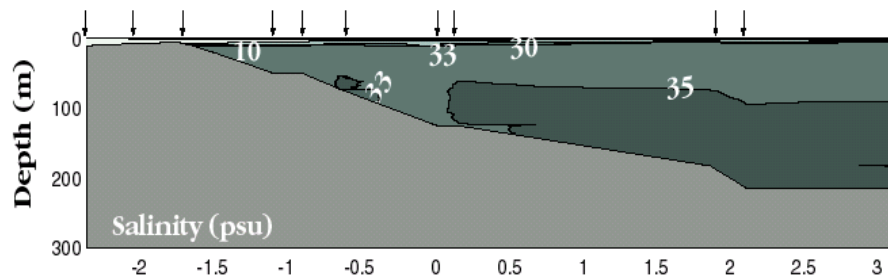
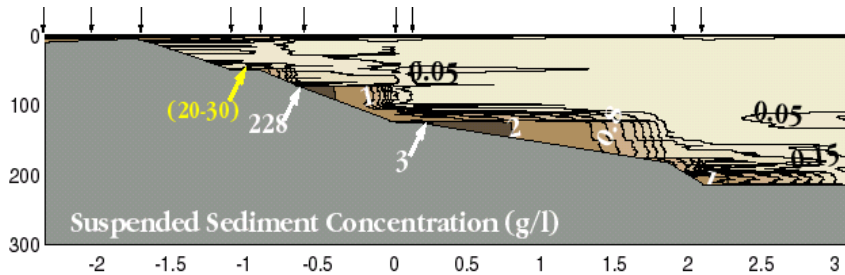
Chezy Equation:

$$h' \rho_p g' \sin \beta = \tau_b + \tau_t = \rho_p u_l^2 (C_D + E)$$

Wright et al. 1990

$$h' \rho_p g' \sin \beta = \tau_b + \tau_t = \rho_p u_l^2 (C_D + E)$$

Wright et al. 1990



Distance from Mouth (km)

How thick for 10g/l?

~ 20 cm

How long to create?

~ 5.6 hrs

How fast?

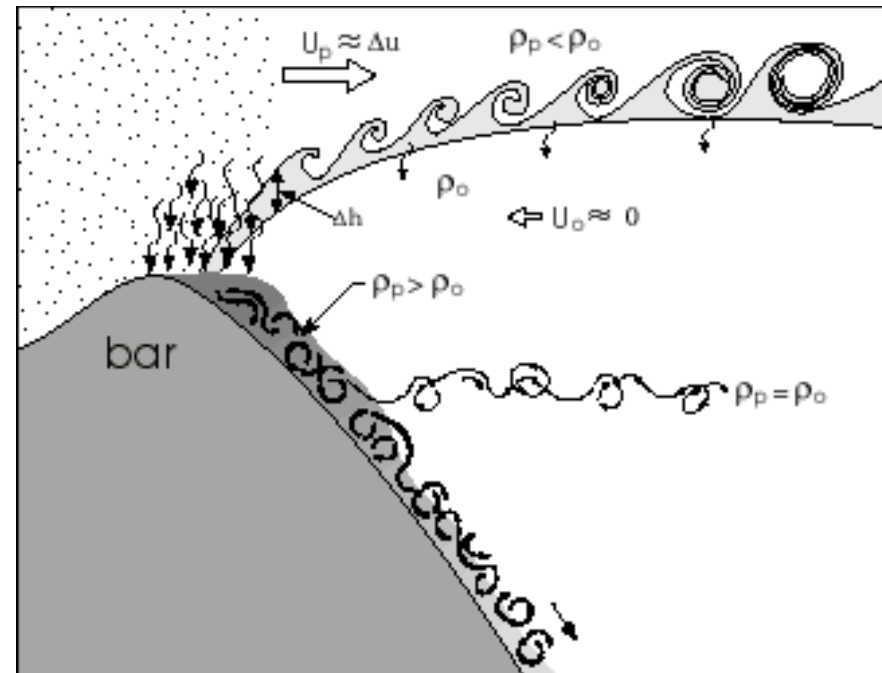
~ 10 cm/s

For 1 g/l, 10 m?

~70 cm/s

# Sepik Summary

- ~15% of the river input of sediment is dispersed via the surface plume with a comparable amount in intermediate layers and the remainder in dense nearbottom flows
- Implications for widespread dispersal of particulates
- Rapid delivery to deep sea





- For these examples, high concentration suspensions formed through settling.
- Requires rapid settling flux,  $w_s C$ , and inhibited mixing (stratification).

### Petitcodiac River experiments

Bay of Fundy area -big tides, strong currents, high concentrations, and a stable working platform

# Instrumentation

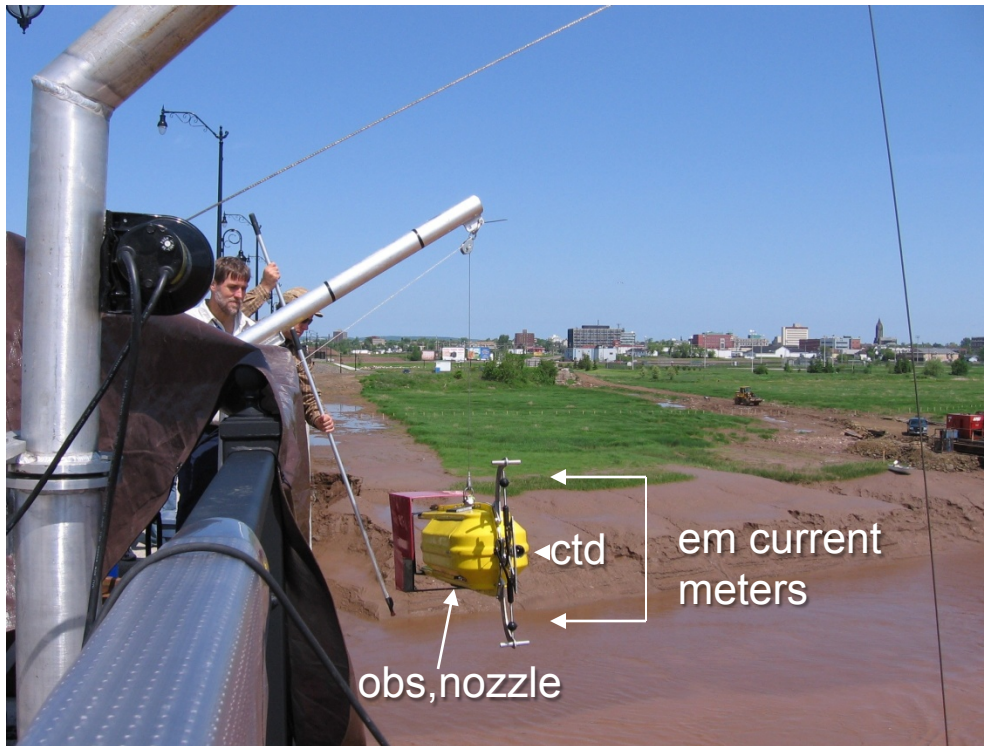
- SUBS profiler
- Surfboard (ADCP, Knudsen)
- ADV profiling tripod



# Instrumentation

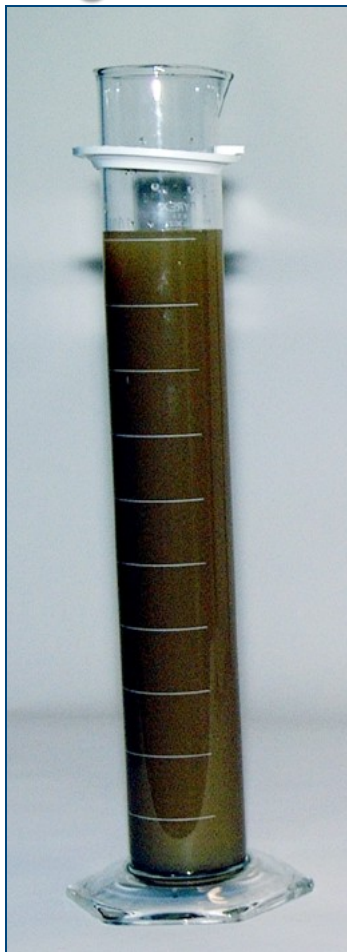
- SUBS profiler

$$\frac{du}{dz}, \rho_z(S, T, C), \text{samples}$$

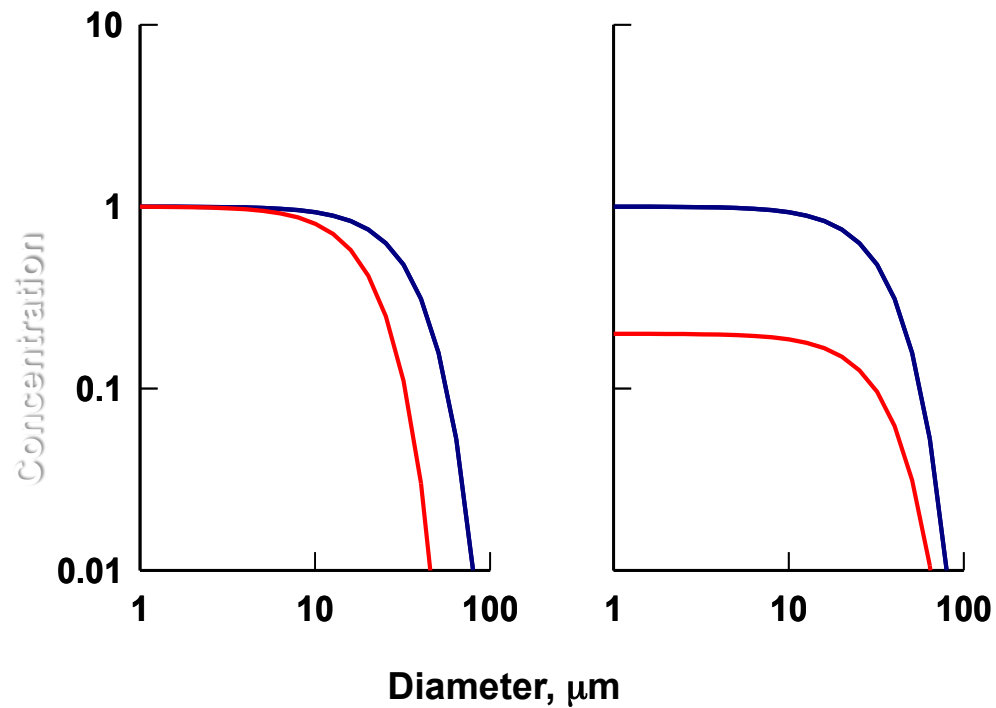


# Floc Settling Necessary for Fluid Mud Formation

Single Grains



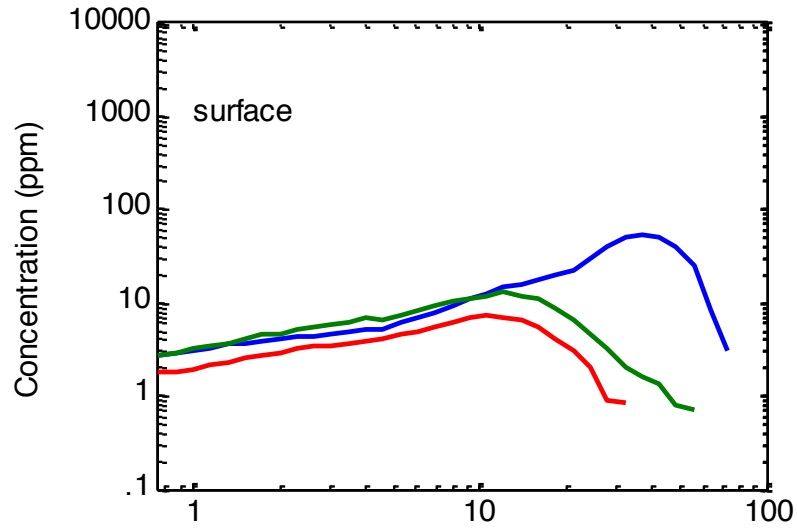
Flocs



Milligan, BIO

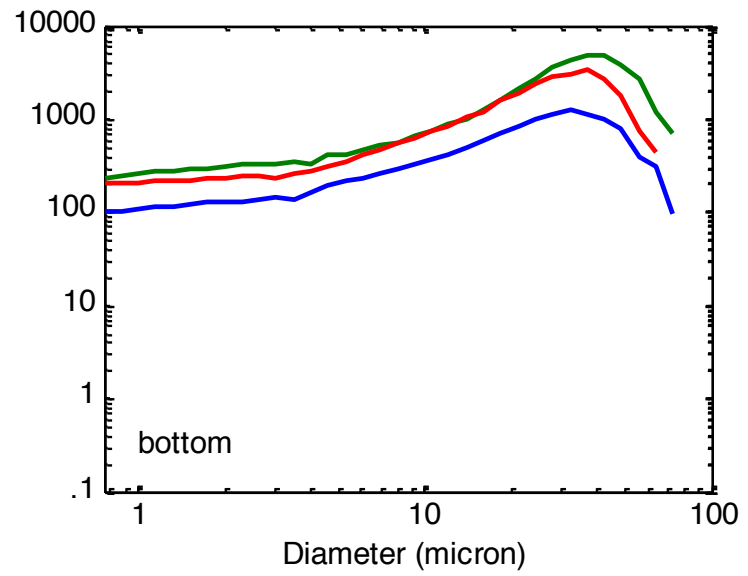
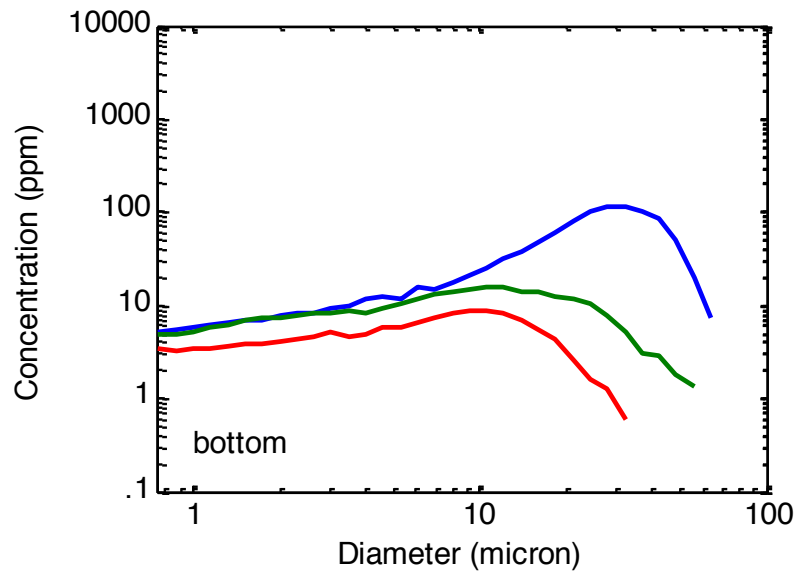
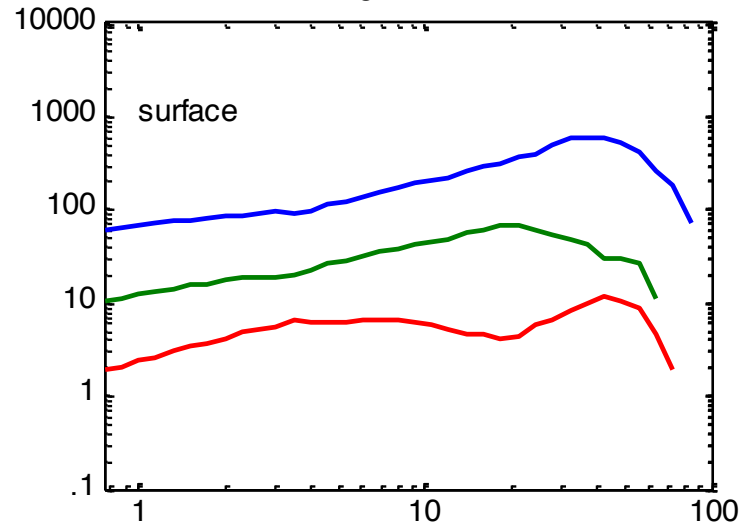
Fresh

June 2006

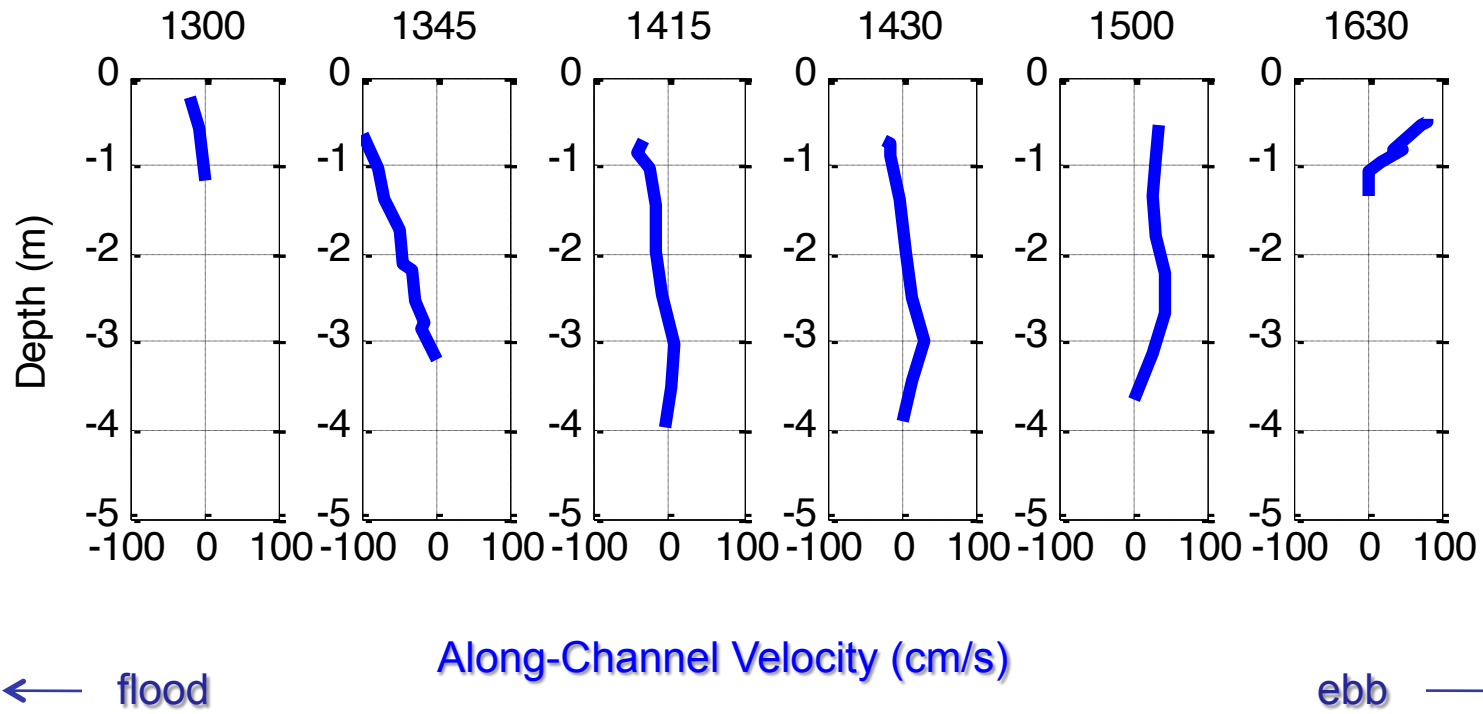
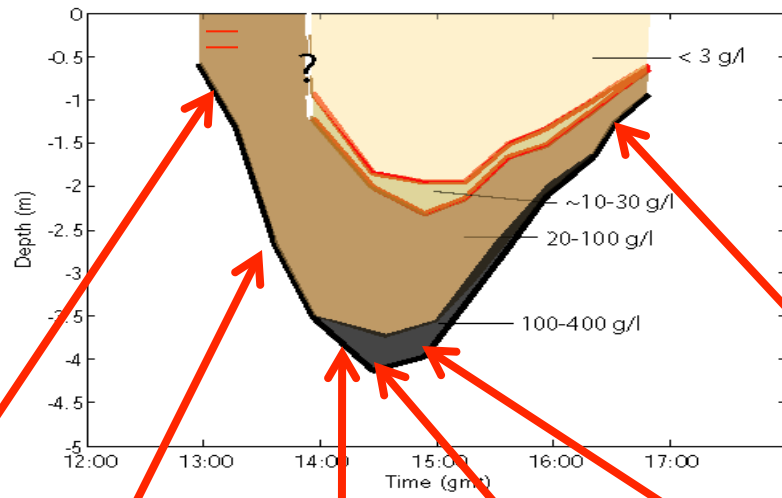


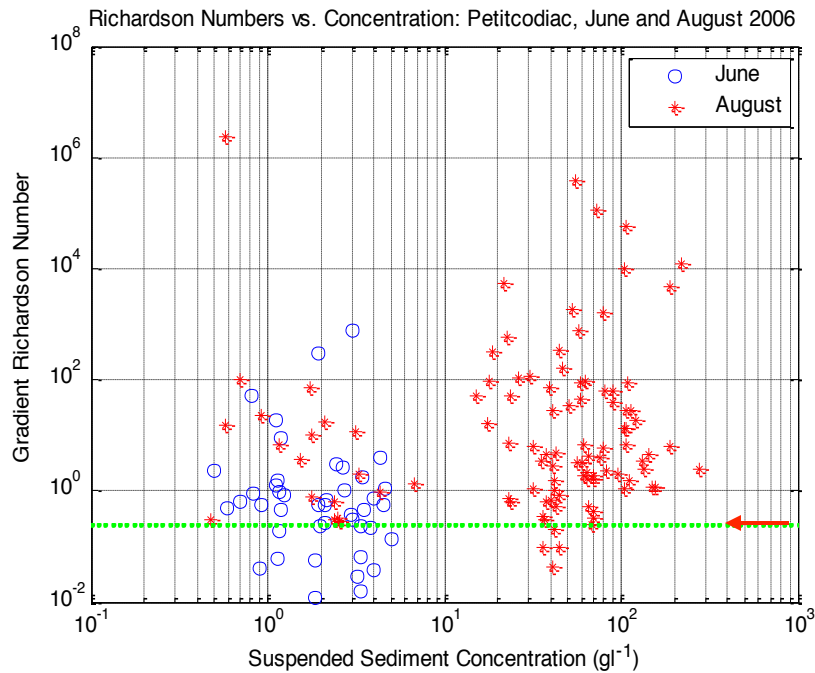
August 2006

Brackish



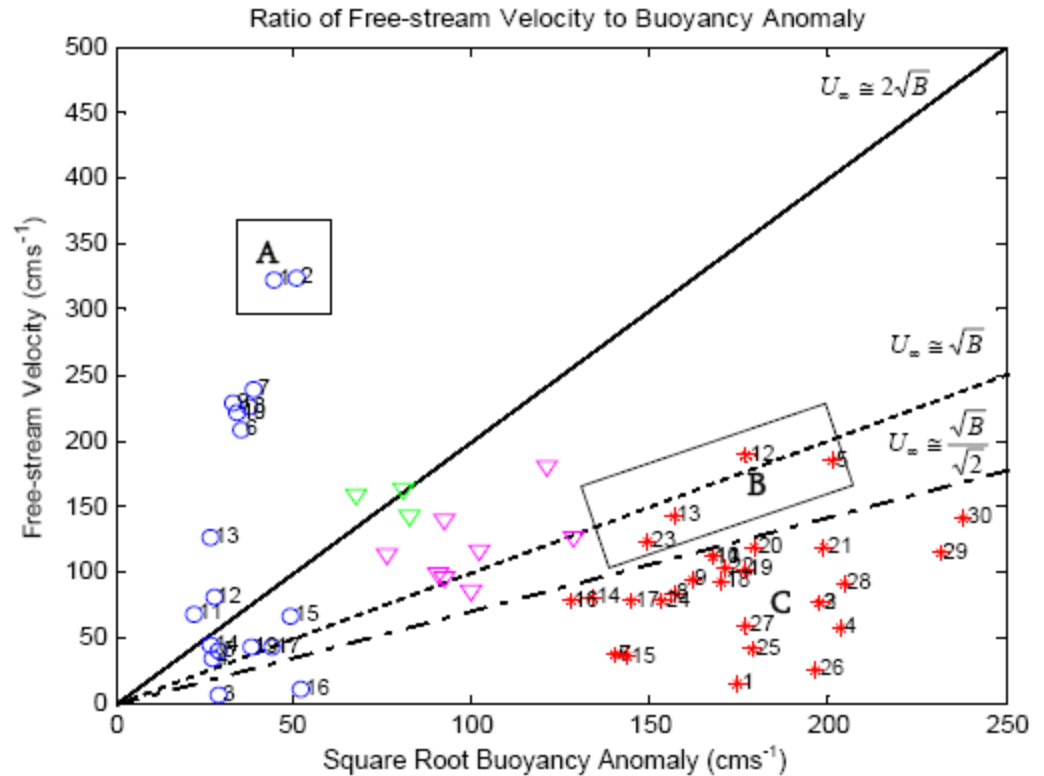
Milligan, BIO





## $Ri_g$ vs Suspended Sediment Concentration

Heath 2009



Carrying capacity threshold condition, dependence on  $Ri_c$

$$\frac{U_\infty}{\sqrt{B}} > \text{or} < 2?$$

# Petitcodiac Summary

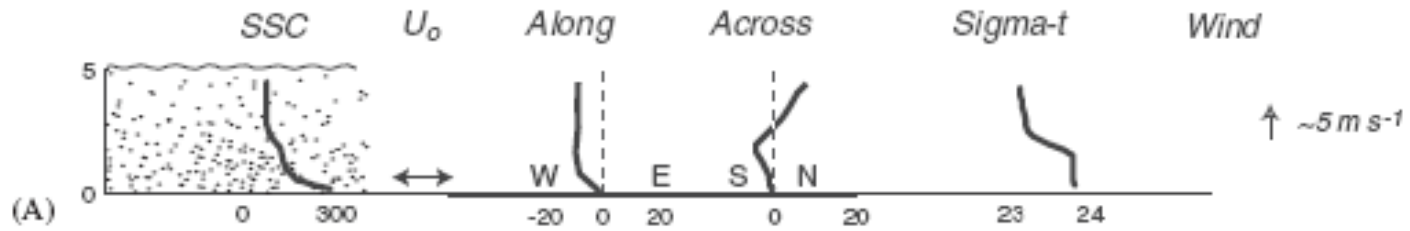
- Floc settling necessary for fluid mud formation
- $Ri_c$  dependence and carrying-capacity threshold noisy but reasonable



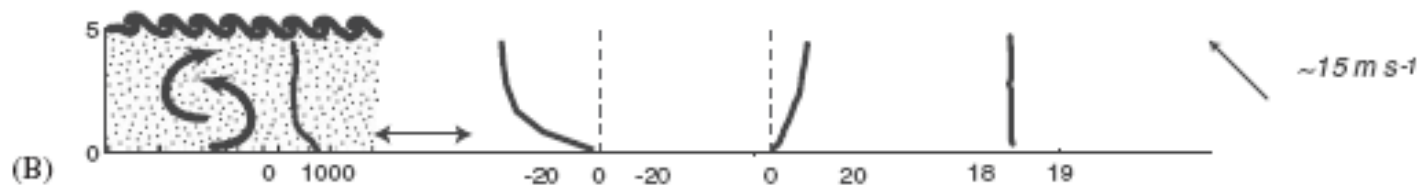
- Most of Louisiana shoreline is retreating, but a section in western LA is prograding
- Prior observations suggested role of cold fronts critical for onshore transport

Source: [http://visibleearth.nasa.gov/data/ev124/ev12470\\_Mississippi.A2002058.1650.1km.jpg](http://visibleearth.nasa.gov/data/ev124/ev12470_Mississippi.A2002058.1650.1km.jpg)

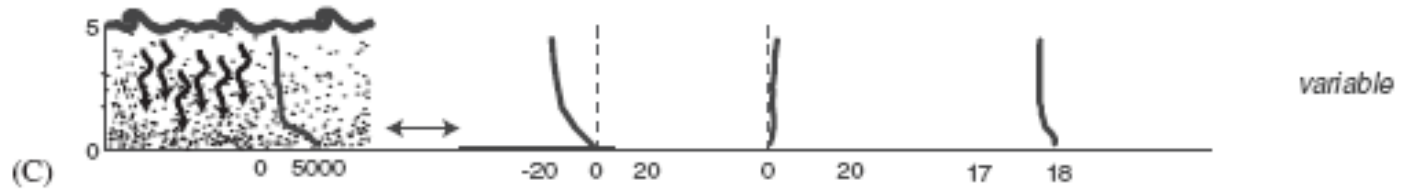
Low Energy



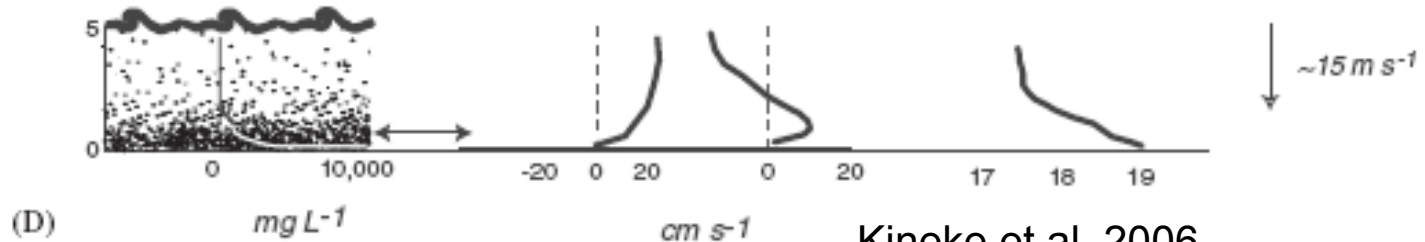
Pre-Front



Front

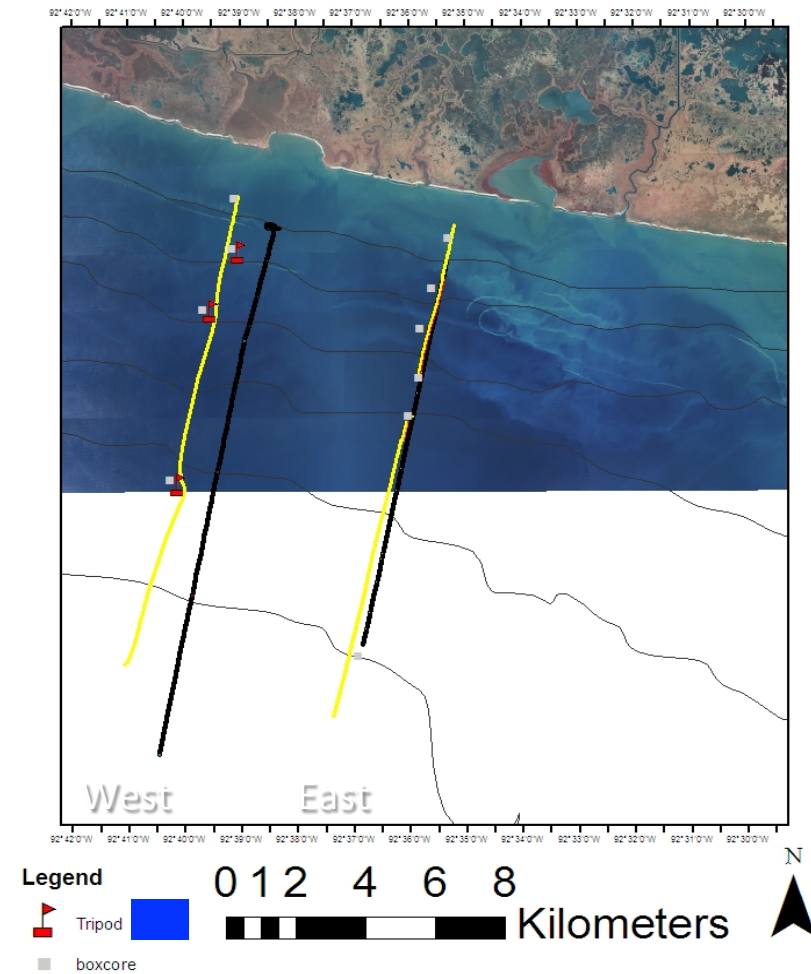
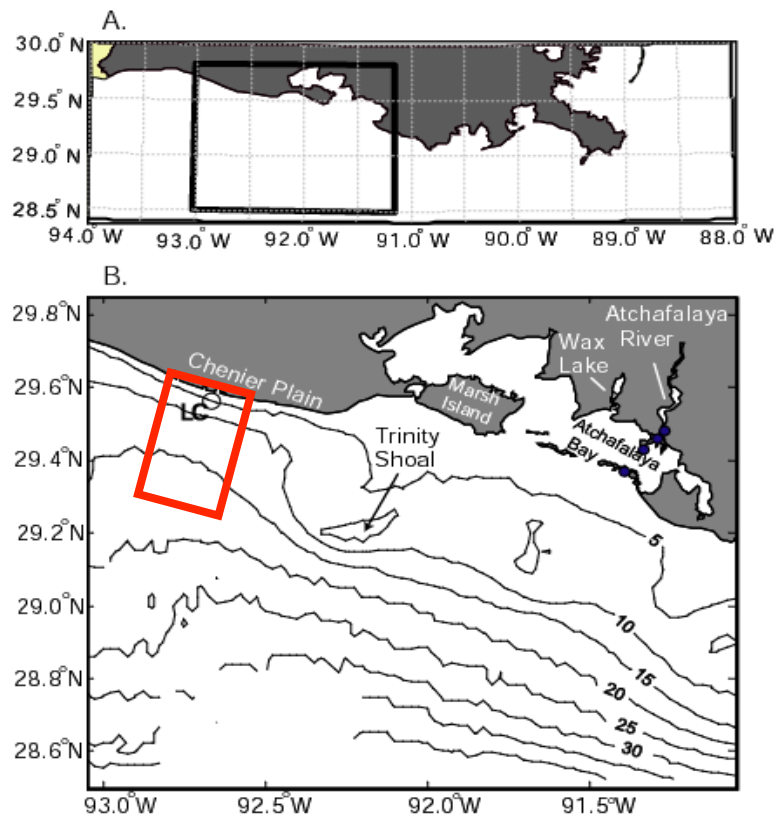


Post-Front



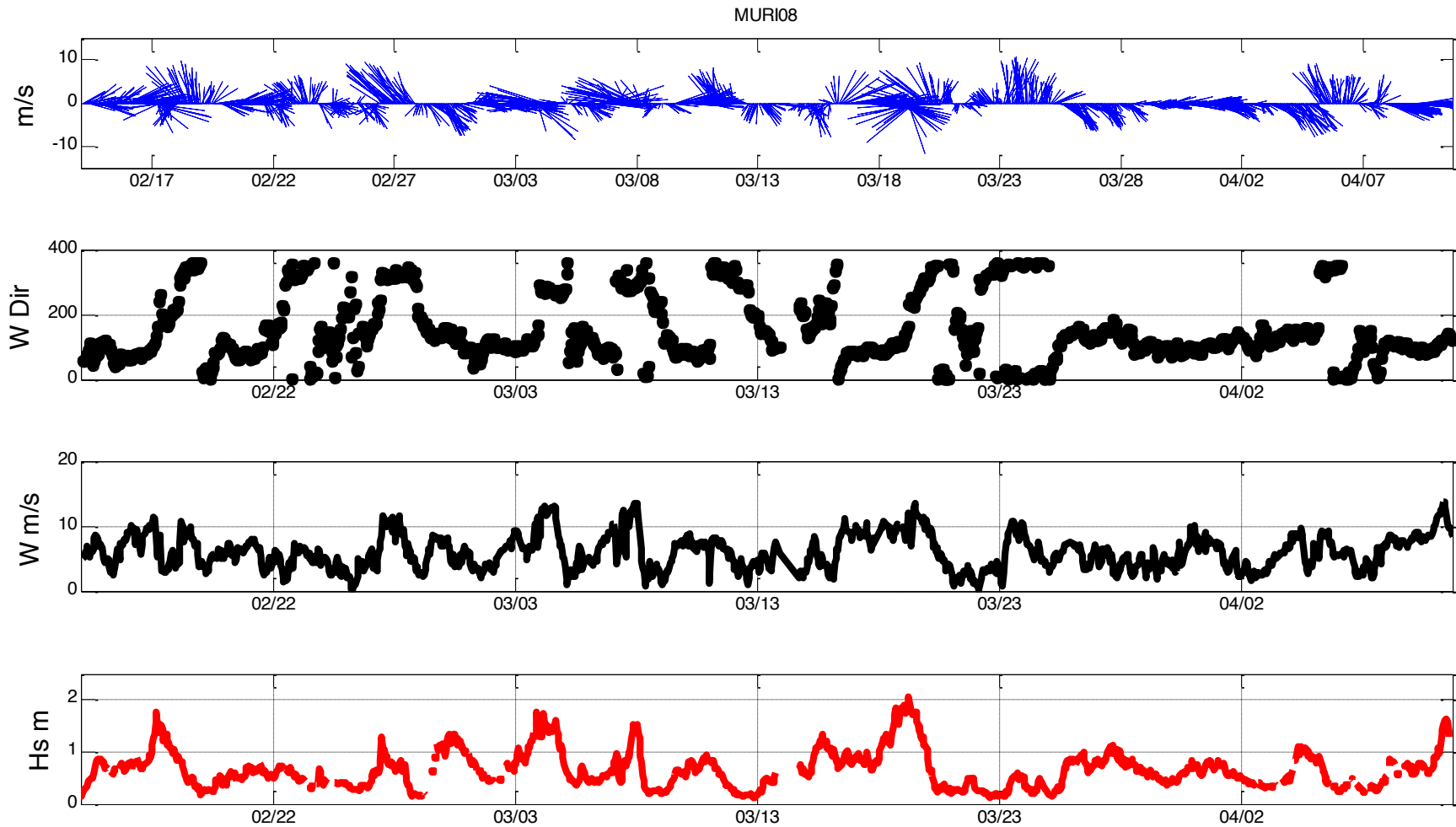
Kineke et al. 2006

# ONR sponsored MURI JHU, WHOI, BC, MU, MIT

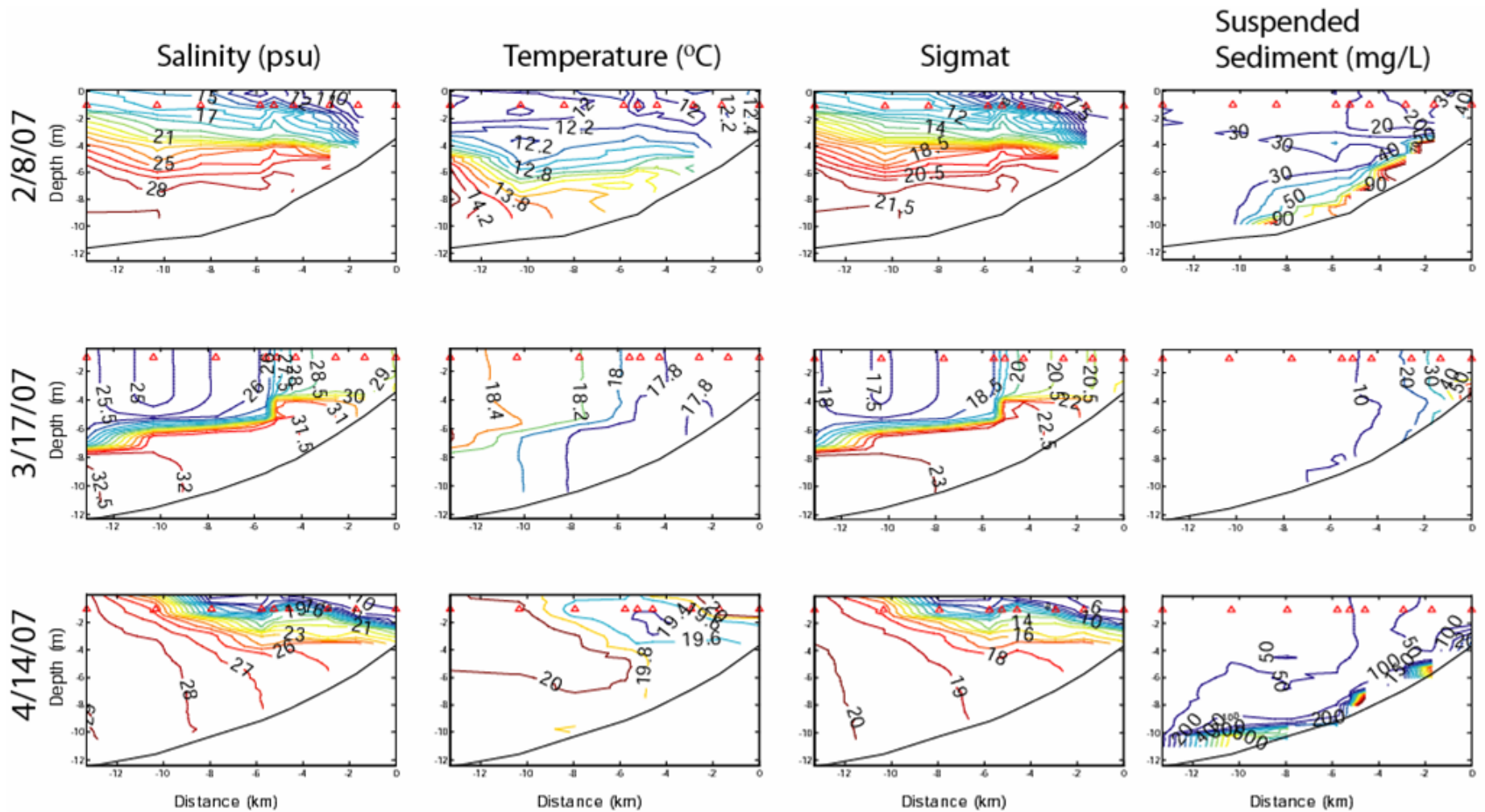


- Water column surveys, coring, bottom boundary layer tripods

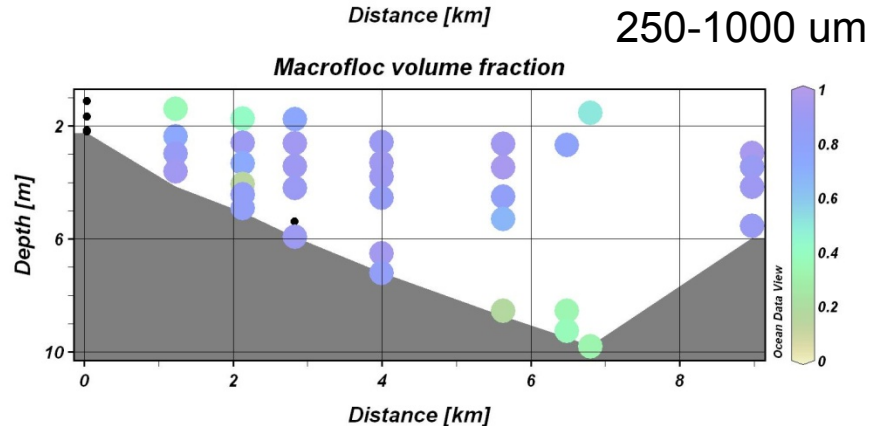
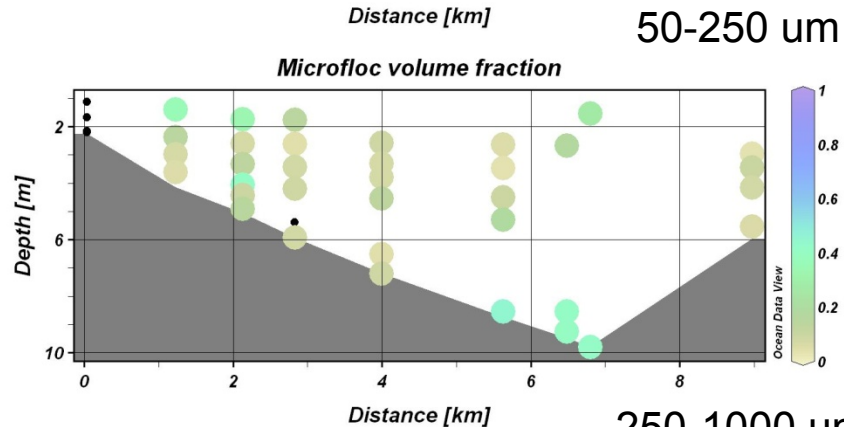
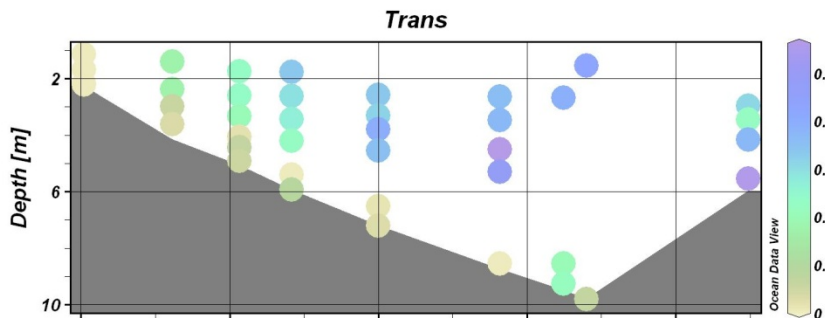
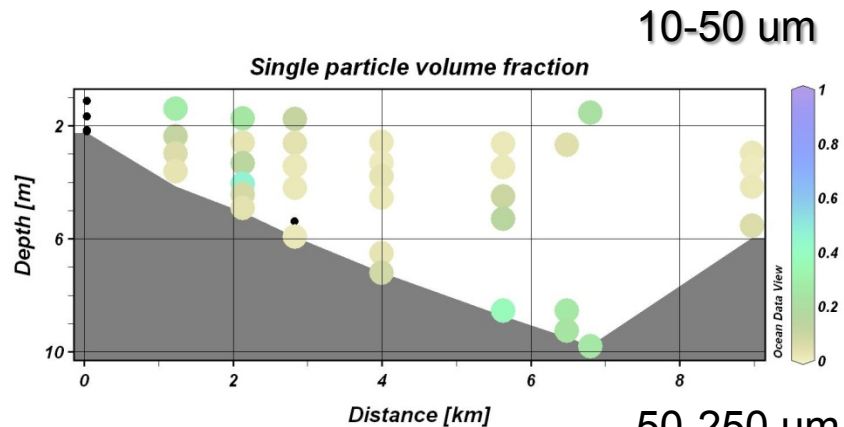
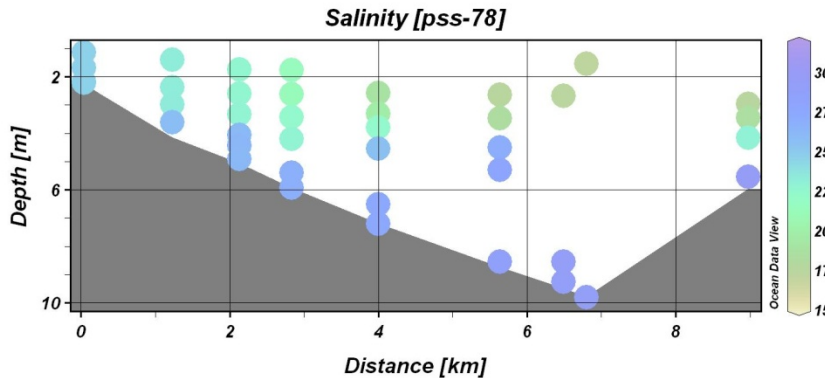
# Cold Fronts supply the energy – typical characteristics



# Eastern Transects 2007



LISST



March 2008 East  
(tripod line)

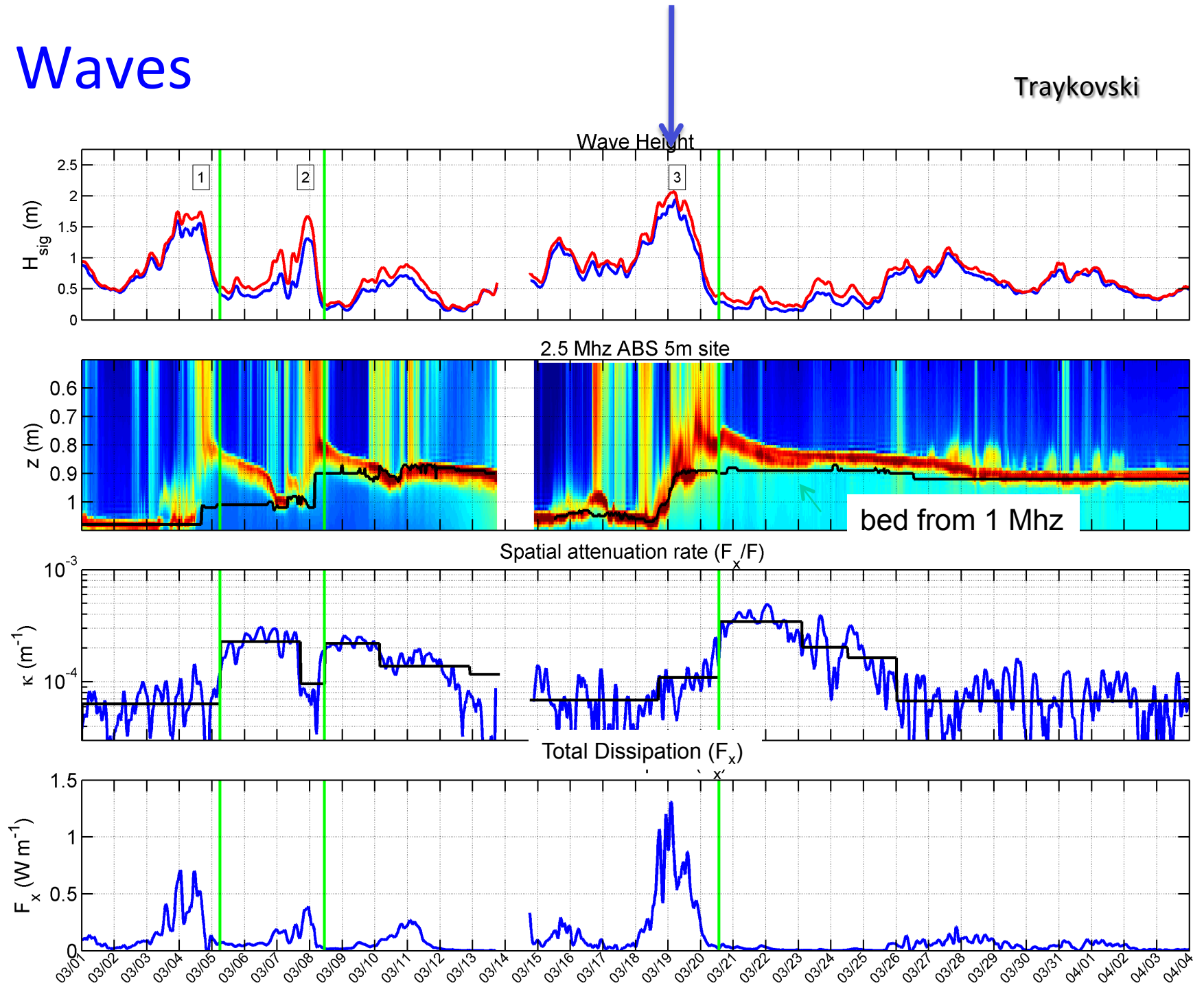
Mean particle size:  
generally 500 – 1000 um

**water can clear in < 1.5 hrs**

Meg Estapa

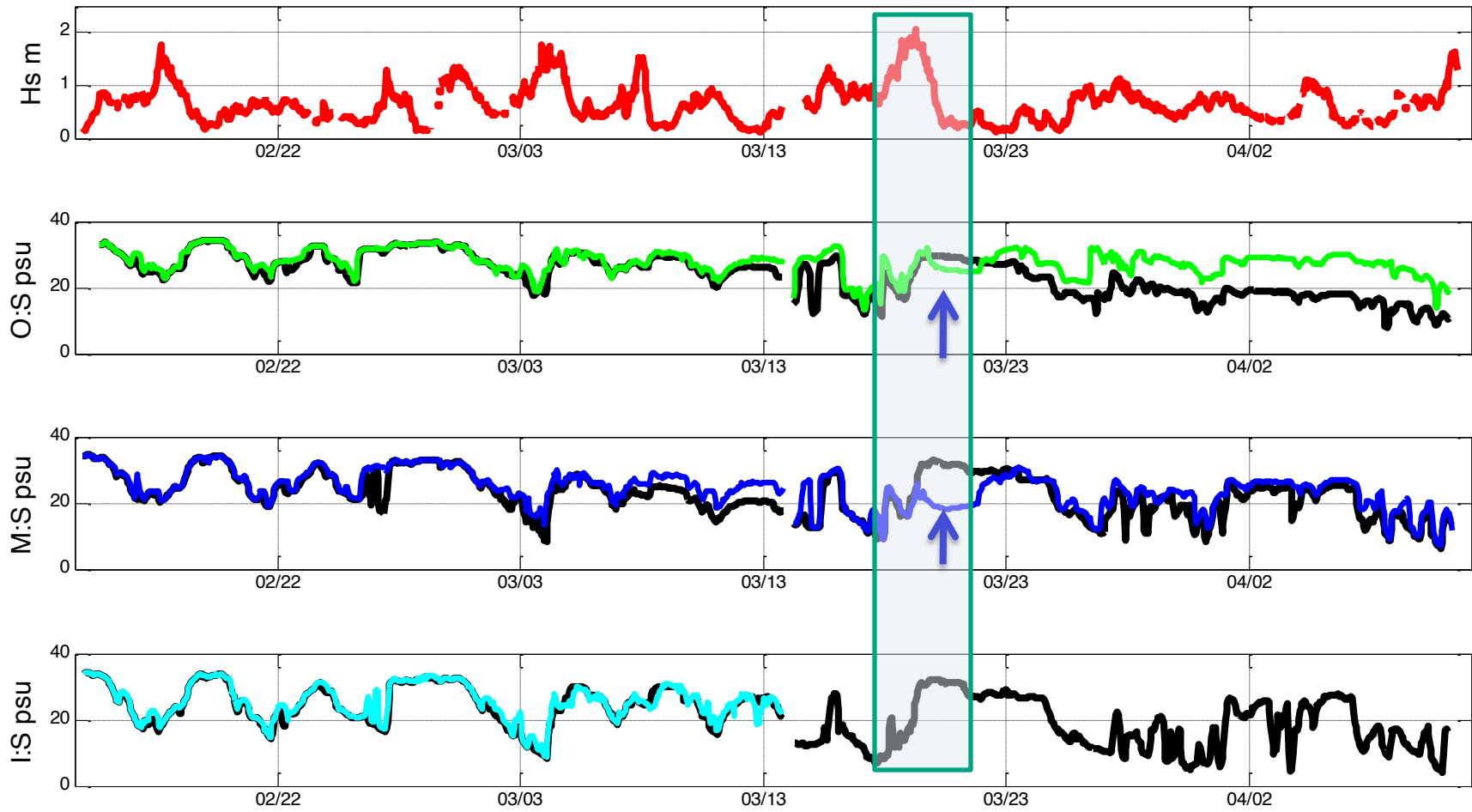
# Waves

Traykovski



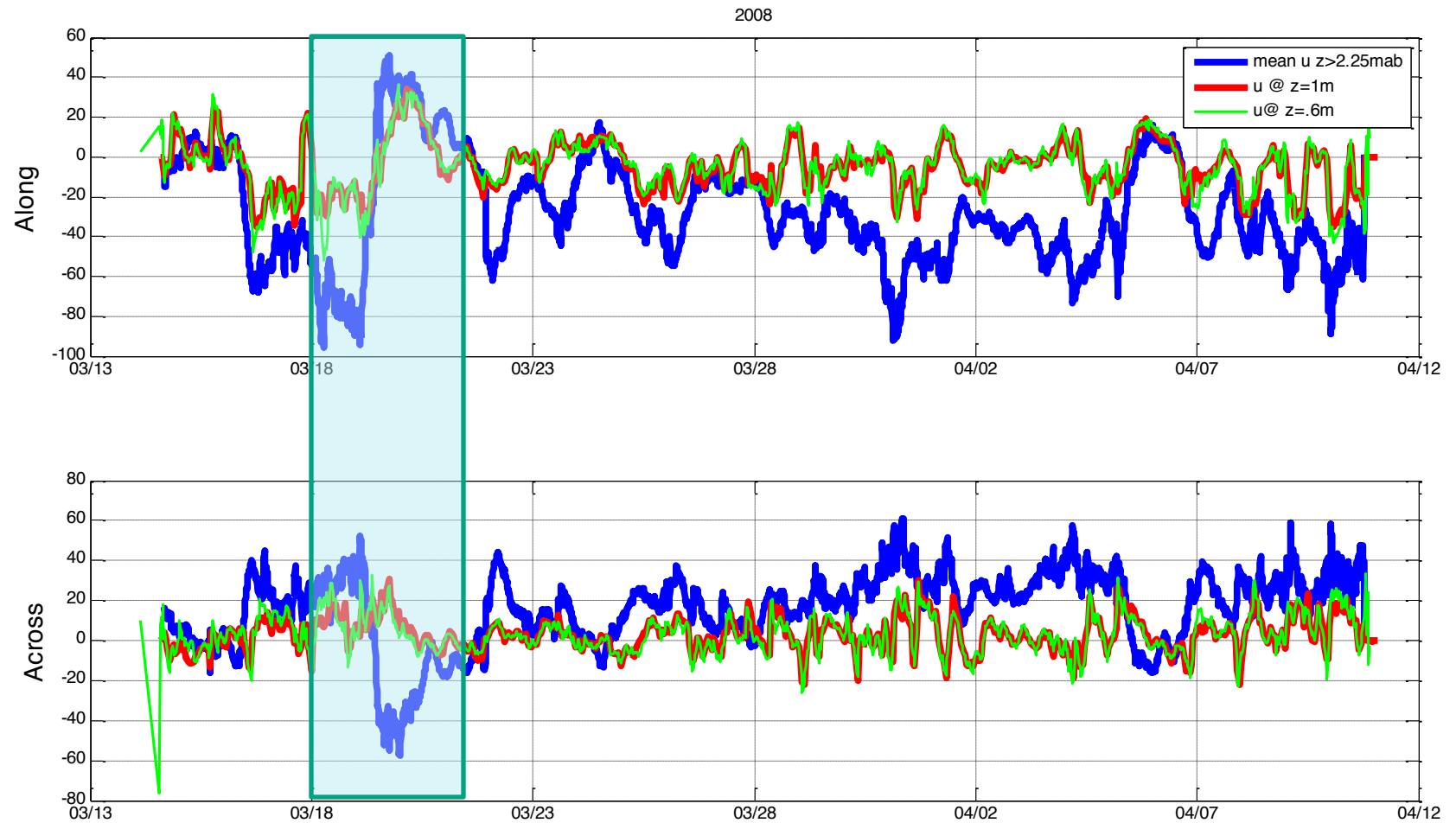
Hs

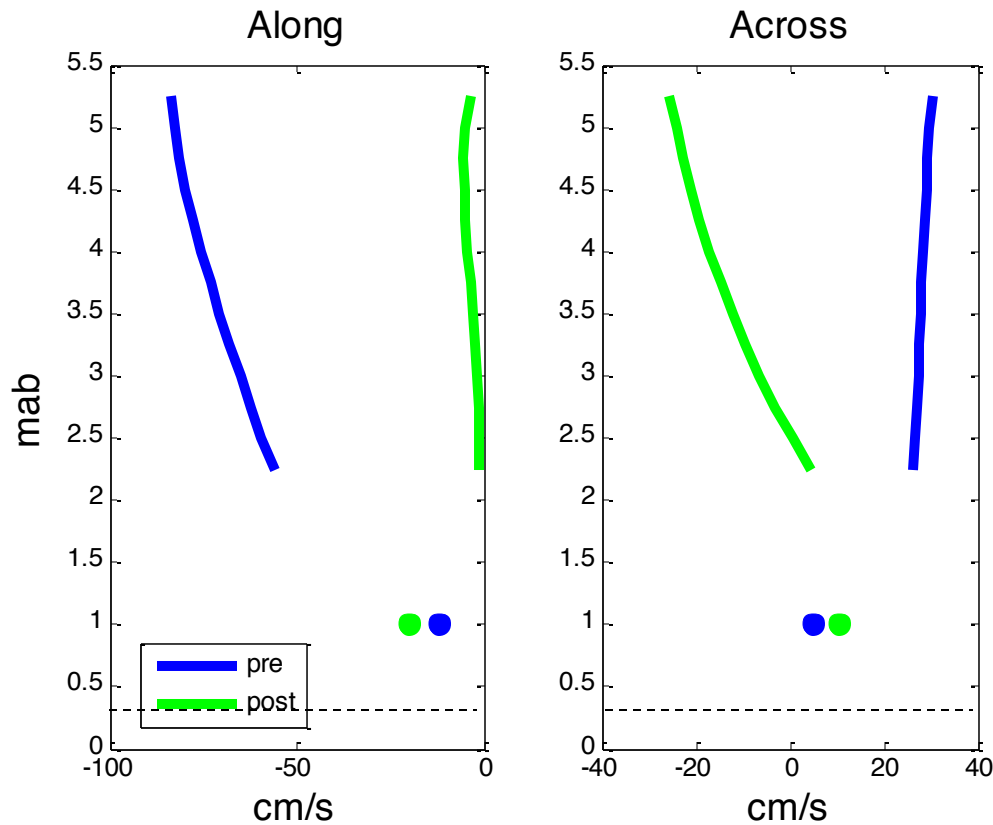
Salinity at 9, 7, 5 m z=1.25 and 0.3 mab





# ADV + ADV + ADCP





Post transport estimates



Order  $20 \text{ cm/s} * .15 \text{ g/l} * 5 \text{ m} = 1500 \text{ g/s cm}$



Order  $10 - 20 \text{ cm/s} * 10 \text{ g/l} * 20 \text{ cm} (?) = 2000-4000 \text{ g/s cm}$

*\*interesting caveat...new analysis of ABS observations by Peter Traykovski indicate very weak intermittent offshore gravity flow of very thin mud below the 'bottom' onshore flow...work in progress*

## Atchafalaya Summary

- The combination of wave dissipation, event-driven mean currents, water column stratification, and particle characteristics all contribute to net onshore transport and accretion.
- Sediment is delivered in late winter/early spring.
- During passage of CFs, waves initially suspend bottom sediment and mix the water column.
- Post-front, stratification is re-established, rapid settling is enhanced and fluid muds form contributing to wave dissipation.
- Onshore sediment transport of bottom waters is greater than offshore sediment transport in waters column above ~1 m.



Photo courtesy of A. Draut