



# Carbonate Fluxes During Earth History

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GEOMAR

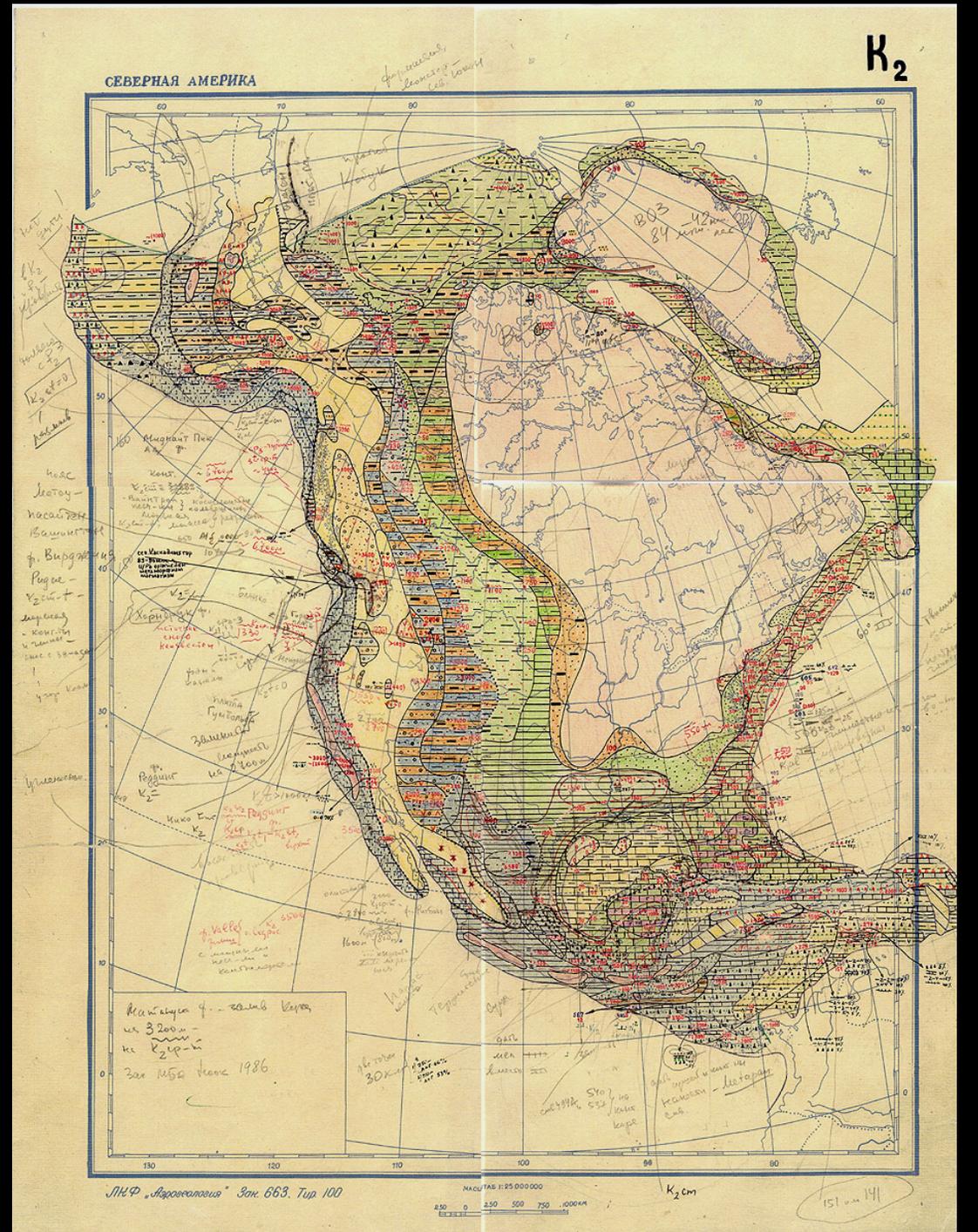
Kiel, Germany

# An Analysis of the Ronov Database

- The database was built under the direction of **Alexander Borisovitch Ronov** (affectionately known as „Sambareesh“) at the Vernadski Institute of Geochemistry, Moscow
- Started in 1947, data on areas, volumes and masses of sediments were actively added and analysed until 1995
- Areg Migdisov and Alex Balukhovsky have been actively working since 1995 to preserve the original data

The database was produced by first constructing Lithologic-Paleogeographical maps of the continents. The original (unpublished) equal-area map for the Upper Cretaceous of North America is shown here.

From the information on the map, the areas and thicknesses of each lithology were measured and the volume and mass calculated.



The compilation includes volumes of the following sediment types for each region:

Continental terrigenous sediment

Coal-bearing terrigenous sediment

Glacial sediment

Molasse

Marine sands

Marine shales

Marine terrigenous sediment

Flysch

Carbonate-terrigenous sediment (Marls)

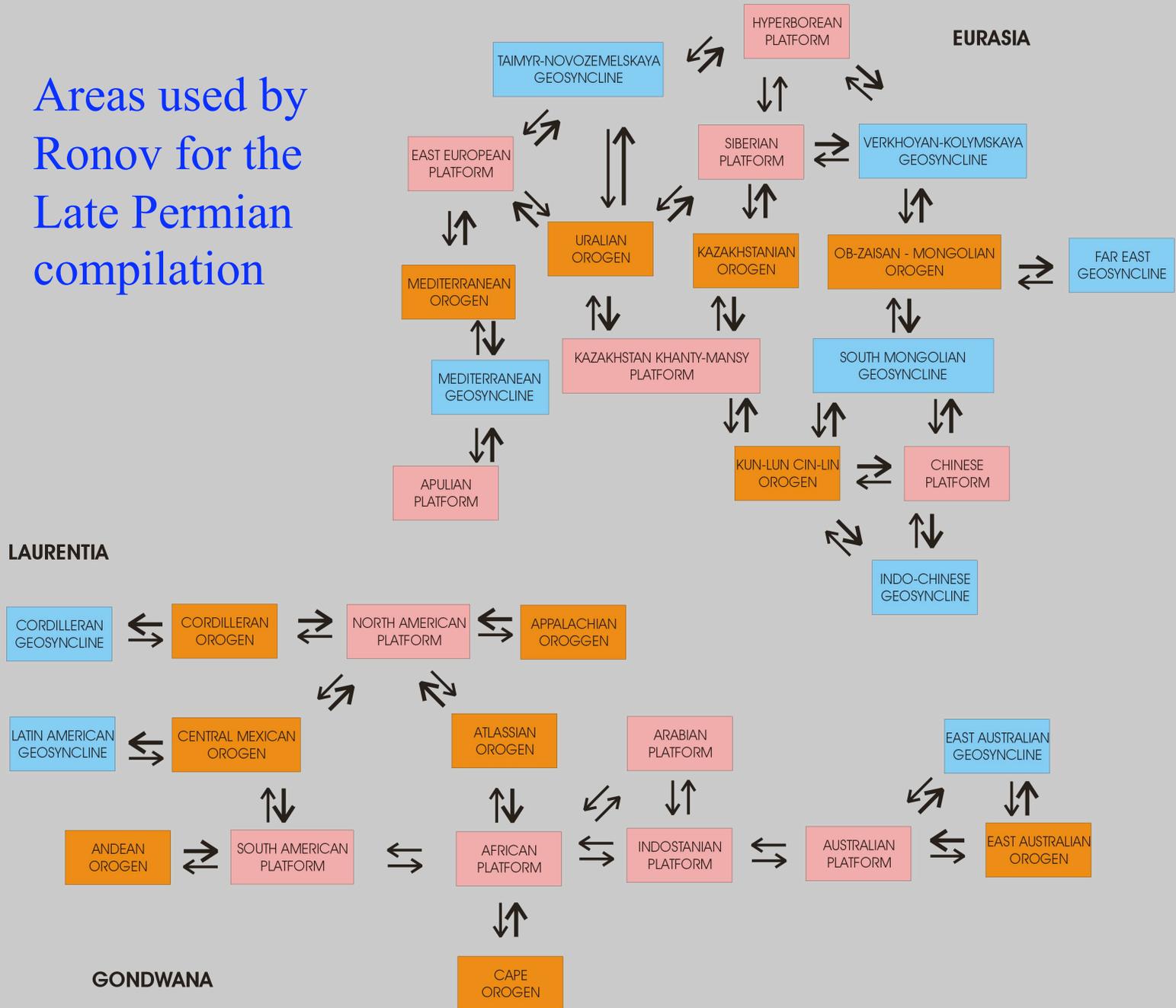
Carbonates

Halite

Gypsum/Anhydrite

Siliceous sediment

# Areas used by Ronov for the Late Permian compilation





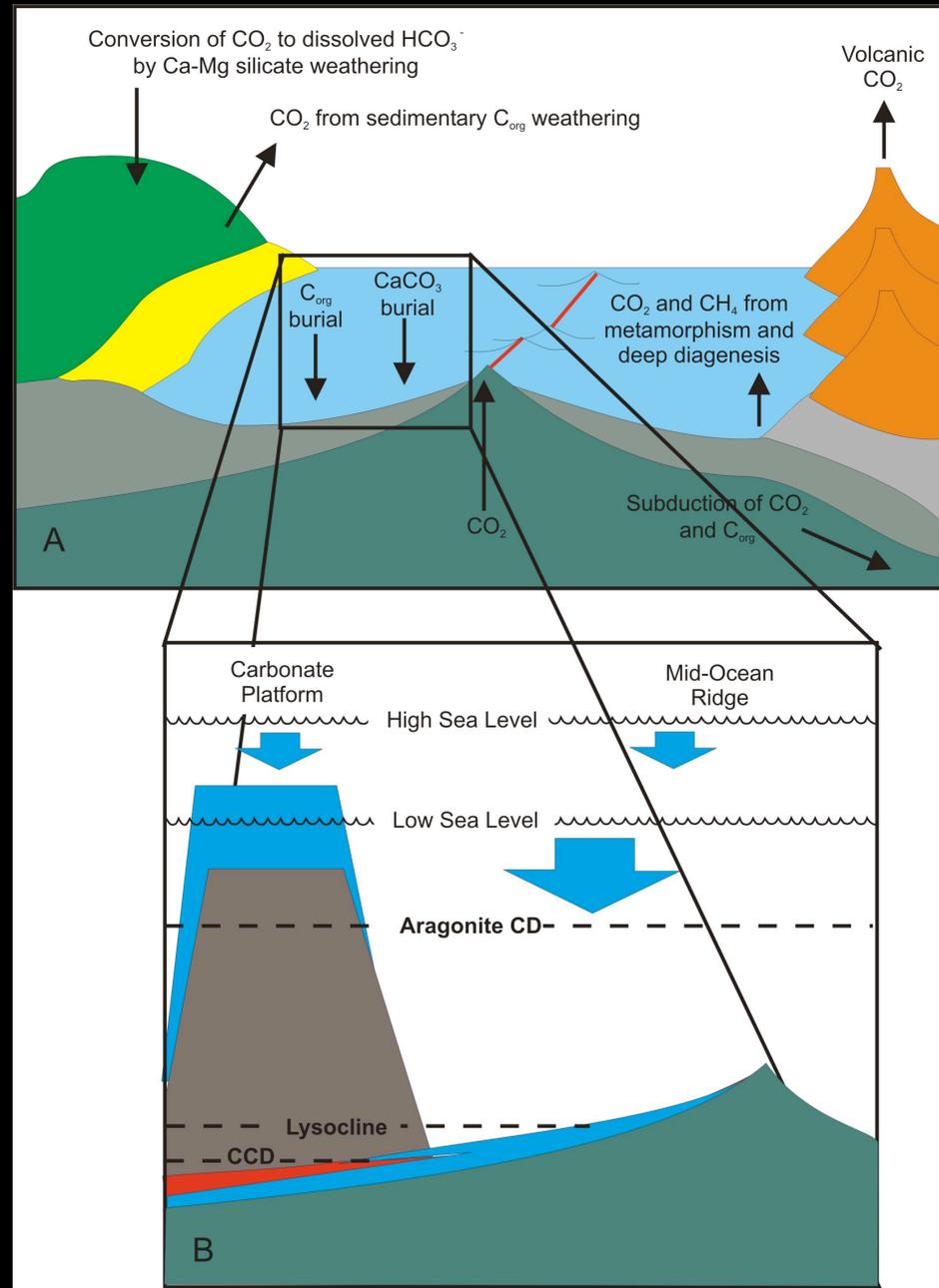
# Detrital and dissolved loads of rivers are deposited in different ways

- The detrital load (sand, silt, clay, etc. settles is deposited in tvalleys, the coastal plain, and settles onto the sea floor
- The dissolved load can be stored in the ocean for short (carbonate) or long (salt) periods of time and its site of deposition is unrelated to where it originally entered the sea

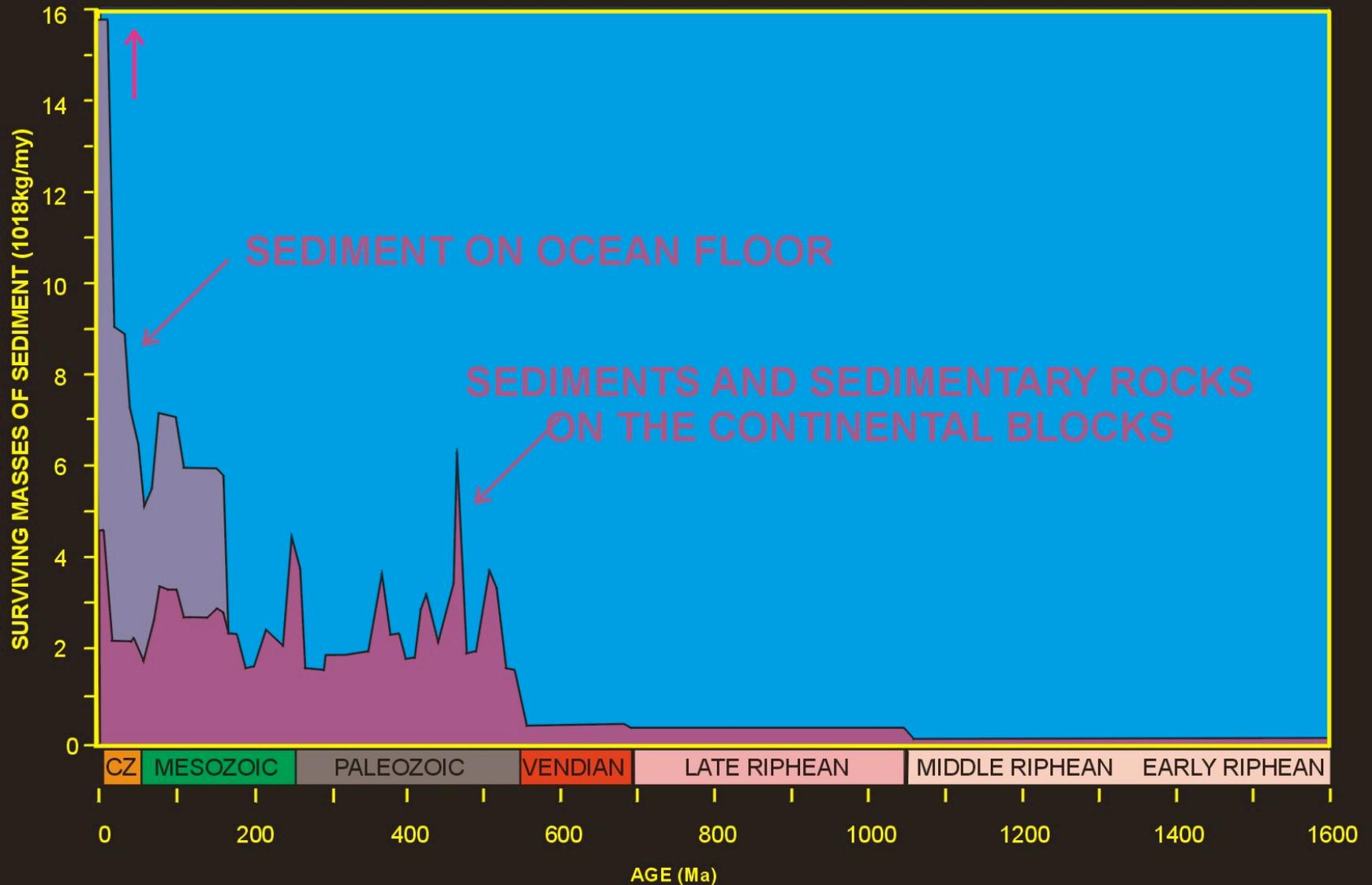
For the dissolved load -  
On the long term: what  
goes in must come out.

But . . .

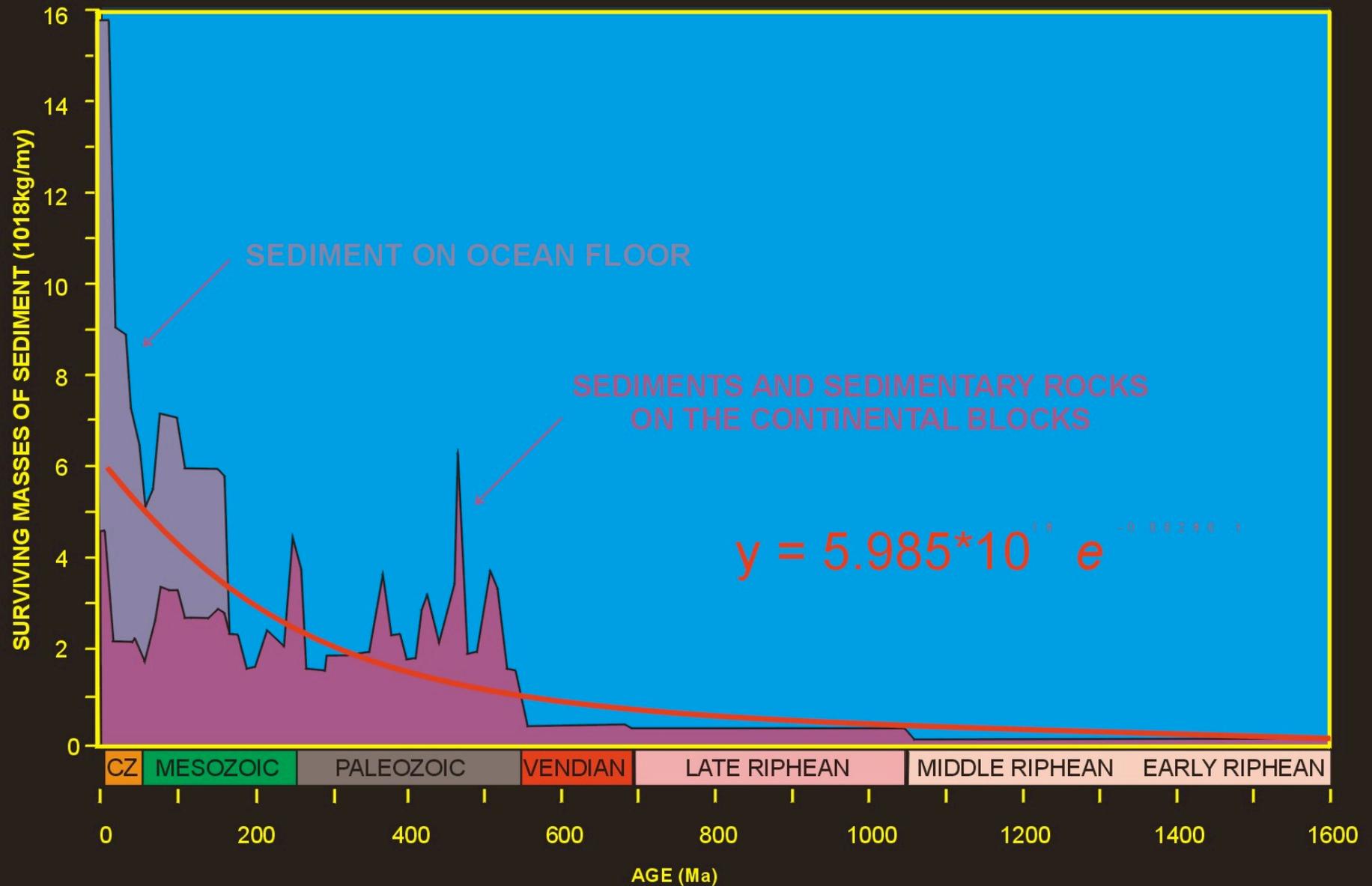
Over the past 100 my  
the calcareous  
nannoplankton have  
become an important  
part of the internal  
workings of the carbon  
cycle



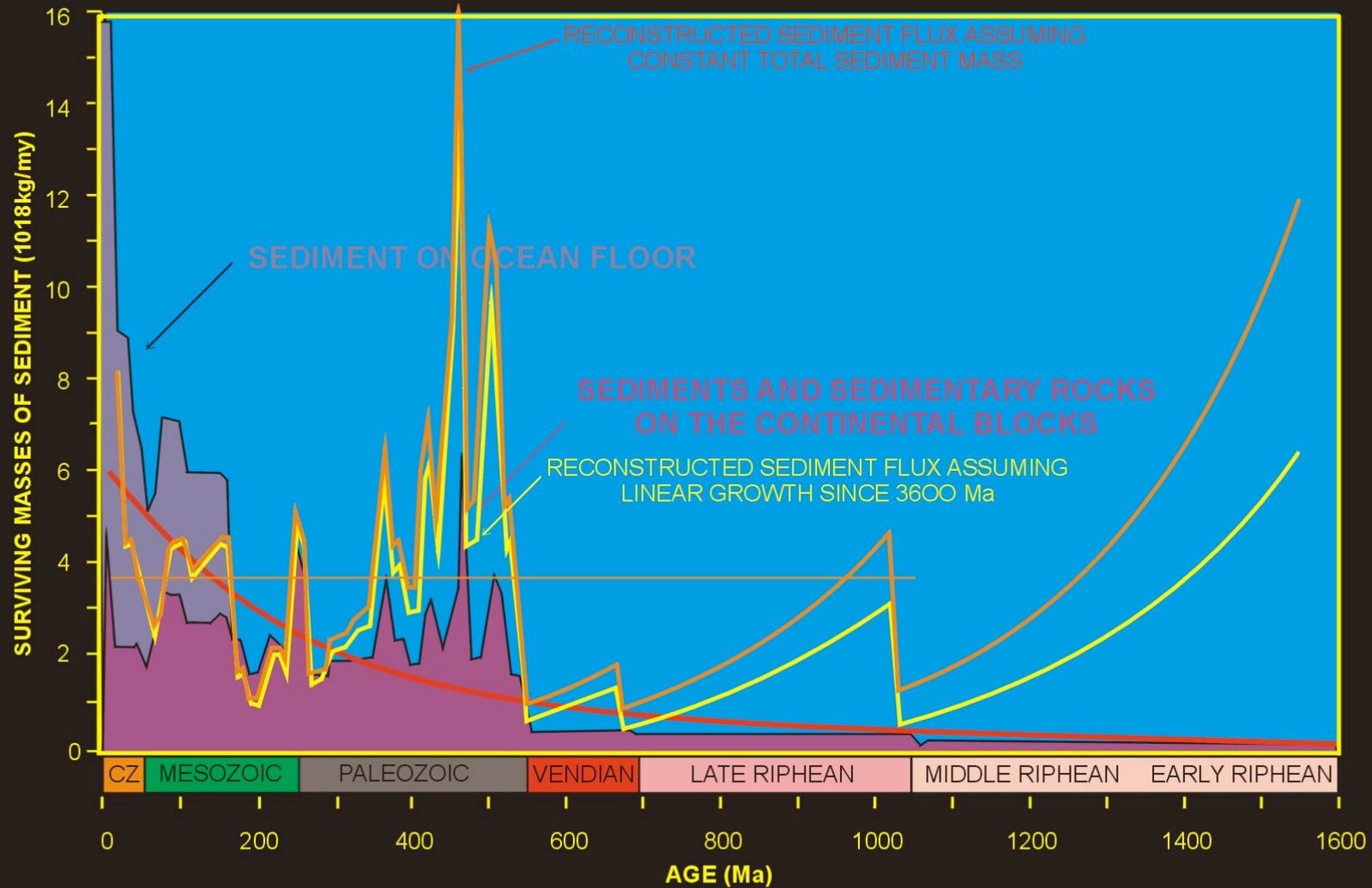
# The Sediment Existing Today



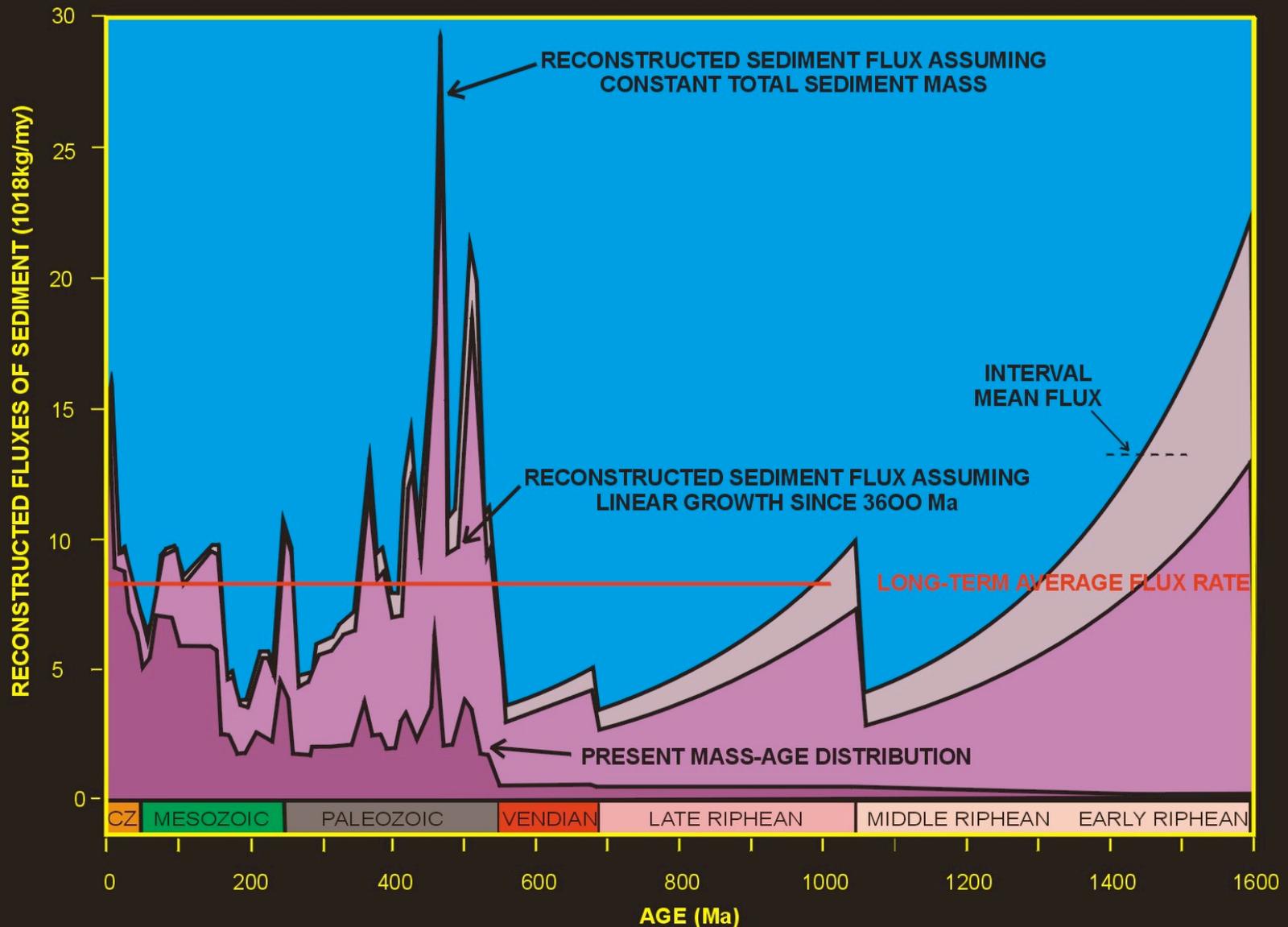
# The Global Recycling Rate of Sediment



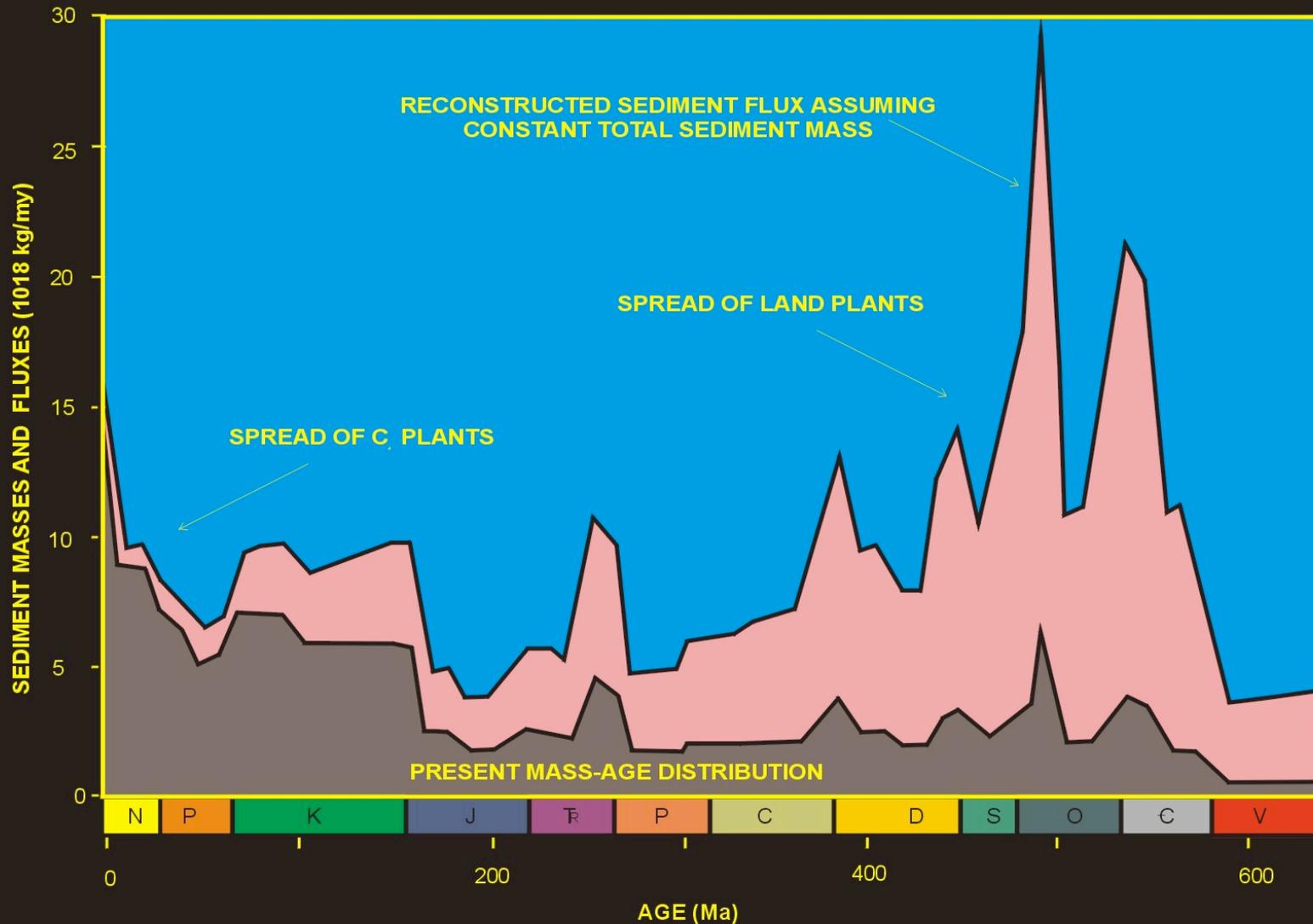
# Reconstructed Sediment Fluxes



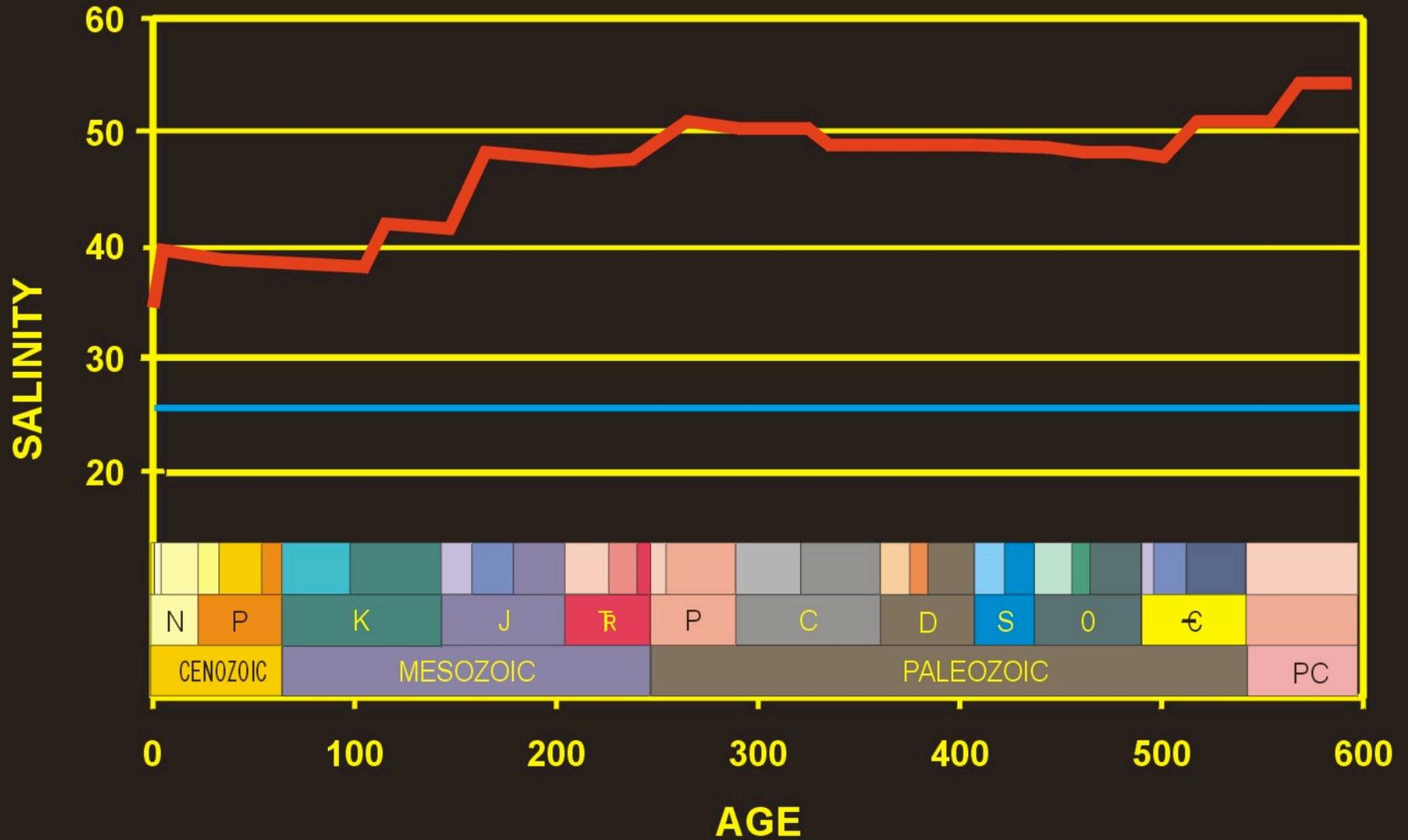
# The Long Term Average Sediment Flux



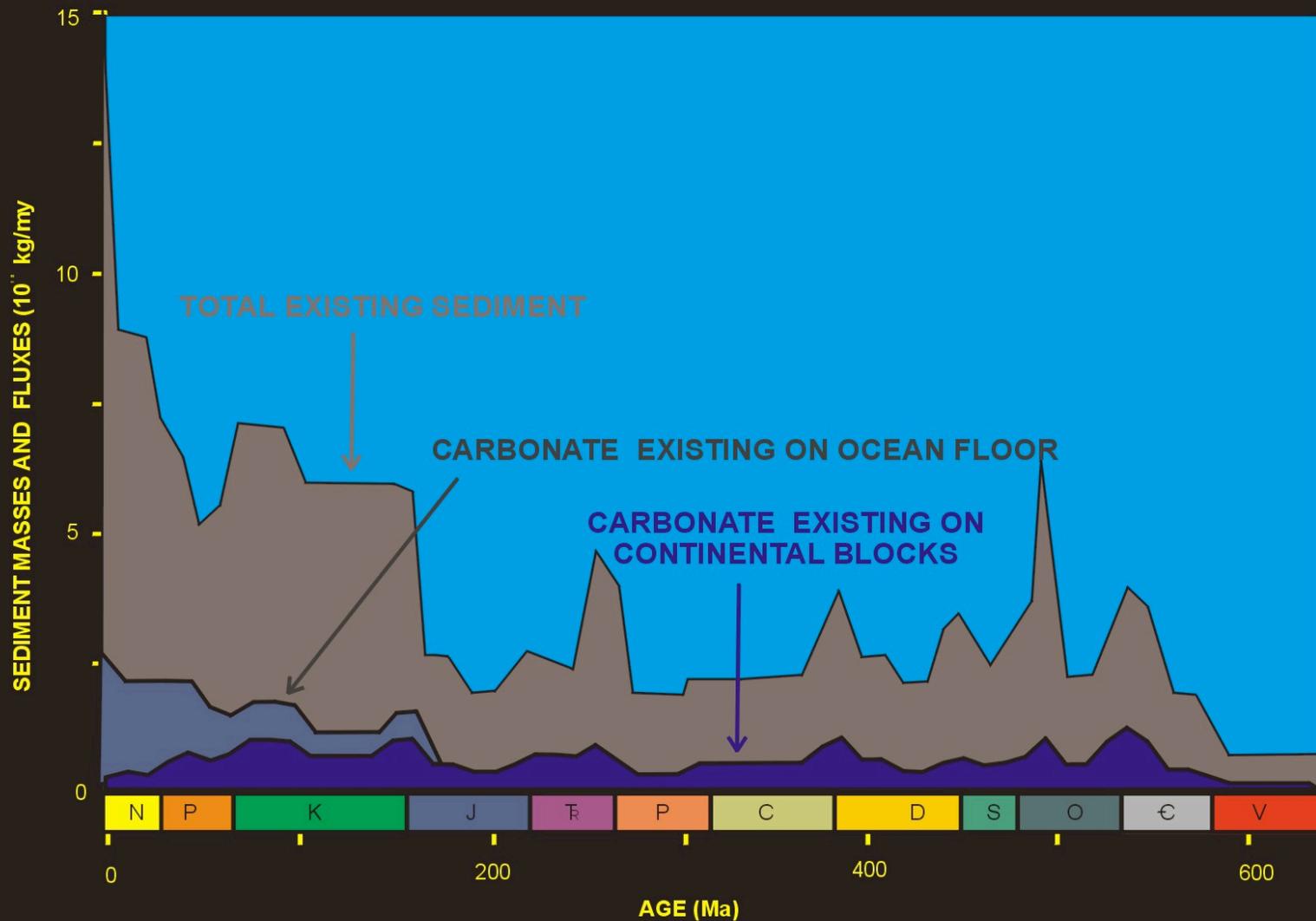
# The Phanerozoic Sediment Flux



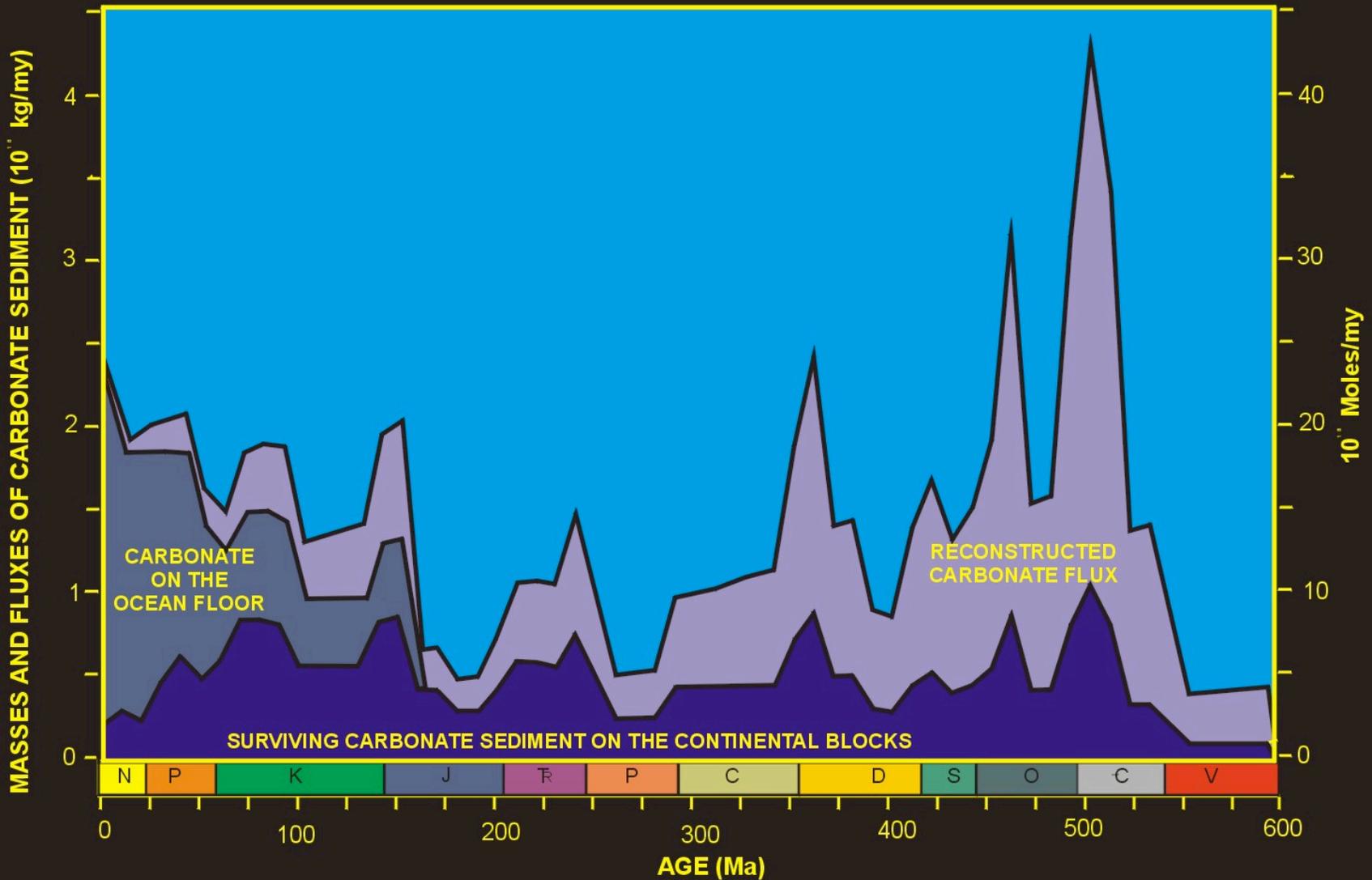
# GLOBAL MEAN SALINITY



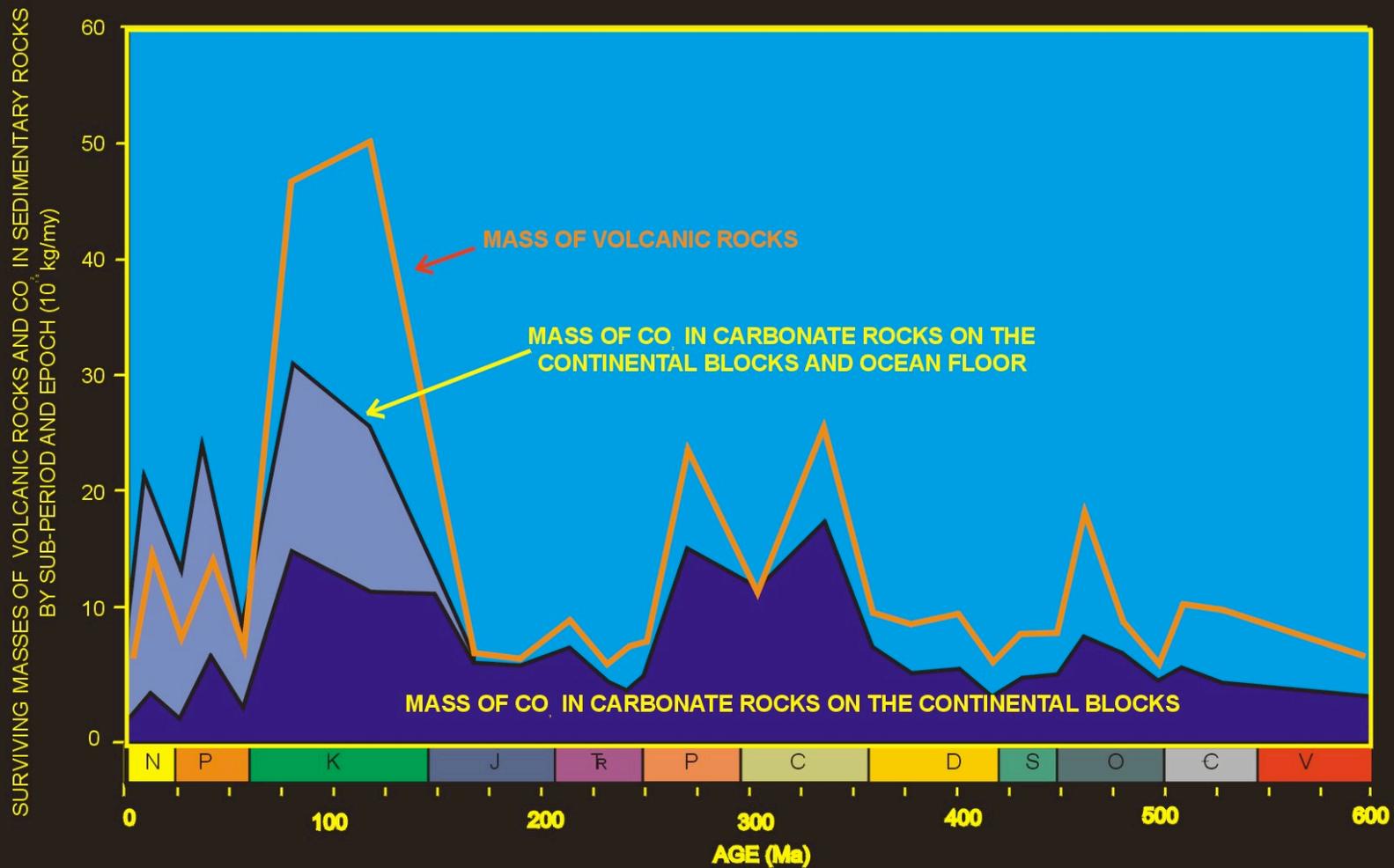
# Carbonate is about 1/4 of the Existing Phanerozoic Sediment



# Phanerozoic Carbonate Flux



# The original idea of Budyko and Ronov – that carbonate deposition is linked to volcanism via CO<sub>2</sub> emission may not be so bad



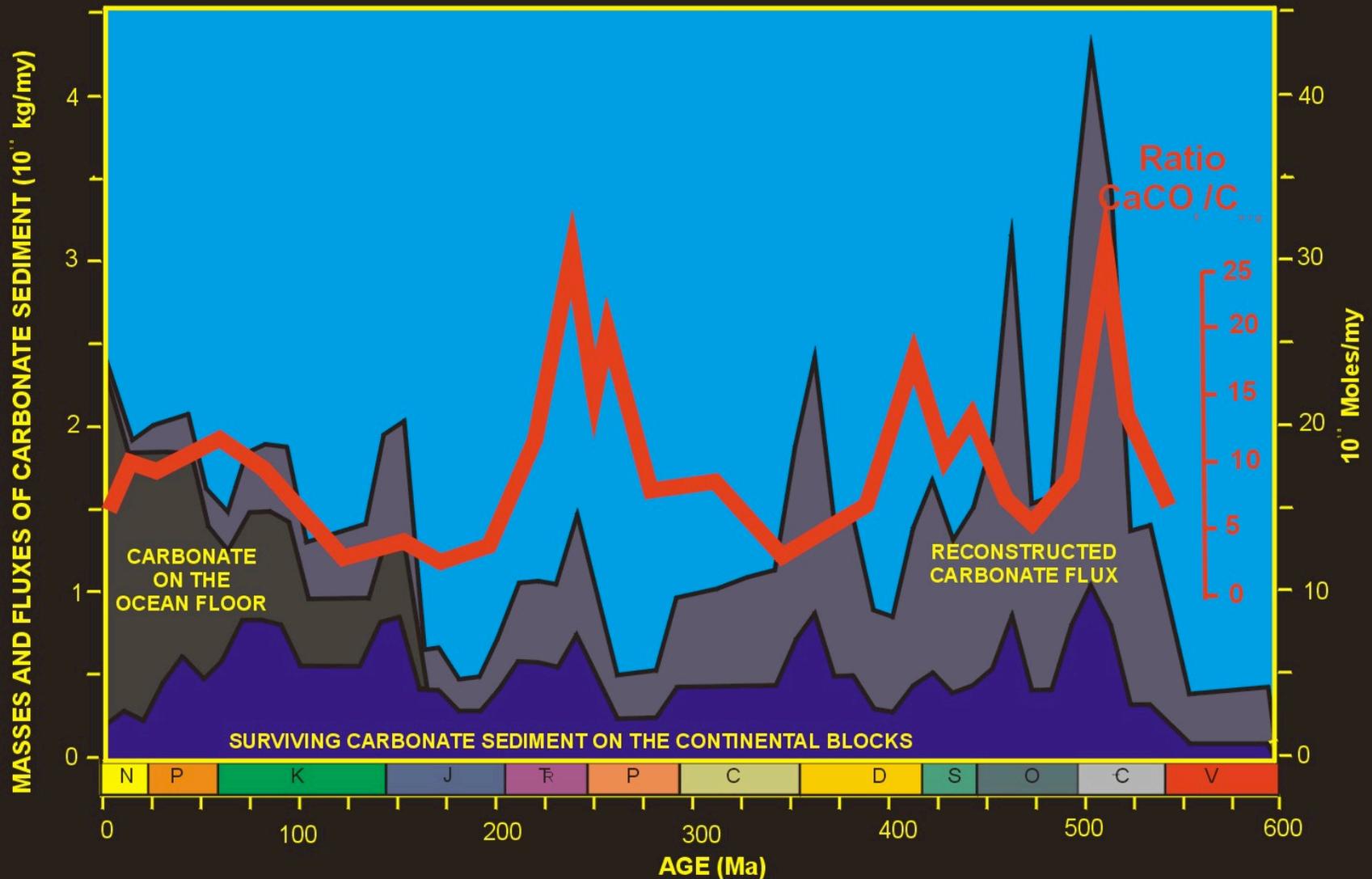
# The relative proportions of $\text{CaCO}_3$ and $\text{C}_{\text{org}}$

- In modern plankton, ( $\gg$ ) 4 moles of C are fixed as  $\text{C}_{\text{org}}$  for every mole fixed as  $\text{CaCO}_3$

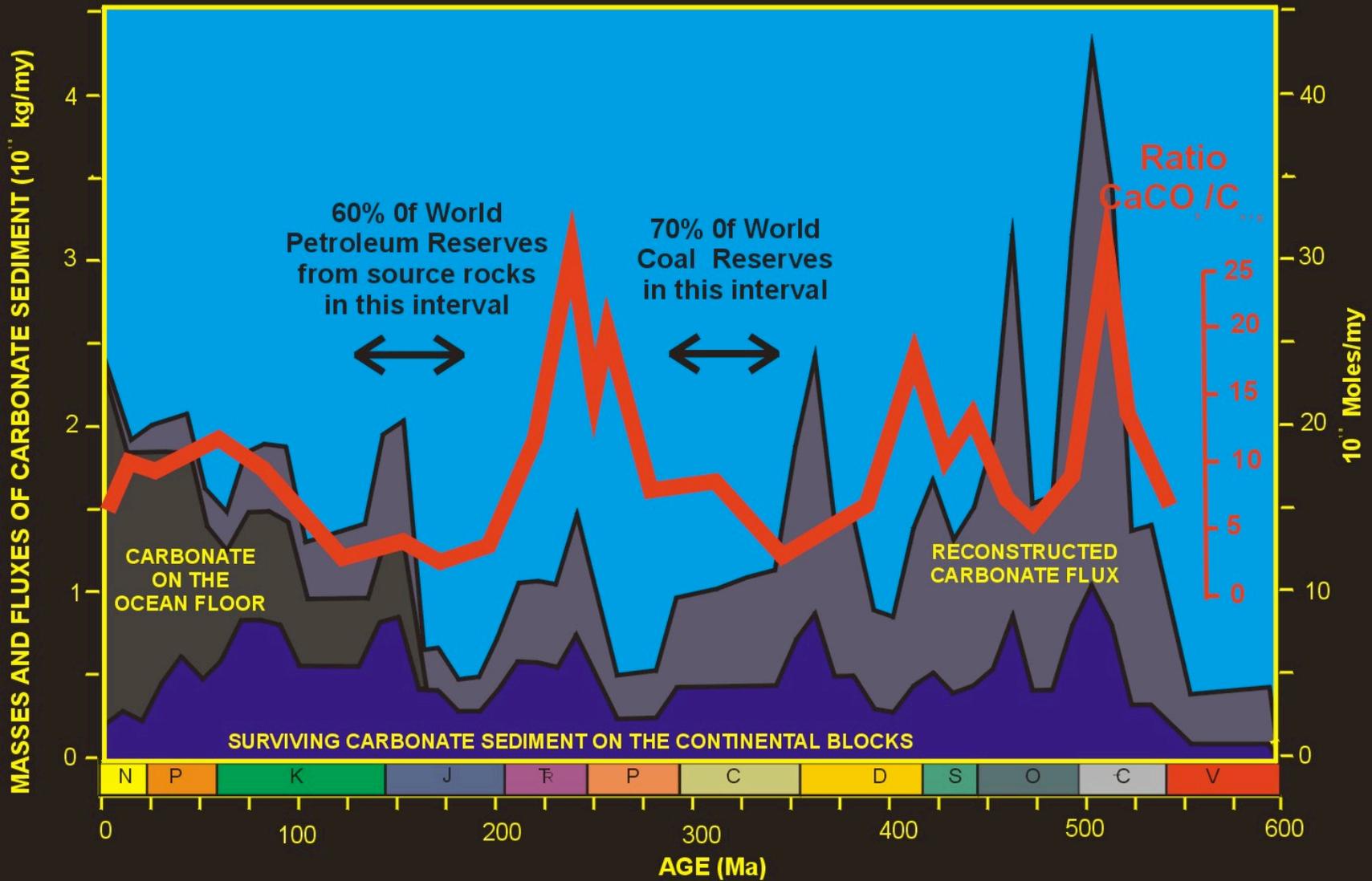
At present about 6 moles of C are buried as  $\text{CaCO}_3$  for every mole buried as  $\text{C}_{\text{org}}$

- This ratio has generally been higher in the past

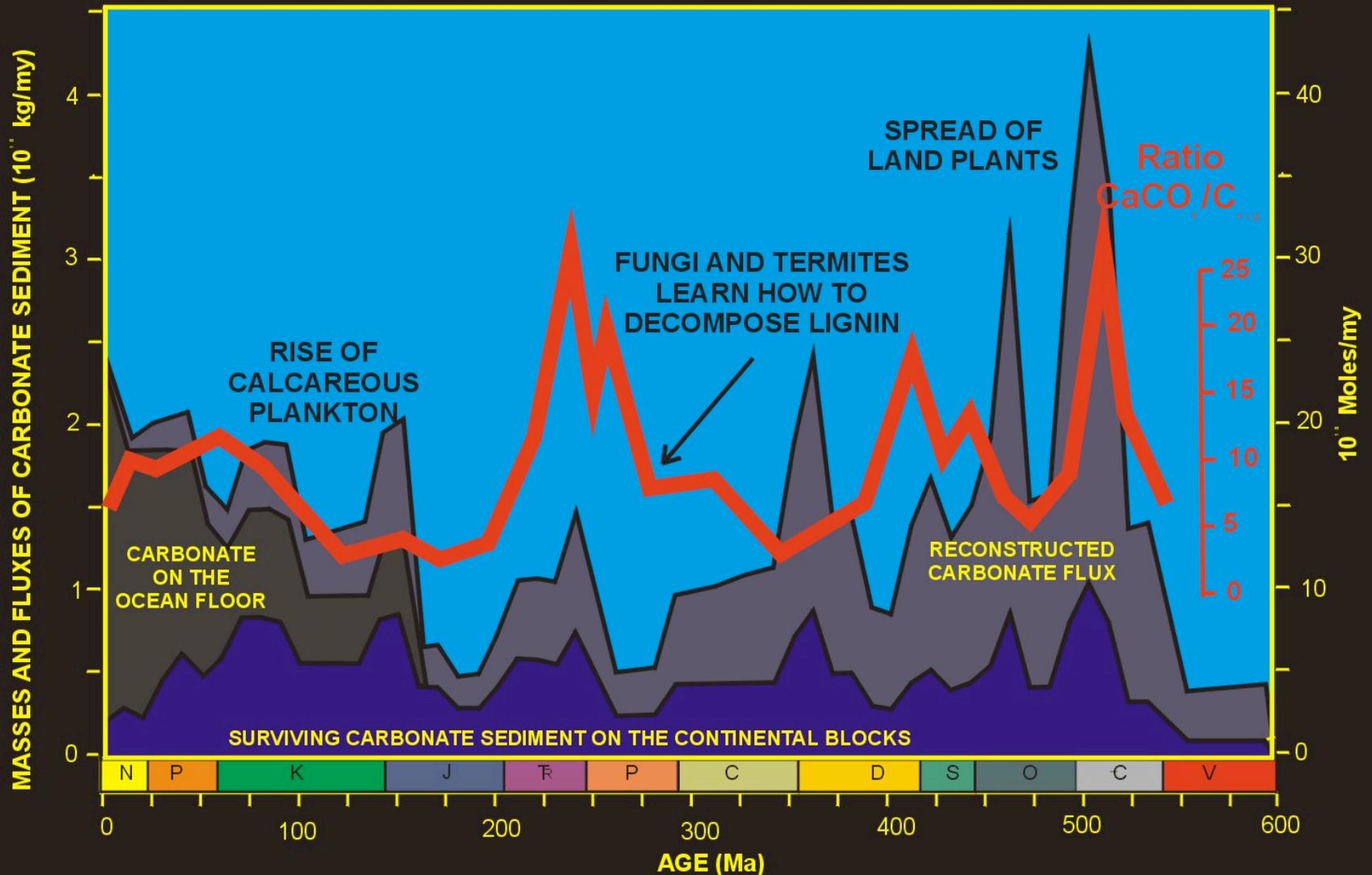
# The ratio of C in $\text{CaCO}_3$ to $\text{C}_{\text{org}}$



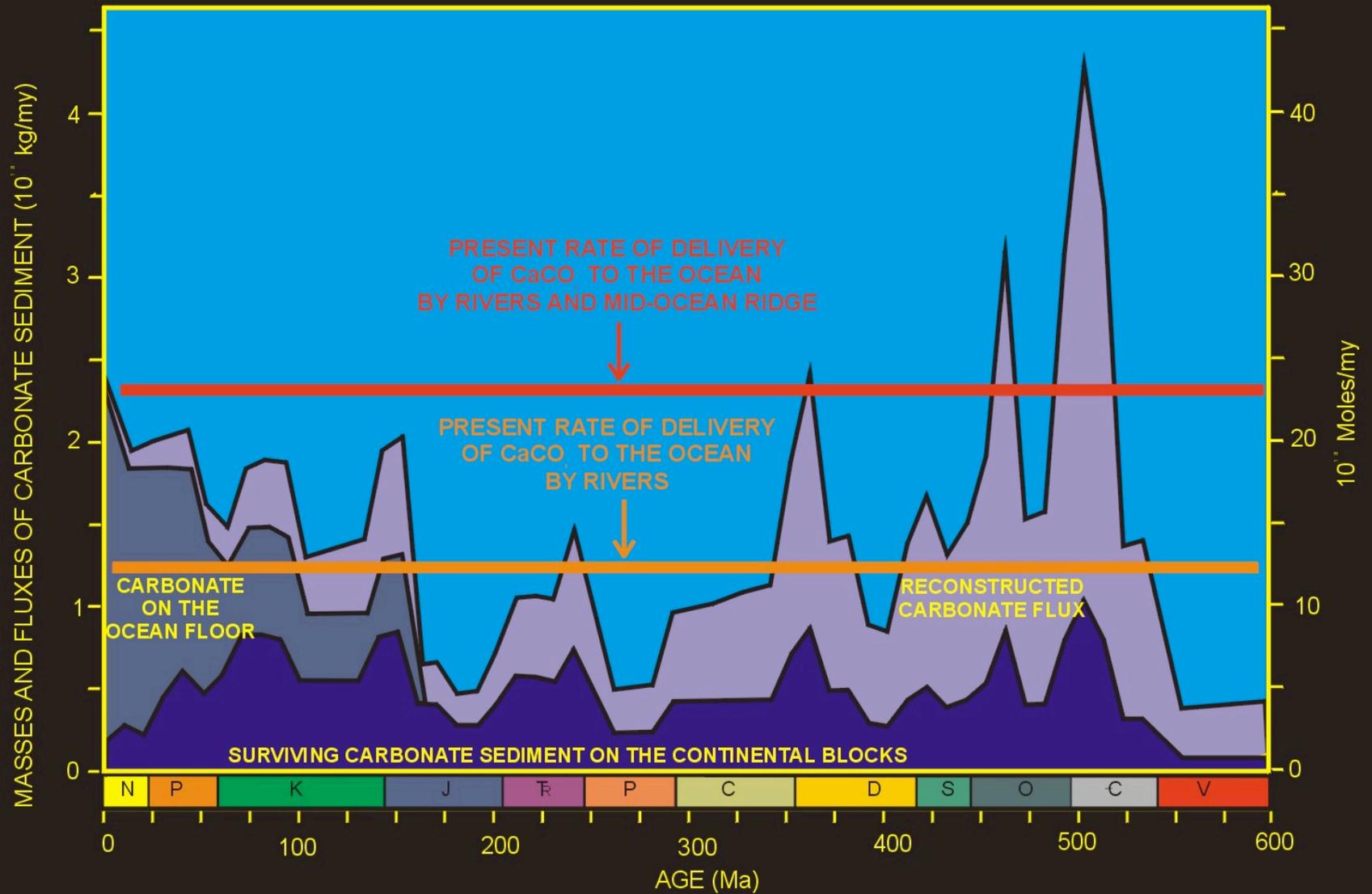
# Deposition of $C_{org}$



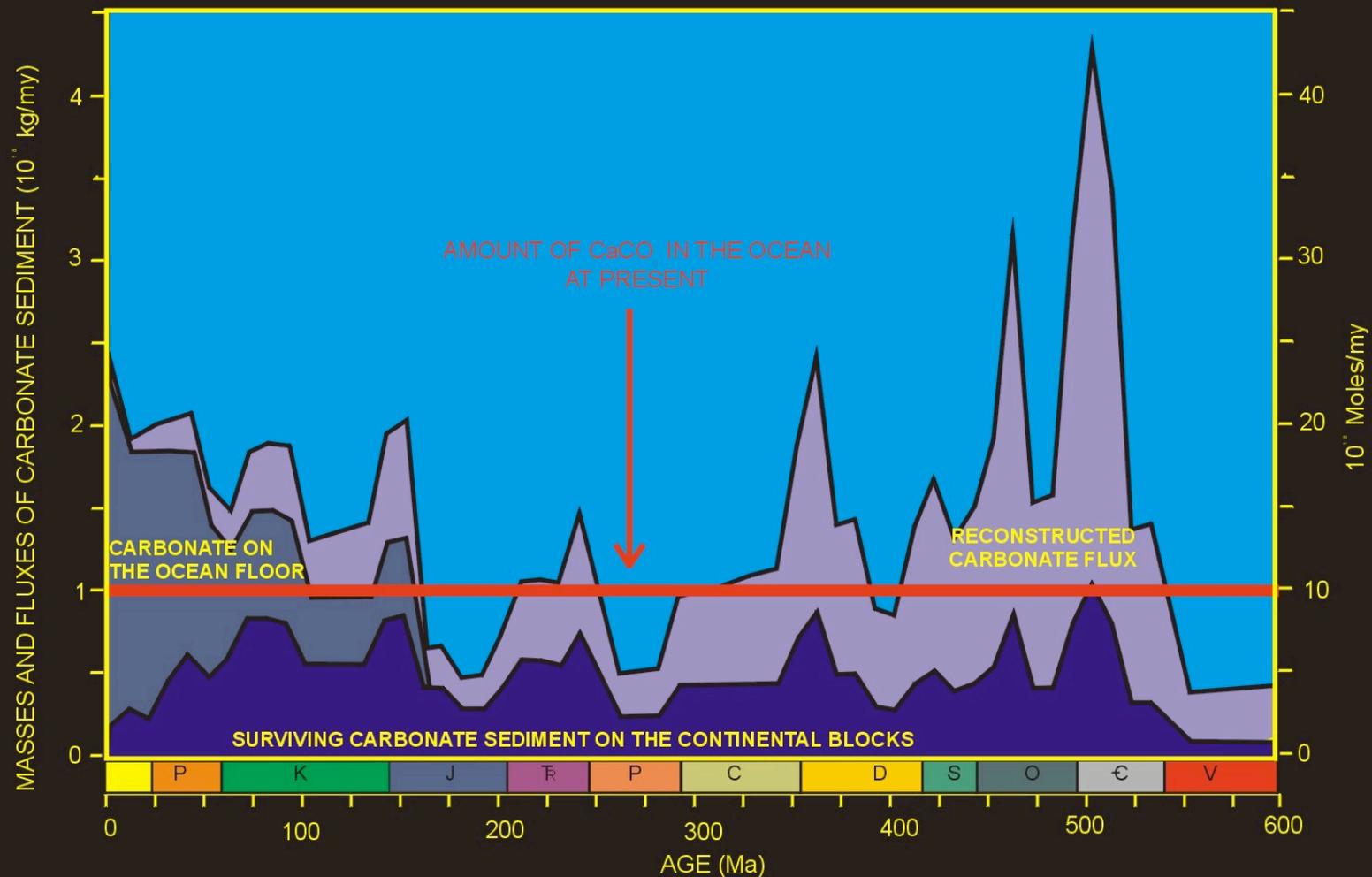
# Major Events in $C_{org}$ History



# Total Carbonate Flux



# Carbonate stored in the ocean is insufficient to maintain output



# The Mesozoic – Cenozoic Story

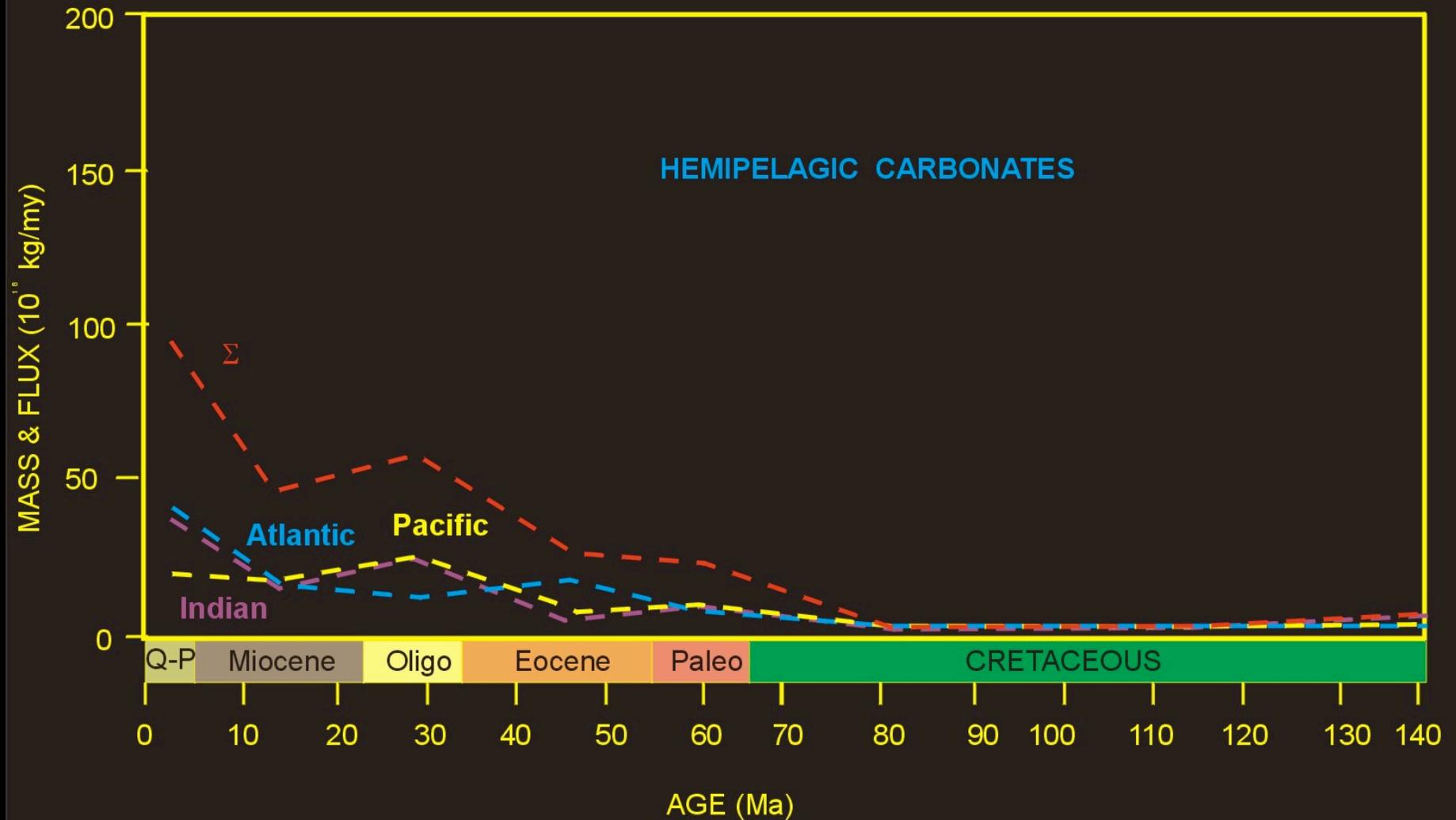
- The major event is a shift of the site of carbonate deposition from the shallow seas to the deep sea
- This was brought on by the development of the calcareous plankton
- The spread of calcareous plankton may be related to declining ocean salinity

# Coccoliths as part of the Carbon Flux

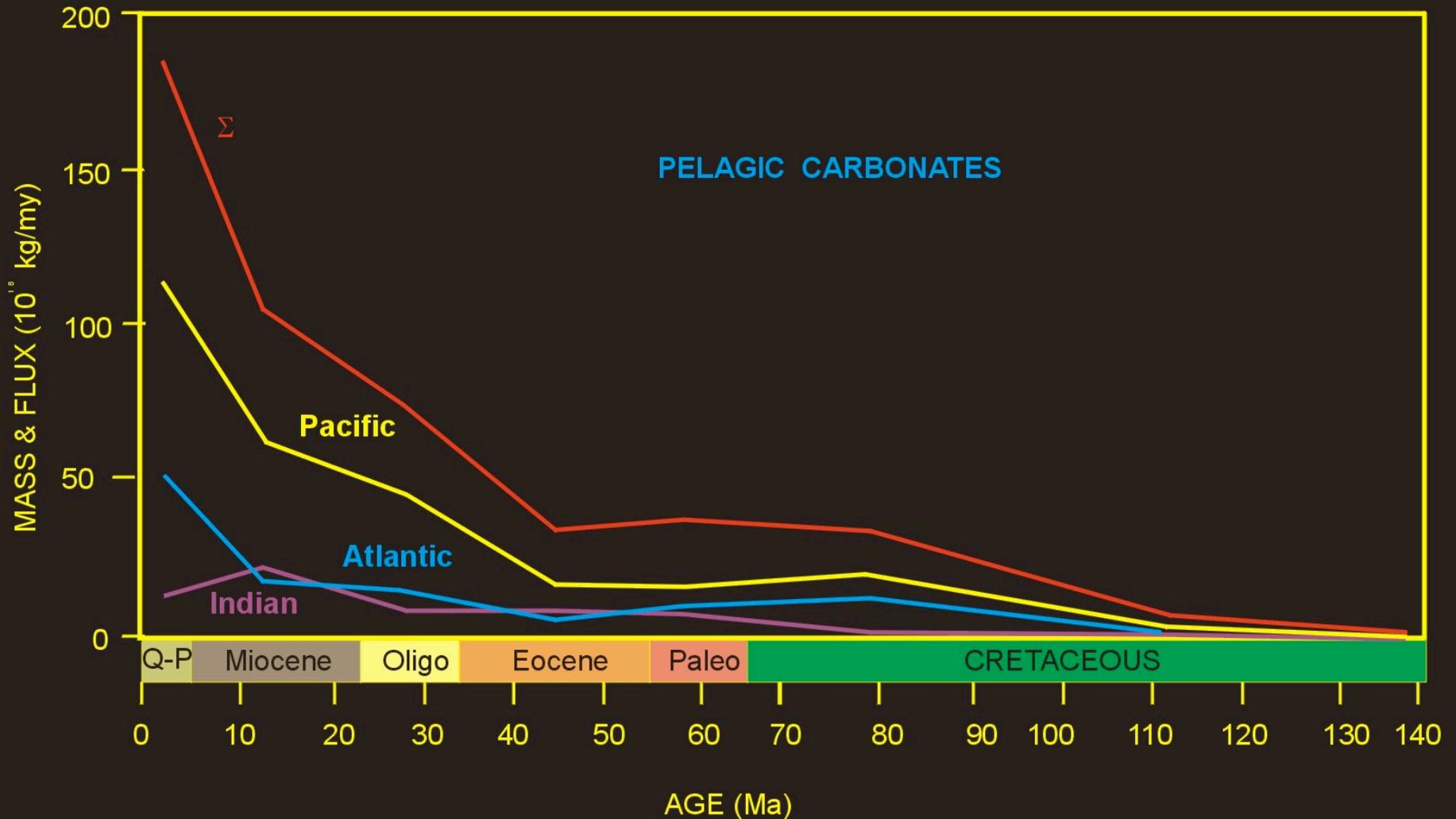
- Coccoliths are part of the pelagic  $\text{CaCO}_3$  Flux
  - Pteropod shells
  - Planktonic foraminiferal tests
  - Coccoliths
- The oceanic  $\text{CaCO}_3$  Flux is part of the  $\Sigma \text{CaCO}_3$  Flux
  - Pelagic  $\text{CaCO}_3$  Flux
  - Benthic  $\text{CaCO}_3$  Flux
- The  $\text{CaCO}_3$  Flux is part of the  $\Sigma$  Carbon Flux
  - $\text{CaCO}_3$
  - Organic carbon

The Carbon Flux is part of the  $\Sigma$  Sediment Flux

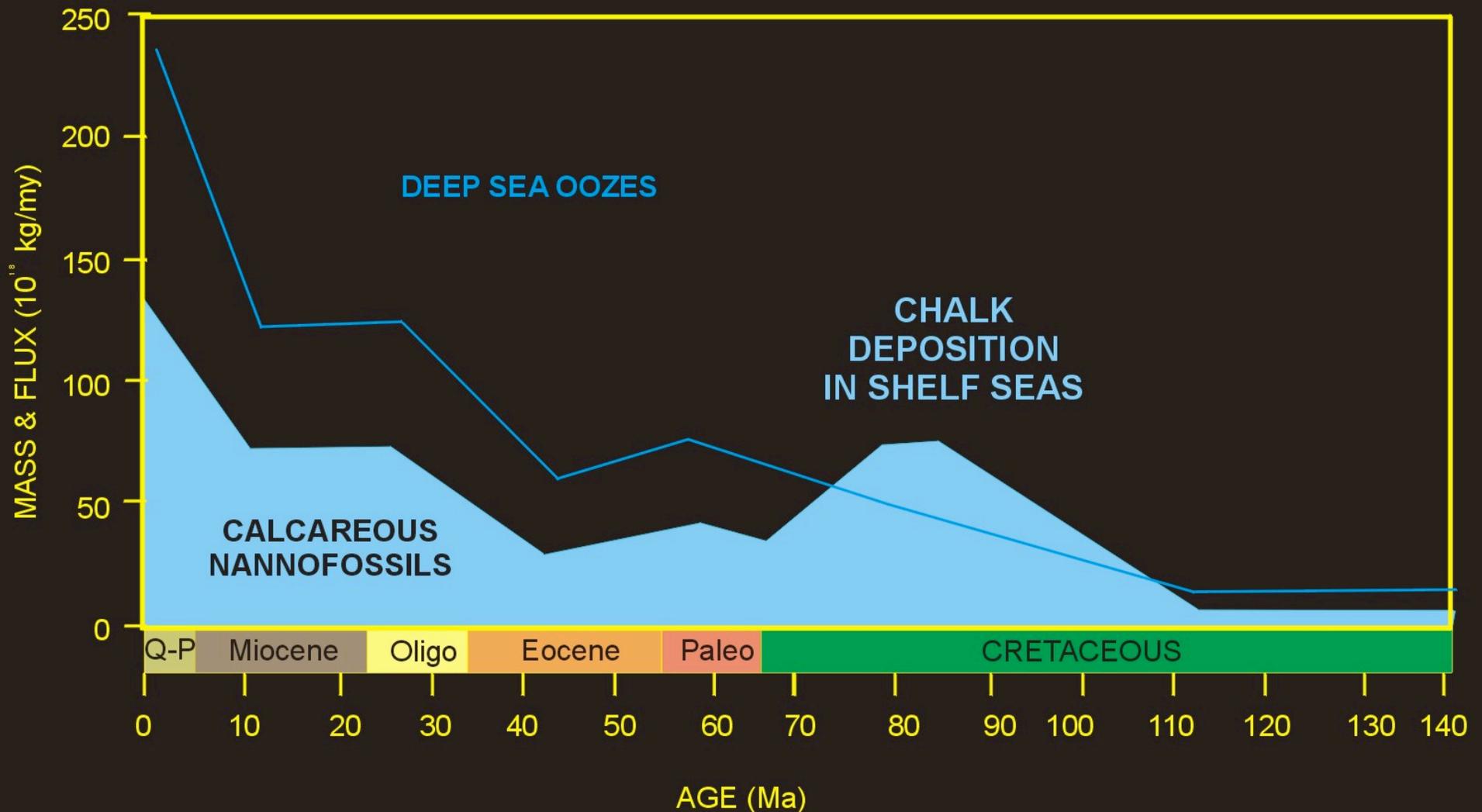
# The Increasing Hemipelagic Flux



# The Increasing Pelagic Flux



# Calcareous Nannoplankton and the Shift from Shallow to Deep Water



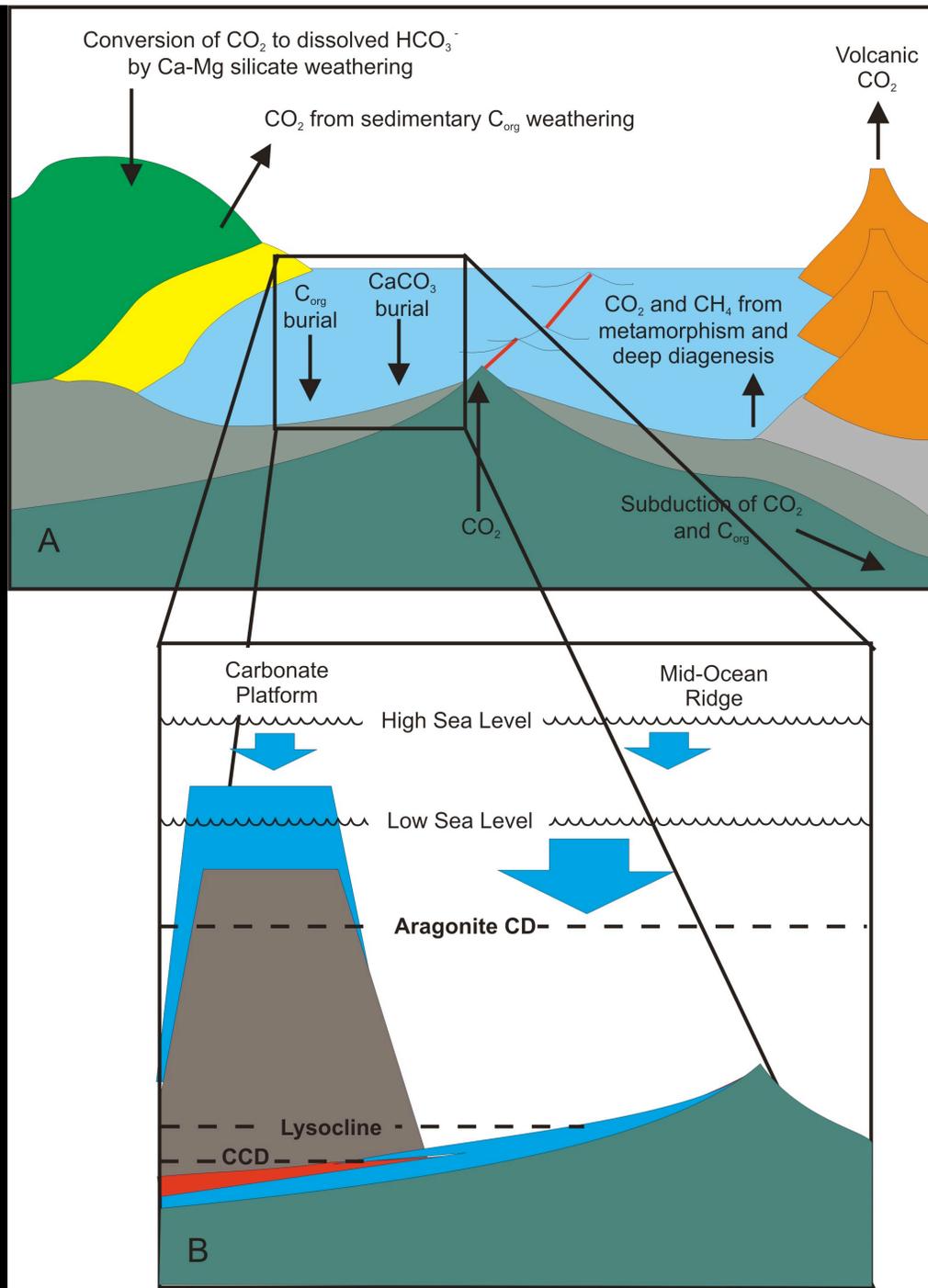
# The Calcareous Plankton

- Have shifted the site of carbonate deposition from shallow seas to the deep sea

During the Late Neogene about 95% of  $\text{CaCO}_3$  deposition has been in the deep sea, 5% in shallow water

But there have been reversals – during the Holocene, after the rapid post-glacial sea level rise, the shallow water  $\text{CaCO}_3$  deposition rate was about 3 times that of the long term average

The modern system is a complex balance with feedbacks to change the  $\text{CO}_3^{2-} - \text{HCO}_3^-$  (ocean alkalinity) proportions to adjust for changes in atmospheric  $\text{CO}_2$ , rapid carbonate output after sea level rise, etc.



# The Paleozoic – Early Mesozoic Story –not well known

- Carbonate fluxes onto the continental shelves and epeiric seas were much larger
- The largest carbonate fluxes were in the Cambrian
- Perhaps very high CO<sub>2</sub> = rapid carbonate recycling
- Perhaps the rise of land plants changes the weathering system

Carbonate was not regularly  
delivered to the deep sea floor  
before 100 Ma

Older ophiolites have dark shales and  
cherts resting on basalt

There are no obducted pelagic  
carbonates in ancient mountain ranges

# Where did all the Early Paleozoic carbonate come from

- It must have already existed in the Precambrian
- What caused the massive recycling into the Cambrian?
- Does this have anything to do with the snowball Earth?
- When did the ocean change from a soda to a sodium chloride solution?

# Carbonate on the Ocean Floor will be ultimately be subducted

- Does the  $\text{CO}_2$  come back through volcanoes?
- Or does it descend into the mantle to form scapolites?
- Scapolites are mantle-stable minerals with the formula



# Have the Calcareous Plankton set in motion a Doomsday Machine?

- Removing C from the Earth's surface and storing it in the mantle through subduction?
- By the way, how did the C get onto the surface of the Earth to begin with?

# Conclusions

- The calcareous plankton are playing a major role in the global carbon cycle
- The calcareous nannoplankton have the largest share
- For the past 100 million years they have been actively altering the global geochemical balance
- In the Early Mesozoic and Paleozoic carbonates were almost entirely restricted to the continental blocks
- Even taking this into account, there were much larger amounts of carbonate deposited in the Paleozoic than since

