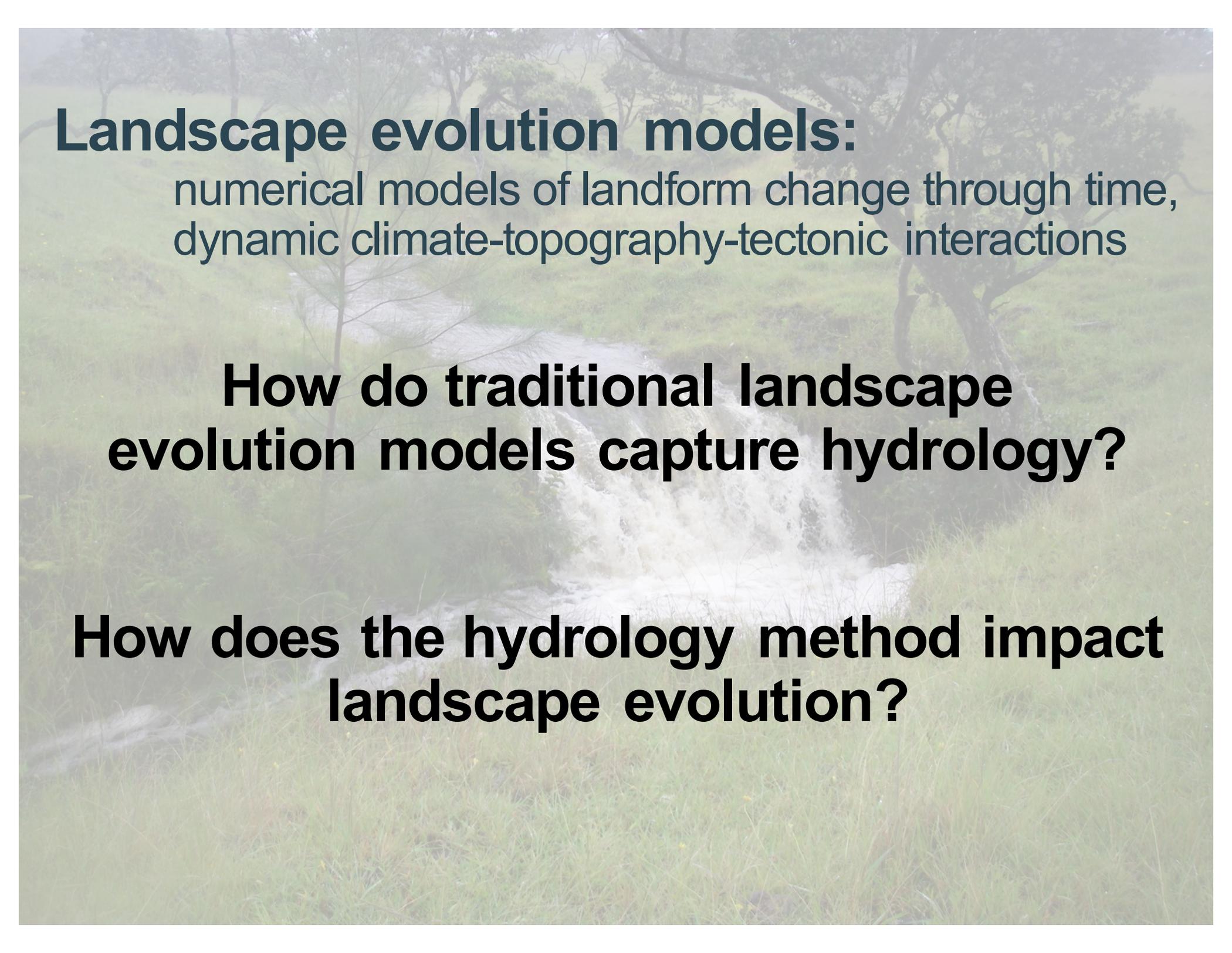


Integrating a 2-D hydrodynamic model into the Landlab modeling framework

Jordan Adams

May 18, 2016



Landscape evolution models:

numerical models of landform change through time,
dynamic climate-topography-tectonic interactions

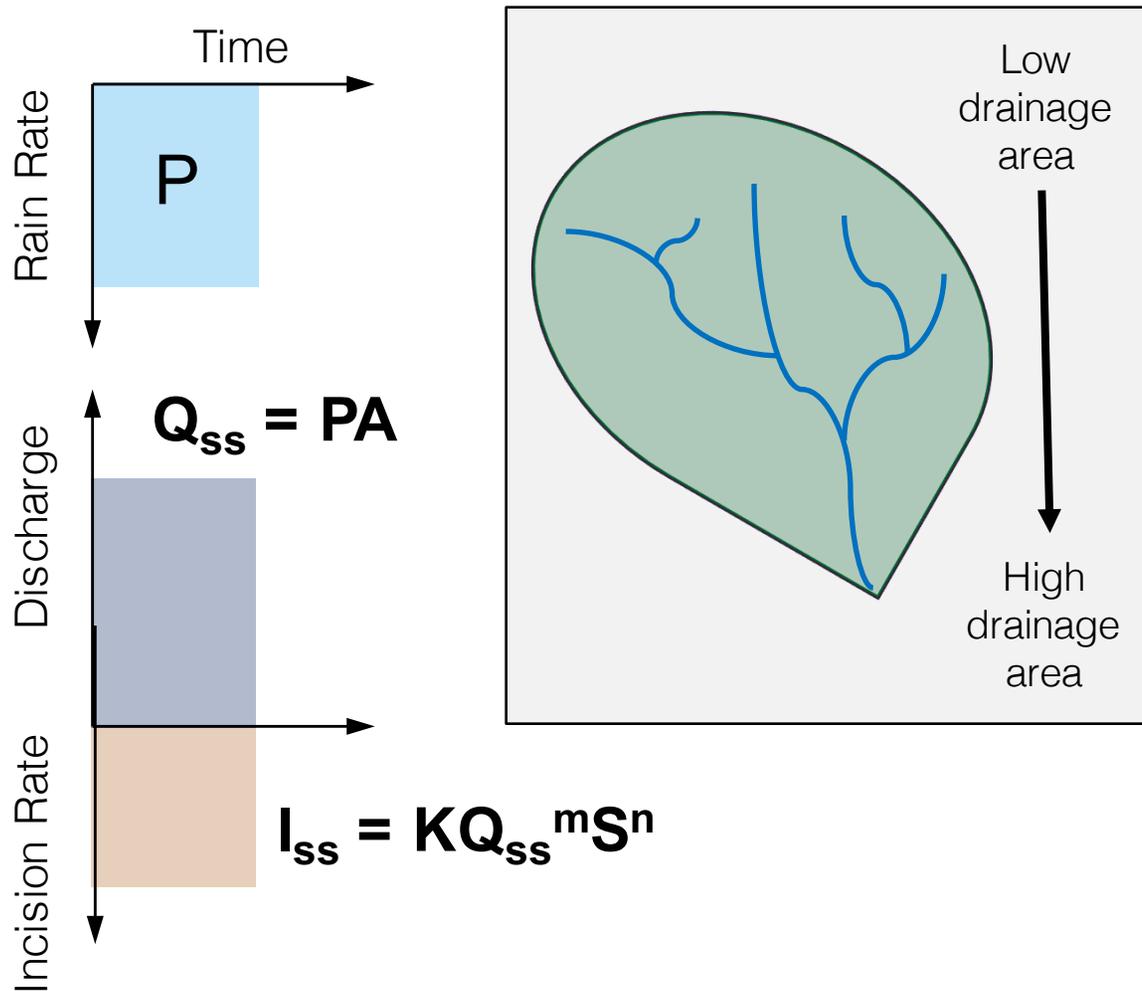
**How do traditional landscape
evolution models capture hydrology?**

**How does the hydrology method impact
landscape evolution?**

Steady-state hydrology

Each event:

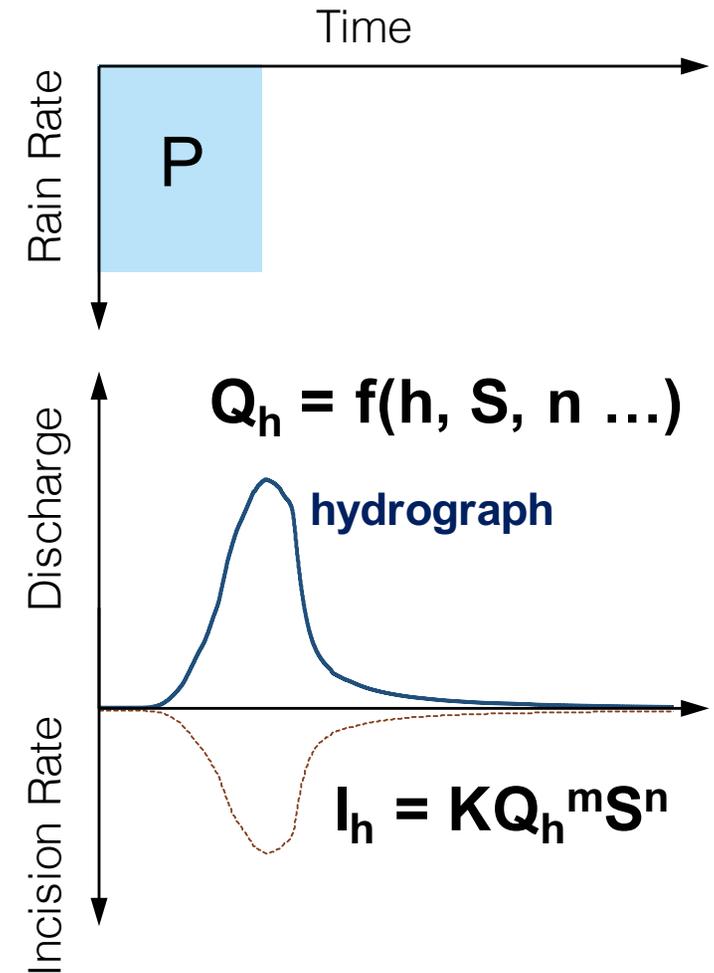
- single rainfall rate
- single discharge at all points in watershed
- single incision rate



Nonsteady hydrology

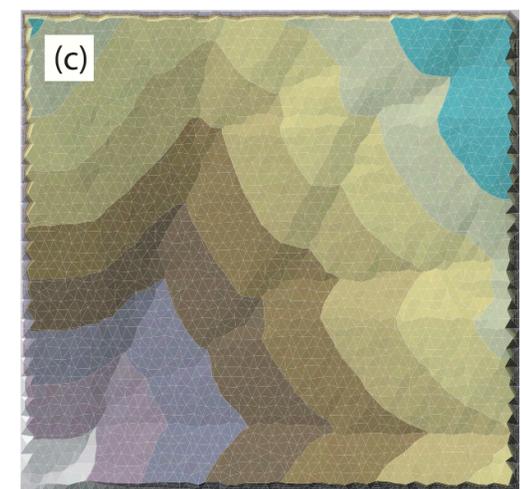
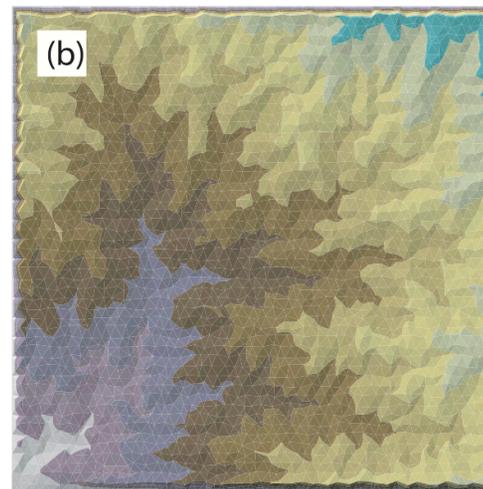
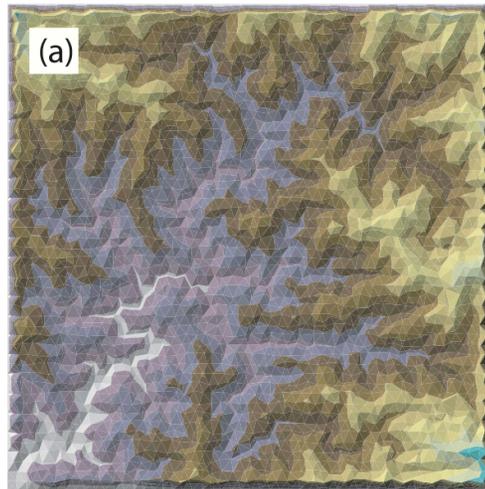
Each event:

- single rainfall rate
- wave propagates across watershed:
 - nonsteady discharge at each point
 - nonsteady incision at each point



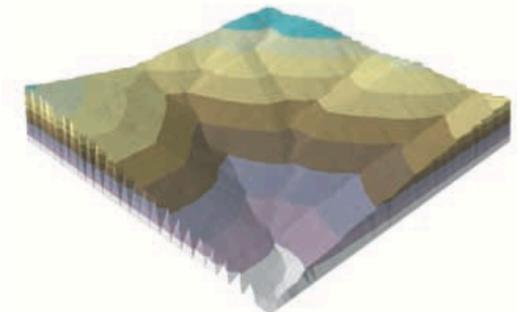
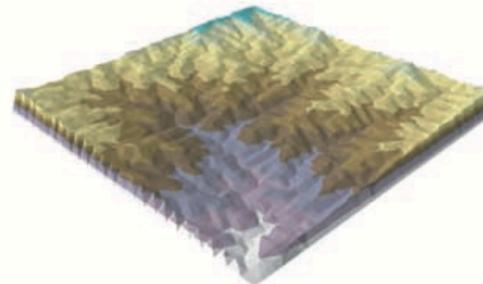
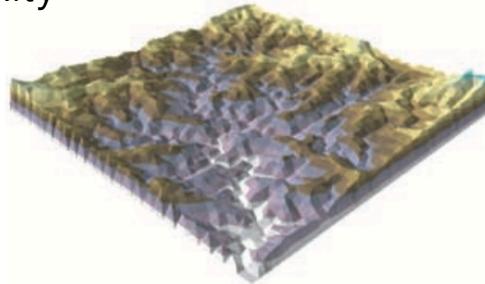
Steady-state assumption can be problematic: large catchments and short-duration precipitation events

Steady-state  nonsteady state



Note:

- greater relief
- increasing convexity
- low valley density

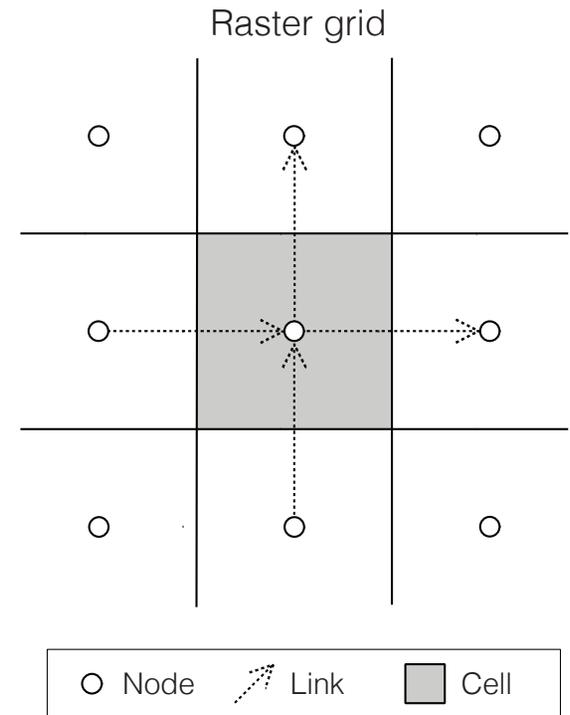


Sólyom and Tucker, 2004, *Journal of Geophysical Research*

Landlab: *A Python toolkit for modeling Earth surface processes*

- Open-source modeling library
 - 2-D gridding libraries
 - Pre-built process components
 - Coupling framework: multi-process models
 - Input / output utilities

**Geared toward (but not limited to!)
Earth-surface dynamics**



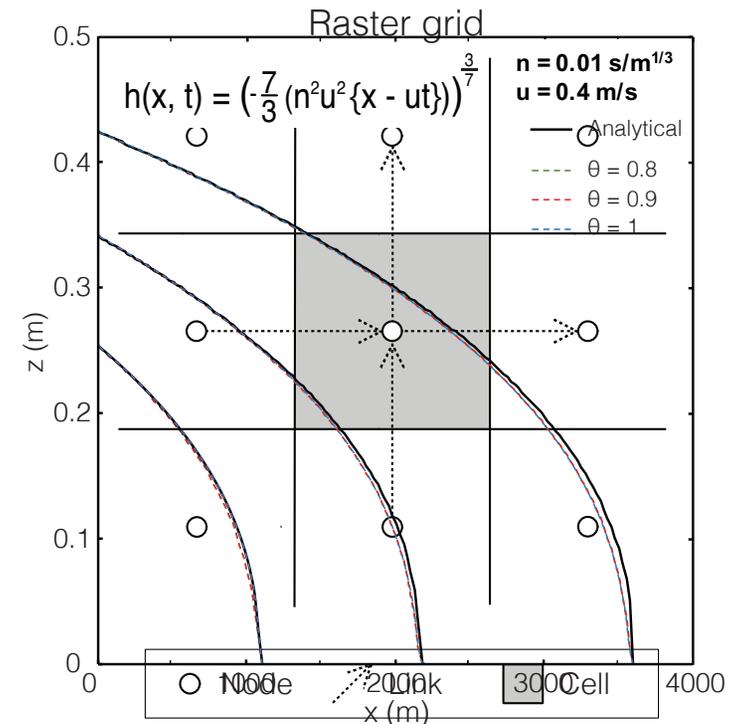
Visit our website: <http://landlab.github.io>



de Almeida overland flow component



- Urban flood inundation model (de Almeida et al., 2012)
- Centered finite-difference, explicit
- Routes hydrograph at all grid locations, flow in D4



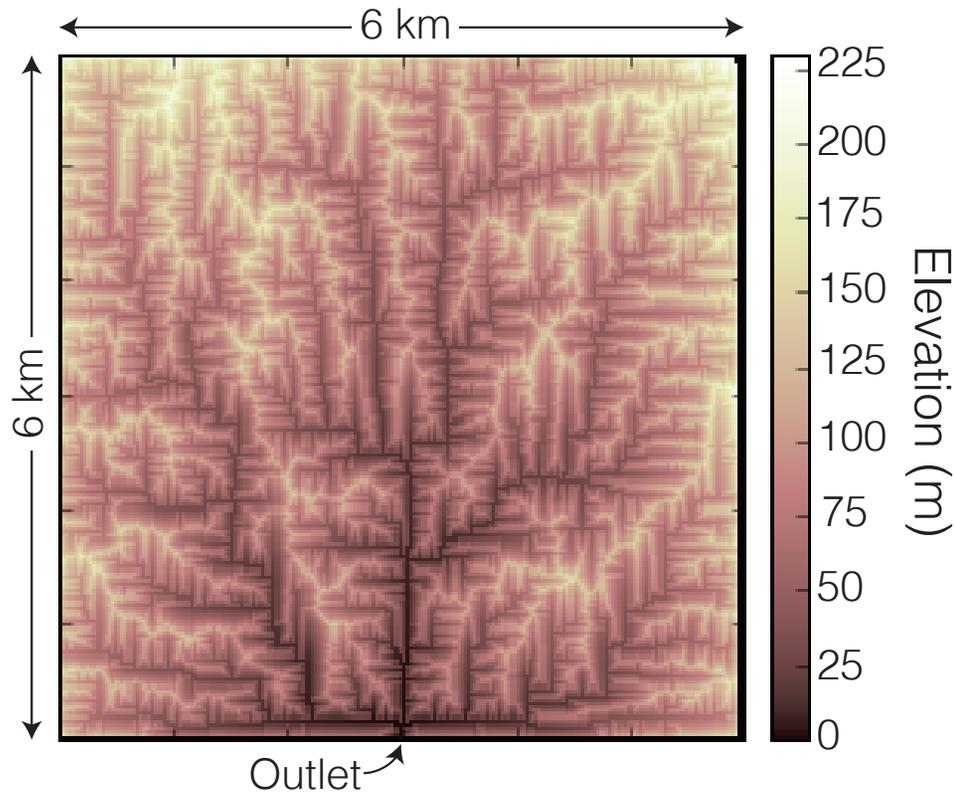
$$q_{t+1} = \frac{[\theta q_t + \frac{1-\theta}{2}(q_{t, \text{left}} + q_{t, \text{right}})] - gh\Delta t S_w}{1 + g\Delta t n^2 |q_t| / h^{7/3}}$$

de Almeida et al., 2012, WRR

$$\Delta t = a \frac{\Delta x}{\sqrt{gh_f}}$$

Hunter et al., 2005,
Advances in Water Resources

Model domain



Simple stream power parameters

$$I = KQ^m S^n$$

Parameter	Value
K (<i>erodibility</i>)	0.0007 m/yr
m	0.5
n	1.0

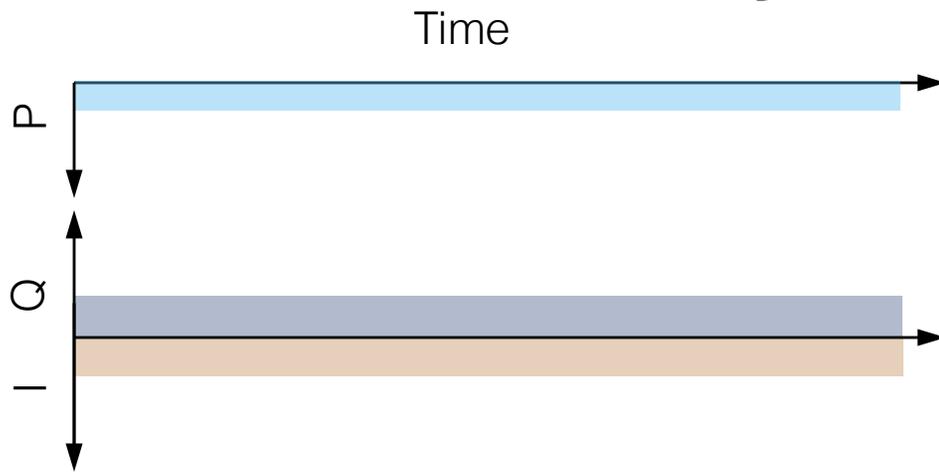
Geomorphic steady-state topography
(**uplift == erosion rate**)

36 km² drainage area

Grid resolution: 30 m x 30 m

Slopes: $\sim 10^{-1}$ to $\sim 10^{-2}$

Steady-State Parameters

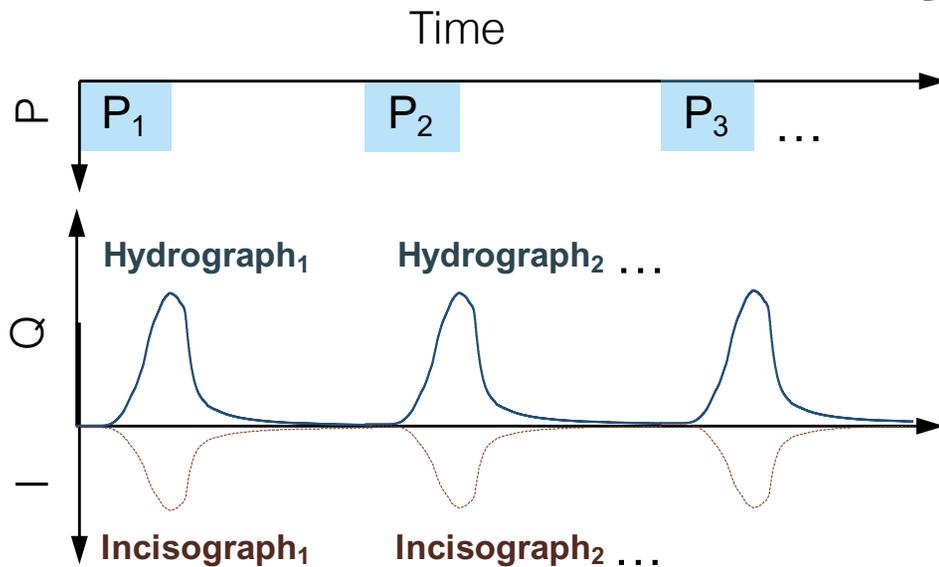


Rainfall characteristics

0.5 m/yr total rainfall for **10 years**

Low, constant rainfall rate

Nonsteady Parameters



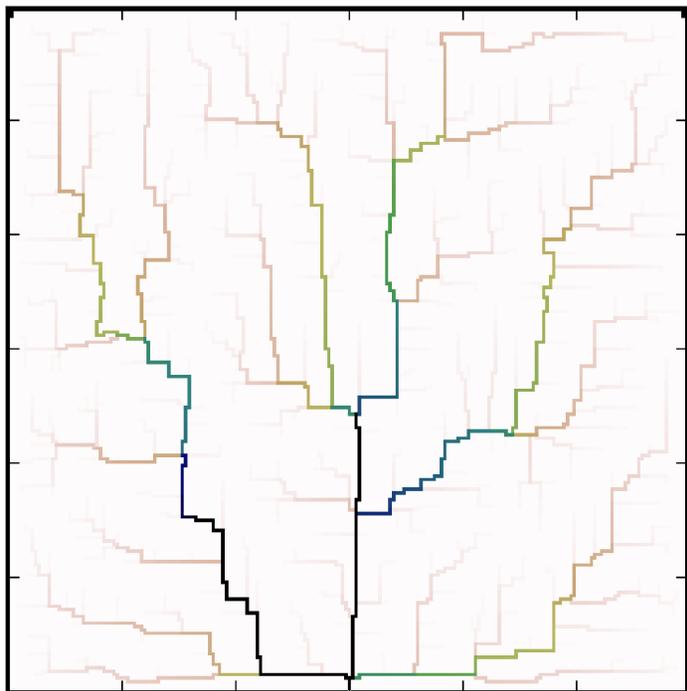
500 hydrograph events

3 different storm types:

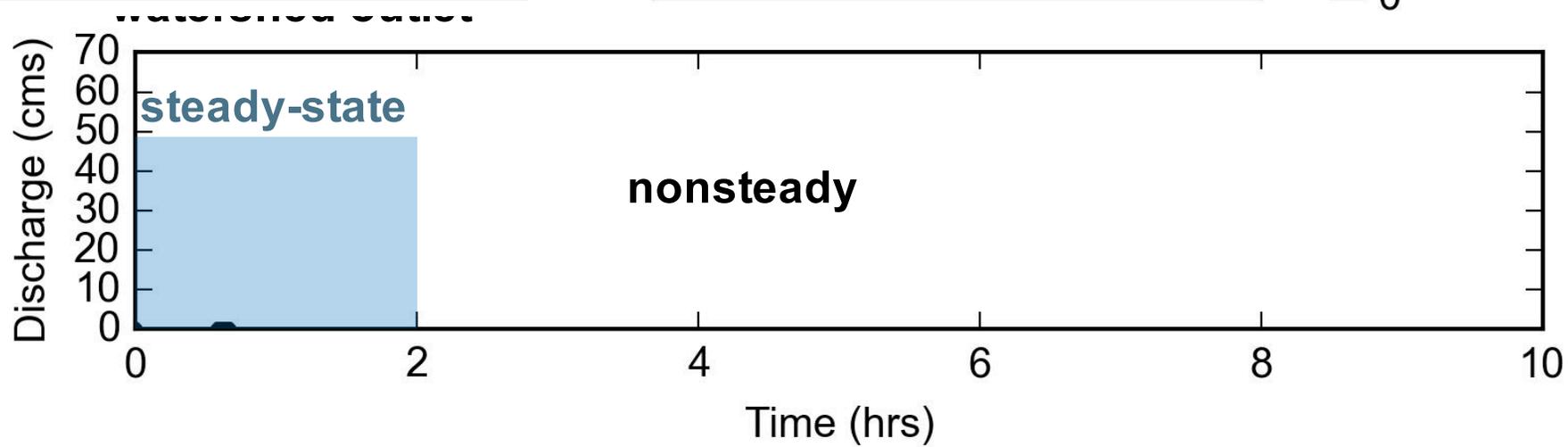
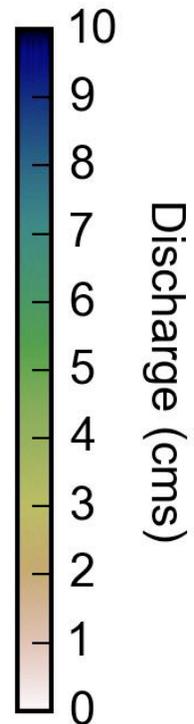
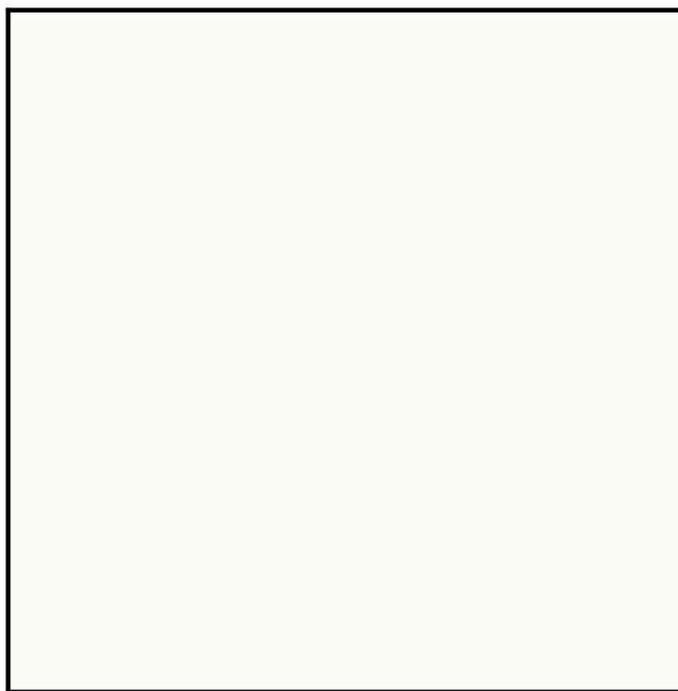
Intensity	Duration
2.5 mm/hr	4 hr
5 mm/hr	2 hr
10 mm/hr	1 hr

Intensity

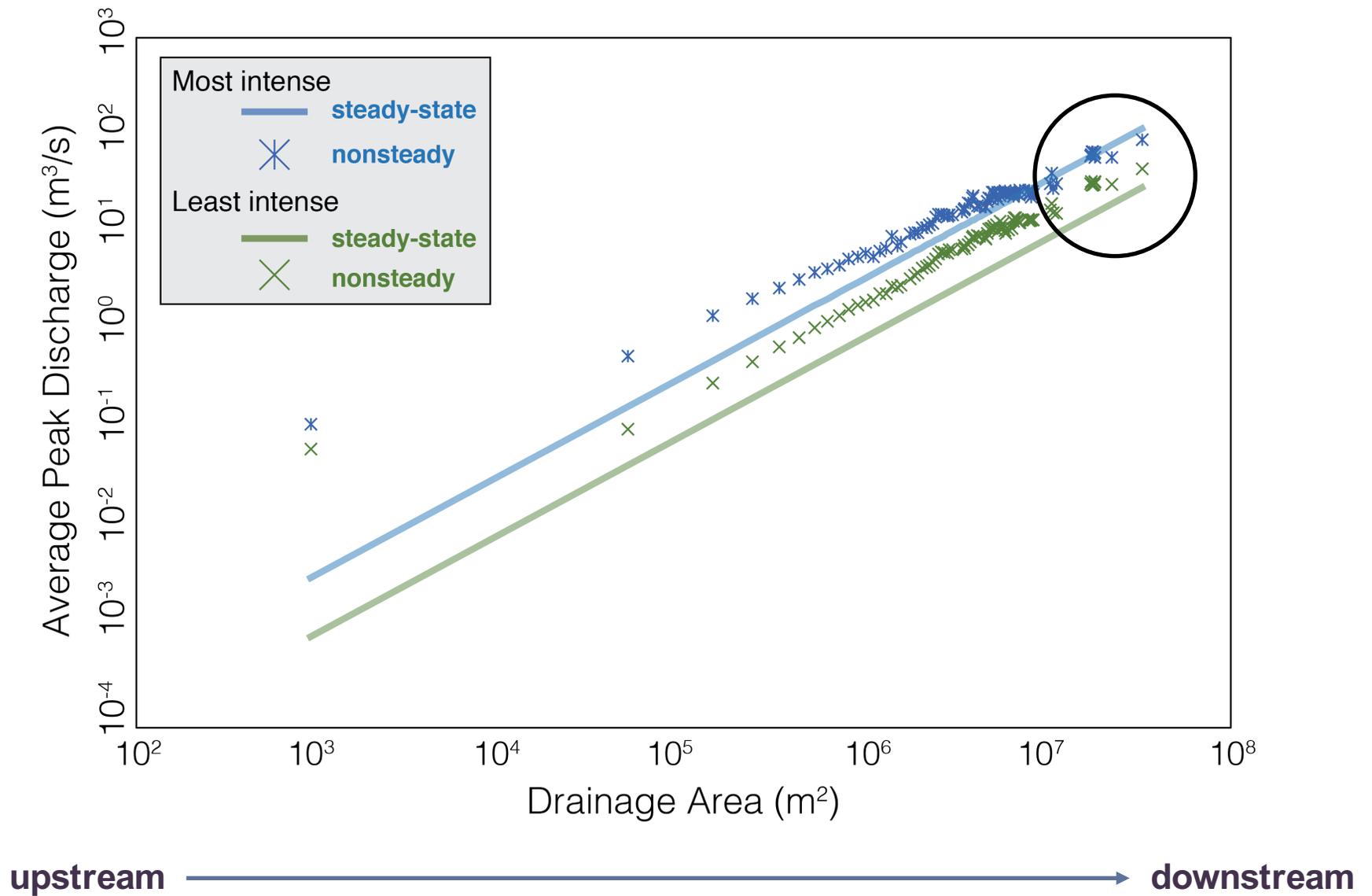
Steady-state



Nonsteady

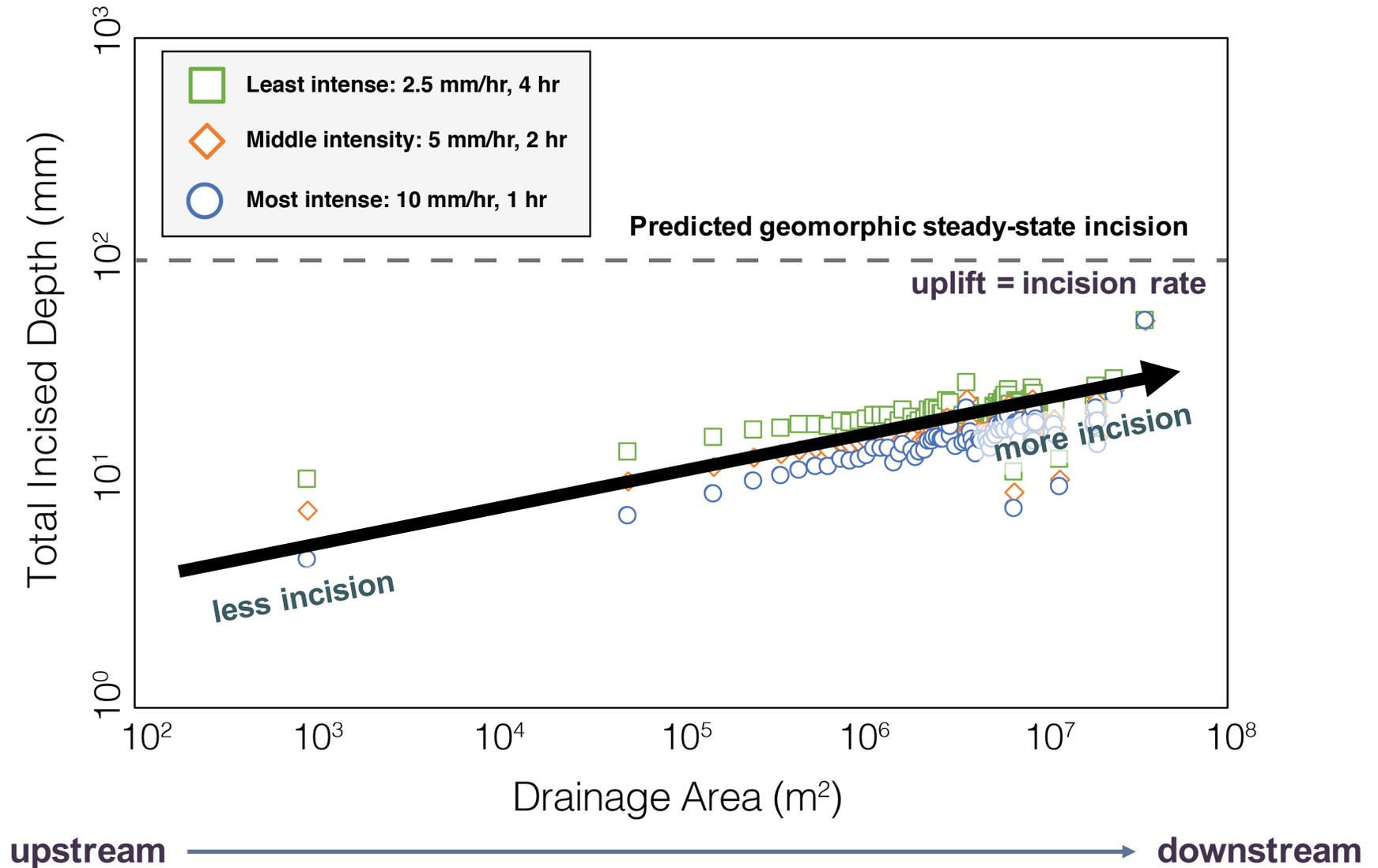


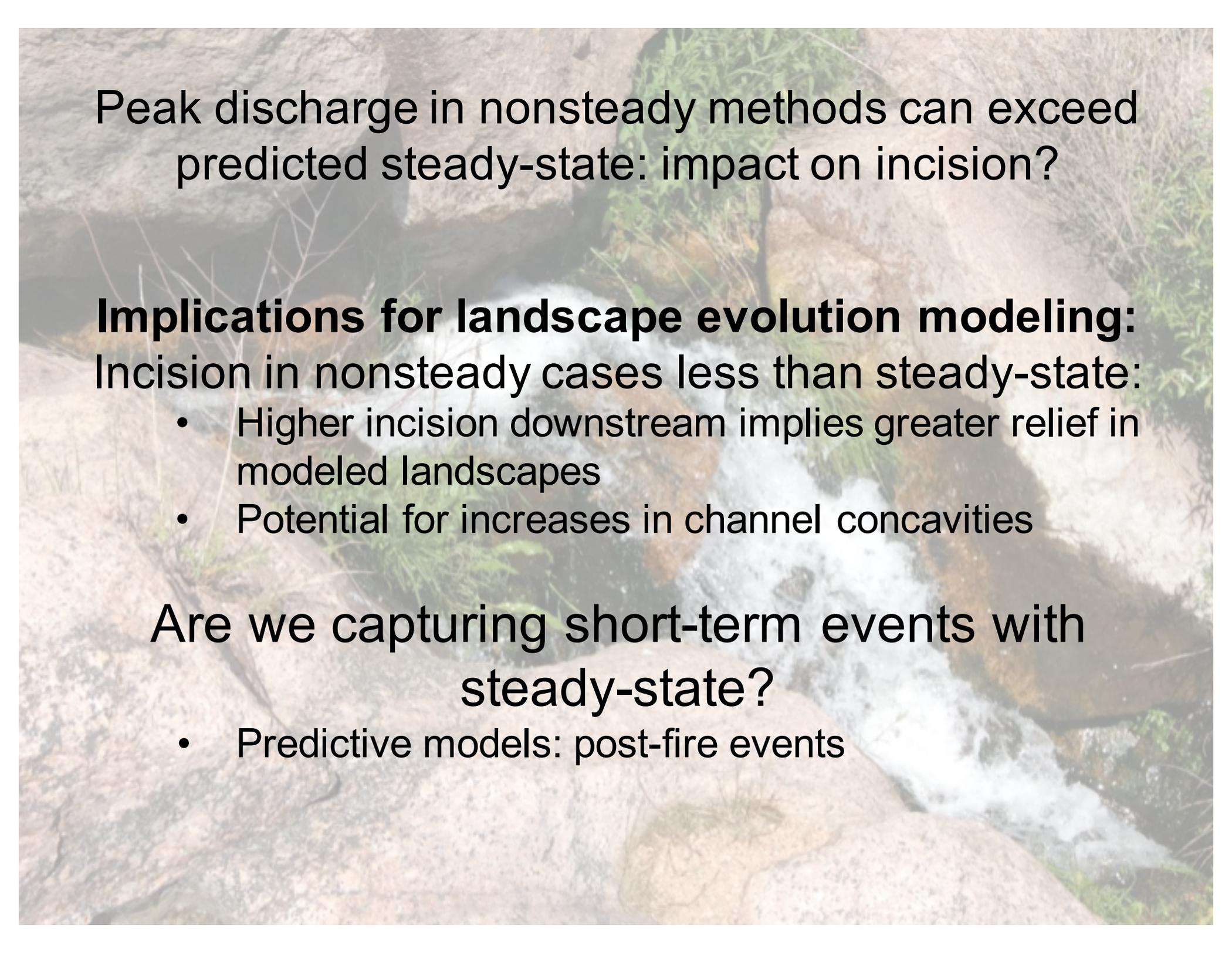
Average peak discharge is greater than predicted steady-state in upstream



Cumulative incised depth by nonsteady method less than predicted by the steady-state hydrology model

In all intensities, higher drainage areas experience greater incision





Peak discharge in nonsteady methods can exceed predicted steady-state: impact on incision?

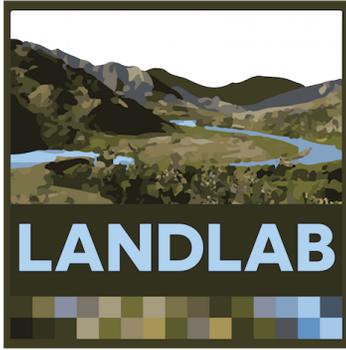
Implications for landscape evolution modeling:
Incision in nonsteady cases less than steady-state:

- Higher incision downstream implies greater relief in modeled landscapes
- Potential for increases in channel concavities

Are we capturing short-term events with steady-state?

- Predictive models: post-fire events

Visit our website: <http://landlab.github.io>



University
of Colorado
Boulder



UNIVERSITY of
WASHINGTON

CSDMS

COMMUNITY SURFACE DYNAMICS MODELING SYSTEM



J. Adams supported by NSF grants ACI-1147519 and ACI-1450338, and a 2016 CSDMS Student Scholarship