

Landscape evolution models and stream piracies phenomena

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Motivations and results:

On earth, landscape morphology is mainly controlled by rivers evolutions and their interactions with hillslopes. But hydrographic network may be re-organized by stream capture and modify deeply the relief. This transition may be induced by several mechanisms (diversion, headward erosion, avulsion, or subterranean filling up). It has interested numerous scientists since a long time (Davis 1895, Blache 1943, Lesson-Quinif 2001 & Le Roux-Harmand 1997-2009...). Here we focus on stream piracies by headward erosion, when an actively eroding low level stream (called the captor) encroaches on the drainage of a nearby stream flowing at a higher level (called the diverter) and diverts part of the water of the higher stream.

During the last decades, several landscapes evolution models (LEM) have been developed to quantify the topography evolution with diffusion and advection equations. These models play an important role in sharpening our thinking to better understand the interaction between landscape evolution processes. LEM were developed basically to simulate erosion, tectonic and climate at different scales of time and space. But, these models were not designed to describe specific mechanisms as the stream capture. It's one of the aims of this work to evaluate LEM for this purpose.

In this presentation, we develop a 1D model based on LEM equations to investigate the stream piracy by headward erosion responses to climatic or tectonic changes. This model incorporates the most common equations used in quantitative geomorphology; diffusion in hillslope, advection in river (detachment-limited mode) and an inequality based on slope and drainage area for the limit between these two domains (Montgomery and Dietrich, 1988).

First, a predictive study with an improved version of GOLEM (software developed by Tucker & Slingerland in 1994) on the Meuse basin shows that several piracies may probably occur in the future. A comparison with the 1D model gives similar results. Then, complex simulations are realized in the Meuse basin taking into account: Meuse deposition tendency and simple

Stream piracies and landscape evolution models:





representation of Karst process etc.

Finally, we present a new approach for parameters calibration based on recorded terraces. The results gives a remarkable differences between this approach and the classical methods (Slope-Area relationship, etc.).

Experimentation on the Meuse basin :

Geomorphologic setting and background:

Eifel Allemagne Luxembourg	The MEUSE river 20km narrow basin (cuestas) valley ~+50m relative elevation 10 to 25m deposit	Captures processes ? Piracy (Terrouin) Avulsion, Karst morphology & lithology control cuesta/depression
-60m -60m -60m -83m -83m -83m -83m -83m -83m -83m -94t -95m -95m -95m -95m -95m -95m -95m -95m	The last My erosion stopped at Monthermé 4 recorded captures: Aisne, Aire ~ -0.9,-0.4Ma Gespunsart ~ -0.9Ma Haute-Moselle ~ -0.3Ma	Imestone/clay <u>Next Capture ?</u> Where & when ? What Impact on the local topography? topography analysis gives limited results1



Reminder of previous results and limitations (EGU 2012):



Illustration of scenarios:

variations:





Impact of Kd (hillslope process)

Smoothed Actual profil

Evolution with mb/nb = 1.2 and Kd = 0.01 m^2/an

Estimation of piracies ages and parameters calibration

First, GOLEM is calibrated according to most cited methods in literature (Odoni, 2007). The simulation shows that the piracies ages are very important (more then 10 My). More complex simulations (Meuse filling, Karst process, etc.) are then realized with a developed 1D model.

The rate of the Meuse avulsion was fixed to 10m/My (ref).

The Karst process had been considered by estimating the ratio between the water flux and drainage area (P=Q/A) for all rivers. Q is based on the hydrologic stations measurement and A is estimated by DEM analysis.



1D simulations tacking in account the Meuse river filling:





The increase of **mb/nb** (fluvial process) promotes the upstream piracies. In our case, if **mb/nb>1** the piracy of the Meuse by the Ingressin does not occur.

The increase of Kd (hillslope process) may promote piracies in regions with weak relative hillslope elevation. In our case, if Kd>0.5m²/y the Aroffe is captured by the Bouvade independently of the fluvial process.



Conclusions and limitations:

The fitted parameters stay in the range of variations given in the literature (Odoni 2007, etc.). In the other hand, the piracies ages are not in agreement with the geologists predictions (Le Roux & Harmand, 1997), when they think that a capture by the Ingressin will occurs in less than 100Ky.

Considering the Meuse avulsion, reduce significantly the piracies ages, but they remain important (more than 3 My for the Ingressin and the Rupt-de-Mad).

The way the Karst process is simulated (impact only the water flux Q) modify the rivers profiles but the ages remain unchanged.

New approach for parameters calibration:

Calibration approach :



Application on the Meuse basin rivers:



The calibration gives similar values for the 4 studied cases: mb = 0.38, nb = 1, Kci = 1, nci = 0.3 and Tci = 0.4 km². Note that Kb varies slightly [26, 40] and Kd varies significantly.