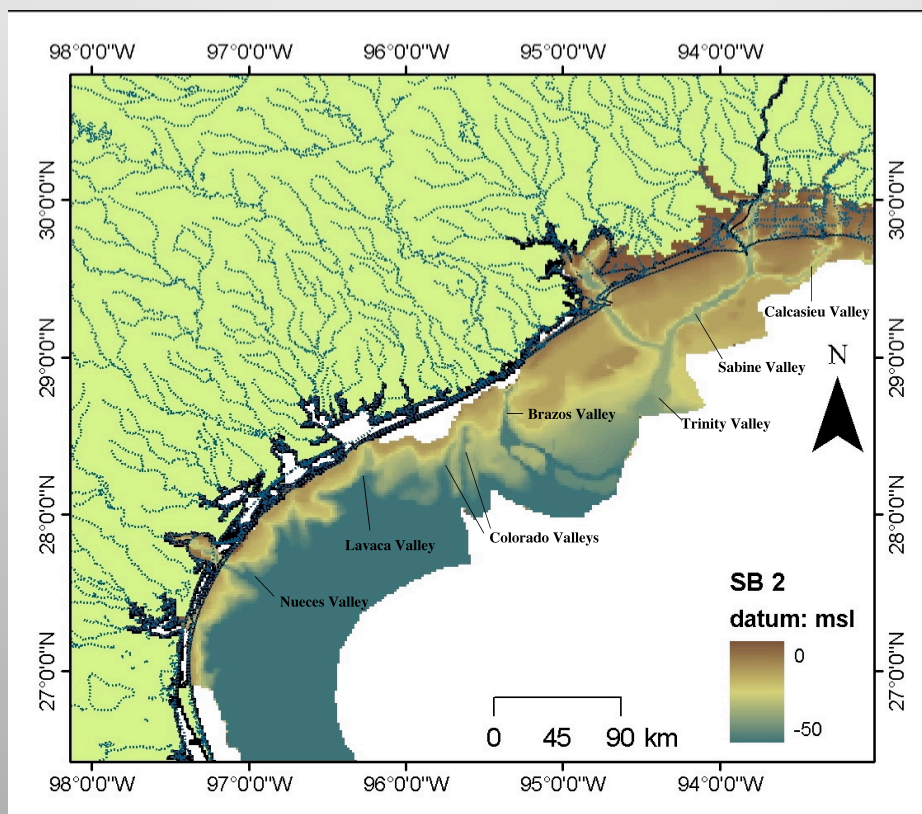


Different Stratigraphic Architectures offshore Texas Reflect Spatial and Temporal Variability in Sediment Supply and Dispersal



Alex Simms
Antonio Rodriguez
Kristy Milliken
Ken Abdulah
Julia Wellner
Heather McKowen
Phil Bart
Lou Bartek
Laura Banfield
Patrick Taha
Robert Weight
Brenda Ecklels
Mark Thomas

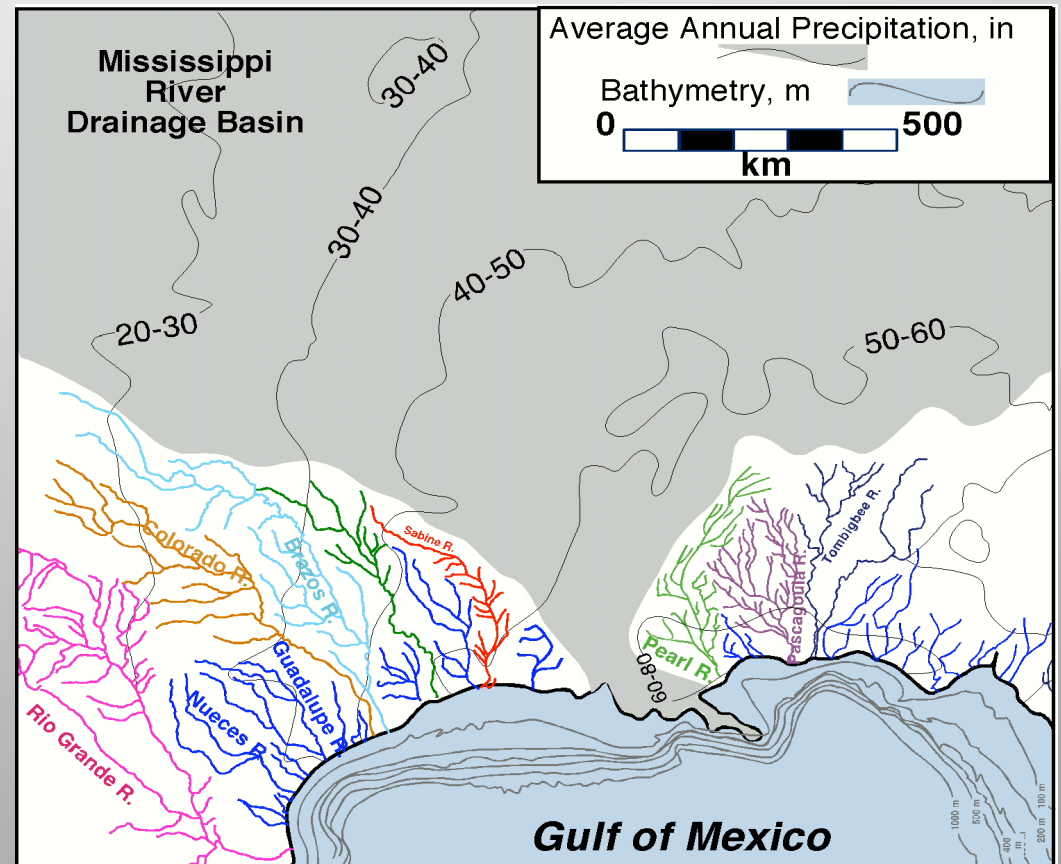
John Anderson, Robert Weight and Rodrigo Fernandez
Rice University

2011 Chapman Conference

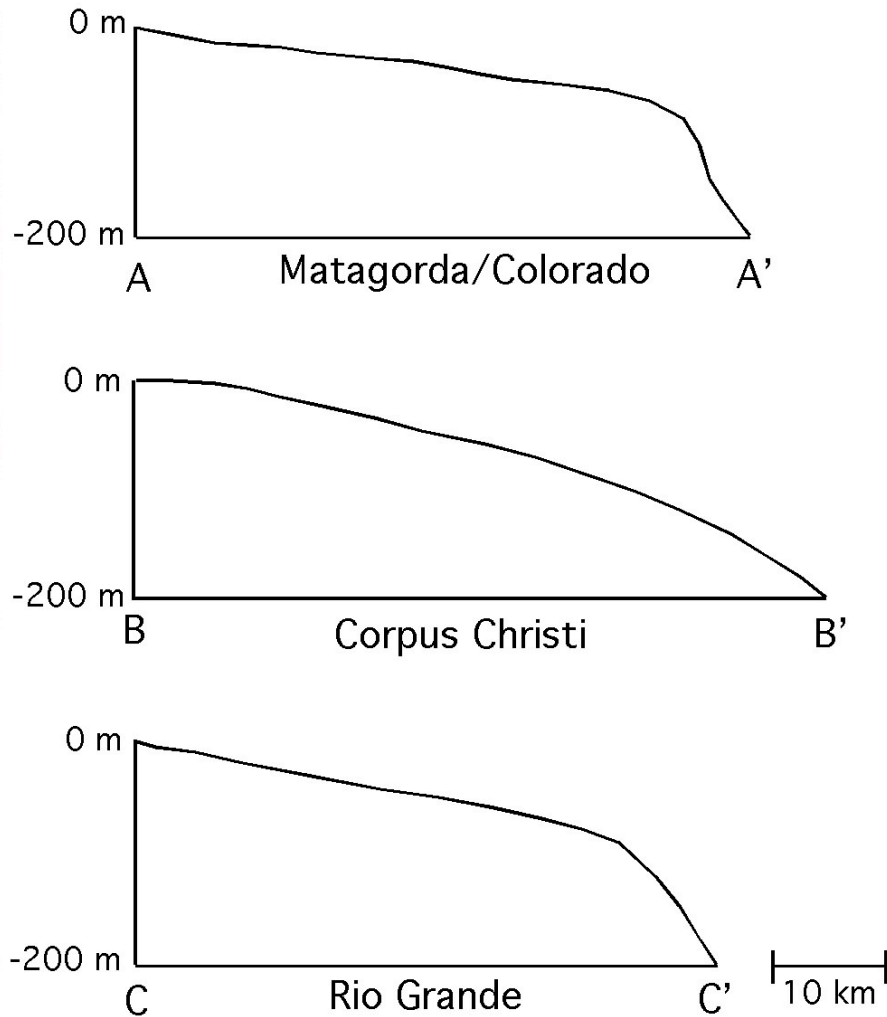
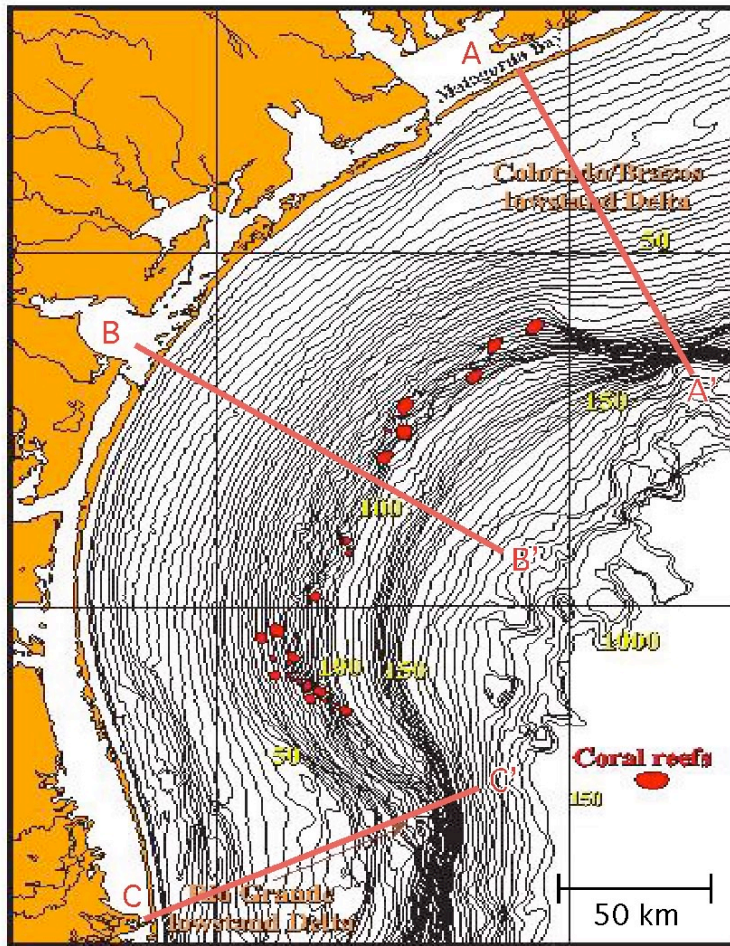


Gulf of Mexico as a Site for Stratigraphic Modeling

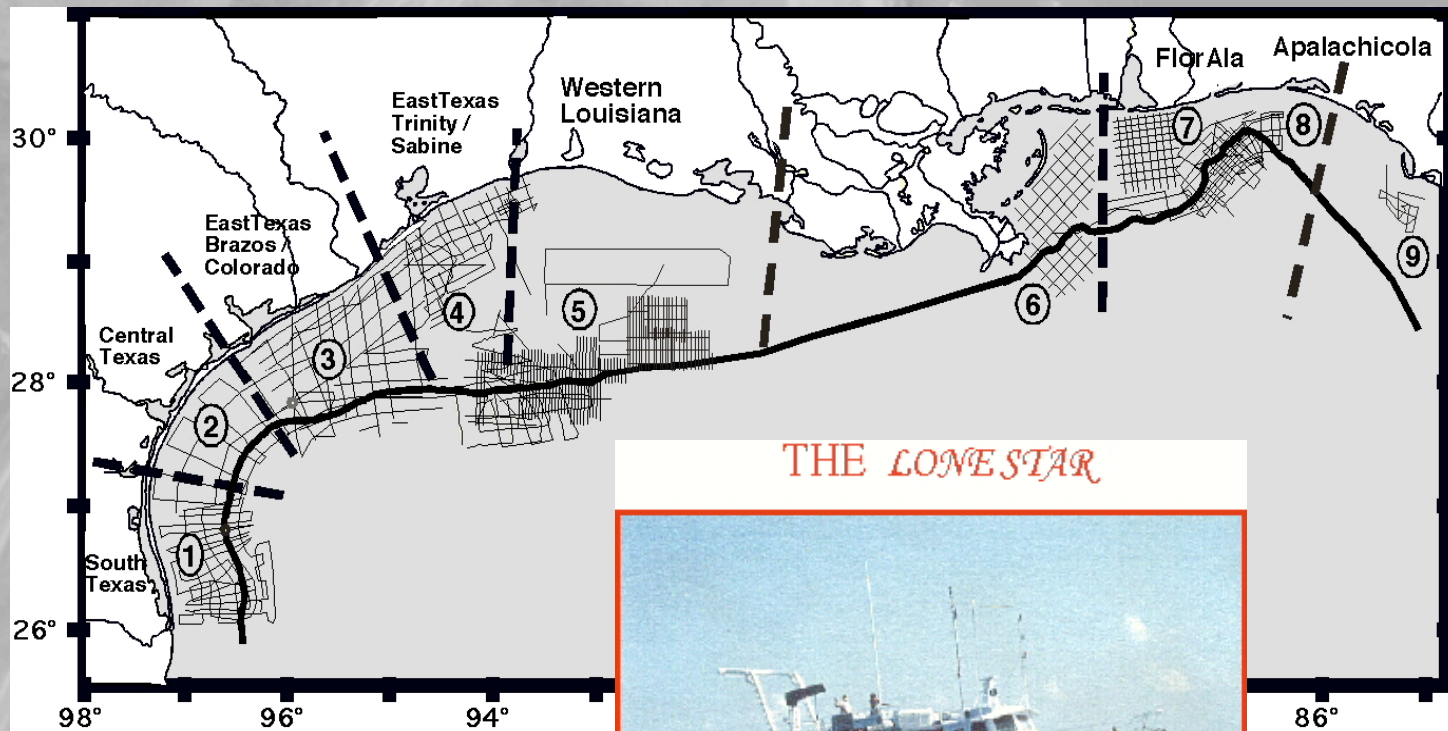
- *Variable climatic settings and rivers with different drainage basin sizes and gradients.*
- *Subsidence rates are well constrained and relatively high compared to most passive margins.*
- *Variable physiography*
- *Abundant drill core*



Long-term depositional patterns reflected in margin physiography.



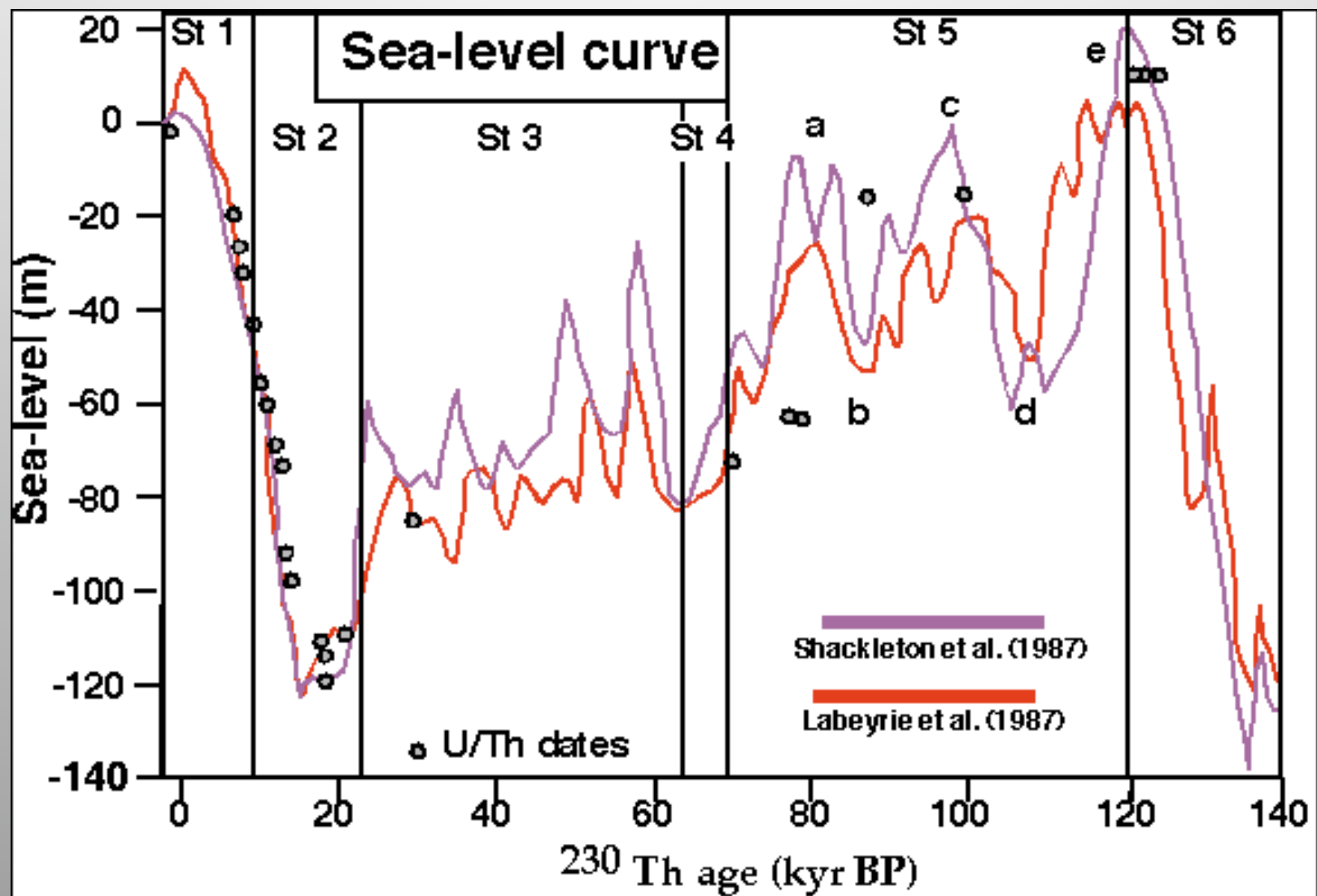
High-resolution seismic data



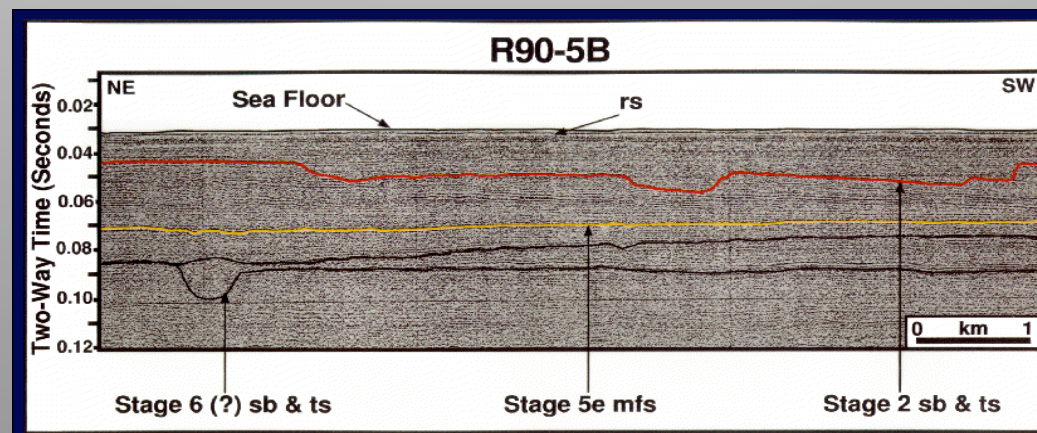
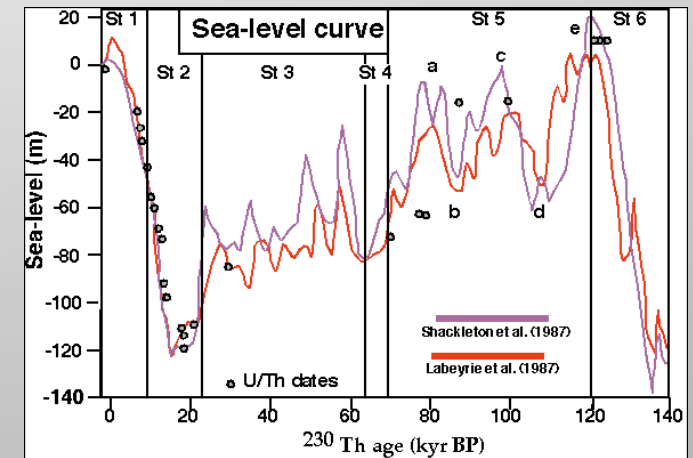
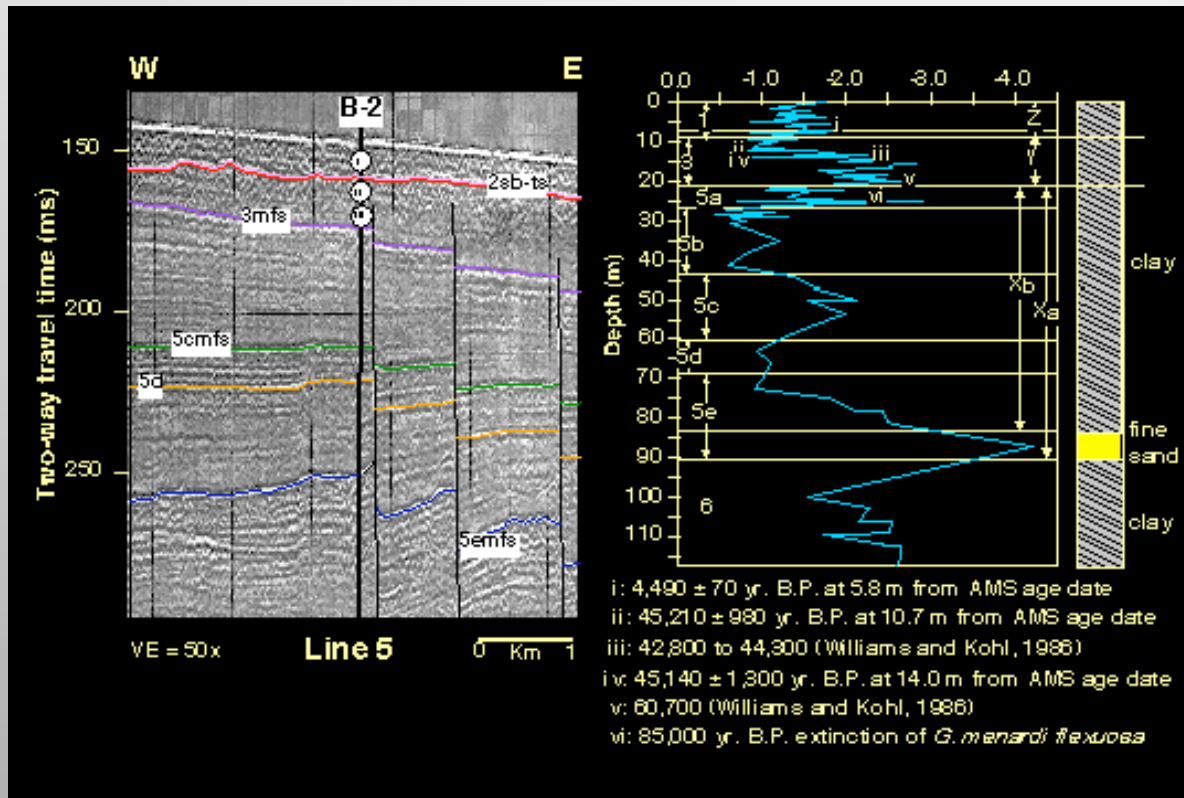
●
(6=LSU)
(7=University of Alabama)



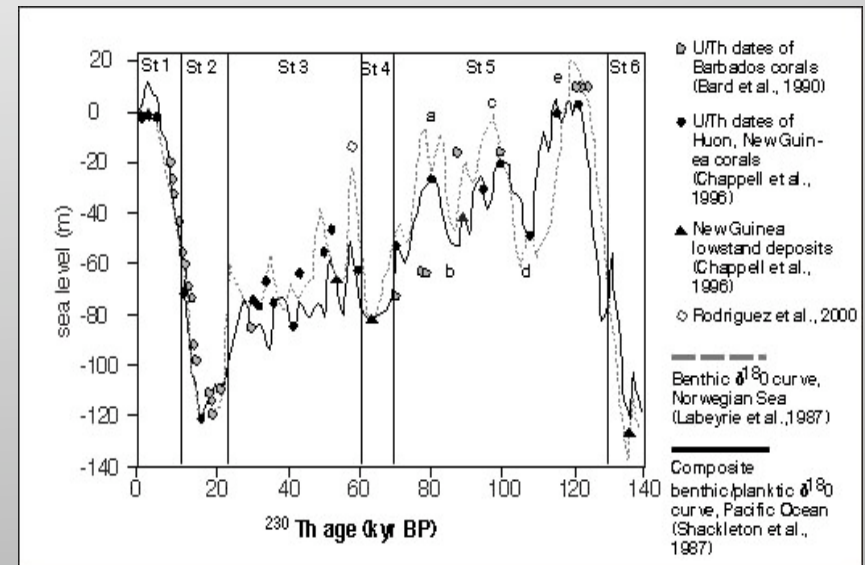
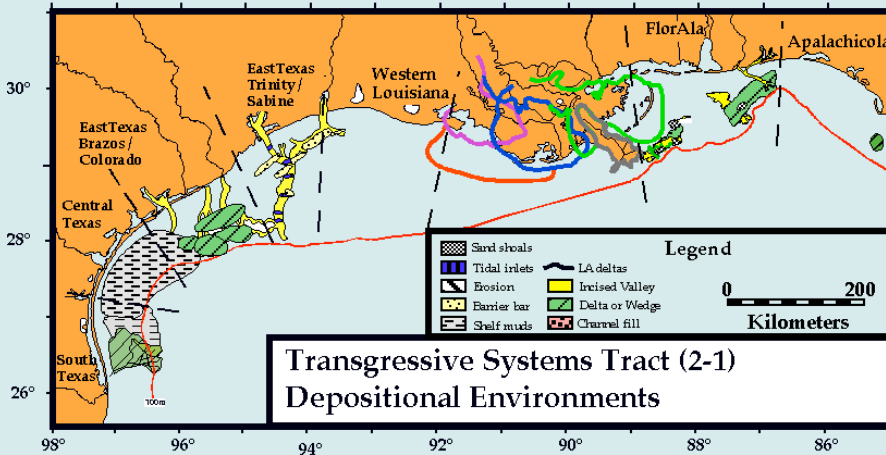
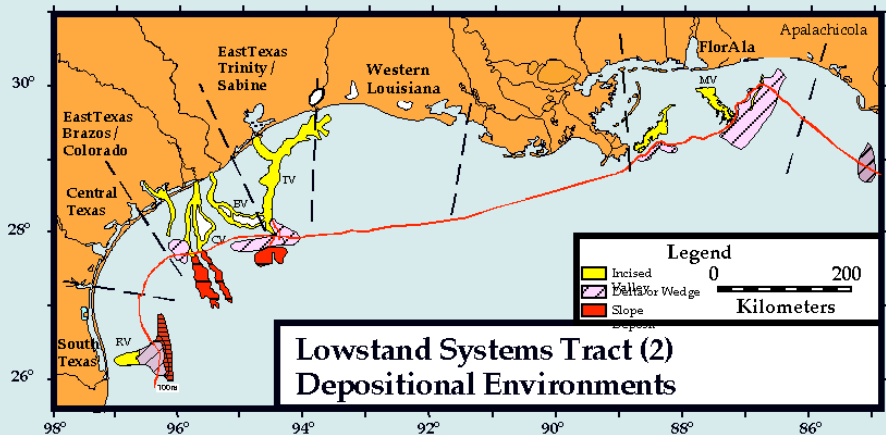
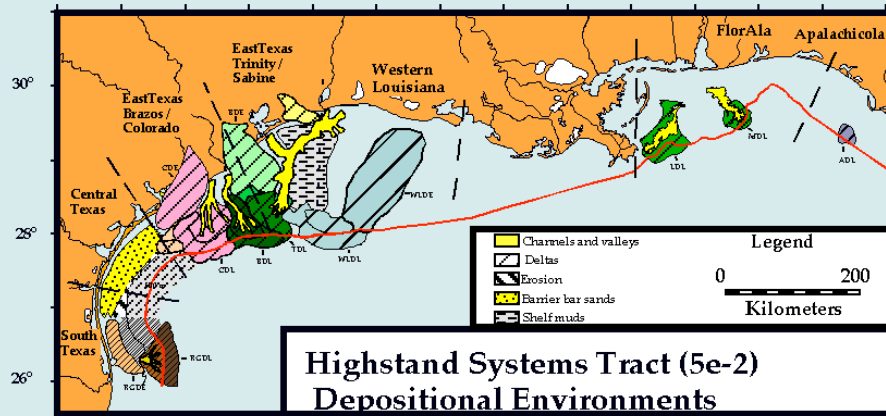
See Anderson and Fillon, 2004., SEPM Spec. Pub. 79)



Seismic and Chronostratigraphy



Paleogeography



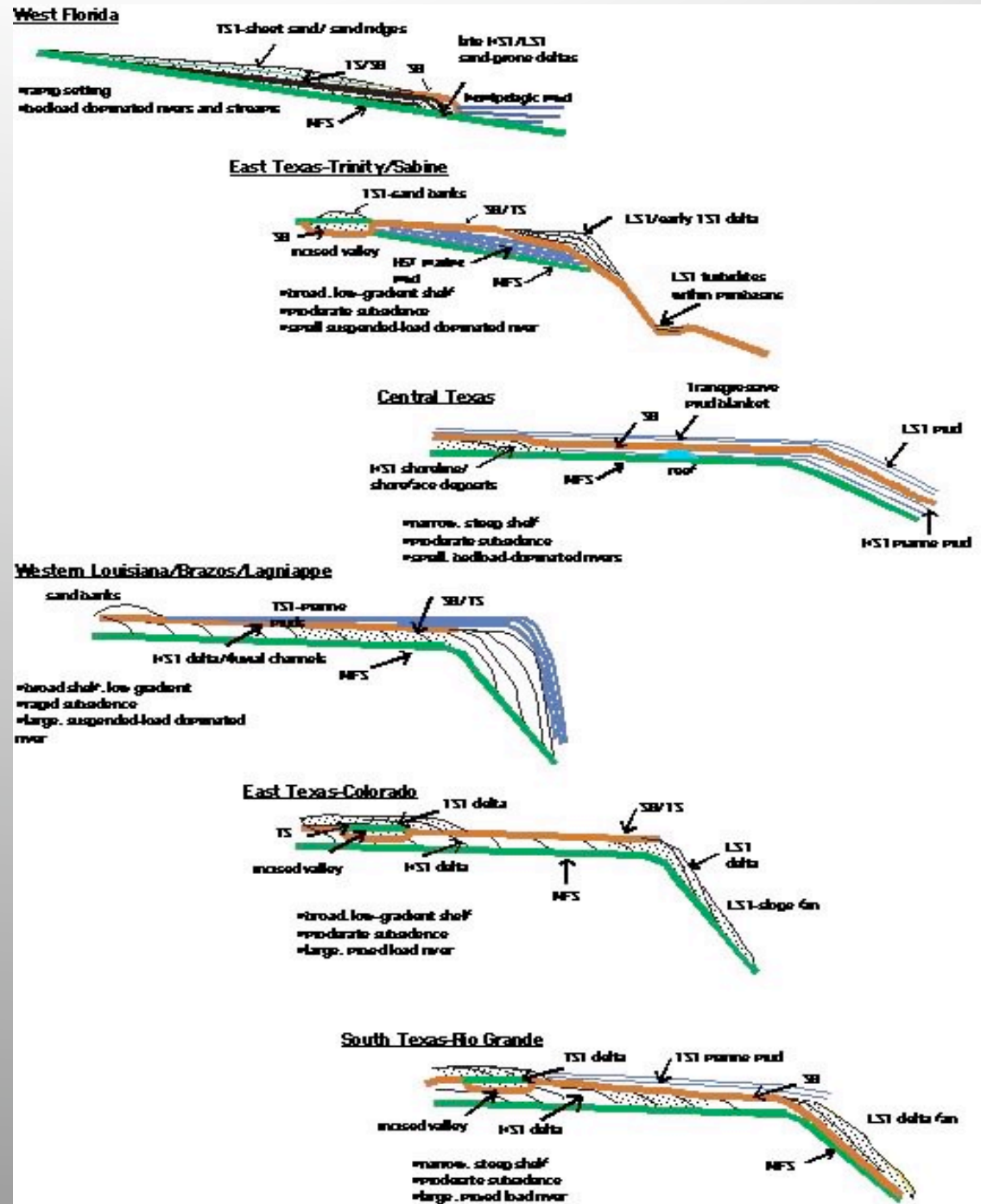
Anderson et al., 2004
SEPM Spec. Pub. 79

Variable stratigraphic architecture across the region

Bedload versus suspended load of main feeders is not reflected by the strata on the shelf and for any given system changes through time

Subsidence is a big factor

One source one sink



	Drainage Area [km ²]	Qw [m ³ /s]	Qs [MT/yr]
Mississippi	3,202,959	15452.0	427.9
Rio Grande	804,792	123.0	36.9
Brazos	115,040	189.0	12.4
Colorado	101,178	76.0	2.8

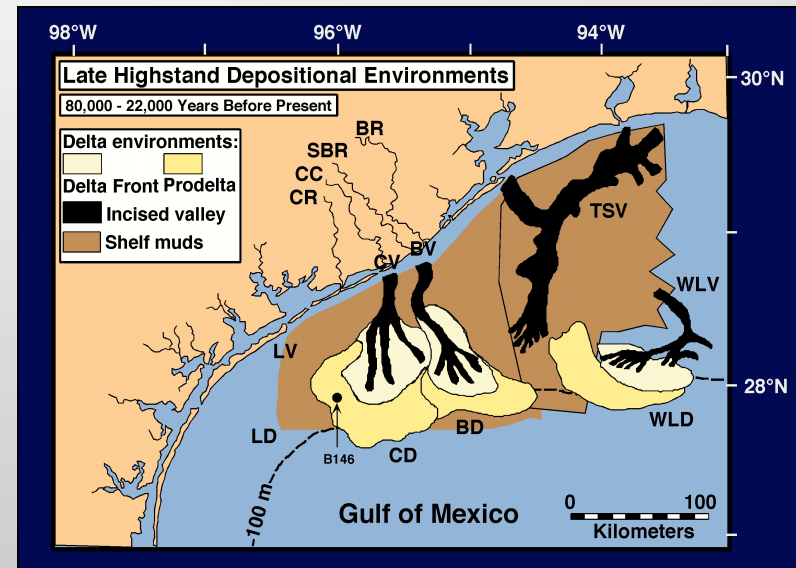
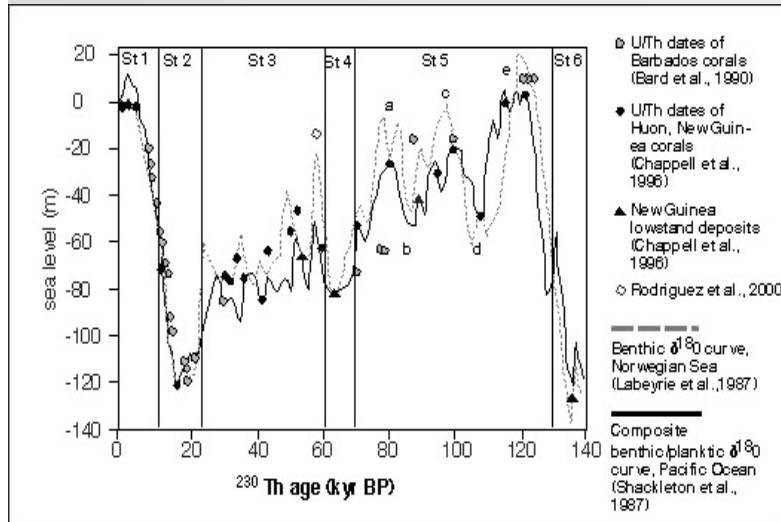
Estimation of Sediment discharge (Qs) using Syvitski and Milliman (2007) QBART mc

Qw	water discharge	Note: all the values of the parameters used to calculate Qs are
R	Relief	$Q_s = w * B * (Q^{0.31}) * (A^{0.5}) * R * T$
T	temperature	
Te	trapping effect	
B	$B = I * L * (1 - Te) * Eh$	
Qs	sediment discharge	

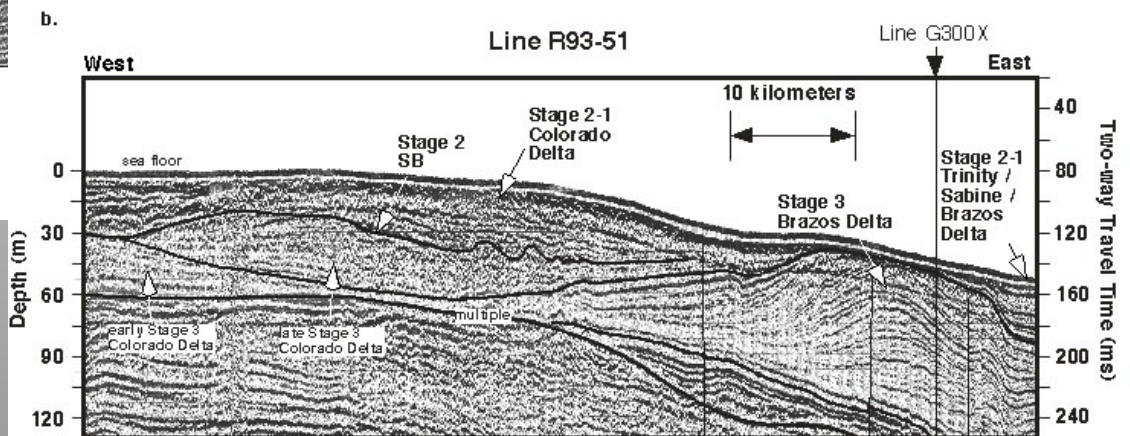
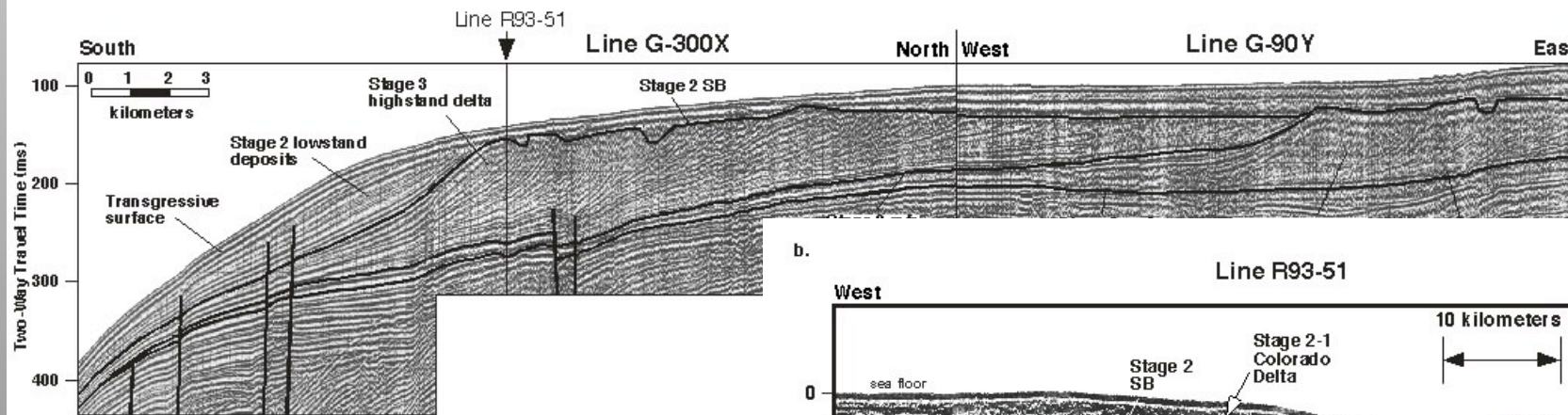
NOTE: the expresion to calculate the bulk sediment discharge from sediment volumes is:

$$Q_s[\text{kg/yr}] = (\text{Vol}[\text{m}^3] * \text{Sed. density}[\text{kg/m}^3]) / \text{time interval}[\text{yr}]$$

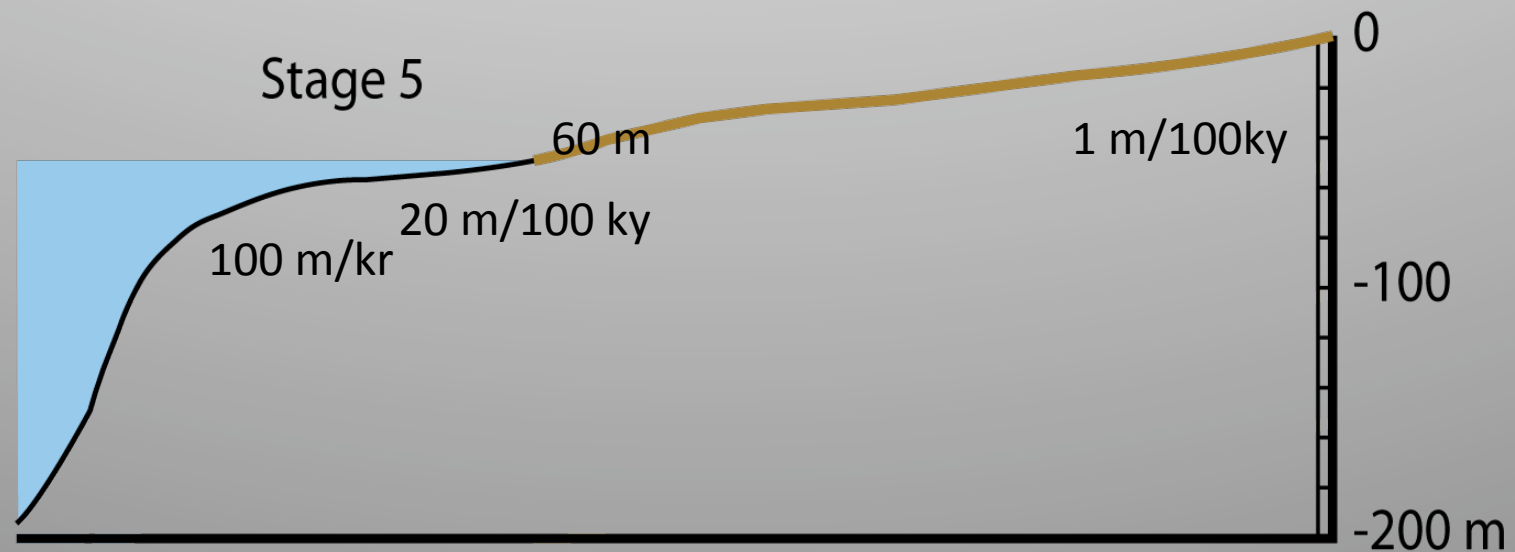
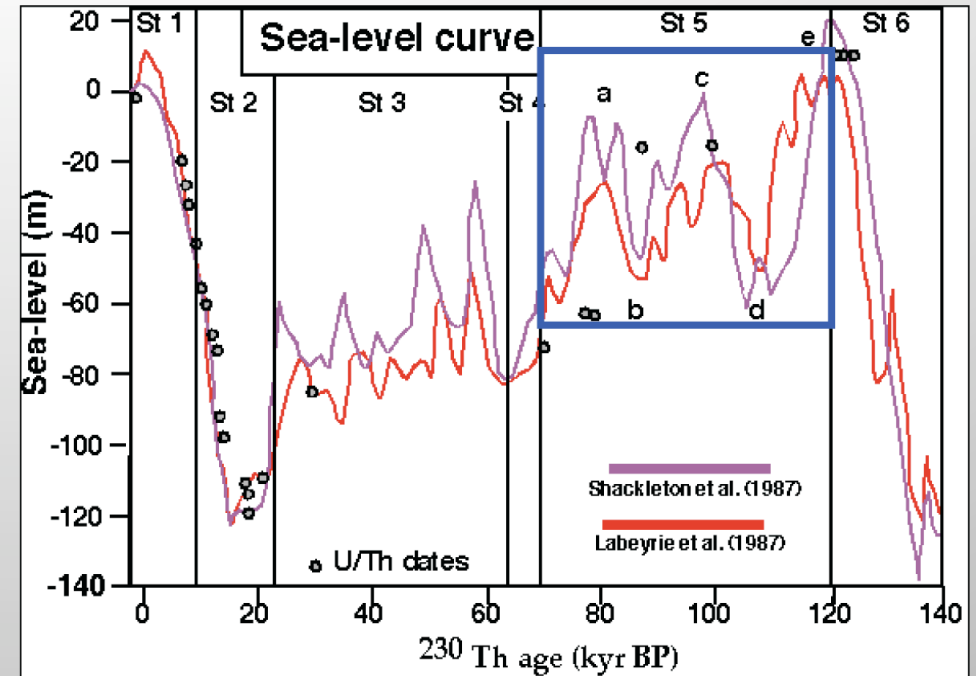
Brazos Delta $\sim 200 \text{ km}^3$ total volume with maximum flux during Stage 3 averaging between $8 \text{ km}^3/\text{ka}$ to $40 \text{ km}^3/\text{yr}$ (discharge of $20 \times 10^6 \text{ t/yr}$ to $100 \times 10^6 \text{ t/y}$, or ~ 2 to $10 \times$ current rate).



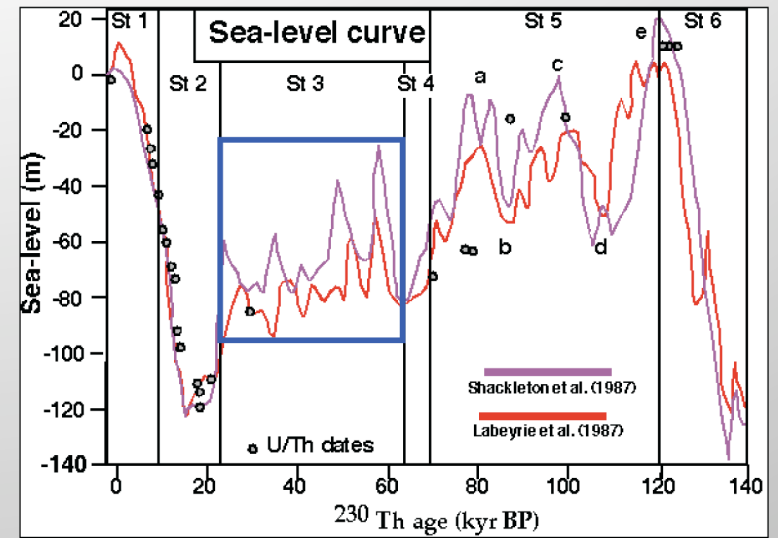
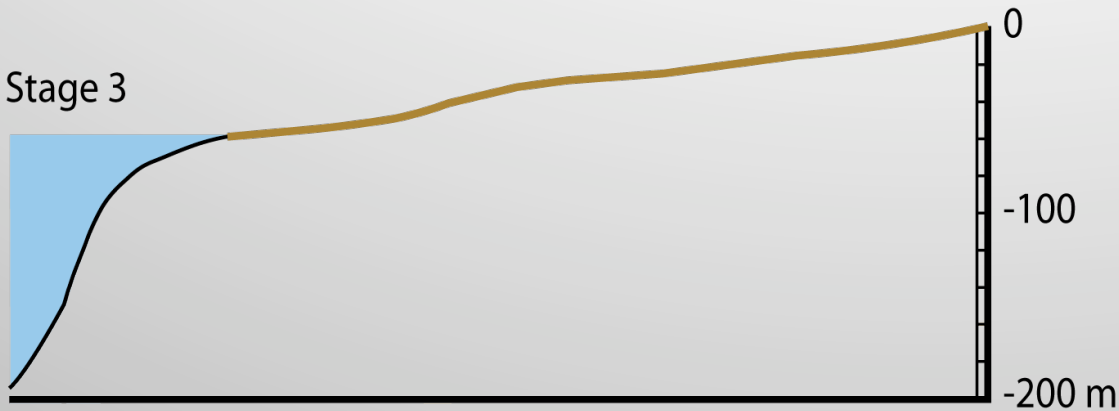
a.



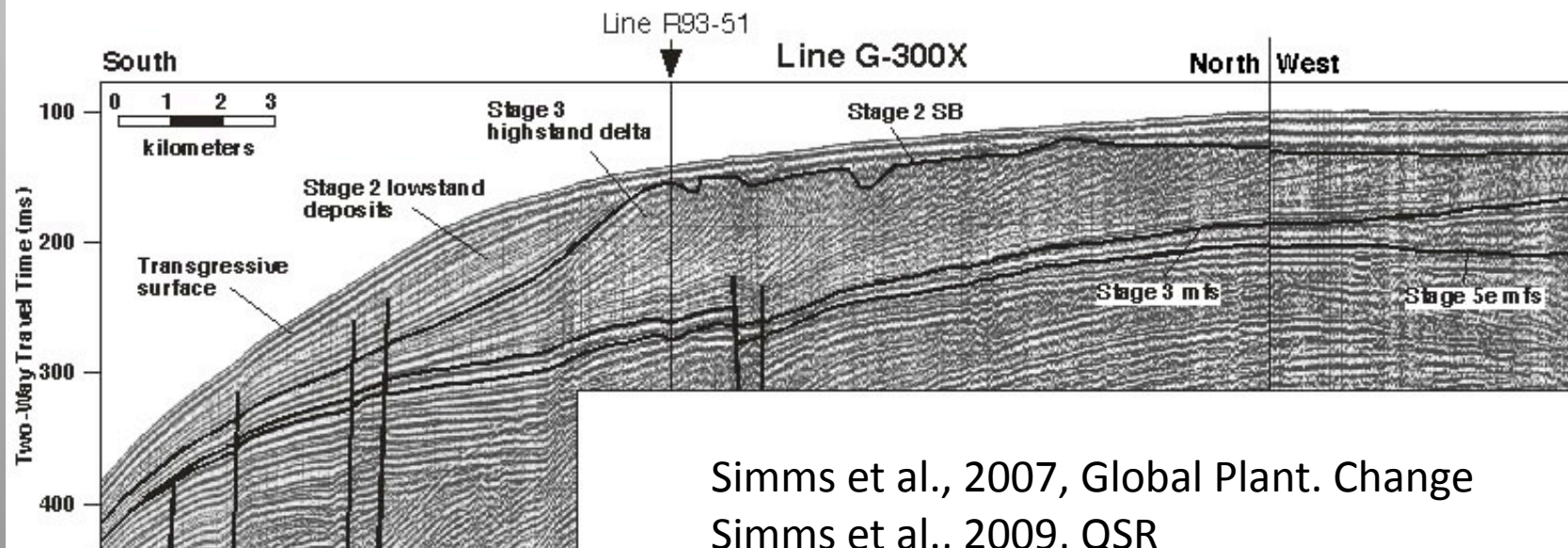
Subsidence versus sea-level



Stage 3



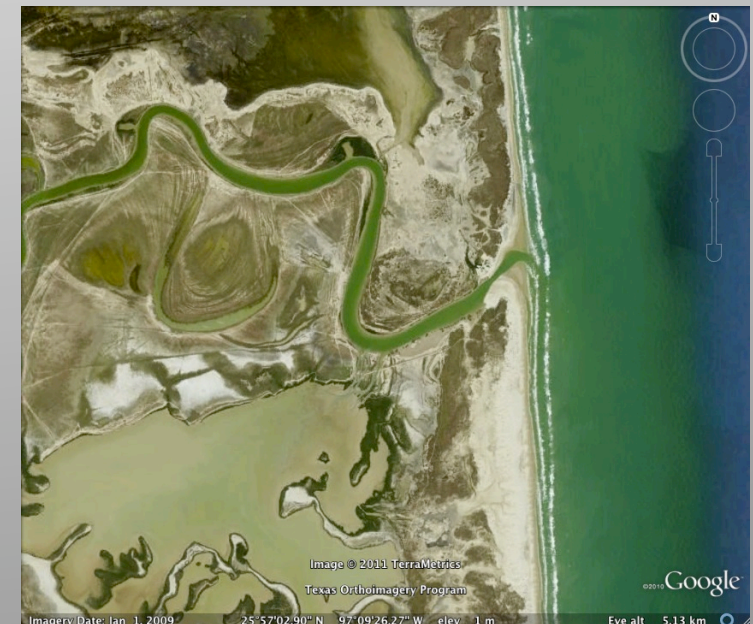
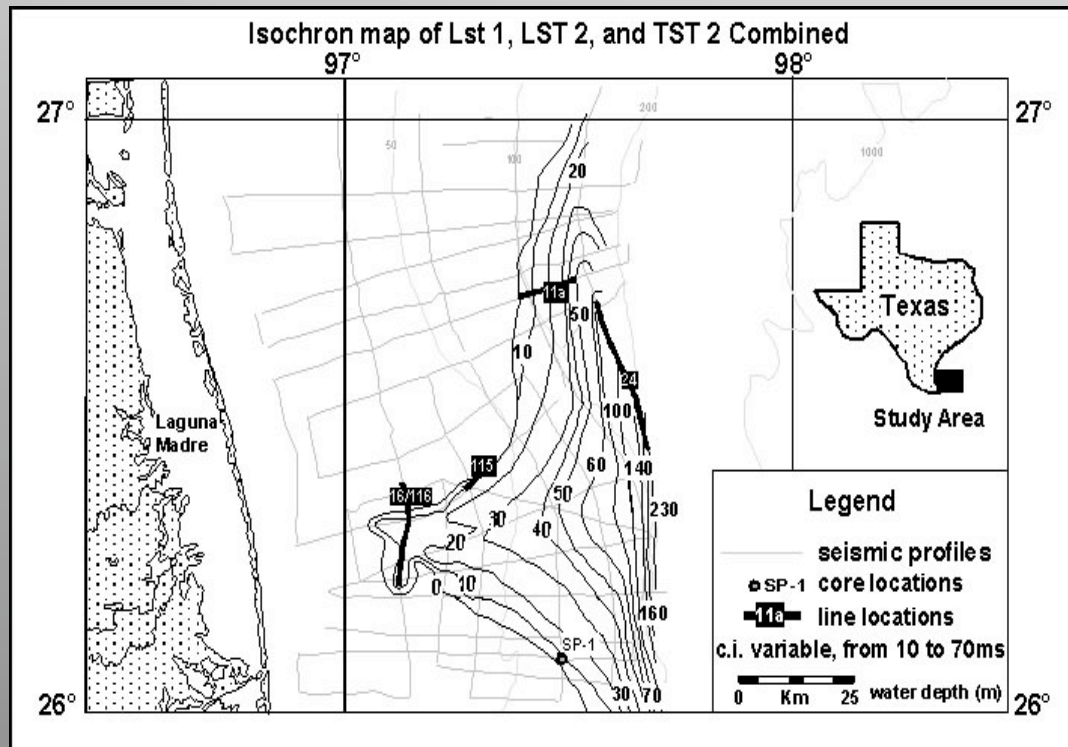
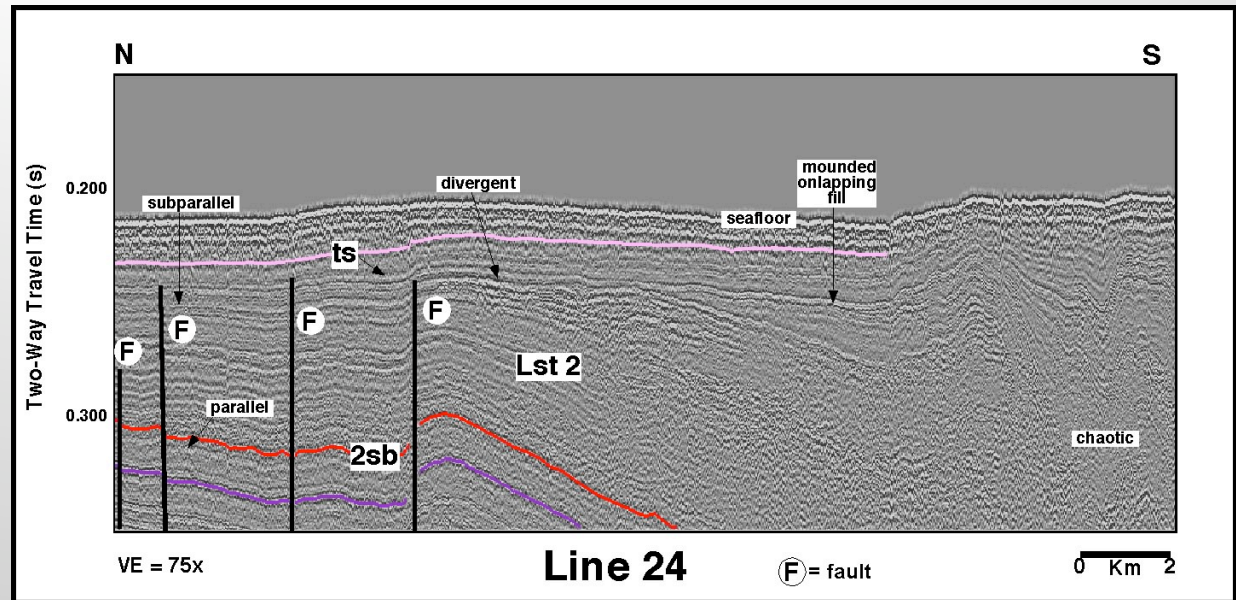
a.



Rio Grande Delta

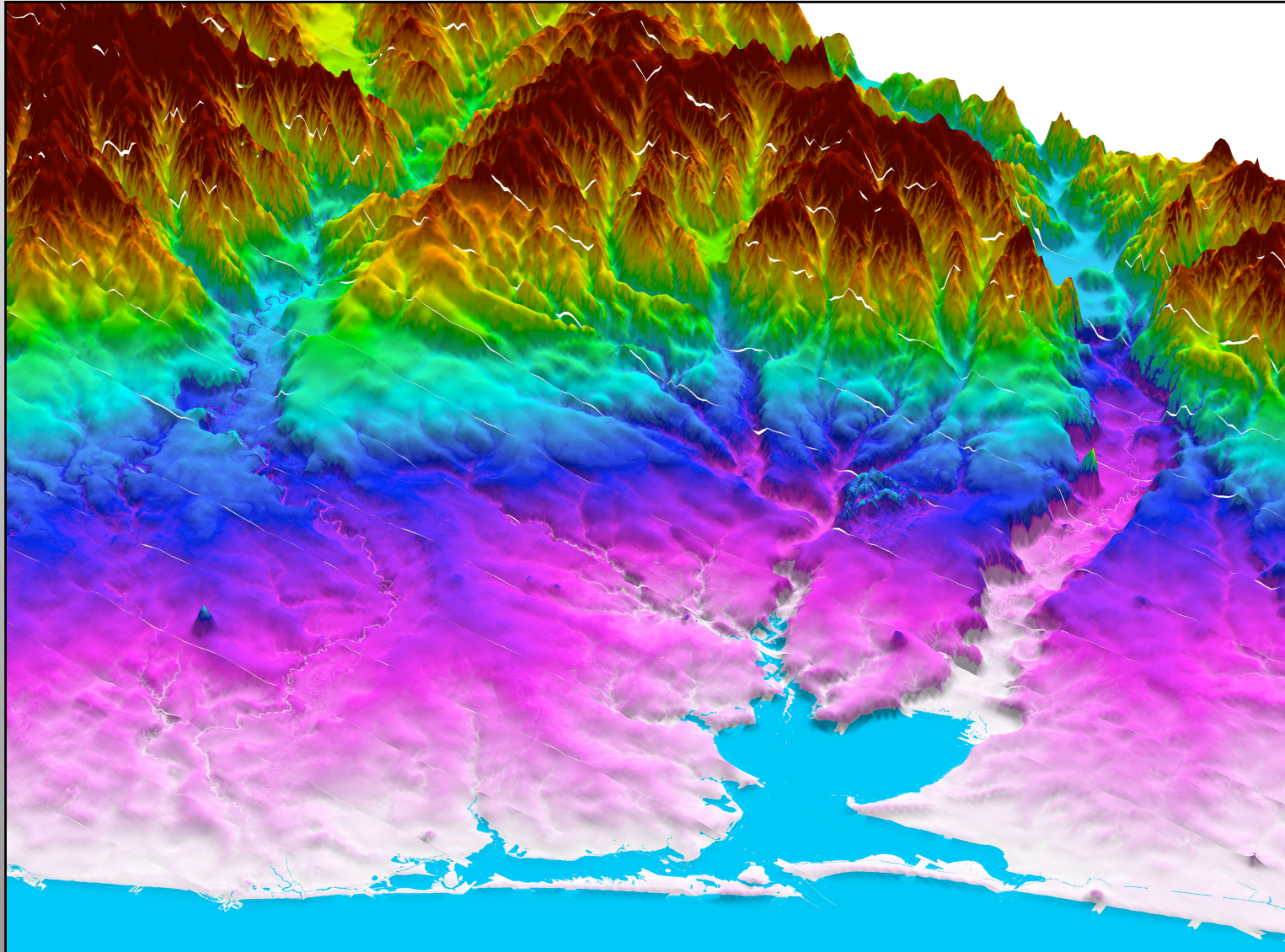
(~220 km³)

Minimum flux during lowstand of 30 km³/ka
(75x10⁶ t/yr or just over twice the current rate)

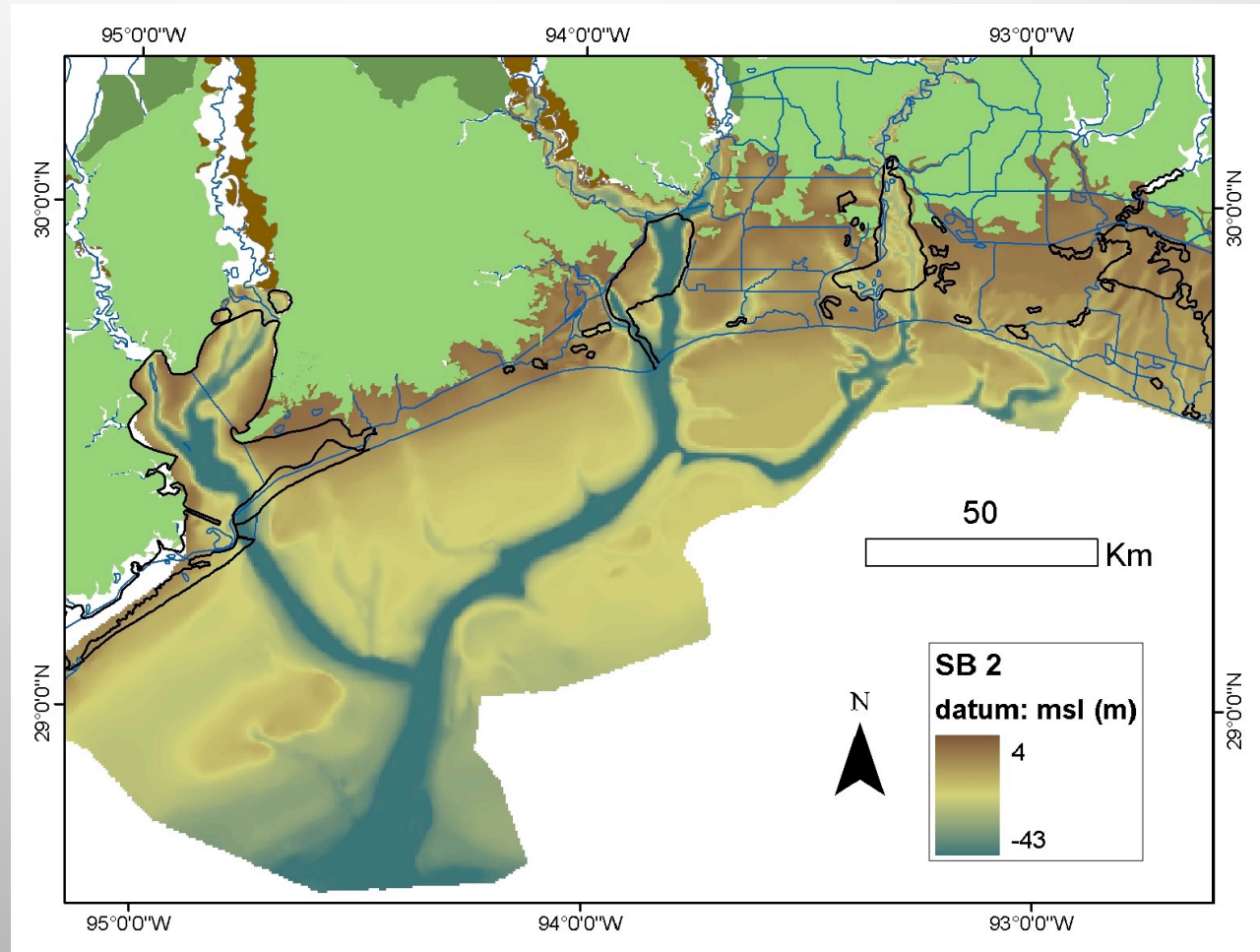


Over-filled versus under-filled valleys

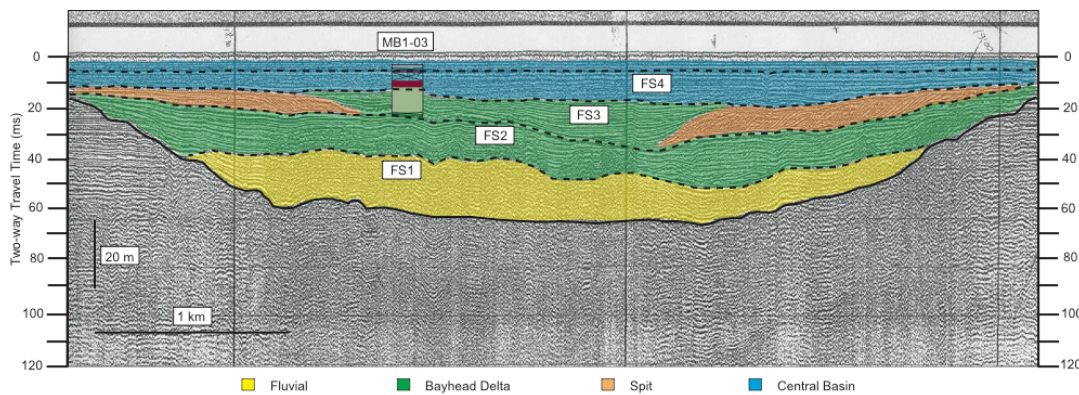
Fill and Purge



Valleys are filled with Holocene sediments



Simms et al., 2007



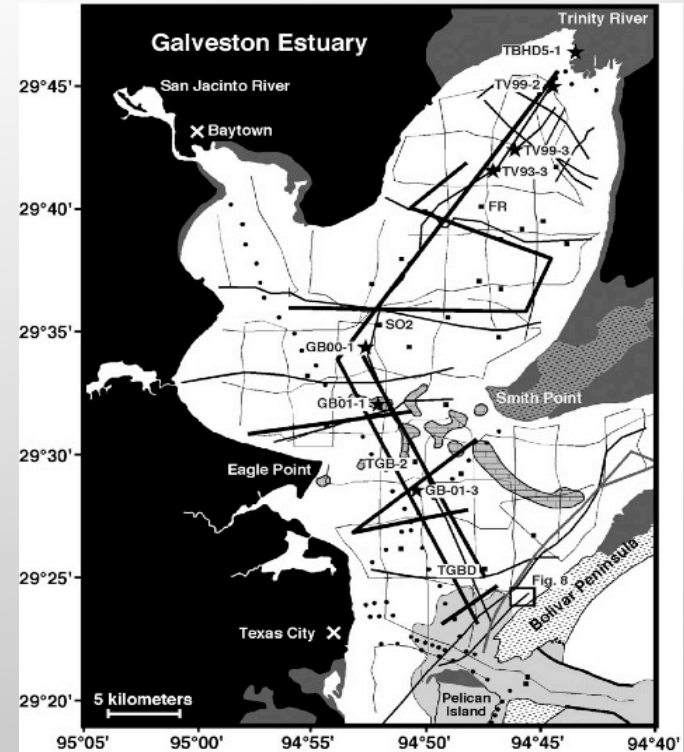
Trinity Valley

An under-filled valley

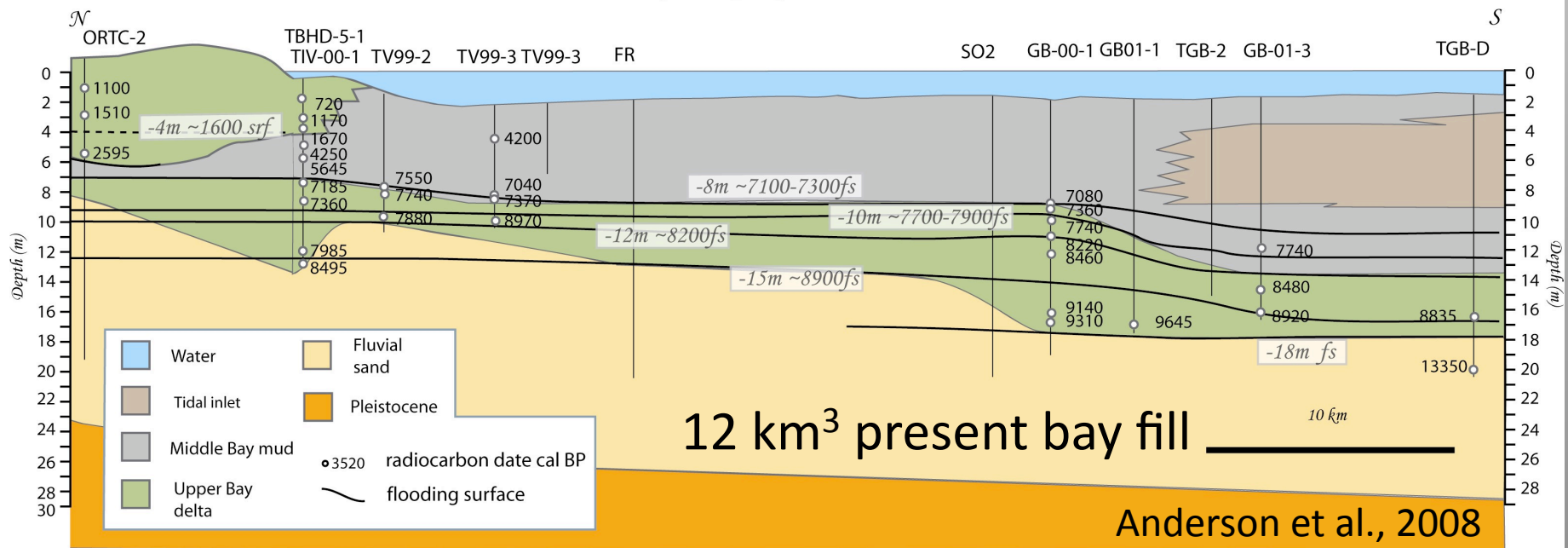


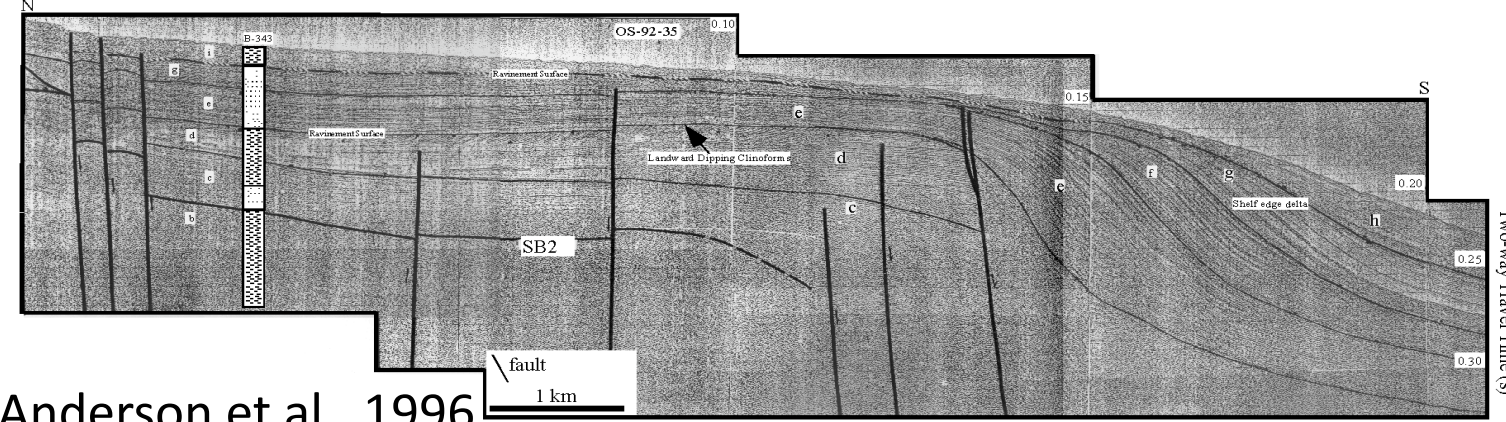
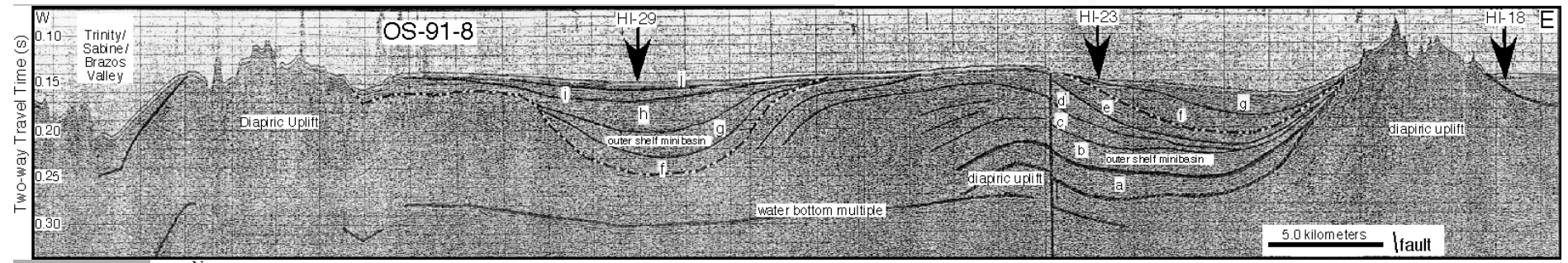
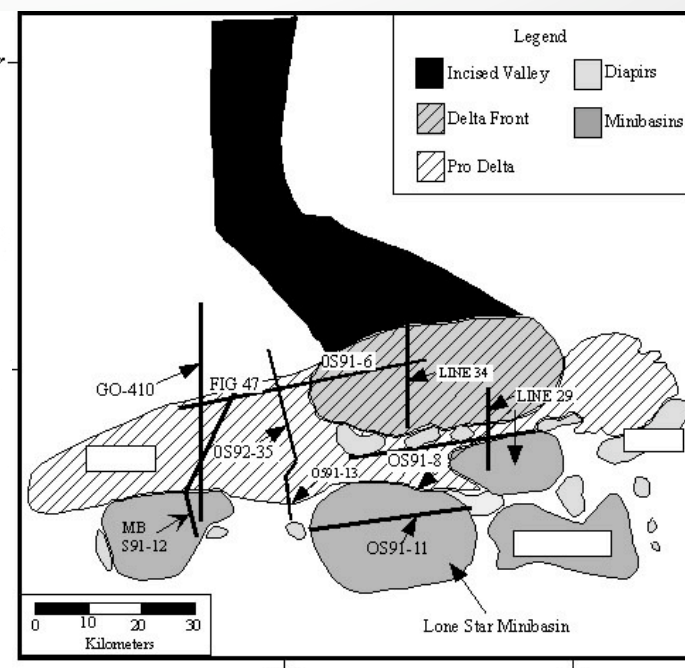
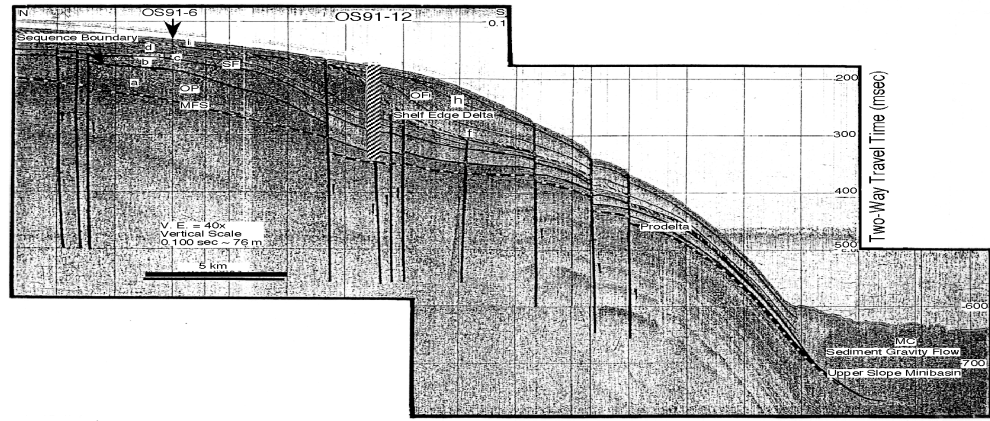
Note that sand is mainly in lower part of valley

We use these data to estimate long-term (Holocene) sediment flux



Trinity Valley Dip Cross section

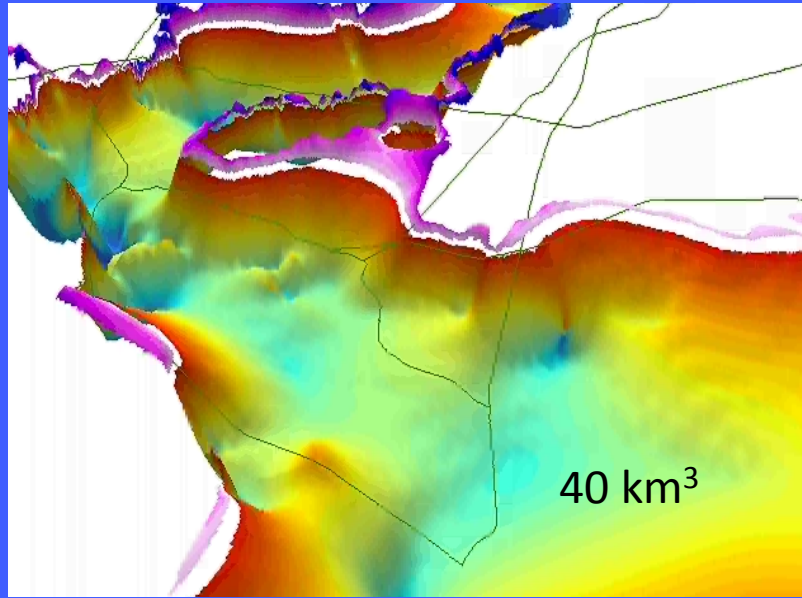
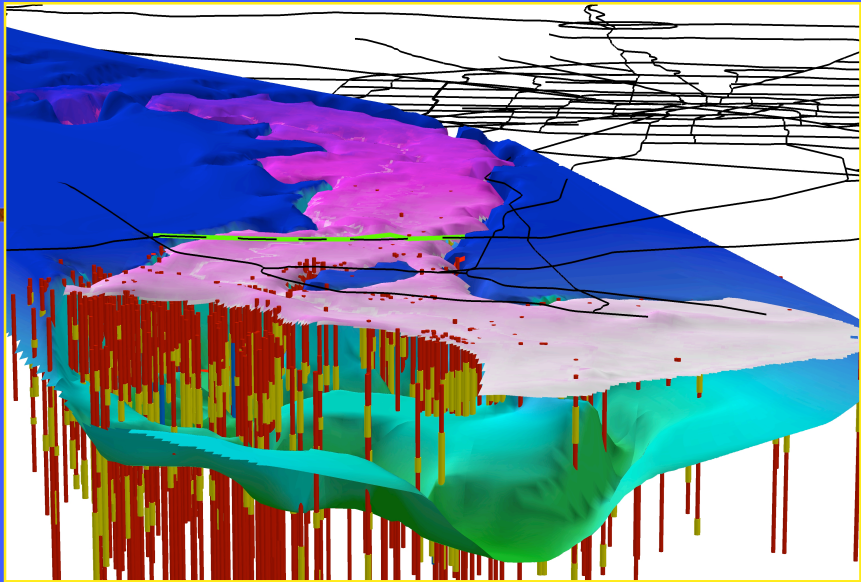




Anderson et al., 1996

Wellner et al., Figure 19

403 Digital Water Well Descriptions



- Clay
- Sand

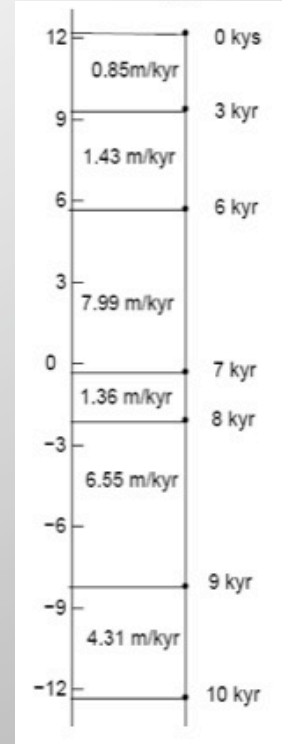
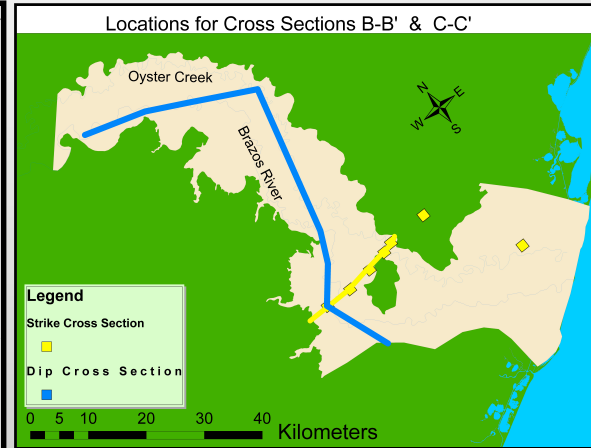
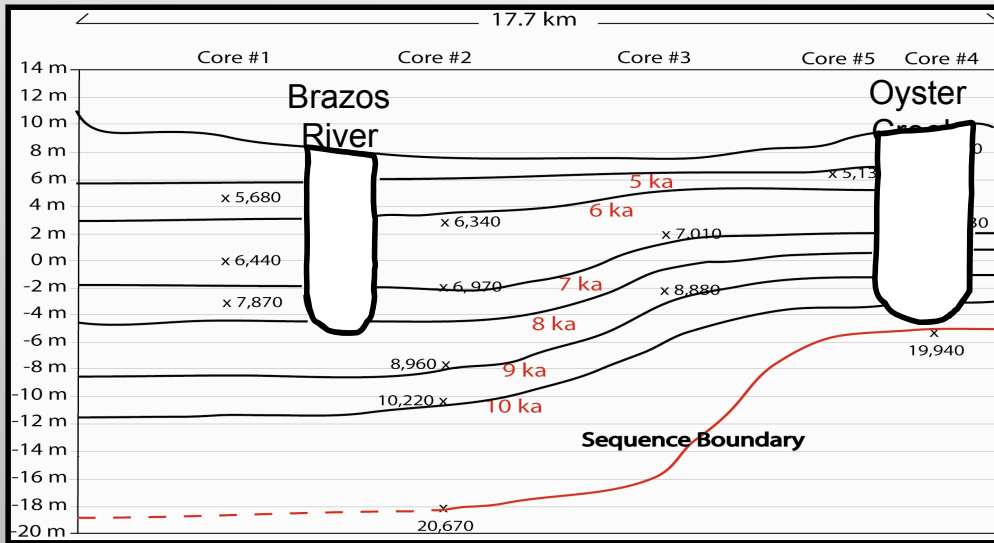
Note the distribution of clay and sand

Taha and Anderson, 2007, Geomorphology

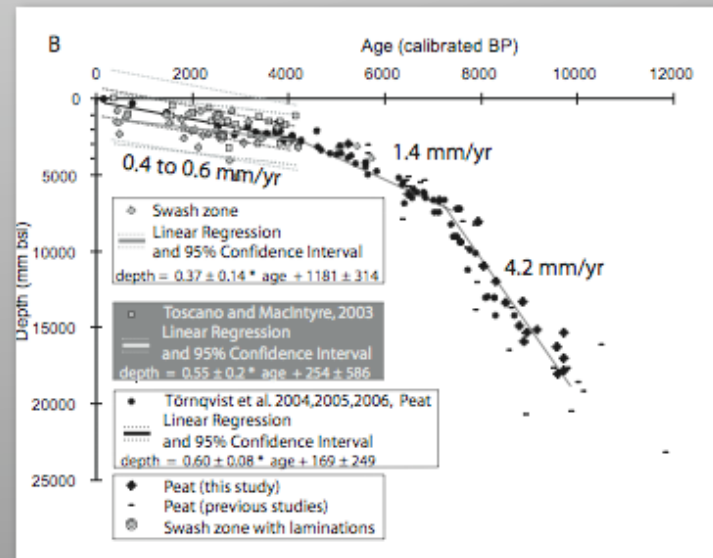
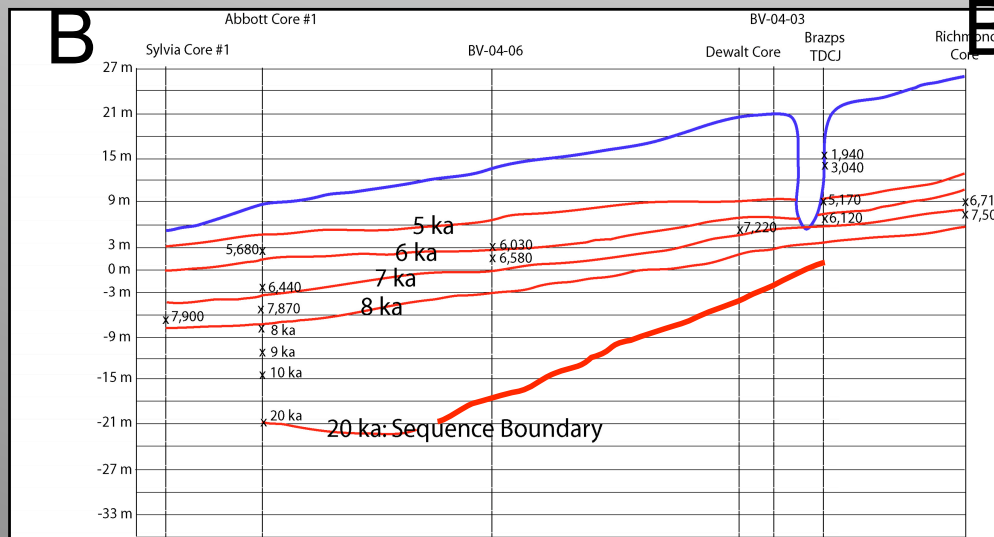
Vertical Exaggeration : 300

In lower valley, aggradation tracks sea-level rise

Lower Colorado Valley

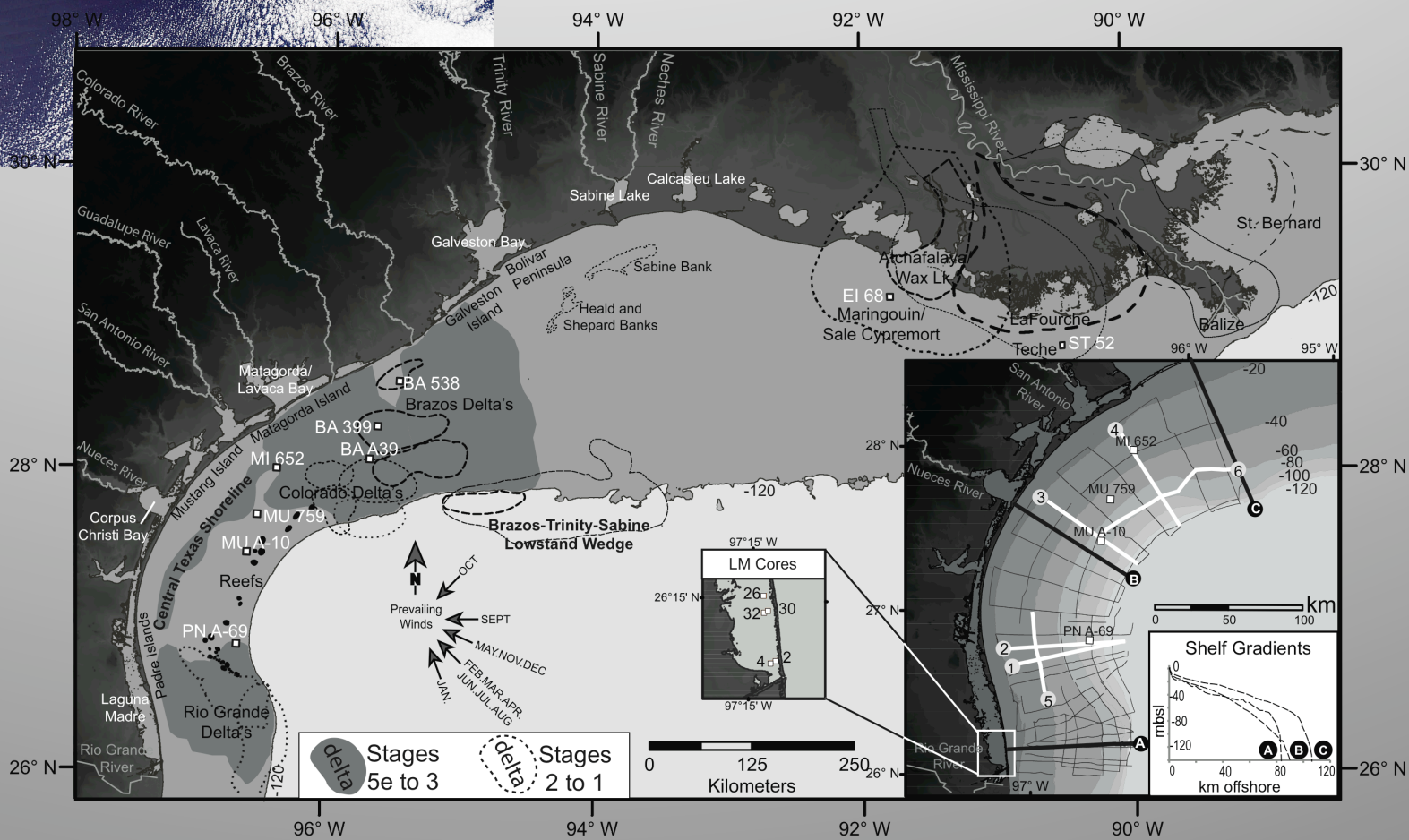
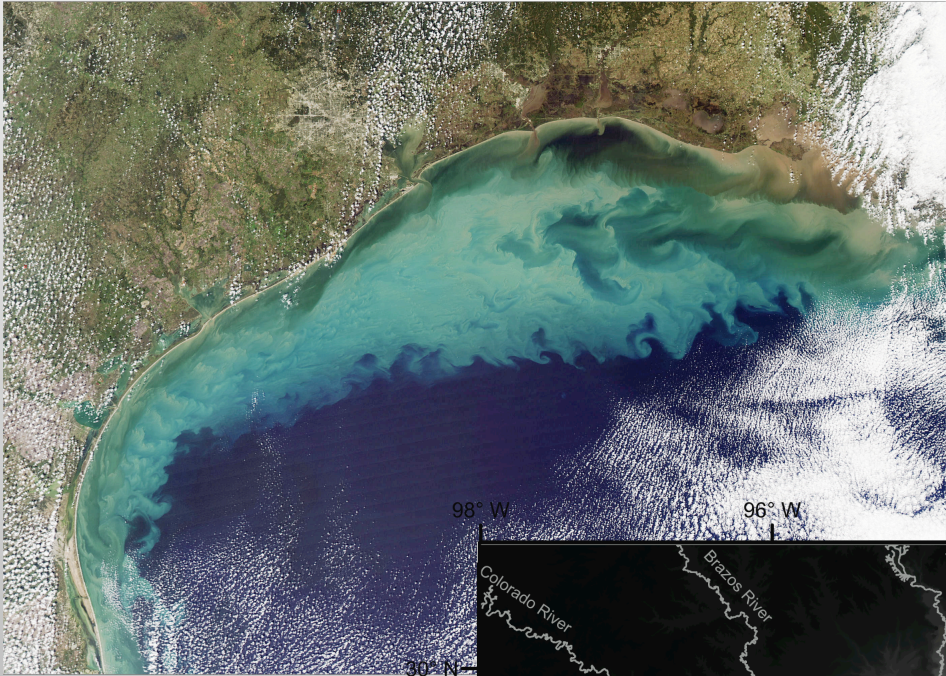


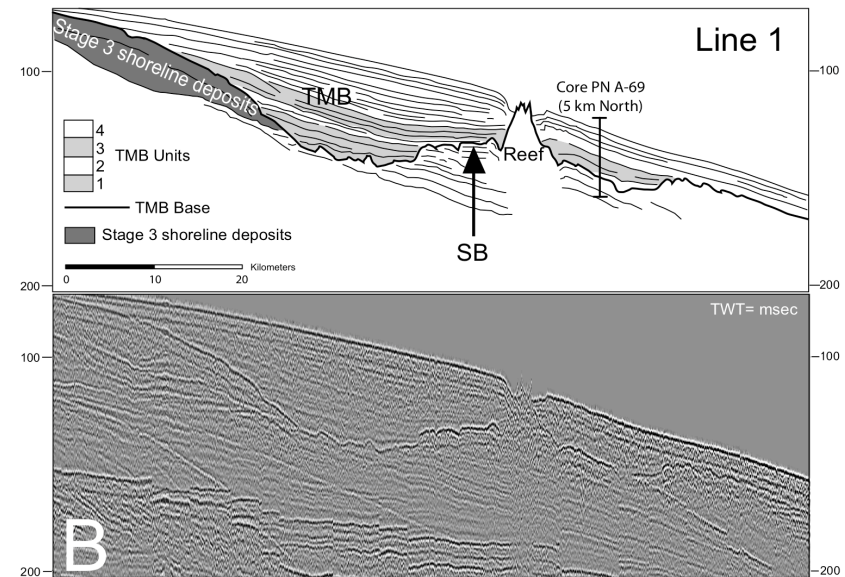
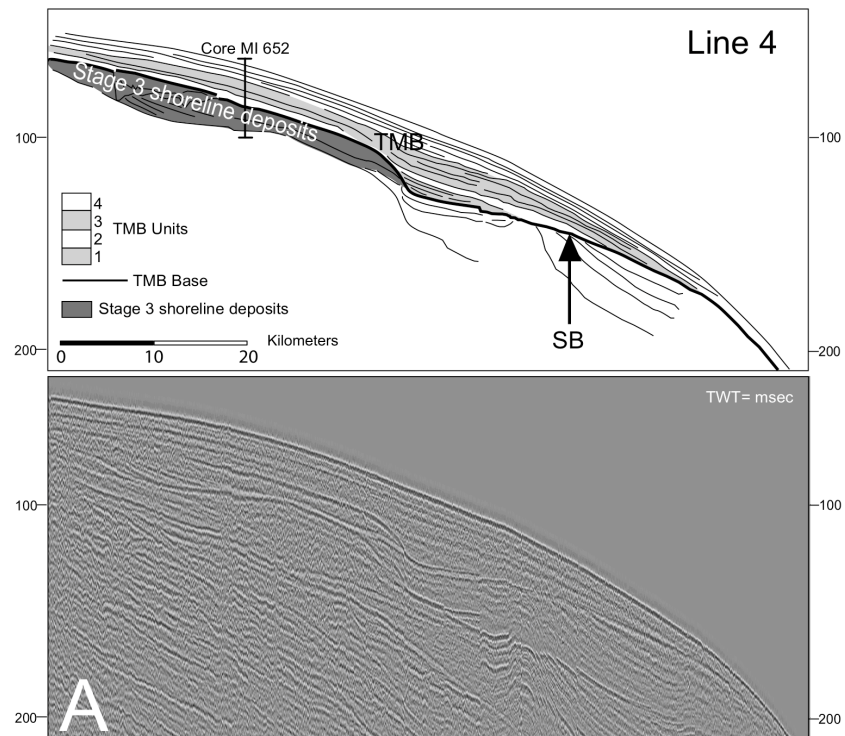
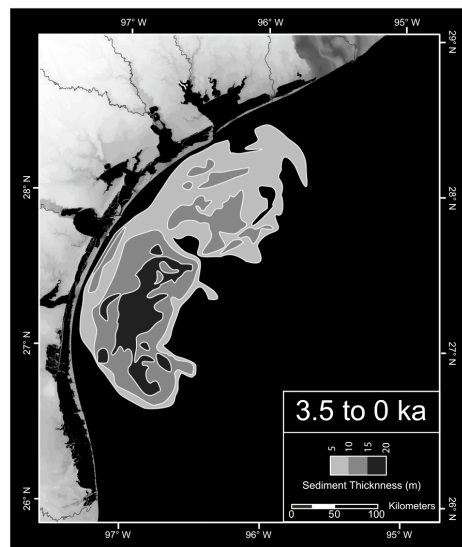
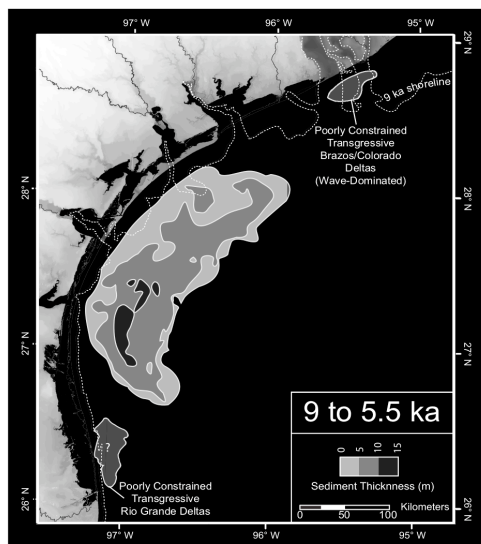
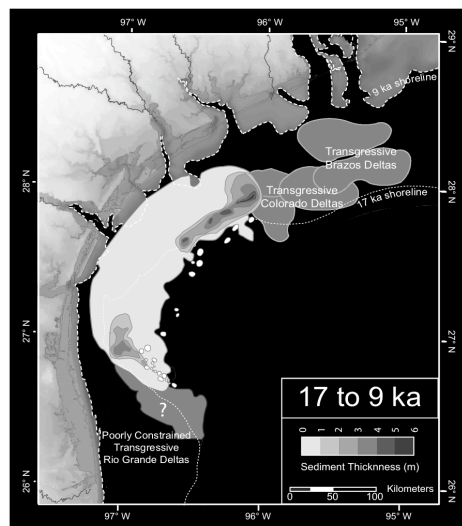
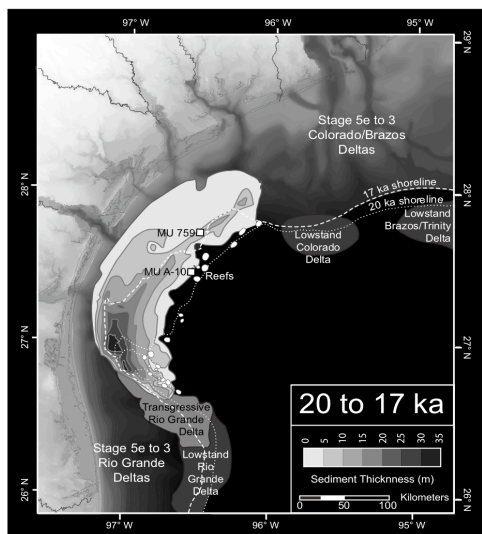
← 80 km →



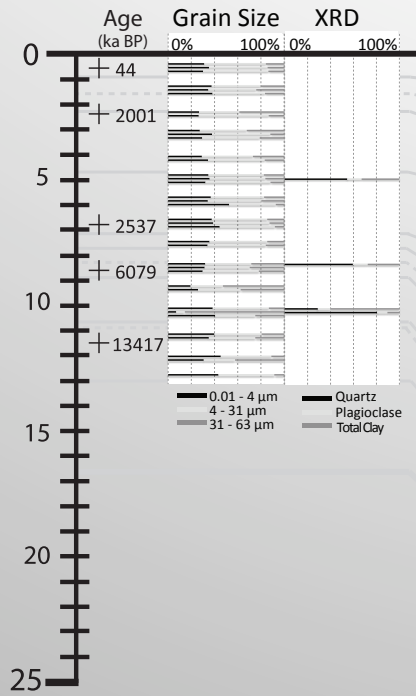
Milliken et al., 2008

The Texas Mud Blanket

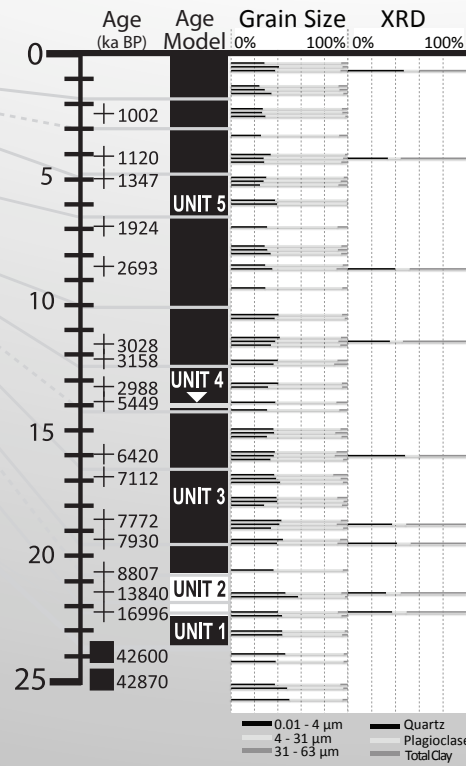




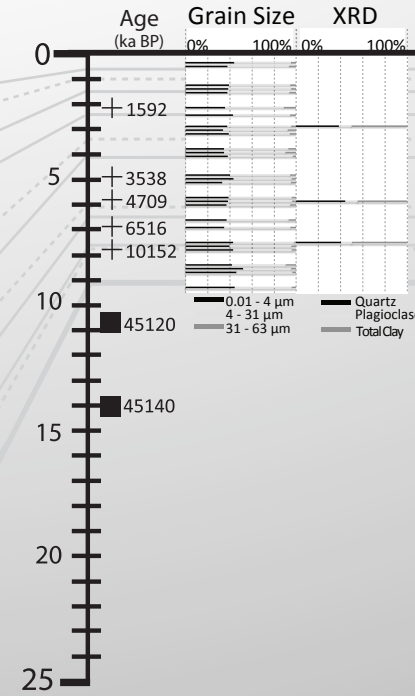
MI 652 (Colorado)



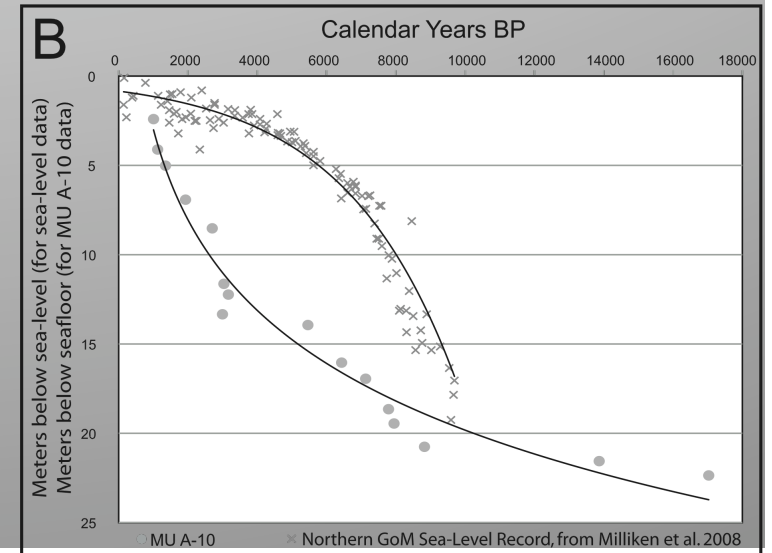
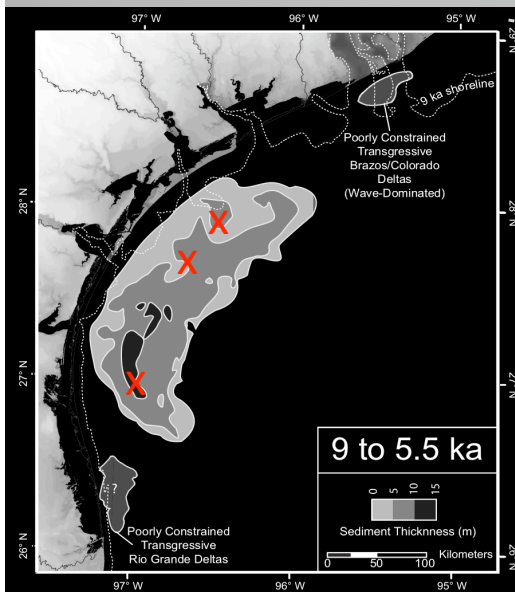
MU A-10



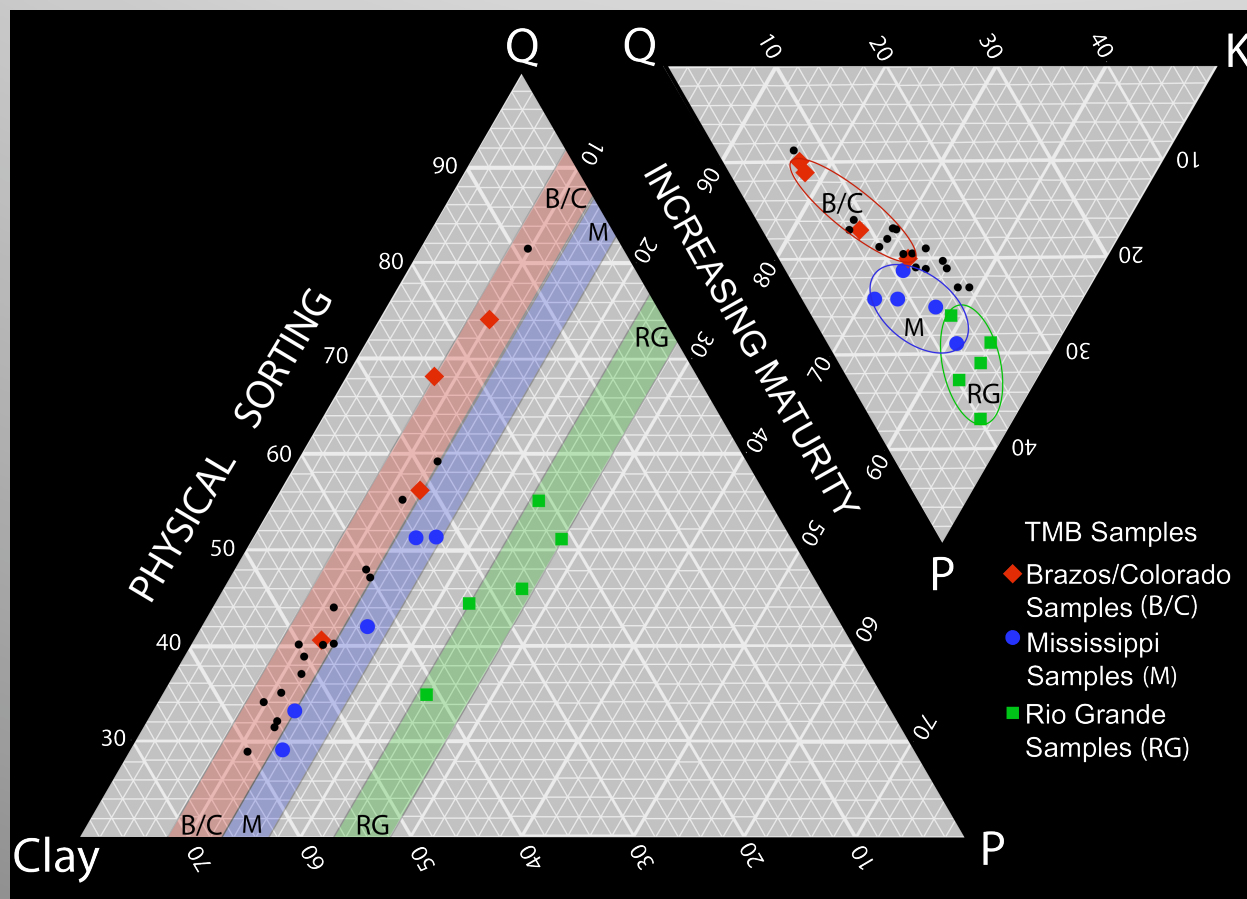
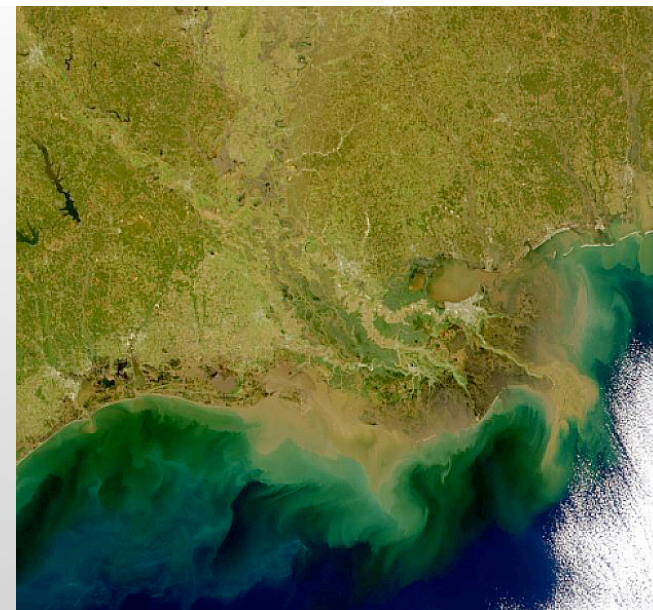
PN A-69 (Rio Grande)

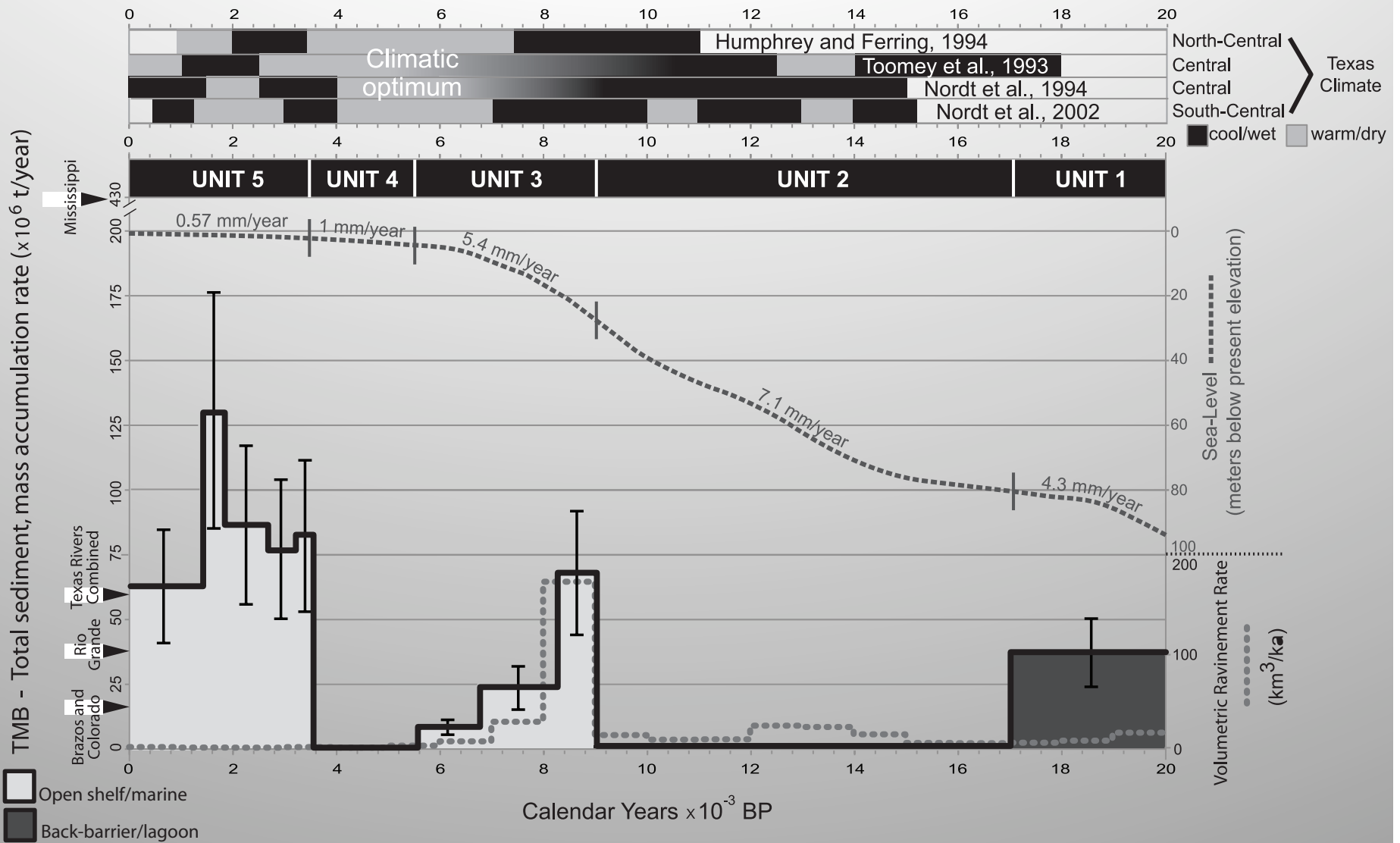


- Seismic Reflections
- Sequence Boundary
- Reflections not seen in periphery locations
- + Viable Radiocarbon Date (ka BP)
- Radiocarbon Dead Date (ka BP)



Sediment Sources





Unit 3 - 58.3 km^3 , = two-order of magnitude increase (from 0.4 to $41.1 \text{ km}^3/\text{ka}$) in flux between ~ 9 and $\sim 8 \text{ ka}$

Unit 5 - 172 km^3 = $\sim 57\%$ of the total TMB, discharge between 10.0×10^7 to 1.50×10^7 metric tons/year. This is equal to 50-75% of its modern Miss. R. discharge. If the sediment sources were solely the Brazos and Colorado rivers, their discharge would need to have been more than eight times their combined present-day discharge

Some Conclusions

1. During the previous fall of sea level, sediment discharge to the outer shelf increased significantly due to recycling of inner shelf sediments.
2. A key factor regulating sediment discharge to the shelf and slope is purging of sediment from fluvial drainage basins.
3. Rivers with greater sediment yield (e.g. Brazos, Lagniappe, and western Louisiana fluvial systems) are prone to avulsion, which means that a significant portion of the sediment they deliver to the basin is sequestered on the shelf. In contrast, smaller rivers tend to occupy the same channel throughout one or more eustatic cycles and act as point sources for slope sediment delivery.
4. The Texas Mud Blanket is the dominant depositional event of the shelf during the transgression and reflects a significant change in longshore sediment delivery and sediment discharge to the shelf in the late Holocene.
5. Transgressive ravinement has played a key role in sediment supply during Transgression.