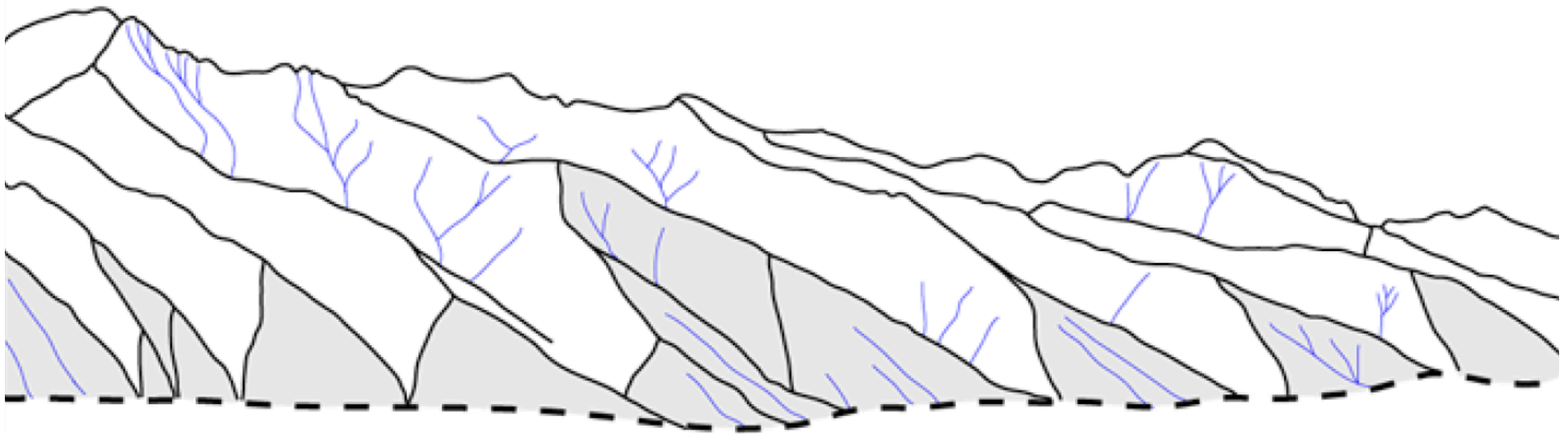


MODES OF EXTENSIONAL FAULTING

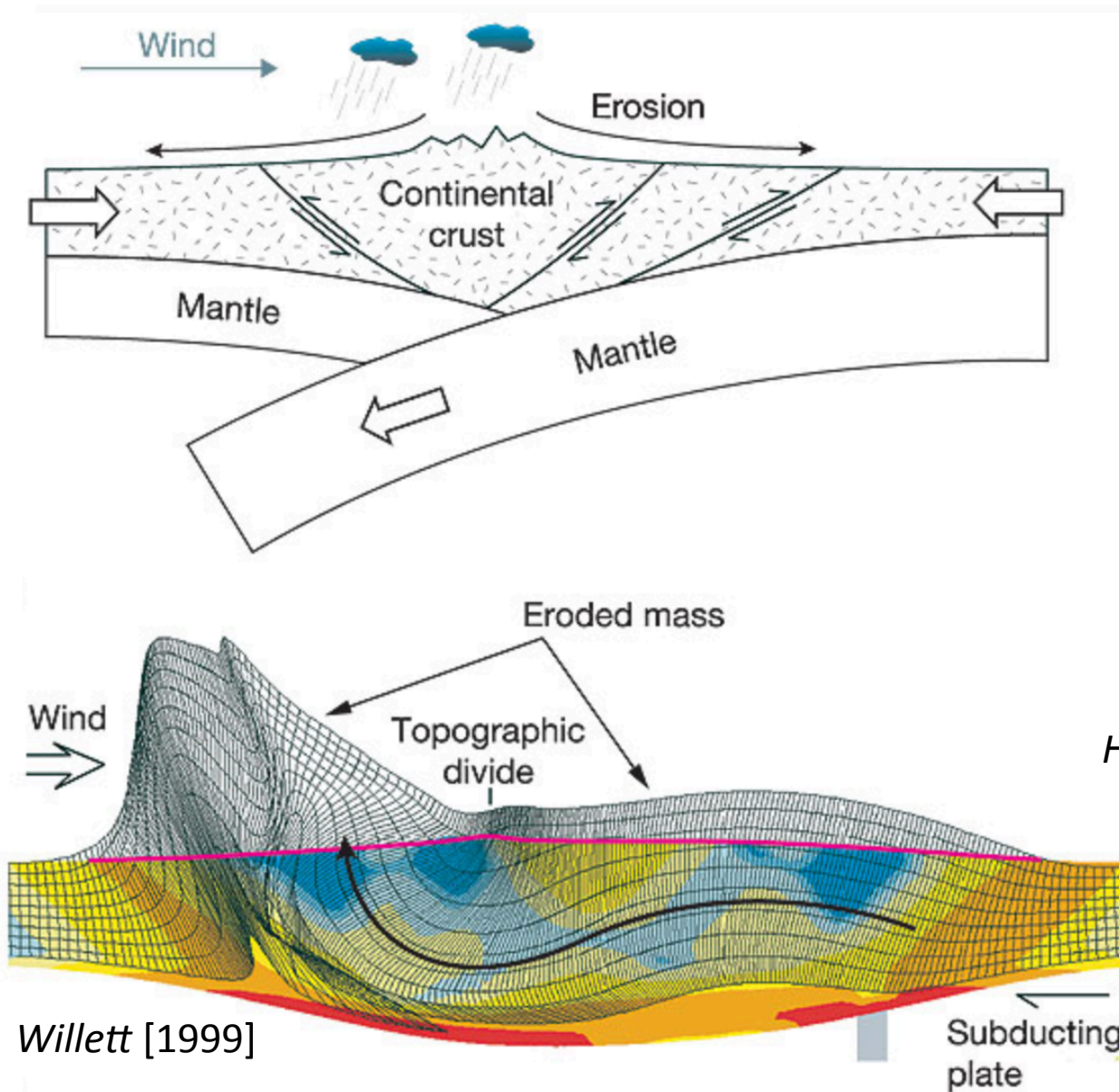
CONTROLLED BY SURFACE PROCESSES



Jean-Arthur Olive^{1,2}, Mark Behn², Luca Malatesta³

1 – MIT/WHOI, now LDEO, 2 – WHOI, 3 – Caltech

Feedbacks between tectonics and erosion



**“Savor the irony
should mountains
owe their
metamorphic muscles
to the drumbeat
of tiny raindrops.”**

Hoffman and Grotzinger [1993]

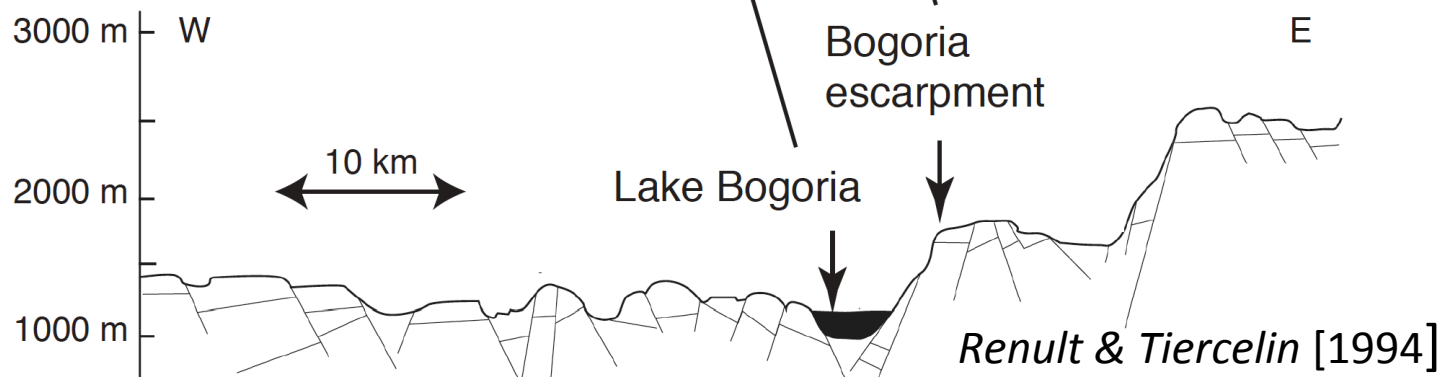
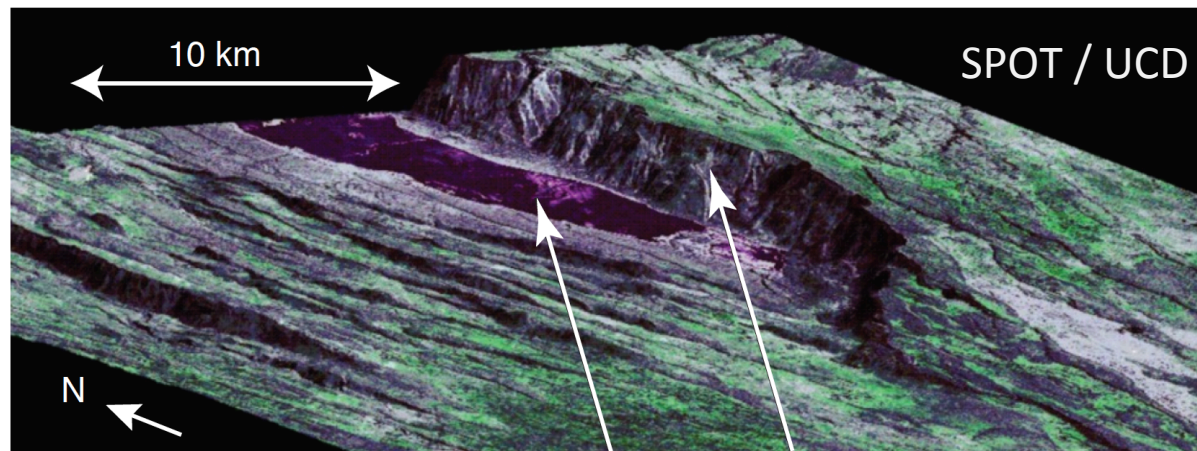
Diversity of (subaerial) normal fault styles

SHORT-LIVED FAULTS

LONG-LIVED DETACHMENTS



Baringo-Bogoria half-graben (Kenya) fault life span
East African Rift - Eastern Branch

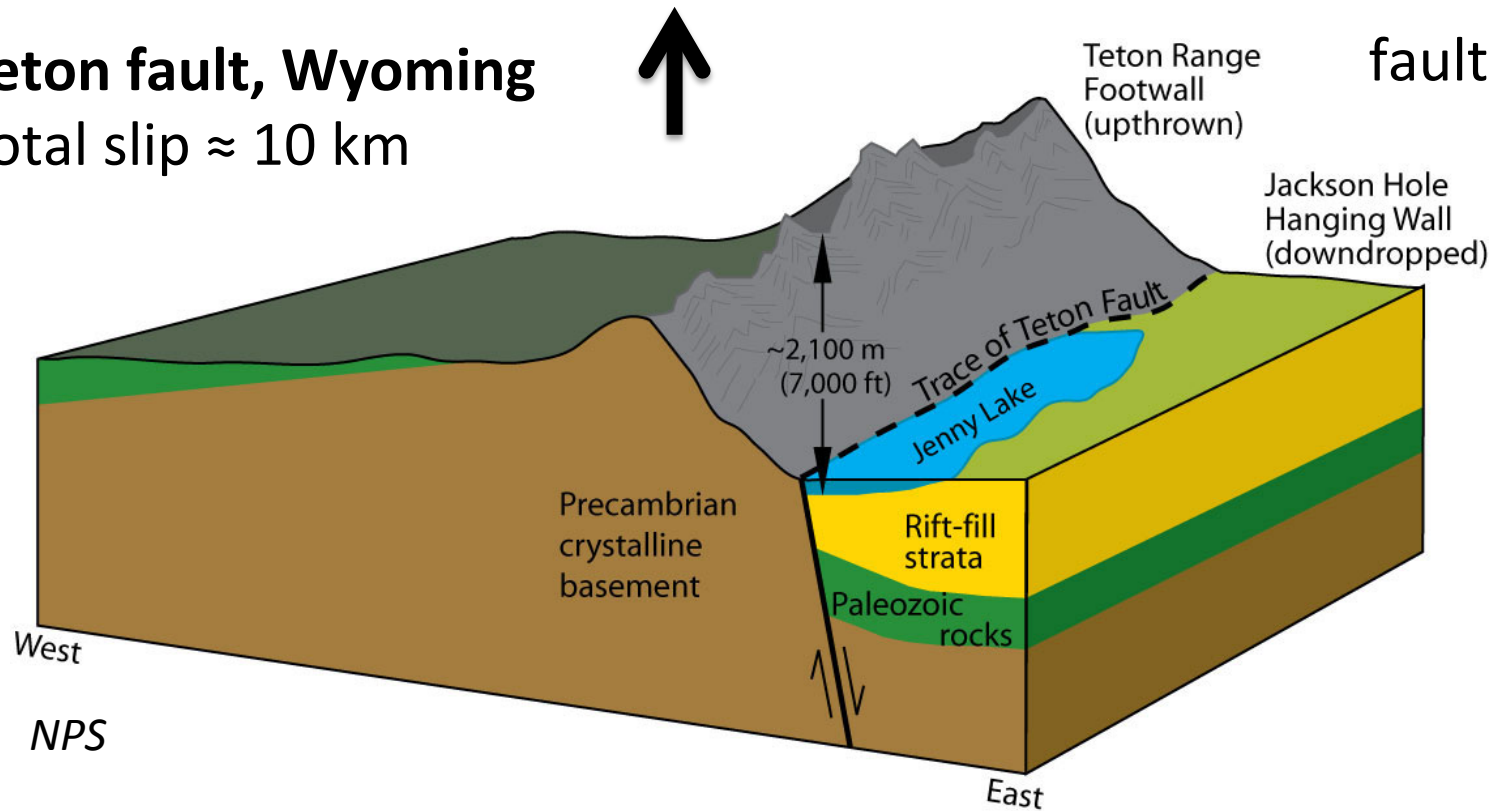


Diversity of (subaerial) normal fault styles

SHORT-LIVED FAULTS

LONG-LIVED DETACHMENTS

Teton fault, Wyoming
Total slip ≈ 10 km



fault life span



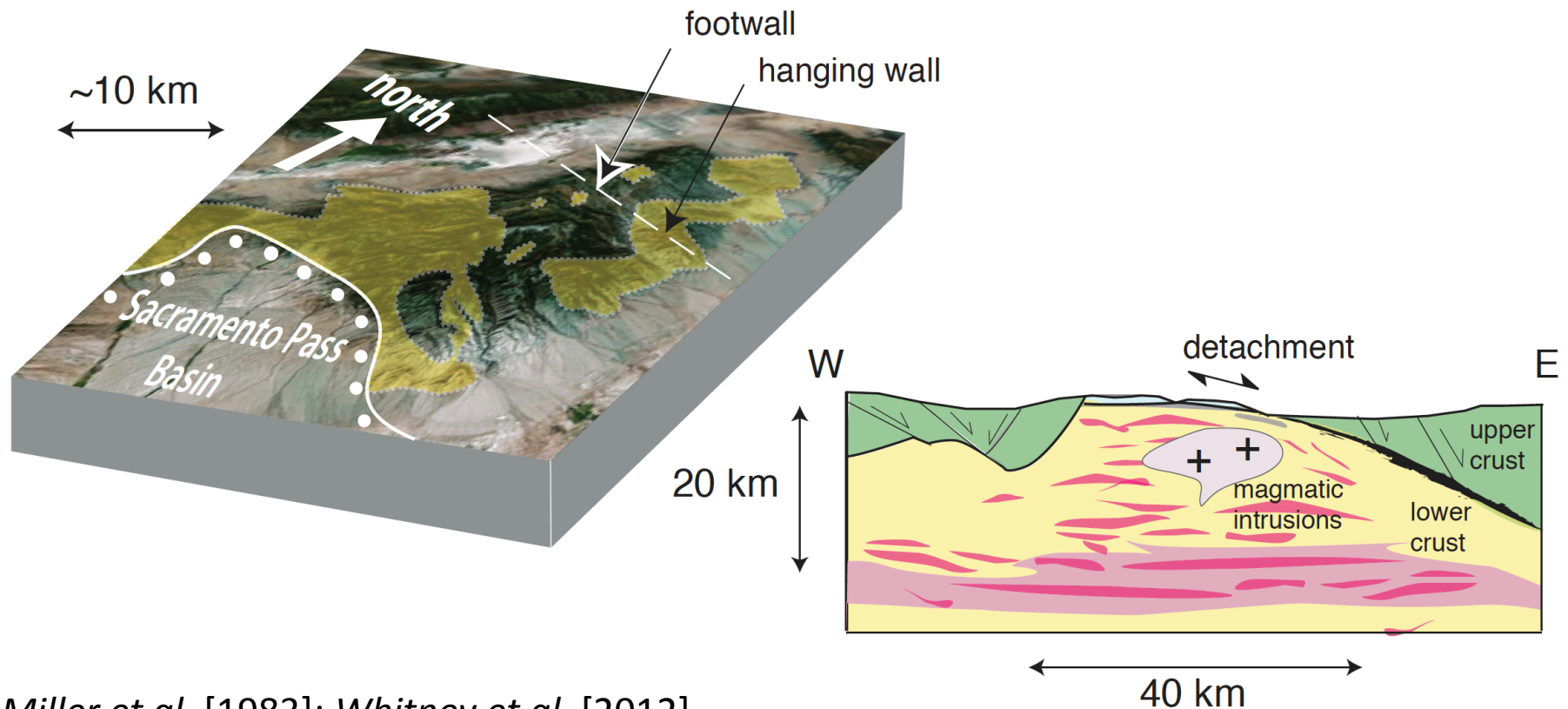
Diversity of (subaerial) normal fault styles

SHORT-LIVED FAULTS

LONG-LIVED DETACHMENTS



Northern Snake Range (Nevada)



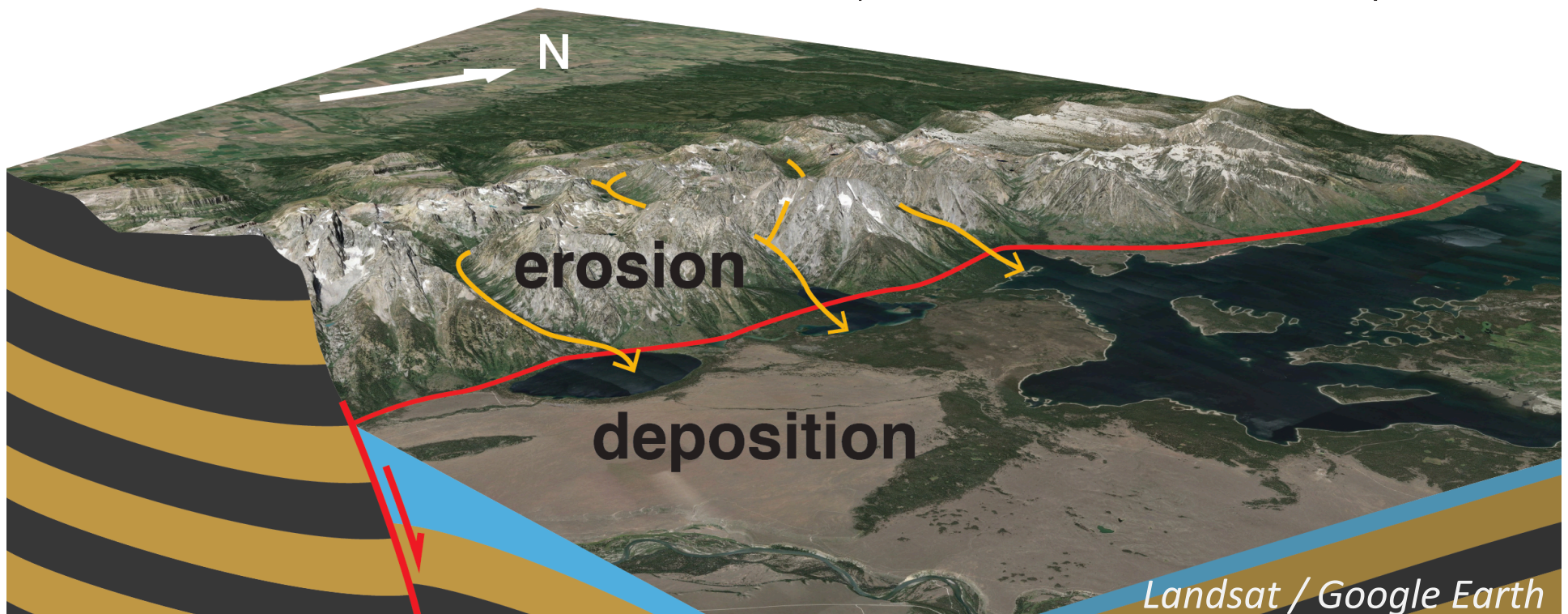
Miller et al. [1983]; Whitney et al. [2012]

Normal faulting vs. surface processes

Surface processes redistribute mass and topographic loads on a growing fault. **Can they affect fault life span ?**

Teton fault, WY

~10 km

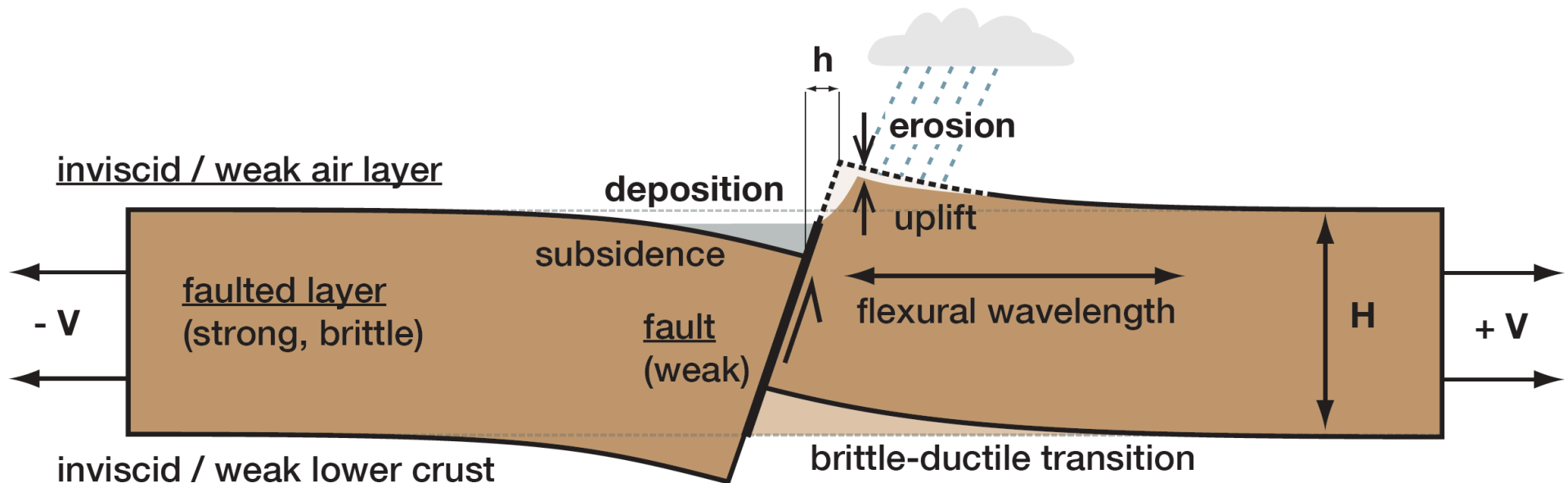


Coupled geodynamics – surface model

Geodynamic model (visco-elastic-plastic flow with FD/PIC):

Strong brittle layer (thickness H) with seeded fault subjected to stretching at half-rate V . Fault evolution monitored as offset h increases.

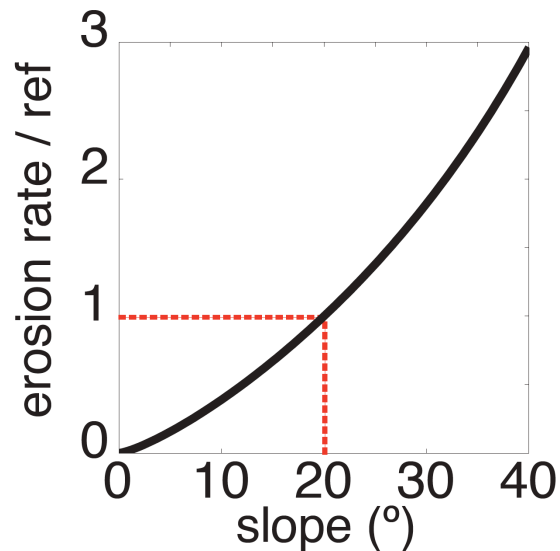
Numerical simulations conducted with the **SiStER** code (*Simple Stokes with Exotic Rheologies*).



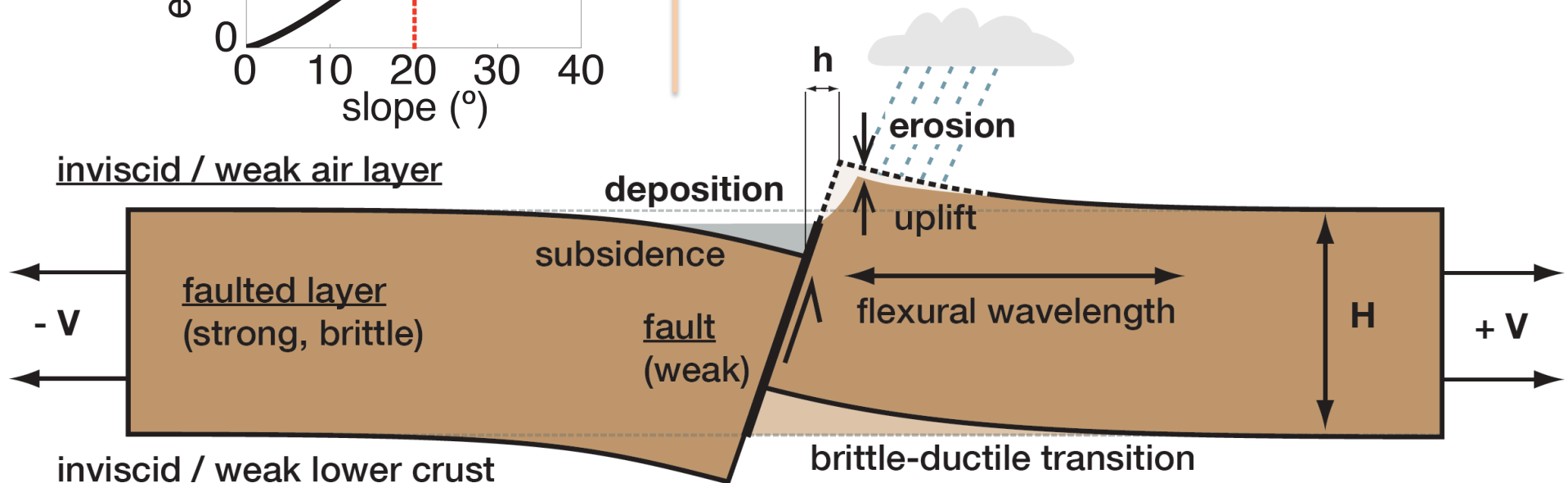
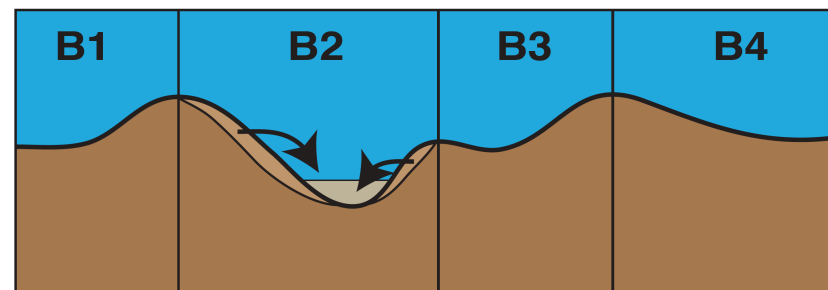
Coupled geodynamics – surface model

Surface evolution model (upper boundary condition):

Step 1: **erode** material

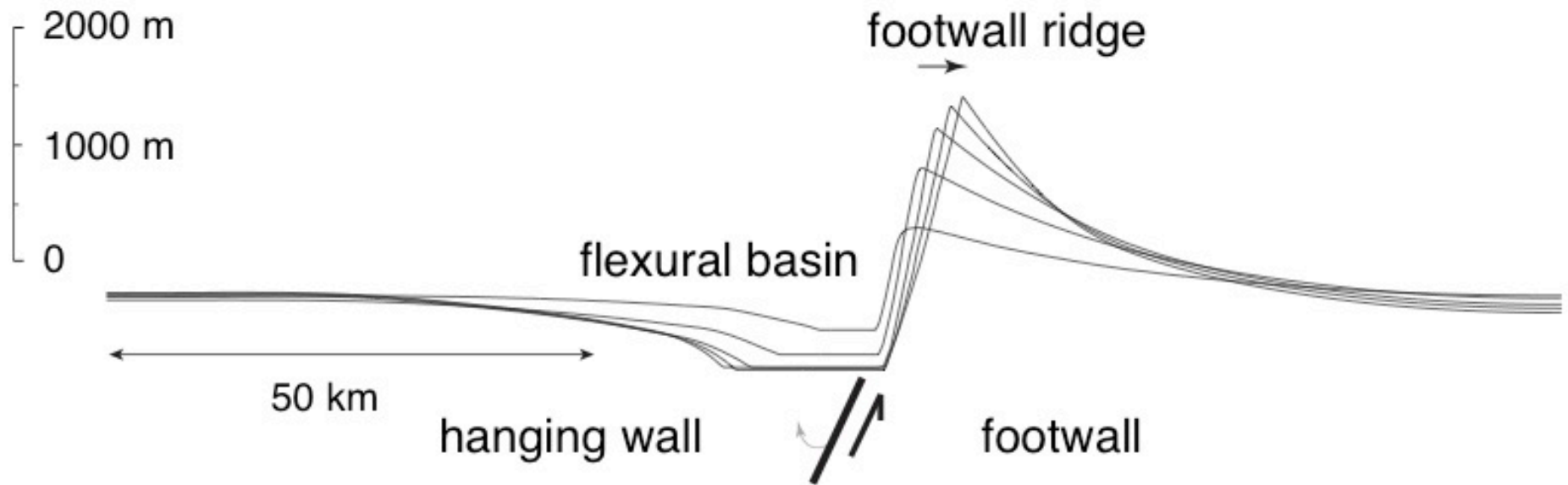


Step 2: **deposit** material flat in corresponding watershed



Fault growth with erosion and sedimentation

Fault growth is accompanied by

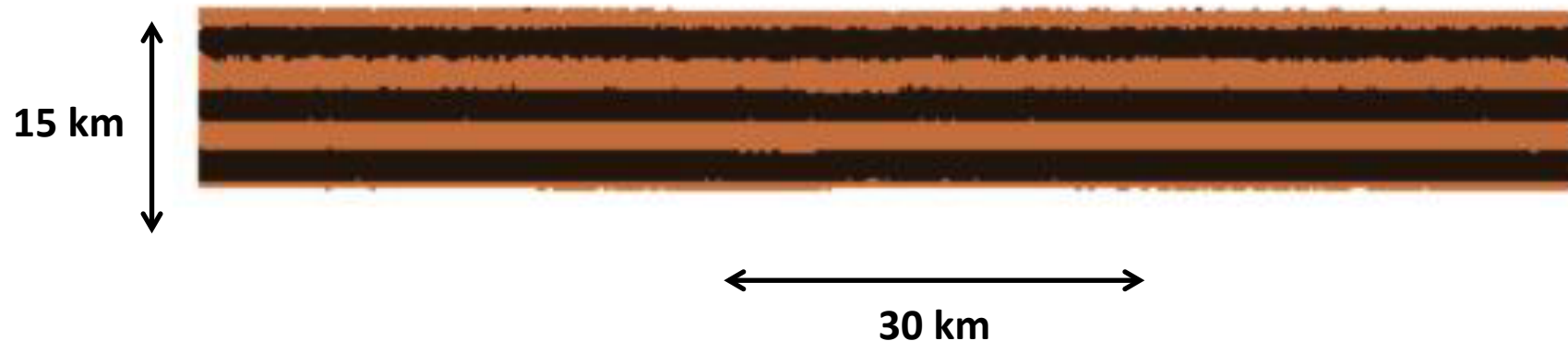


- Flexure of the footwall and hanging wall
- Rapid rotation of the fault plane down to 30-45°.
- Erosion of the footwall block, ridge migration, deposition in flexural basin.

Sequence of faulting with **slow** erosion

Extending a **15 km thick** layer at **2 mm/yr** (full rate).
Ref. erosion rate = **0.05% of fault slip rate**.

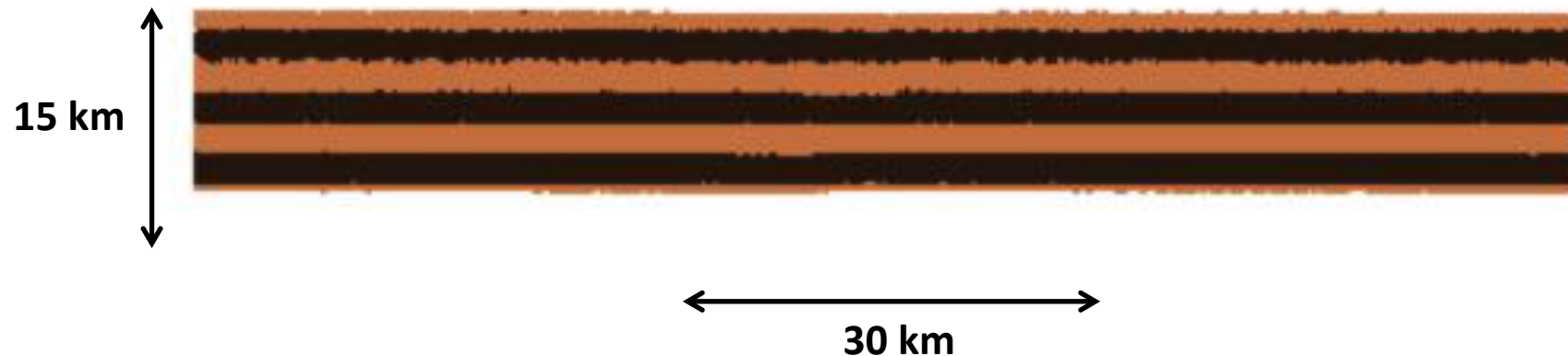
total extension = 0.74857 km



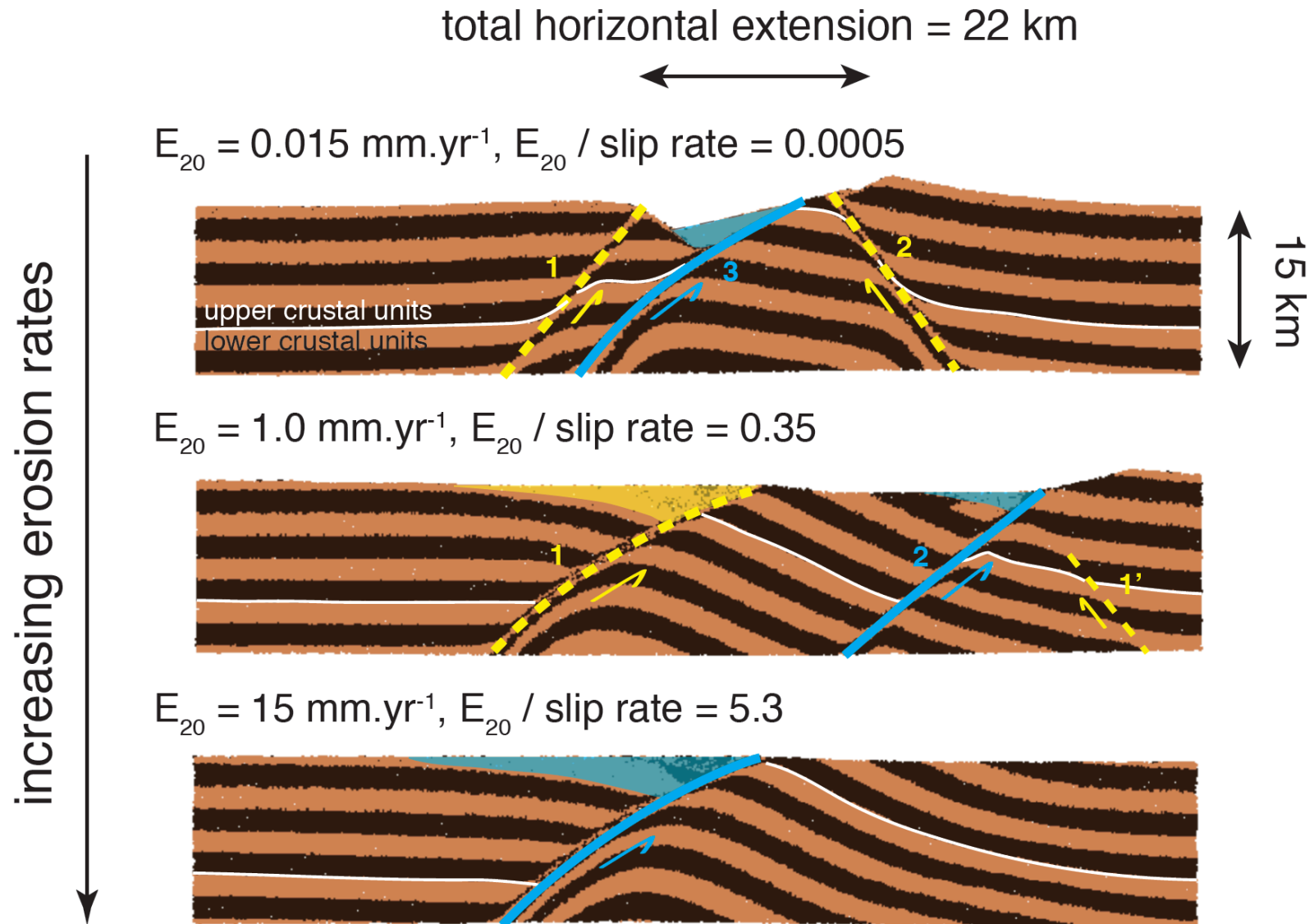
Sequence of faulting with **fast** erosion

Extending a **15 km thick layer** at **2 mm/yr** (full rate).
Ref. erosion rate = **35% of fault slip rate**.

total extension = 0.74856 km

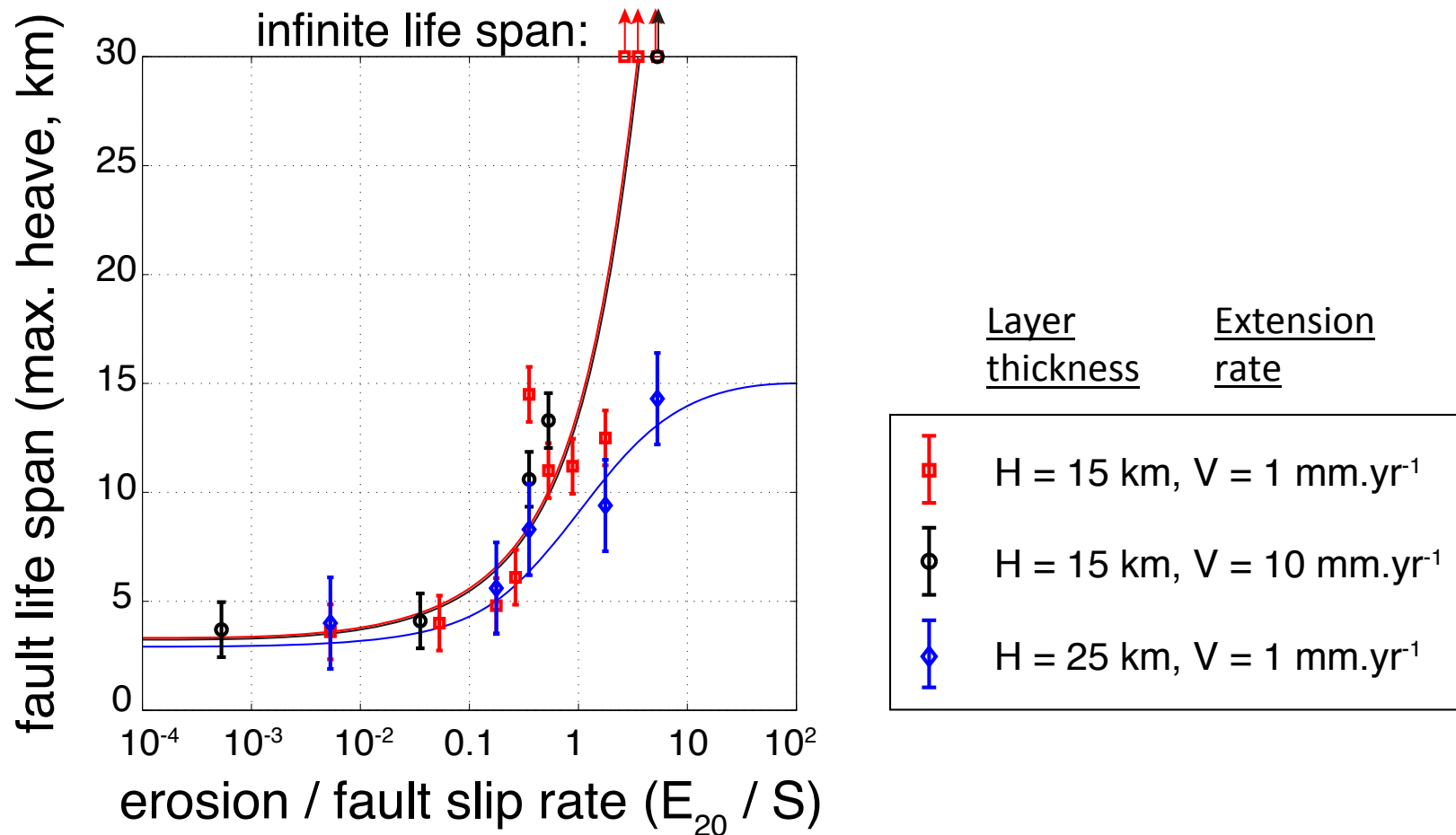


Faster surface processes = longer-lived faults



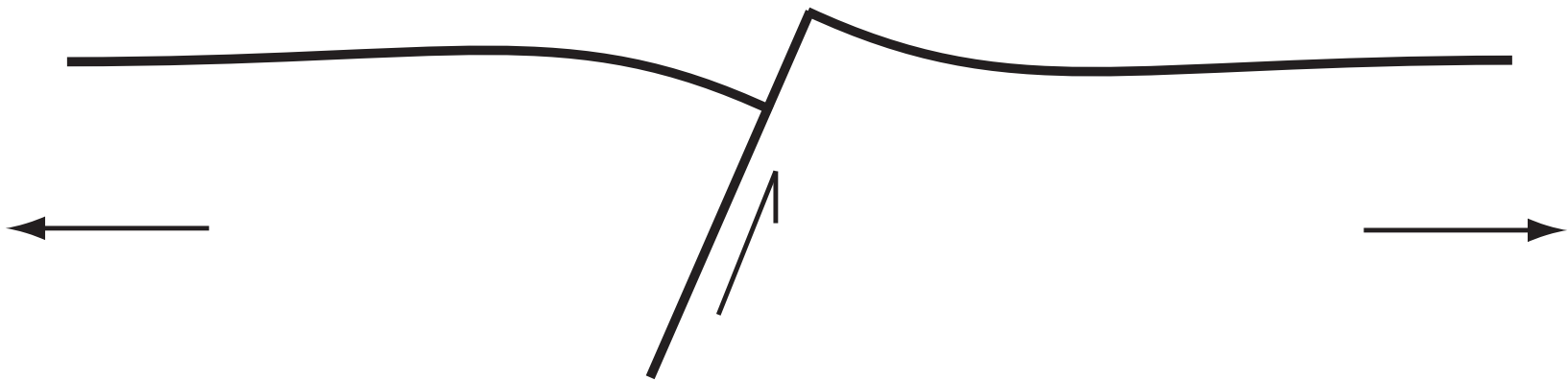
Faster surface processes = longer-lived faults

Surface processes **further enhance** fault life span in **thinner** faulted layers for a given erosion / slip rate.



Force balance on a growing normal fault

How to keep a normal fault active ? [*Forsyth, 1992*]



Force balance on a growing normal fault

How to keep a normal fault active ? [*Forsyth, 1992*]

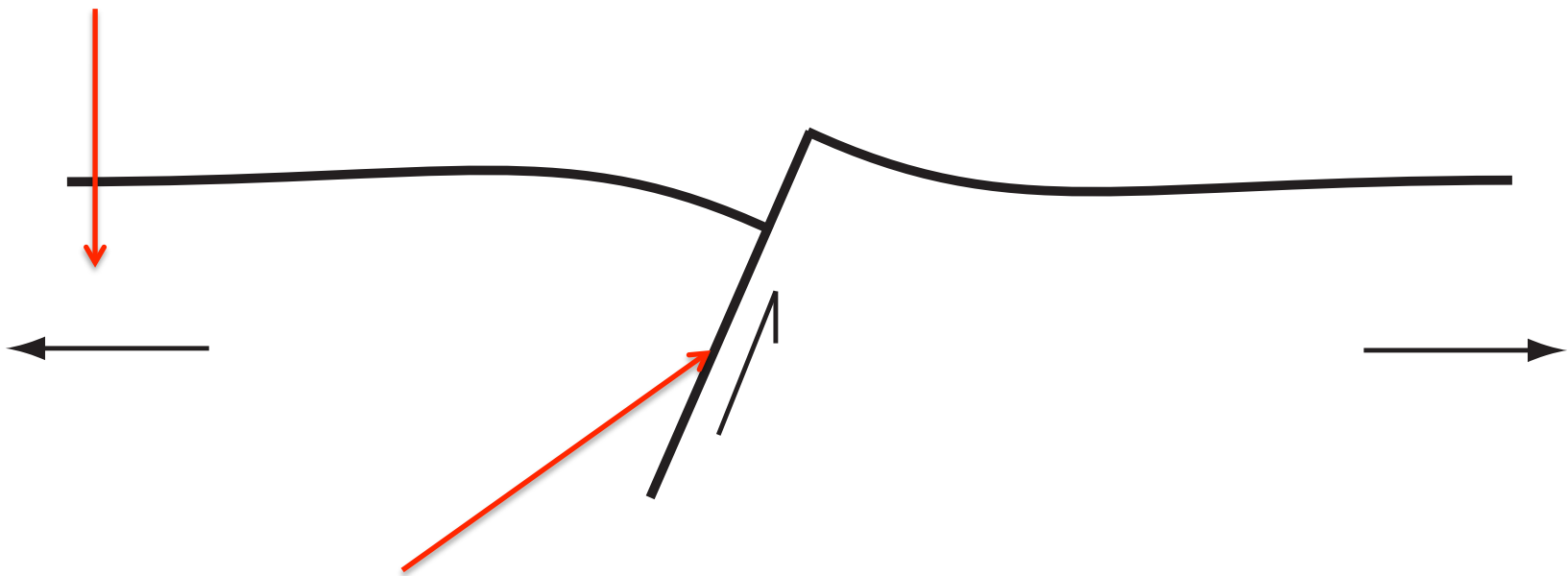
Supply extensional work



Force balance on a growing normal fault

How to keep a normal fault active ? [*Forsyth, 1992*]

Supply extensional work

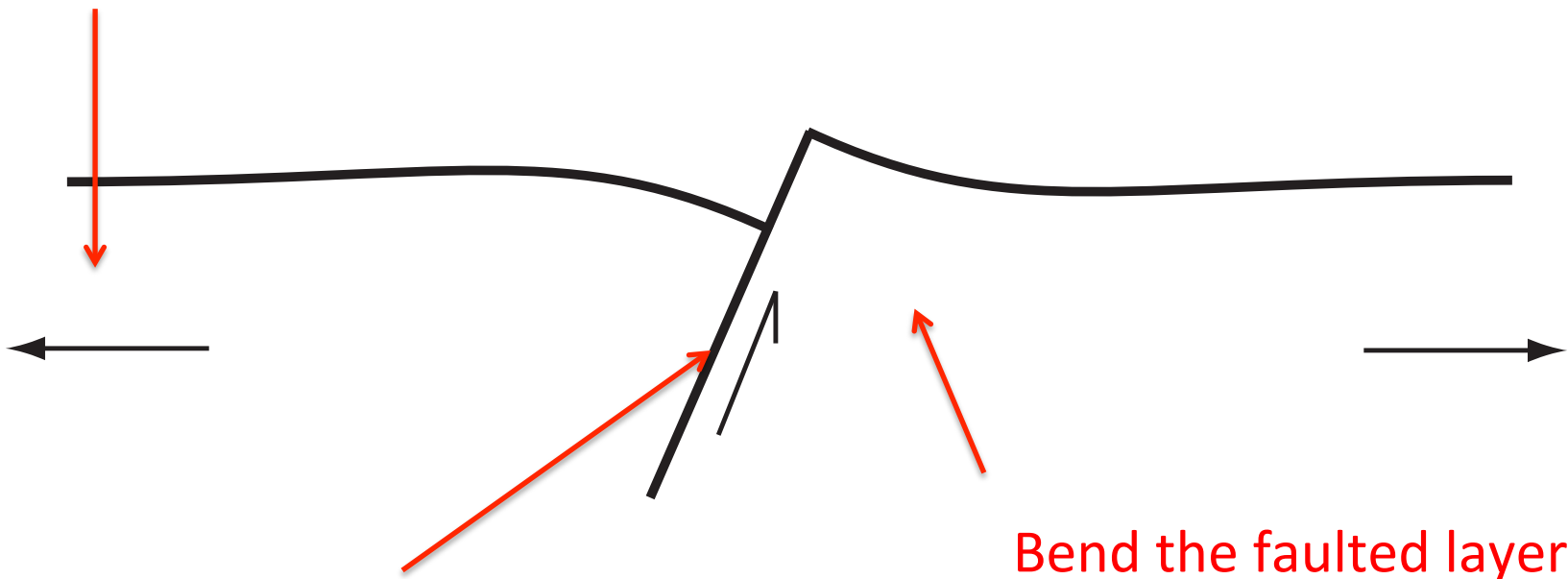


Overcome frictional
resistance on the fault

Force balance on a growing normal fault

How to keep a normal fault active ? [*Forsyth, 1992*]

Supply extensional work

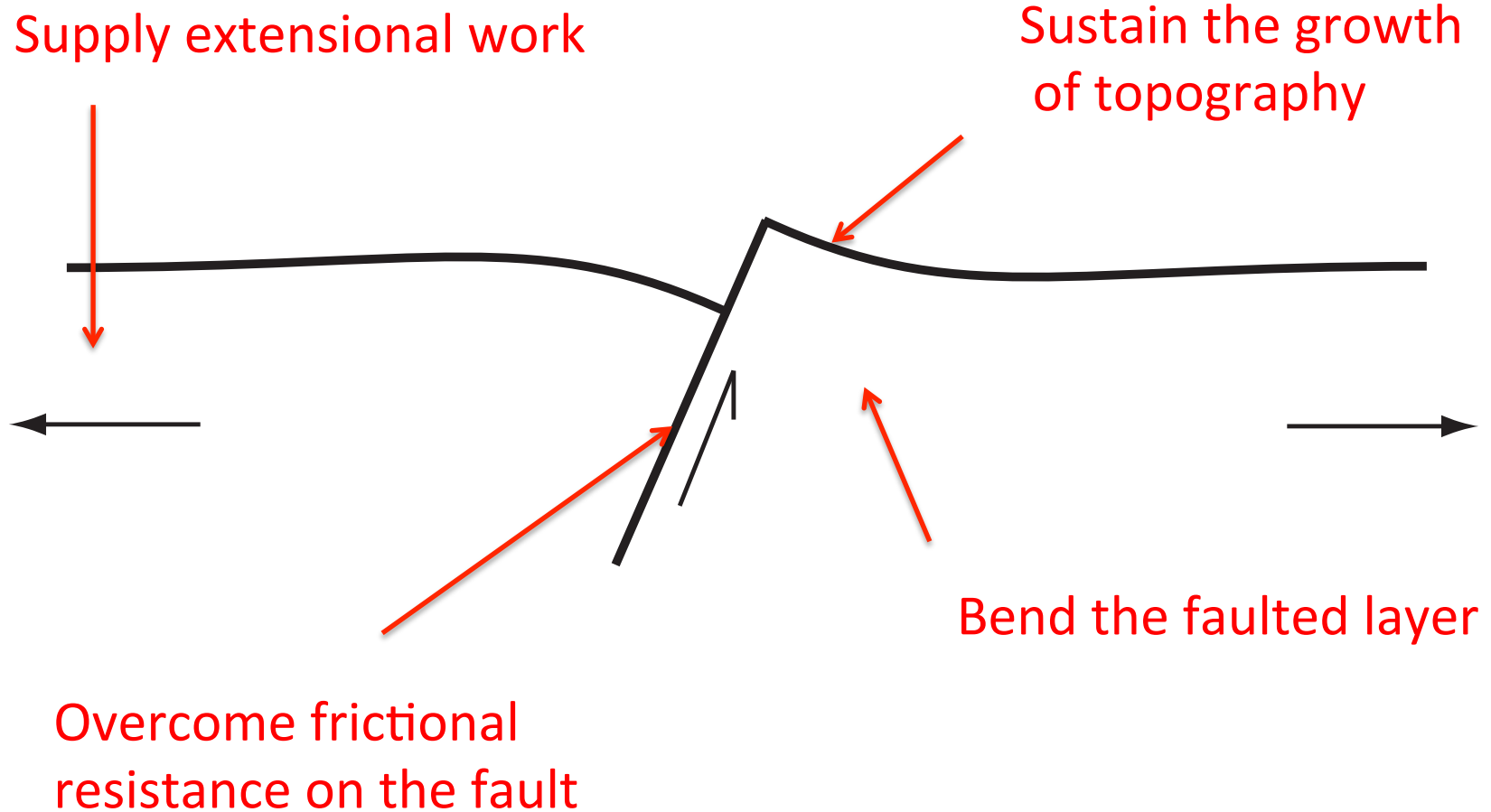


Overcome frictional
resistance on the fault

Bend the faulted layer

Force balance on a growing normal fault

How to keep a normal fault active ? [*Forsyth, 1992*]



Topographic forcing on fault life span

FORCE REQUIRED TO KEEP A FAULT ACTIVE

=

OVERCOMING FRICTION ALONG THE FAULT

+

BENDING THE FOOTWALL AND HANGING WALL

+

SUSTAINING THE GROWTH OF TOPOGRAPHY

⊕

⊖

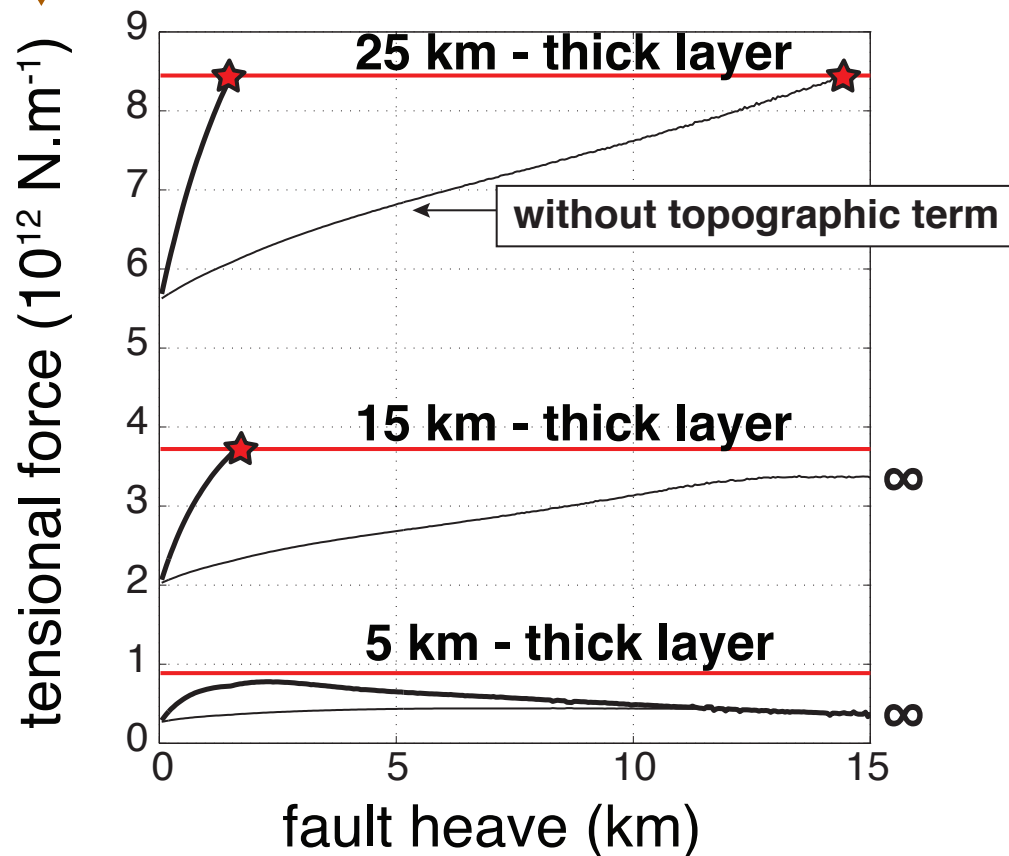
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fault rotation

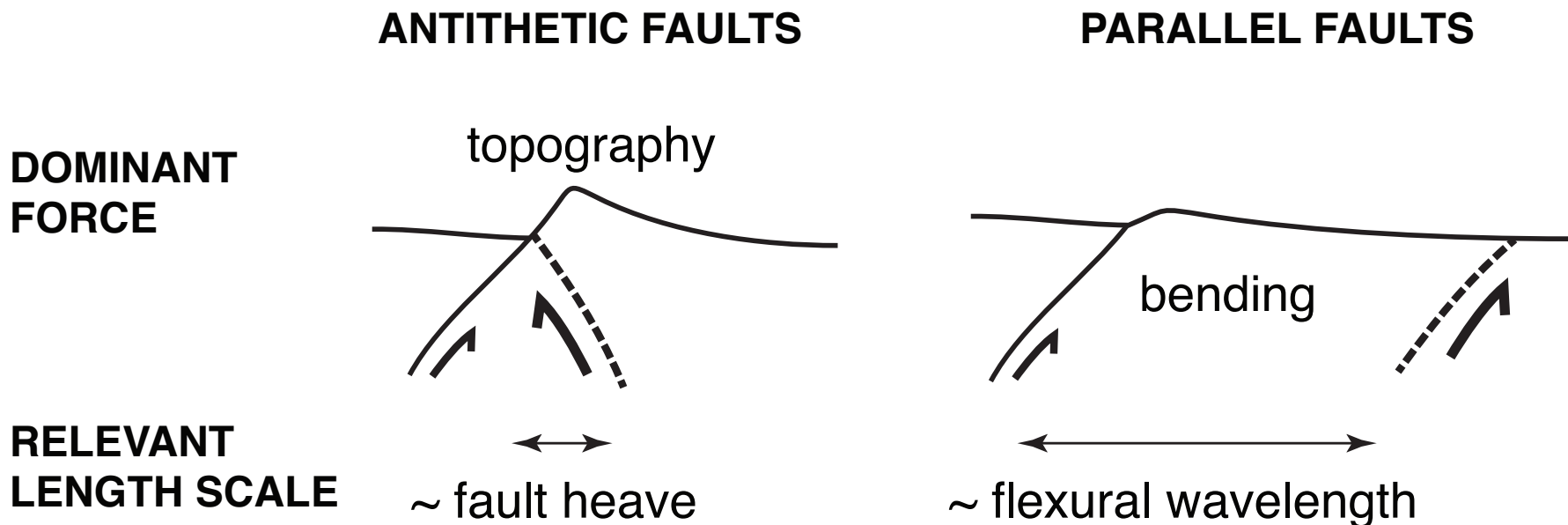
surface processes



Surface processes enhance fault life span by **relieving the energy cost** of topography build-up.

Implications for the geometry of the 2nd fault

Topography build-up favors the formation of an **antithetic second fault** in the vicinity of the initial fault.



When **surface processes suppress topography**, internal **bending** controls the location and geometry of the 2nd fault.

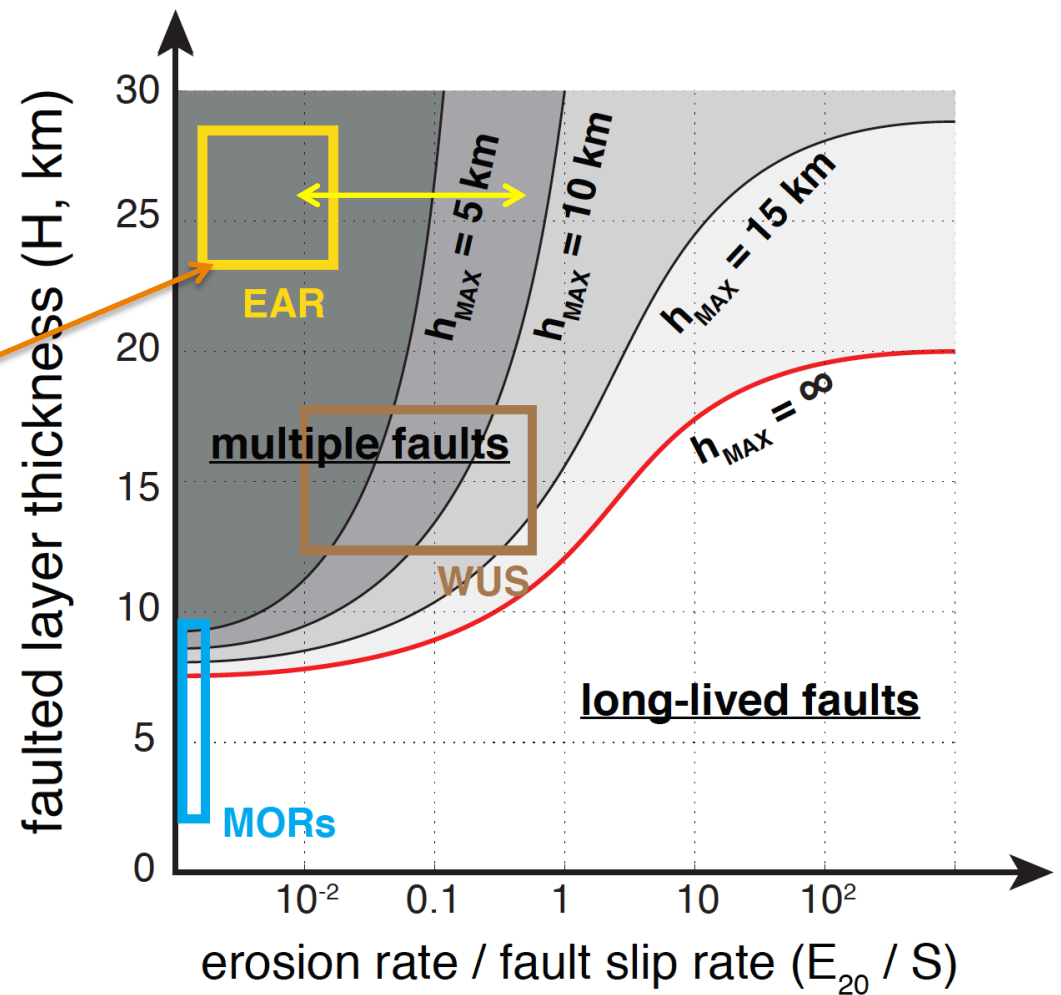
Implications for rift dynamics

Surface processes **enable the large offsets (10+ km)** observed on major range-bounding normal faults.

(These would otherwise be difficult to achieve in 15 km-thick brittle upper crust)

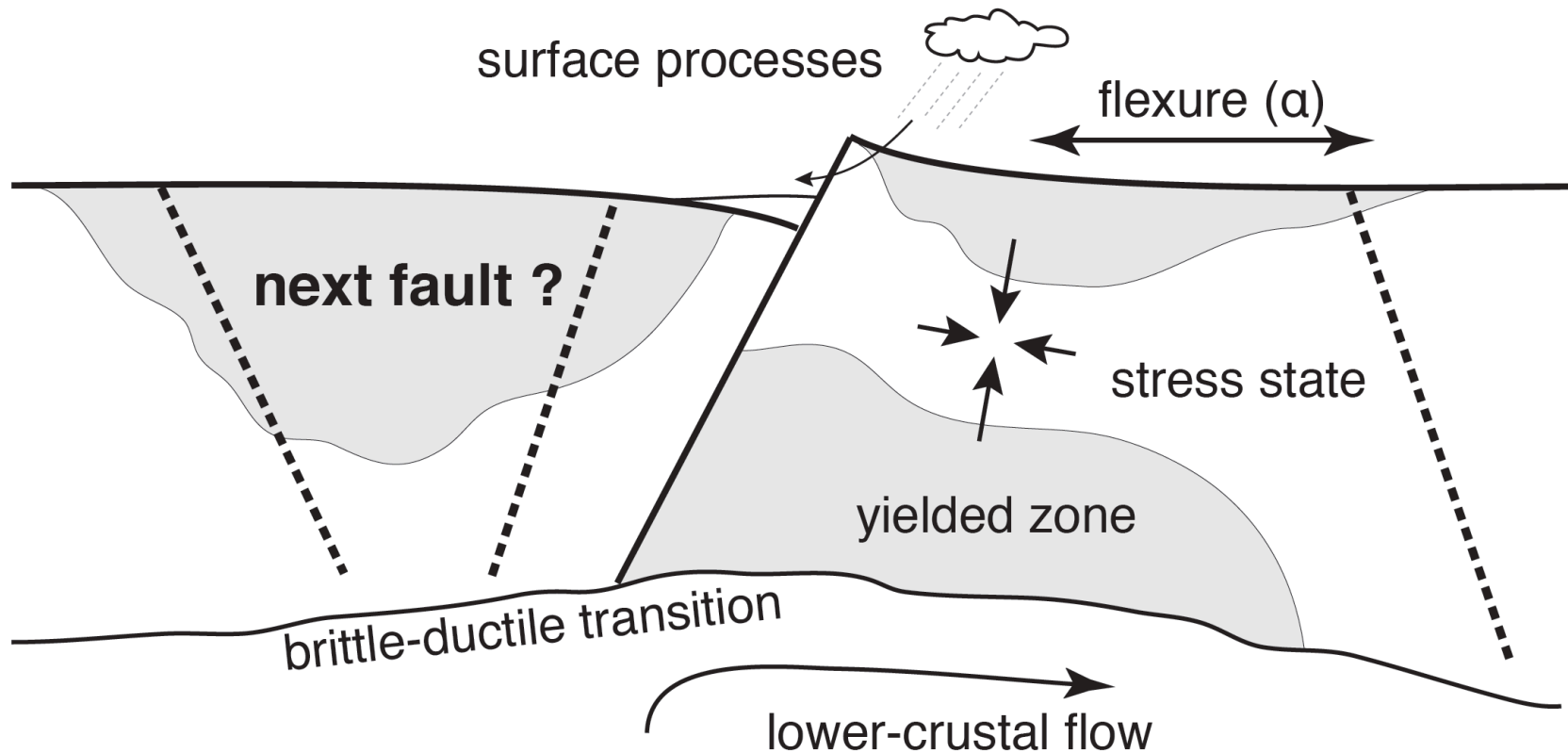


Rwenzori Mountains (Uganda)



Future directions

- **Top-down** vs. **bottom-up** controls on rift dynamics ?



- Improved models of **sediment deposition** for load redistribution
- **Model, meet data**: mechanical vs. geomorphological observables